

An occurrence of a fully-oxidized natural titanomaghemite in basalt

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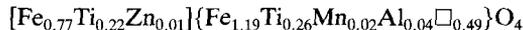
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Abstract

Titanomaghemite occurs in a relatively fresh doleritic intrusion in an area of Precambrian gneiss in Minas Gerais, Brazil. It hosts exsolution lamellae of ilmenite and contains more than 90% of the iron in the ferric form. It is more resistant to weathering than the ilmenite and is inherited virtually unaltered by the resulting soils. Titanomaghemite, extracted as grains from a weathered rind of the rock, has lattice parameter $a_0 = 0.8348(3)$ nm and has a canted spin structure due to substitution of non-magnetic ions on tetrahedral and octahedral sites of the spinel structure. The average canting angle is $32 \pm 3^\circ$ and canting occurs predominantly on the octahedral iron sublattice. Its formula, based on microprobe analysis and Mössbauer spectroscopy may be expressed as:

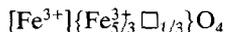


where [] and { } denote ions on tetrahedral and octahedral sites, respectively. The spontaneous magnetization of the mineral is $36(3)$ J/T/kg.

KEYWORDS: titanomaghemite, ferric iron, gneiss, Minas Gerais, Brazil.

Introduction

MAGHEMITES are cation-deficient iron oxides with a structure related to that of spinel, which contain little or no ferrous iron. The ideal formula, $\gamma\text{-Fe}_2\text{O}_3$ may be written



where [] and { } denote, respectively, tetrahedral (A) and octahedral (B) sites of the spinel structure, and \square represents a vacancy. There is a solid-solution series between magnetite, Fe_3O_4 , and maghemite, which converts to hematite, $\alpha\text{-Fe}_2\text{O}_3$, on heating to about 300°C . Both cation (e.g. Al, Ti) and anion (OH) substitutions are common in the structure, and a range of lattice parameters ($0.830 \leq a_0 \leq 0.835$ nm) is found in the literature (Lindsley, 1976). The term titanomaghemite refers to cation-deficient spinels in the Fe^{3+} -rich field defined by $\text{Fe}_3\text{O}_4\text{-Fe}_2\text{TiO}_4\text{-FeTiO}_3\text{-}\gamma\text{-Fe}_2\text{O}_3$.

Titanomaghemite occurs extensively in sea-floor basalts, where it usually forms by gradual undersea alteration of titanomagnetite produced at the mid-ocean ridge, with a time-constant of about 10^6 years (Johnson and Atwater, 1977). Partial oxidation of titanomagnetite may also take place at high temperatures and be retained as the basalt is quenched (Hauptmann, 1974). Occurrences are also reported in seamounts (Kono *et al.*, 1980) and subaerial volcanic rocks (Akimoto and Kushiro, 1960).

Maghemite in soils can be formed from goethite by burning in the presence of biomass. Such goethite is frequently aluminous, and aluminium will then be incorporated into the structure of maghemite (Schwertmann and Fechter, 1984; Fitzpatrick, 1988). There is also some evidence of the possible formation of soil maghemite by a solution-precipitation process (Kopp and Lee, 1987). Highly magnetic soils, often reported to

contain maghemite, are widespread in Brazil (Resende *et al.*, 1988, Curi and Franzmeier, 1987). They frequently form on basalt, and magnetization is used there as an aid in soil classification.

Here we report the presence of fully-oxidized titanomaghemite in a relatively fresh, intrusive rock of basaltic composition from an area of Precambrian gneiss in Minas Gerais, Brazil. The rock is the parent for a soil with a high concentration of magnetic material, as described by Resende *et al.* (1986). We have determined the composition, intrinsic magnetization and spin structure of this titanomaghemite, which will facilitate its quantitative determination by magnetic measurements.

Experimental methods

Electron microprobe measurements were made with a wavelength dispersive ARL-SEM. The excitation potential was 20 kV, the beam current 20 nA and the take-off angle was 52.5°. Natural standards were used and the data were corrected using the modified program Magic IV (Colby, 1968).

The samples were analysed by X-ray diffraction (XRD) with Co-K α radiation using a Philips powder diffractometer. An internal silicon standard was used and the step size was 0.02° 2 θ per step. The data were computer fitted using standard XRD fitting programs.

