

Criddleite, $\text{TlAg}_2\text{Au}_3\text{Sb}_{10}\text{S}_{10}$, a new gold-bearing mineral from Hemlo, Ontario, Canada¹

DONALD C. HARRIS, ANDREW C. ROBERTS

Geological Survey of Canada, 601 Booth Street, Ottawa, Ontario, Canada K1A 0E8

J. H. GILLES LAFLAMME

CANMET, 555 Booth Street, Ottawa, Ontario, Canada K1A 0G1

AND

CHRIS J. STANLEY

Department of Mineralogy, British Museum (Natural History), Cromwell Road, London, England SW7 5BD

Abstract

Criddleite, ideally $\text{TlAg}_2\text{Au}_3\text{Sb}_{10}\text{S}_{10}$, is a rare constituent within the Hemlo gold deposit, Hemlo, Ontario, Canada. The mineral occurs as 20 to 50 μm -sized lath-like, tabular or anhedral grains usually surrounding or penetrating aurostibite, or associated with native antimony, native gold and stibnite. Criddleite is opaque with a metallic lustre and a black streak. It has been synthesized by reacting TlSbS_2 and high purity Ag, Au, Sb and S in an evacuated silica glass tube at 400 °C. The measured density of the synthetic material is 6.86; the calculated density is 6.57 g/cm^3 . The difference is due to minor admixed aurostibite, native antimony and a dyscrasite-like phase within the charge. VHN_{25} is 94–129. Mohs hardness (calc.) = 3–3½. In reflected plane-polarized light in air, natural criddleite is weakly bireflectant with a discernible reflectance pleochroism from grey-blue to slightly greenish grey-blue. The mineral has a distinct to moderate anisotropy with rotation tints in shades of buff to slate grey. Reflectance spectra and colour values for both natural and synthetic criddleite are given. X-ray study showed that synthetic criddleite is monoclinic (pseudotetragonal) with refined unit-cell parameters $a = 20.015(2)$, $b = 8.075(2)$, $c = 7.831(2)$ Å, $\beta = 92.01(2)^\circ$, $V = 1264.9 \pm 1.0$ Å³ and $a:b:c = 2.4786:1:0.9698$. The space group choices are $A2/m(12)$, $A2(5)$ or $Am(8)$, diffraction aspect $A^*/*$. The seven strongest lines in the X-ray powder diffraction pattern [d in Å (I) (hkl)] are: 5.63(90) (011), 3.91(50) (002), 3.456(50) (320), 2.860(70) (700), 2.813(100) (022), 2.018(60) (040) and 1.959(70) (004). Electron microprobe analyses are reported of natural criddleite in five polished sections of drill core from four holes. The averaged empirical formulae, based on 26 atoms, are $\text{Tl}_{0.92}\text{Ag}_{1.99}\text{Au}_{2.93}\text{Sb}_{9.87}\text{S}_{10.28}$ (natural) and $\text{Tl}_{0.94}\text{Ag}_{2.03}\text{Au}_{2.89}\text{Sb}_{9.76}\text{S}_{10.38}$ (synthetic).

KEYWORDS: criddleite, new mineral, Hemlo gold deposit, Hemlo, Ontario, Canada, reflectance data, X-ray data, electron microprobe analyses, synthesis.

Introduction

CRIDDLEITE, ideally $\text{TlAg}_2\text{Au}_3\text{Sb}_{10}\text{S}_{10}$, is a newly described gold-bearing mineral from the Hemlo gold deposit, Hemlo, Ontario, Canada. The mineral is named for Alan J. Criddle (1944–), of the British Museum (Natural History), London, England, for his substantial contributions to the

quantitative reflectance of ore minerals and for his research on opaque minerals. The mineral and the mineral name have been approved by the Commission on New Minerals and Mineral Names, International Mineralogical Association. Type material is preserved at: the British Museum (Natural History) as polished section E.1230, BM 1987,351; within the Systematic Reference Series of the National Mineral Collection, housed at the Geological Survey of Canada, Ottawa, Ontario,

¹ Geological Survey of Canada Contribution Number 10288.

