SHORT COMMUNICATIONS

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An occurrence of akaganéite (β -FeOOH · Cl) in Recent oxidized carbonate concretions, Norfolk, England

THE natural occurrence of β -FeOOH was first reported by M. Nambu (cited by Mackay, 1962). and named akaganéite after the locality, Akagané mine, Iwate Prefecture, Japan. Subsequent reports of natural occurrences of this mineral have been few. Logan et al. (1976) suggested on the basis of Mössbauer data that akaganéite is present in some New Zealand soils. Johnston (1977) identified akaganéite coating andesite boulders at White Island volcano, New Zealand. Johnston and Glasby (1978) reported akaganéite as a secondary oxyhydroxide mineral in marine and fossil manganese nodules. Rozenson et al. (1980) found the mineral in Jurassic evaporite deposits at Makhtesh Ramon, Israel. Akaganéite also occurs in Red Sea hot brines (Holm et al., 1983; Holm, 1984).

The author recently identified akaganéite as a significant component in oxidized siderite-calciteiron-sulphide concretions at Warham, north Norfolk, England. The concretions are actively forming in reduced intertidal marsh and sandflat sediments (Pye, 1981, 1984). Fresh, *in situ* concretions are black, but where they have been exposed by wave action or creek bank collapse they are partially oxidized and show brown, yellow and orange mottling. When sectioned many oxidized concretions were found to contain zones of relatively

Table 1. XRD data for Norfolk akaganéite compared with synthetic #-FeOOH (Mackay, 1960) and natural akaganéite from the Red Sea hot brines (Holm <u>et al</u>., 1983)

ASTM 13-157 (Mackay, 1960)		Norfolk (this paper)		Red Sea hot brine (Holm et al. 1983)	
d _{obs} Å	1/1	d _{obs} Å	I _{rel}	d Å	Irel
7.40	100	7.48	medium	7.32	medium
5.25	40	5.23	weak	5.21	v. diffuse
3.70	10	3.72	v. weak		
3.31	100	3.32	v. strong	3.31	strong
2.61	40	2.63	weak	2.59	v. weak
2.54	80	2.54	strong	2.54	medium
2.34	20	2.33	v. weak		
2.28	40	2.29	medium		

pure iron oxyhydroxide, up to several cm^3 in volume, close to the centre. This material appeared to be replacing wood and/or metal nucleii around which the concretions grew.

X-ray powder diffraction analysis indicated the oxyhydroxide is akaganéite. The diffraction patterns obtained showed a close similarity to that for synthetic β -FeOOH reported by Mackay, 1960 (Table 1).

Energy-dispersive X-ray microanalysis performed in the scanning electron microscope showed the akaganéite contains a significant amount of Cl⁻, with traces of Cr, Mn, Si and S (Fig. 1). Reliable quantitative chemical data could not be obtained by this method due to the very small size of individual akaganéite crystals (< 0.5 μ m) and the difficulty of preparing satisfactory polished sections. Cl⁻ is known to favour the formation of β -FeOOH (Weiser and Milligan, 1935), and may be essential to stabilize its structure which is of the hollandite type, comprising square-sided tunnels (each side 5 Å in dimension) parallel to the *c*axis. The tunnels can accommodate large anions, especially Cl⁻ but also OH⁻, F⁻, SO₄²⁻ and NO₃⁻

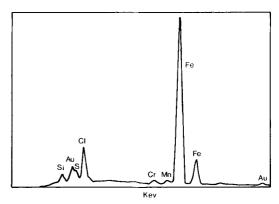


FIG. 1. Energy-dispersive X-ray spectrum obtained from a flat surface of aggregated akaganéite (specimen is gold coated).

(Gallagher, 1970; Murray, 1979). It is uncertain whether some Cl^- also substitutes for structural OH^- .

The Cl⁻-rich nature of the Norfolk akaganéite is consistent with its formation in waters of nearmarine composition. Surface waters in the lower marsh-sandflat system, which is covered by about 65% of tides, typically contain 18000–19000 ppm Cl⁻; porewaters from 10–40 cm below the surface of the lower marsh contain 20000–24000 ppm Cl⁻ (average seawater = 18986 ppm Cl⁻).

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Department of Earth Sciences, Cambridge University, Downing Street, Cambridge CB2 3EQ K. Pye

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Spencerite from Kabwe, Zambia, and the infrared spectroscopy of the Kabwe zinc phosphates

A MONG a suite of specimens from Kabwe, Zambia (previously known as Broken Hill, Northern Rhodesia) sent to Mr L. Clark and the author for investigation was a small piece consisting mainly of a portion of a very pale greenish hemimorphite nodule with 2–3 mm thick solid walls. The matrix within this hemimorphite nodule had been dissolved away and partly replaced by tiny colourless spiky smithsonite crystals, while colourless thick blades of hopeite to about 5 mm sit on the hemimorphite. Spanning spaces between hemimorphite and hopeite crystals, and growing around the hopeite are colourless pearly thin lath-shaped crystals to about 7 mm in length, representing the latest generation of mineralization present. The infrared spectrum of these laths indicates that they are the rare zinc phosphate spencerite, $Zn_4(PO_4)_2(OH)_2 \cdot 3H_2O$. Spencerite has only been reported before from two localities, Hudson's Bay mine, Salmo, British Columbia, the type locality (Walker, 1916, 1918), and Turf Pits mine, Grassington, Yorkshire (British Museum (Natural History) specimen, collected by A. W. G. Kingsbury in 1963; see Embrey, 1978). The rarity of spencerite is due to the fact that it is