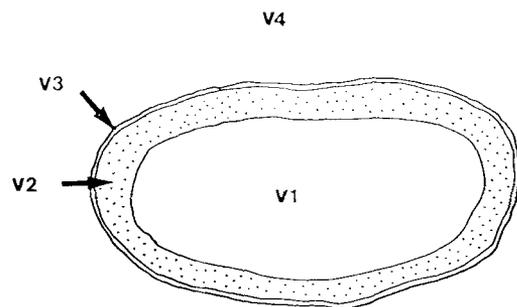


FIG. 1. Schematic diagram of the doleritic cobble (≤ 10 cm in length) showing: V1, rock of basaltic composition containing magnetic grains V1m; V2, yellow rind of weathered material containing magnetic grains V2m; V3, red outer rind; V4, surrounding red magnetic soil.

Room-temperature magnetization measurements were made in fields up to 1.5 Tesla using

a vibrating sample magnetometer. ^{57}Fe Mössbauer spectra at room temperature and 4.2 K were collected in 512 channels using a conventional constant acceleration spectrometer and a source of ^{57}Co in Rh.

Results

Petrographic examination. The intrusion is located in the hilly to mountainous region of the Zona da Mata, Minas Gerais, Brazil. The area is formed of Precambrian gneiss with minor mafic intrusions and has a deep weathering mantle (≥ 10 m). Cobble-sized (≤ 10 cm maximum dimension) rounded pieces of dolerite (see Fig. 1), which are distributed throughout the weathering mantle, were examined. The cobbles are surrounded by rinds of yellow weathered material (containing quartz, gibbsite, and goethite) several millimetres thick.

In thin section, the dolerite is mainly fresh with occasional weathered patches associated with fractures ($\approx 5\%$ of total area). The fresh rock contains approximately 65% plagioclase feldspar, 30% pyroxene, and 5% opaque minerals, and is non-porphyrific. The plagioclase forms laths 1–1.5 mm long and falls within the labradorite range, with 50–70% of the calcic end-member (anorthite). The pyroxene crystals (grain size ≤ 0.5 mm) are of pale augite and show some strain deformation into sub-grains; they form polycrystalline clusters (≤ 1.5 mm) and are generally interstitial to the plagioclase. The opaques (grain size ≤ 1 mm) are irregularly shaped and partly interstitial, providing evidence that they may have formed later than the other rock components. The opaque grains are of ilmenite and of magnetite/maghemite (a spinel-structure phase) with ilmenite exsolution parallel to $\{111\}$. A characteristic lamellar structure (Putnis and McConnell, 1980) is clearly seen in the dolerite and in grains extracted from the weathered rind (Fig. 2), with approximately 20% of the area of the opaque grains being ilmenite and the rest magnetite/maghemite. We show later that the latter phase is actually maghemite.

In the weathered patches, near fractures, the rock has been heavily altered to an orange and brown birefringent material which may include goethite and oxidized chlorite.

Electron microprobe. Opaque mineral grains from both the basalt (V1m) itself and from the surrounding yellow rind (V2m) were analysed. Five to eight points were examined in the ilmenite and maghemite regions and the results averaged (see Table 1).

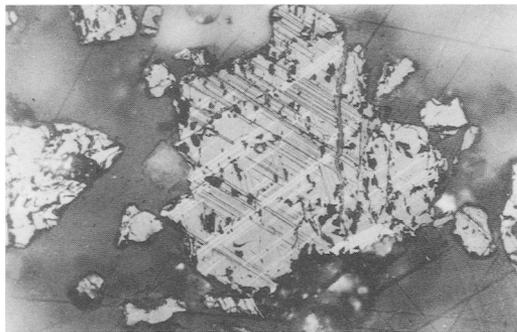


Fig. 2. Photograph of a grain of V2m, under slightly uncrossed polarizers, showing exsolution lamellae of ilmenite (scale: horizontal length of photograph = 840 μm).

Table 1. Electron microprobe analyses for samples V1m and V2m (wt%)

Oxide	V1m		V2m	
	Maghemite	Ilmenite	Maghemite	Ilmenite
FeO (± 0.06)	69.83	46.84	72.29	48.12
TiO ₂ (± 0.02)	21.74	51.51	20.32	50.36
Al ₂ O ₃ (± 0.06)	1.74	0.08	1.86	<0.05*
MgO (± 0.08)	<0.03*	0.59	<0.02*	0.83
MnO (± 0.05)	0.64	0.64	0.64	0.62
ZnO (± 0.04)	0.24	0.24	0.34	<0.02*
Total*	94.22	99.68	95.47	100.00

* Indicates that the value is below the detectable limit

* Maghemite oxide wt% totals fall below 100 because iron is presented as FeO rather than as Fe₂O₃

The degree of titanium substitution is high for the maghemite regions of both samples and a little aluminium, manganese and zinc is also present. No chromium substitution was detected.