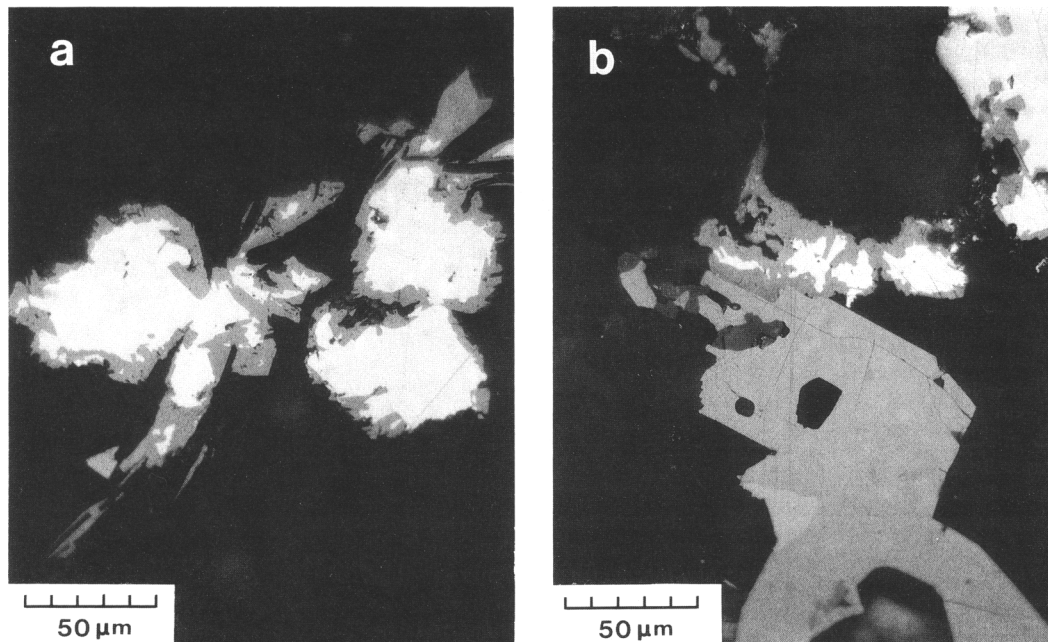


FIG. 1. (a) Photomicrograph of criddleite (grey) surrounding and penetrating austrobitite (white) within quartz gangue. Oil immersion. Polished section GG23, 651.6. (b) Photomicrograph of criddleite surrounding and partially replacing austrobitite (white) in contact with a coarse grain of chalcostibite (light grey) within quartz gangue. Oil immersion. Polished section GG23, 651.6.

Canada as a polished section, and as drill core specimens under catalogue number NMC 65186; and as a polished section at the Royal Ontario Museum, Toronto, Ontario, Canada.

Occurrence

The Hemlo gold deposit is located near the northeast shore of Lake Superior, 35 km east of Marathon, Ontario, Canada, adjacent to the Trans-Canada Highway # 17, at latitude 48° 40' N and longitude 86° 00' W. The deposit was discovered in 1982 and, within the last few years, three mines have been brought into production. They are, from east to west, the David Bell mine, the Golden Giant mine and the Page-Williams mine. In 1983, a mineralogical study of the deposit was initiated and preliminary reports on the mineralogy have been published by Harris (1986*a, b, c*). The gold deposit is substantially enriched in Mo, Sb, As, Hg, Tl, V and Ba and contains a diverse assemblage of minerals. The most common ore minerals are pyrite, molybdenite, stibnite, tetrahedrite-tennantite, native gold (containing up to 22 wt. % Hg), sphalerite (containing up to 27 wt. % Hg), zinkenite, realgar, arsenopyrite, cinnabar

and aktashite. Minor-to-trace ore minerals are austrobitite, boulangerite, bournonite, twinnite, geocronite, chalcostibite, stibarsen, native arsenic, native antimony, arsenopyrite, orpiment, berthierite, routhierite, parapiroterite, dufrénoyite, baumhauerite, gersdorffite, ullmannite, barian tomichite, magnetite, ilmenite and chromite. The gangue minerals are principally quartz, baryte, sericite, vanadian muscovite (containing up to 8.5 wt. % V_2O_5), rutile (with as much as 5.6 wt. % V_2O_5 , 6.5 wt. % Sb_2O_3 , 2.3 wt. % WO_3) and barian microcline (containing up to 16.6 wt. % BaO).

