Analyses of some of the opaque minerals (Table I) show that there is a considerable range already present in *in situ* Fe–Ni alloys at Red Mountain. The patchy distribution of iron described by Challis has not been observed in Red Mountain alloys, but compositions ranging from Ni_{86.1} Fe_{33.9} to Ni_{90.9} Fe_{9.1} have been determined. Cobalt is also present in significant amounts (Table I).

Compositional relations among analysed Ni-Fe-S minerals are shown in fig. 2. Tie lines connect analyses of mineral grains that occur within single samples but which do not necessarily represent equilibrium assemblages. However, the arrangement of tie lines is consistent with equilibrium at temperatures of less than 550 °C (Kullerud et al., 1969). The data also suggest that Fe/Ni of the alloys is mainly controlled by bulk composition (i.e. rock plus fluid) during serpentinization. The most iron-rich alloys occur with pyrrhotine (po), whereas iron-poor alloys occur with pentlandite (pn).

Although the mechanism proposed by Challis (1975) may be valid, the natural refining process is unnecessary to explain the compositional range of Fe-Ni alloys in the north-west Otago-south-Westland region, since a considerable range is represented in the source serpentinites of the Red Mountain area. This range probably reflects the Fe-Ni range of bulk compositions during hydrothermal alteration.

Acknowledgements. Professor D. S. Coombs made helpful suggestions to the manuscript; the University of Otago Grants Committee provided financial support for research in the Red Mountain area.

Geology Dept., University of Otago, Dunedin, New Zealand

JOHN M. SINTON

REFERENCES

CHALLIS (G. A.), 1975. Min. Mag. 40, 247.

KULLERUD (G.), YUNG (R. A.), and MOH (G. H.), 1969. Econ. Geol. Mon. 4, 323.

LINDGREN (M.), 1928. Mineral Deposits. 3rd edn. New York (McGraw-Hill).

SWEATMAN (T. R.) and LONG (J. V. P.), 1969. Journ. Petrol. 10, 323.

[Manuscript received 12 December 1975]

MINERALOGICAL MAGAZINE, SEPTEMBER 1976, VOL. 40, PP. 794-6

Volborthite in Liguria, Northern Italy

VOLBORTHITE Cu₃(VO₄)_{2.3}H₂O, so far described in few localities (Guillemin, 1956; Witkind, 1961; Leonardsen and Petersen, 1974), has been found in two parts of Liguria. In the Cassagna Mine (Eastern Liguria), where Mn-ores occur within the M. Alpe Chert Formation (Passerini, 1965) that lies on the top of ophiolite sequences volborthite is mainly associated with the Cu-bearing minerals chalcosine, cuprite, malachite, azurite, and native copper, in a mineralization related to fossil araucaria trunks (Cortesogno and Galli, 1974). In the upper Neva Valley (Albenga, Western

Liguria) volborthite occurs within a chalcopyrite-azurite-malachite-galena-aurichal-cite-sphaerocobaltite mineralization (Pelloux, 1926) in Triassic dolomitic limestones (Vanossi, 1971).

h k l	I	I	2	3	h k l	I	1	2	3
001	100	7·17 Å	7·17 Å	7·16 Å	222/402	25	2·047 Å	2·045 Å	2·045 Å
200	5	5.293	5.262	5.262	4 2 Ï	5	1.9277		1.9239
I I O	10	5.112	5.097	5.112	13 <u>1</u>	10	1.8521		I ·8557
2 O Ī	5	4.429		4.418	4 0 3∫	10	1 0 3 2 1		1 0551
ΙΙΊ	25	_	4.241	4.241	5 I 2				
111/201	25	4.102	4.124	4.087	004	20	1.7942	1.8041	1.7958
002	15	3.576	3.583	3.583	223	20	1.7876	1.7876	1.7876
202	30	3.092	3.090	3.090	422	5		_	1.7683
3 I O		_	-	_	223/114	20	1.7204	-	1.7144
I I Ž	45	2.993	3.008	2.998	132	20			1.7056
020		-		_	5 1 2	20	1.6725		1.6810
I I 2	40	2.885	2.885	2.885					
311/202	35	2.858		2.858	$2V\gamma$		63°		
021/311	30	2.715	_	2.711	а		1.820	-	_
400	30	2.648	2.626	2.637	β		1.835		
220)		A	0.540	2.564	γ		1.920		
40 Ī j	50	2.571	2.543	2.564	a		10·643 Å	10·671 Å	10·583 Å
2 2 I	45	2.392	2.399	2.383	b		5.874	6.037	5·866
402	10	2.238	2.222	2.217	c		7.202	7.275	7.208
2 2 2		_	_		β		95·17°	92·78°	94·90°

TABLE I. Physical data on Ligurian and synthetic volborthite

In the Cassagna Mine volborthite appears as 'rose flowers' up to 1 mm in diameter in veins and small geodes: the mineral shows lamellar shape and pseudohexagonal habit, with a {001} cleavage, micaceous brightness, and variable colour from yellow to greenish-yellow. In thin section the lamellae show an intense yellow colour, weak pleochroism ($\alpha = \beta = \text{yellow}$; $\gamma = \text{yellow-green}$), straight extinction and fast along the length related to the {001} cleavage; strong dispersion. The volborthite of Neva Valley is present as yellow-green dusty coats and veins without a defined habit, replacing Cu-bearing minerals, mainly malachite and azurite.

Physical data are listed in Table I together with those of volborthite synthesized by the method of Guillemin (1956), starting from Na-metavanadate.

Istituti di Mineralogia e Petrografia, University di Genova, Corso Europa—Genova, Italy. G. GHISO

B. MESSIGA

REFERENCES

CORTESOGNO (L.) and GALLI (M.), 1974. Ann. Mus. Civ. St. Nat. Genova, 80, 142-56. GUILLEMIN (C.), 1956. Bull. Soc. franç. Min. Crist. 79, 219-75.

^{1.} Volborthite of Cassagna Mine.

^{2.} Volborthite of Neva Valley.

^{3.} Synthetic volborthite.

 $[\]alpha$, β , γ and a, b, c all \pm 0.002.

LEONARDSEN (E. S.) and PETERSEN (O. V.), 1974. Amer. Min. 59, 372–3. PASSERINI (P.), 1965. Boll. Soc. Geol. Ital. 84, 93–176. PELLOUX (A.), 1926. Ann. Mus. Civ. St. Nat. Genova, 52, 1–13. VANOSSI (M.), 1971. Atti Ist. Geol. Univ. Pavia, 21, 89–114. WITKIND (I. J.), 1961. U.S. Geol. Surv. Bull. 1107–C, 247–92.

[Manuscript received 21 October 1975]

© Copyright the Mineralogical Society.