Lattice parameters. The lattice parameter for the spinel-structure phase, from X-ray diffraction, is 0.8359(1) nm for the sample from the basalt (V1m) and 0.8348(3) for the grains extracted from the weathered rind (V2m). These values are both smaller than expected for a titanomagnetite with the observed Fe/Ti ratio of ≈ 3.8 (≈ 0.849 nm; O'Reilly and Readman, 1971) but they are some-

what larger than the accepted average value for end-member maghemite of 0.834 nm. The decrease in lattice parameter may reflect further oxidation of any residual Fe²⁺ in the spinel as the rock weathers from the rind.

Magnetization. The saturation magnetization (σ_s) is 18 J/T/kg of sample for the opaque grains extracted from the basalt after crushing (V1m) and 25 J/T/kg for those in the rind (V2m). These figures fall within the expected range for soil maghemites (≤ 60 J/T/kg; Coey, 1988).

Mössbauer spectroscopy. Room temperature Mössbauer spectra of V1m and V2m are shown in Fig. 3. They consist of three principal components; a magnetic sextet M with isomer shift (relative to α -Fe) $\delta = 0.32$ mm/s, quadrupole shift $\Delta = -0.01$ mm/s and hyperfine field $B_{\text{hf}} = 49.7$ T and two quadrupole doublets, D1 with $\delta = 0.28$ mm/s and $\Delta = 0.50$ mm/s and D2 with $\delta = 1.06$ mm/s and $\Delta = 0.65$ mm/s. Fit parameters and relative intensities of the three components are shown in Table 2. The ferrous doublet D2 is readily identified as ilmenite, for which $\delta = 1.07$ mm/s and $\Delta = 0.68$ mm/s (Murad and Johnston, 1987). D1, which represents 22% of the relative absorption area in V1m, but only 10% in V2m is identified as ferric iron in associated minerals. In addition there is a little ferrous absorption in V1m ($\delta = 1.12(1)$ mm/s, $\Delta = 2.05(1)$ mm/s) which is attributed to ferrous iron in the associated minerals. The magnetic sextet has a hyperfine splitting close to that of pure maghemite (49.9 T; Murad and Johnston, 1987). There is no sign of a separate ferrous pattern, or of the characteristic mixed-valence Fe²⁺/Fe³⁺ pattern found in stoichiometric or non-stoichiometric magnetite (Coey *et al.*, 1971). The magnetically-ordered iron in the spinel-structure phase is apparently all Fe³⁺; from the quality of the fits, it can be affirmed that *at most* 10% of the iron in this phase could be Fe²⁺.

From the microprobe analyses in Table 1, the composition of the ilmenite is estimated to be (Fe_{0.99}Ti_{0.98}Mg_{0.02}Mn_{0.01})O₃ in V1m and (Fe_{1.01}Ti_{0.95}Mg_{0.03}Mn_{0.01})O₃ in V2m, i.e. there is practically no difference from the ideal formula FeTiO₃. The relative areas of the M and D2 components of the Mössbauer spectra are 61:17 for V1m and 78:10 for V2m, which correspond to a maghemite: ilmenite ratio of approximately 4:1 for V1m and 8:1 for V2m (assuming identical recoil-free fractions for maghemite and ilmenite). This ratio is in agreement with the optical estimates. The weathering process in the rind appears to remove ilmenite in preference to maghemite.

Taking all the iron in the maghemite to be Fe³⁺, this leads to the following compositions:

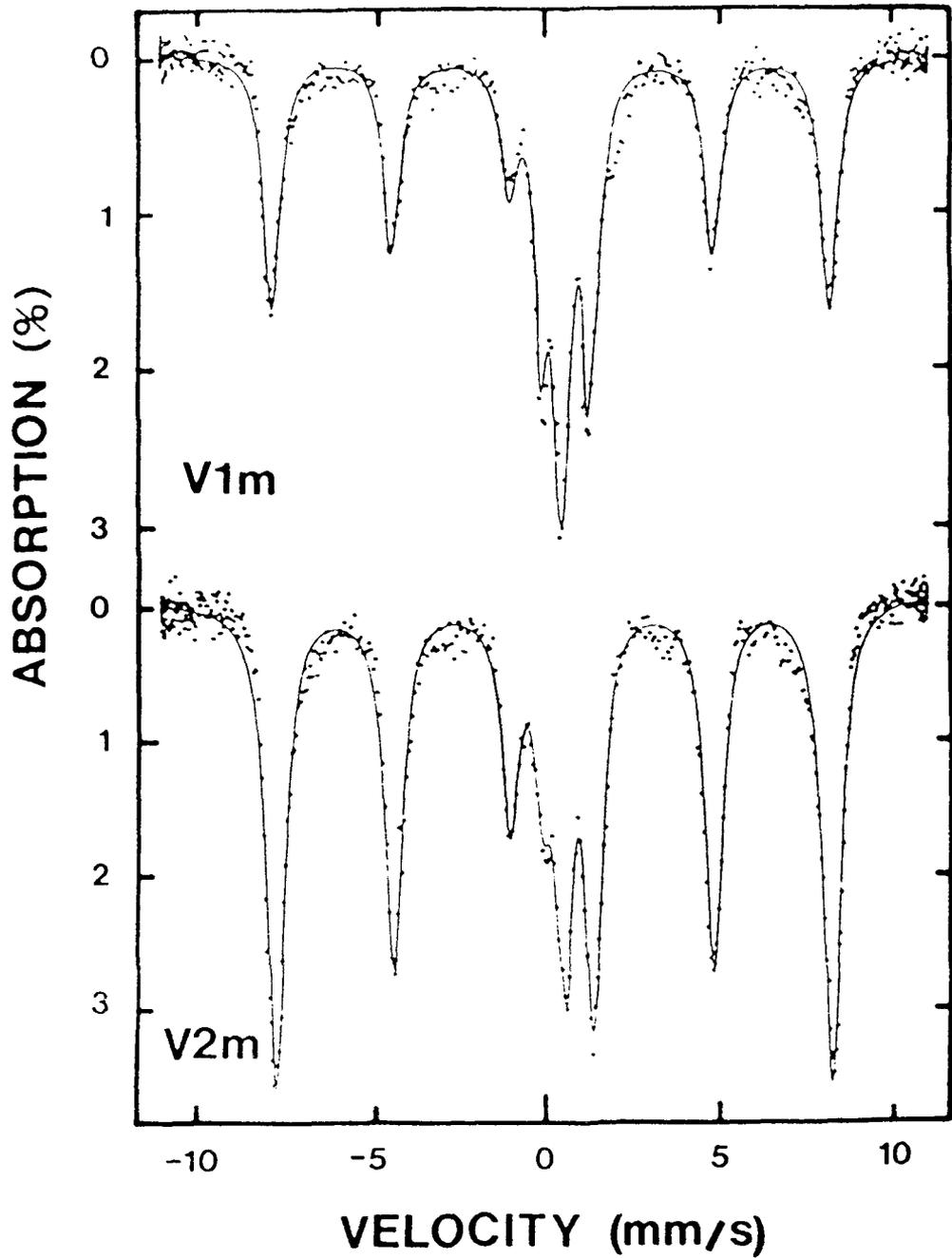


FIG. 3. Room-temperature Mössbauer spectra of the magnetic grains V1m (from the rock) and V2m (from the weathered rind).

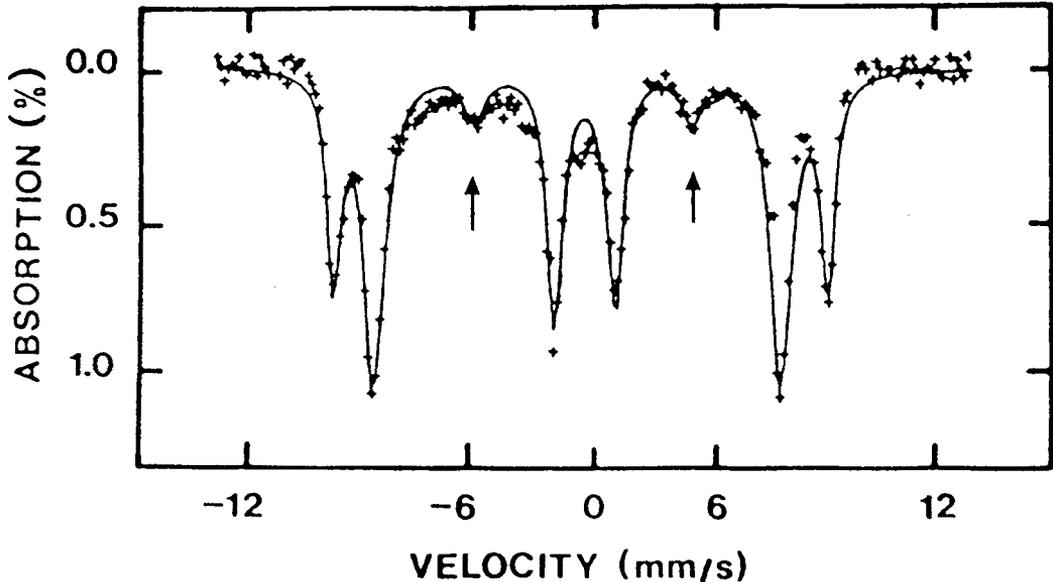
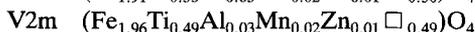
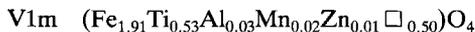