Criddleite is very rare within the deposit and has been found in polished sections of drill core specimens from only four holes: GG23, 651.6 m; GG25, 324.9 m; W70, 758.6 m and W92, 642.9 m. In polished section GG23, 651.6 m, it occurs as 40 μ m-sized anhedral rims around austrobitite (Fig. 1*a* and *b*), associated with chalcostibite and parapiroterite. In polished section GG25, 324.9 m, it occurs as an intergrowth with native gold and austrobitite in gangue. The largest composite grain is 20 \times 60 μ m. In polished section W70, 758.6 m, it occurs as an intergrowth with native antimony in gangue. The largest composite grain

Table 1. Electron microprobe analyses of criddleite

Polished Section	Tl	Ag	Au	Sb	S	Total
GG23, 651.6	7.1	8.8	23.1	48.2	13.1	100.3
GG23, 651.6 - 2a	7.8	8.7	22.8	47.9	13.3	100.5
GG25, 324.9	7.5	8.7	23.0	47.4	13.1	99.7
W92, 642.9	7.5	8.7	23.2	47.5	13.1	100.0
W70, 758.6	7.4	7.8	22.3	47.3	12.8	97.6
Atomic proportions total = 26						
GG23, 651.6	0.87	2.04	2.94	9.92	10.23	
GG23, 651.6 - 2a	0.95	2.01	2.89	9.81	10.34	
GG25, 324.9	0.92	2.03	2.94	9.81	10.29	
W92, 642.9	0.92	2.03	2.96	9.81	10.28	
W70, 758.6	0.93	1.86	2.91	10.01	10.28	
Average	0.92	1.99	2.93	9.87	10.28	

is $30 \times 70 \mu\text{m}$. Other ore minerals observed in this polished section are native gold, stibnite, molybdenite, aurostibite, zinkenite and pyrite. In polished section W92, 642.9 m, it occurs as $20 \mu\text{m}$ -sized grains in gangue associated with, but not in contact with, stibnite, tetrahedrite, native antimony and aurostibite. In all sections, criddleite commonly forms lath-like or anhedral grains and is a primary mineral formed from the same hydrothermal fluids that have resulted in the diverse assemblage of ore minerals.

Electron microprobe analyses

The quantitative analyses were performed with a CAMEBAX electron microprobe operated at 20 kV and with a beam current of 30 nA. The analytical results are given in Table 1. The X-ray lines and standards used were Au- $M\alpha$, Ag- $L\alpha$ (synthetic $\text{Au}_{0.678}\text{Ag}_{0.322}$); Tl- $M\alpha$, Sb- $L\alpha$, S- $K\alpha$ (synthetic TlSbS_2). The averaged empirical formula of the five analyses, based on 26 atoms, is $\text{Tl}_{0.92}\text{Ag}_{1.99}\text{Au}_{2.93}\text{Sb}_{9.87}\text{S}_{10.28}$ or, ideally, $\text{TlAg}_2\text{Au}_3\text{Sb}_{10}\text{S}_{10}$.

Synthesis

Owing to the small size of the natural criddleite grains, syntheses were undertaken: (a) to verify the composition of the natural material as determined by the electron microprobe analyses, (b) to obtain synthetic material suitable for single-crystal X-ray studies, and (c) to use the synthetic material as a potential electron microprobe standard. A 200 mg charge, equivalent to the formula $\text{TlAg}_2\text{Au}_3\text{Sb}_{10}\text{S}_{10}$, was synthesized by reacting TlSbS_2 and high-purity Au, Ag, Sb and S in an evacuated silica glass tube. The starting materials

were heated at 350°C for three days and then heated at 400°C for an additional three days. The charge was then examined megascopically and, although it appeared to have melted, it did not appear homogeneous; a small amount of unreacted material (Au?) was present on the surface of the charge. It was, therefore, decided to increase the temperature to 450°C and examination of the charge after one day at this higher temperature indicated a more complete reaction. Thereafter, the charge was annealed for a period of two months at 400°C and then subjected to grinding, pelletizing and reannealing at 400°C for an additional period of two months. This synthesis was successful because microscopic and electron microprobe examination of a polished section containing a portion of the synthetic charge showed approximately 96% criddleite, 2% aurostibite, 1% native antimony and 1% of a dyscrasite-like phase which has the composition Ag 35.2, Au 44.0, Sb 22.8 wt. %.

Electron microprobe analyses of synthetic criddleite, using the same standards and operating conditions that were applied to the analyses of the natural grains, gave Tl 7.7, Ag 8.8, Au 22.9, Sb 47.8, S 13.4, total 100.6 wt. %, which corresponds to $\text{Tl}_{0.94}\text{Ag}_{2.03}\text{Au}_{2.89}\text{Sb}_{9.76}\text{S}_{10.38}$. Subtracting the extra phases from the starting composition of Tl 8.0, Ag 8.5, Au 23.2, Sb 47.7, S 12.6, total 100.0 wt. %, gives a calculated composition for the synthetic criddleite of Tl 8.3, Ag 8.5, Au 22.7, Sb 47.5, S 13.0, total 100.0 wt. %, in reasonable agreement with the electron microprobe results.

Optical properties

The polished sections containing the natural criddleite and the synthetic equivalent were buffed with a paste of MgO in distilled water on Lamplan 'velvet' immediately before making observations and measurements.

Criddleite takes an excellent polish, although its polishing hardness is low. In reflected plane-polarized light (colour temperature $\sim 3200\text{K}$) in air it is weakly birefractant with a discernible reflectance pleochroism from grey-blue to slightly greenish grey-blue. In oil, this subtle reflectance pleochroism is enhanced. Criddleite has a distinct to moderate anisotropy with rotation tints in shades of buff to slate grey. Simple twinning is common. The synthetic criddleite has identical optical characteristics but, in addition to simple twinning, polysynthetic twinning is also common.

Reflectance measurements were made relative to a WTiC standard (Zeiss no. 314) at intervals of 10 nm from 400 to 700 nm in air and in oil (Zeiss N_D 1.515, DIN 58.884) at 21°C . A plane glass

TABLE 2. Reflectance data and colour values in air and oil for natural and synthetic criddleite.

λ_{nm}	GG 23 651.6		synthetic		GG 23 651.6		synthetic	
	R_1	R_2	R_1	R_2	im_{R_1}	im_{R_2}	im_{R_1}	im_{R_2}
400	40.6	41.85	41.4	42.5	26.3	28.3	27.0	29.8
410	40.4	41.4	41.05	42.1	25.95	27.5	26.6	28.75
420	40.1	41.0	40.7	41.75	25.4	26.7	26.15	27.9
430	39.7	40.5	40.15	41.3	25.1	26.2	25.7	27.2
440	39.4	40.1	39.7	40.9	24.7	25.9	25.2	26.5
450	39.0	39.8	39.3	40.55	24.4	25.6	24.7	26.0
460	38.6	39.4	39.0	40.25	24.1	25.4	24.3	25.7
470	38.4	39.3	38.6	39.9	23.8	25.2	23.9	25.5
480	38.1	39.2	38.2	39.75	23.6	25.1	23.6	25.5
490	37.9	39.2	37.9	39.65	23.3	25.0	23.3	25.45
500	37.65	39.1	37.6	39.6	23.0	24.8	23.1	25.4
510	37.5	39.0	37.5	39.6	22.8	24.7	22.9	25.3
520	37.3	38.9	37.2	39.45	22.6	24.6	22.6	25.1
530	37.0	38.7	37.0	39.3	22.3	24.3	22.35	24.9
540	36.7	38.4	36.8	39.1	22.1	24.2	22.2	24.6
550	36.4	38.2	36.4	38.9	21.8	23.9	21.55	24.4
560	36.1	38.0	36.15	38.8	21.6	23.7	21.7	24.2
570	35.9	37.9	35.8	38.55	21.3	23.5	21.4	24.0
580	35.55	37.65	35.6	38.4	21.1	23.3	21.1	23.8
590	35.3	37.4	35.2	38.1	20.85	23.1	20.8	23.6
600	35.0	37.2	34.8	37.8	20.6	22.9	20.5	23.3
610	34.7	36.95	34.45	37.5	20.35	22.6	20.2	23.0
620	34.35	36.6	34.1	37.2	20.0	22.3	19.85	22.6
630	34.0	36.3	33.7	36.8	19.8	22.0	19.5	22.2
640	33.7	36.0	33.3	36.3	19.6	21.7	19.2	21.8
650	33.4	35.6	33.0	36.0	19.3	21.4	18.9	21.4
660	33.1	35.3	32.7	35.6	19.1	21.1	18.7	21.1
670	32.8	35.0	32.45	35.1	18.9	20.75	18.5	20.8
680	32.65	34.6	32.3	34.8	18.9	20.5	18.5	20.4
690	32.6	34.4	32.3	34.5	18.8	20.3	18.5	20.25
700	32.5	34.1	32.35	34.3	18.9	20.1	18.6	20.05