FIG. 4. Mössbauer spectrum of sample V2m at 4.2 K in an applied field of 5 T (the arrows indicate the positions of lines 2 and 5).

Table 2 Mössbauer parameters for V1m and V2m at room temperature

	δ^* (mm/s)	Δ (mm/s)	B_{hf} (T)	Relative area	
V1m	M	0.33(1)	-0.02(1)	49.7(1)	0.61
	D1	0.30(1)	0.51(1)	-	0.22
	D2	1.07(1)	0.63(1)	-	0.17
V2m	M	0.32(1)	-0.01(1)	49.7(1)	0.78
	D1	0.26(1)	0.48(2)	-	0.10
	D2	1.05(1)	0.68(1)	-	0.12

* δ quoted relative to α -Fe



where \square represents a vacancy.

A Mössbauer spectrum of V2m was also obtained at 4.2 K in an applied field of 5 T (Fig. 4). The main features are

- (i) a splitting of the ferric sextet, separating the A site and B site iron sublattices and
- (ii) reduction in the intensities of lines 2 and 5 of this pattern.

The ilmenite is also magnetically ordered at this

temperature, with $B_{hf} = 4.8$ T (Murad and Johnson, 1987), but the applied field broadens the lines from the randomly-oriented and magnetically anisotropic antiferromagnetic ilmenite grains, so that its absorption contributes to the background in the range -1 to $+1$ mm/s.

The presence of lines 2 and 5 of the maghemite spectrum (indicated by arrows in Fig. 4) shows that spin canting occurs. From computer fitting, the canting is found to be associated with the spectrum from the octahedral sublattice and an average canting angle of $32 \pm 3^\circ$ is deduced. The ratio of occupancies of tetrahedral to octahedral sites is derived from the intensities of the magnetically split subspectra (assuming equal recoilless fractions) to be 0.65(5).

Using this information, the formula for titanomaghemite in V2m may finally be rewritten as



Here we assume an A-site preference for Zn and a B-site preference for Al and Mn.

The calculated saturation magnetization from this formula is 33 J/T/kg with the magnetic structure indicated in Fig. 5. This value agrees well with the magnetization of the titanomaghemite deduced from the experimental value corrected for the presence of ilmenite (36 J/T/kg).

The spin canting appears mostly on B sites, because of the relatively greater non-magnetic substitution on A sites in the above formula, and

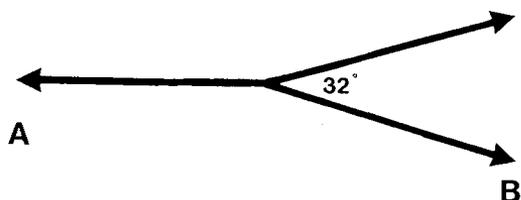


Fig. 5. The magnetic structure of titanomaghemite in V2m showing the directions of spins in the A and B sublattices.

the presence of a maximum of only four A site superexchange neighbours of any B site in the spinel structure (Coey, 1987). Spin canting will be significant for iron on those B sites with only one or two A site iron neighbours. The proportion of reversal spins associated with B site iron with zero A site neighbours will be negligible.

Conclusion

Petrographically fresh rock of basaltic composition has been found to contain highly-oxidized grains of titanomaghemite, which host exsolution lamellae of ilmenite.

The spinel has a canted spin structure, with canting mainly on the octahedral iron sublattice caused by substitution of non-magnetic ions, mainly titanium. The magnetization, 36 J/T/kg, is only about half that of pure maghemite. The titanomaghemite is more resistant to weathering than ilmenite, and persists far into the soils formed from the basalt. Magnetic measurements may be useful to quantify it in this context.

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