COLOUR VALUES

Illuminant C

x	.300	.303	.299	.303	.295	.299	.293	.298
y	.308	.312	.307	.312	.303	.308	.302	.308
Y%	36.2	38.0	36.1	38.7	21.6	23.7	21.7	24.2
λ_d	481	483	481	482	480	482	480	482
$F_{\theta}^{\%}$	4.5	3.0	5.0	3.0	7.0	4.9	7.7	5.1

Illuminant A

x	.437	.440	.436	.440	.432	.436	.430	.435
y	.407	.408	.406	.408	.405	.407	.405	.407
Y%	35.8	37.7	35.7	38.3	21.3	23.4	21.3	23.9
λ_d	493	495	493	495	492	494	492	495
$F_{\theta}^{\%}$	2.5	1.8	2.8	1.8	3.8	2.9	4.2	3.0

reflector was used with $\times 40$ objectives, the effective numerical apertures of which were adjusted to 0.35. As the extinction positions could not be established with certainty, the maximum and minimum positions for the measured grains were determined photometrically at 560 nm. Otherwise, the equipment used and the measurement procedures were as described by Criddle *et al.* (1983).

Owing to the small grain size and paucity of natural material, the measurements (Table 2 and Fig. 2) were obtained on a single lath-shaped ($25 \times 15 \mu\text{m}$) grain included in aurostibite in polished section GG23, 651.6. Other grains of criddleite were examined but gave less reliable results because of twinning or because of difficulties encountered in levelling grains adjacent to the

much harder quartz gangue. Check measurements on these grains, however, showed no significant differences in either the magnitude or dispersion of the spectra.

The most strongly anisotropic and bireflectant grain of the synthetic preparation was selected for measurement. Table 2 shows that the differences in both reflectance and bireflectance between this and the most bireflectant grain in section GG23, 651.6 are minimal. All spectra in Fig. 2 were smoothed graphically between 400 and 450 nm.

The mineral is monoclinic and, because the grains could not be oriented, the reflectance spectra are labelled R_1 and R_2 (Criddle, 1980). One sectioned grain of the synthetic mineral is isotropic, and its measured reflectance R_0 is very close to R_1 .

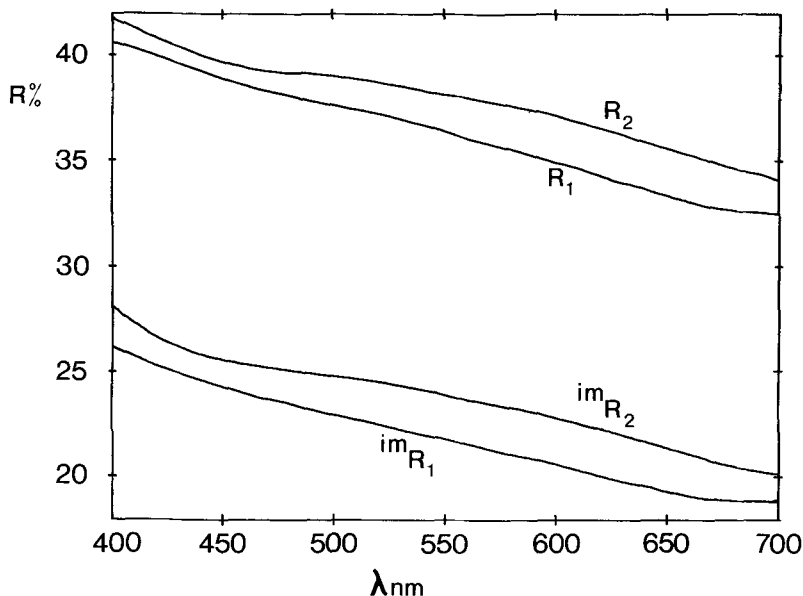


FIG. 2. Reflectance spectra for criddleite in air and in oil. (Specimen GG23, 651.6.)

It follows that the mineral is effectively uniaxial, as well as being pseudotetragonal.

Dominant wavelength values (λ_d in Table 2) for the CIE A illuminant (close to the colour temperature of the light source used for the qualitative description) confirm the overall grey-blue colour of the mineral. The perceived slightly more greenish grey-blue hue for R_2 is a function of its lower excitation purity and slightly higher luminance ($Y\%$) compared with R_1 .

Microhardness and physical properties

Synthetic criddleite has a microhardness VHN_{25} in the range 94–129 based on 10 indentations. This corresponds to a calculated Mohs hardness of 3–3½. In shape, these indentations were perfect to slightly deformed with minor fracturing at the corners. The grain size of the natural material was too small for VHN measurements. The mineral is opaque with a metallic lustre and a black streak.

X-ray powder and single-crystal study

The small grain size and twinning characteristics of natural criddleite prevented extraction of a suitable fragment from a polished section for precession single crystal study. However, a 57.3 mm Debye-Scherrer powder film of natural criddleite showed a sufficient number of observable diffraction lines for comparison to be made with

a powder pattern of synthetic material obtained with the same diameter camera; the two patterns were for one and the same phase. The measurable reflections for natural criddleite (in Å), 5.66 (weak), 2.85 (strong), 2.81 (strong), 2.02 (medium) and 1.96 (medium) are consistent with the strongest reflections presented in Table 3 for synthetic material. Other weak reflections were identified but, because of the spotty nature of the film, were unmeasurable. Therefore, two fragments of synthetic material were mounted and studied by precession single-crystal methods employing Zr-filtered Mo radiation. One fragment was orientated such that a^* was coincident with the dial axis and the other fragment mounted with 011^* coincident with the dial axis. The levels collected were: $hk0 \rightarrow hk2$, $h0l \rightarrow h2l$, $0kl \rightarrow 4kl$ and $011^* \wedge a^*$. No evidence of twinning was observed on the precession single-crystal films. Criddleite is pseudotetragonal, monoclinic with measured (and calculated) cell parameters derived from zero-level precession films of: $a = 19.96$, $b = 8.057$, $c = 7.809$ Å, $\beta = 92.08^\circ$. The space-group absence conditions; hkl with $k+l \neq 2n$, $hk0$ with $k \neq 2n$, $h0l$ with $l \neq 2n$, $0kl$ with $k+l \neq 2n$ and $0k0$ with $k \neq 2n$, indicate that $A2/m$ (12), $A2$ (5) or Am (8) are the permissible space-group choices. The diffraction aspect is $A^*/*$. A fully indexed Guinier-DeWolff powder pattern of synthetic material is presented in Table 3. Admixed end-member aurostibite was used as an internal standard.

Table 3. X-ray powder data for synthetic criddleite

lest.	dÅmeas.	dÅcalc.	hkl	lest.	dÅmeas.	dÅcalc.	hkl
5	10.01	10.00	200	60	2.018	2.019	040
30	6.67	6.67	300	3	2.001	2.000	1000
90	5.63	5.62	011	35	1.980	1.979	722
10	5.45	5.45	111	70	1.959	1.957	004
40	5.00	5.00	400	1	1.932	1.932	340
5	4.94	4.95	211			1.911	731
20	4.85	4.85	211	5b	1.909	1.906	713
30	4.34	4.35	311	10	1.891	1.890	731
40	4.25	4.25	311	10	1.873	1.872	440
40	4.03	4.04	020			1.844	404
15	4.00	4.00	500	10	1.847	1.844	713
3	3.95	3.96	120	1	1.815	1.815	333
50	3.91	3.91	002	1	1.802	1.802	540
1	3.75	3.74	220			1.792	1020
10	3.69	3.69	202	15	1.793	1.792	333
50	3.456	3.454	320			1.765	922
10	3.431	3.428	302	10	1.761	1.761	024
			600	10b	1.741	1.739	342
10	3.331	3.334	302	10	1.726	1.725	342
1	3.298	3.299	311	10	1.715	1.716	324
3	3.225	3.221	511	3	1.697	1.697	442
5	3.142	3.141	420	3	1.689	1.689	324
5	2.896	2.899	611			1.680	533
70	2.860	2.858	700	3	1.679	1.680	442
100	2.813	2.810	022			1.677	424
3	2.774	2.773	122	1	1.662	1.662	604
1	2.689	2.688	222	30	1.648	1.649	740
3	2.612	2.613	322	5	1.640	1.641	704
			711	3b	1.610	1.611	1022
30	2.569	2.573	322	5	1.588	1.589	704
40	2.545	2.545	031			1.585	733
15	2.522	2.522	711	5	1.580	1.582	051
10	2.500	2.500	800	3	1.567	1.568	804
30	2.480	2.482	013	3	1.548	1.549	733
			422	10	1.536	1.537	015
20	2.473	2.474	113	5	1.486	1.486	315
			231			1.429	1400
3	2.426	2.424	422	15	1.425	1.421	1040
3	2.392	2.391	213			1.414	615
5	2.385	2.387	331	3	1.415	1.405	044
10	2.366	2.369	331			1.404	144
3	2.351	2.352	313	40	1.399	1.388	751
			720				
1	2.329	2.327	322				
5	2.305	2.302	313				
3	2.278	2.273	522				
5	2.225	2.223	900				
			820				
10	2.125	2.123	622				
30	2.030	2.029	722				

- Guinier-DeWolff camera employing Fe filtered Co radiation (λ Co $K\alpha_1 = 1.78892 \text{ \AA}$)
- b = broad line
- intensities estimated visually
- pattern run at CANMET by Mr P. Carrière
- pattern corrected for shrinkage
- indexed on $a = 20.015$, $b = 8.075$, $c = 7.831 \text{ \AA}$, $\beta = 92.01^\circ$

Powder lines solely ascribable to admixed aurostibite, the dyscrasite-like phase and native antimony have been deleted. The refined unit-cell parameters for synthetic criddleite, based on 36 reflections between 4.04 and 1.548 Å for which unambiguous indexing, aided by visual inspection of precession single-crystal films, was possible, are: $a = 20.015(6)$, $b = 8.075(2)$, $c = 7.831(2) \text{ \AA}$, $\beta = 92.01(2)^\circ$, $V = 1264.9 \pm 1.0 \text{ \AA}^3$ and $a:b:c = 2.4786:1:0.9698$. With $Z = 2$, the calculated density for the synthetic empirical formula is 6.57 g/cm³, about 4.4% lower than the measured Berdan balance density of 6.86(7) g/cm³. The difference is ascribable to minor admixed aurostibite, a dyscrasite-like phase and native antimony in the synthetic charge.

Acknowledgements

The authors thank Mr P. Carrière (CANMET) for running the Guinier-DeWolff camera film of synthetic criddleite and Drs A. C. Bishop, P. Henderson, and R. F. Symes, all of the British Museum (Natural History), and Mr P. G. Embrey, for their comments on the manuscript.

References

- Criddle, A. J. (1980) Editorial policy for the second issue of the IMA/COM quantitative data file. *Can. Mineral.* **18**, 553-8.
- Stanley, C. J., Chisholm, J. E., and Fejer, E. E. (1983) Henryite, a new copper-silver telluride from Bisbee, Arizona. *Bull. Mineral.* **106**, 511-17.

- Harris, D. C. (1986a) The minerals in the main Hemlo gold deposit, Ontario. In *Current Research, Part A, Geol. Survey of Canada, Paper 86-1A*, 49-54.
- (1986b) Mineralogy of the main Hemlo gold deposit. In *The Hemlo gold deposits, Ontario*. Geological Association of Canada, Mineralogical Association of Canada, Canadian Geophysical Union, Joint Annual Meeting, Ottawa '86, *Field Trip 4: Guidebook* (Harris, D. C., ed.), 74 pp.; also in *Gold '86 Excursion Guidebook: Toronto* (Pirie, J. and Downes, M. J., eds.), 158-65.
- (1986c) Mineralogy and geochemistry of the main Hemlo gold deposit, Hemlo, Ontario, Canada. In *Proceedings of Gold '86* (Macdonald, A. J., ed.). International Symposium on the Geology of Gold, Toronto, 1986, 297-310.

[Manuscript received 25 January 1988;
revised 7 April 1988]