Dufrenite in iron-formation on the Kangnas farm, Aggeneys district, Bushmanland, South Africa

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Abstract

An occurrence of yellowish green to green dufrenite in iron-formation from Proterozoic metamorphites on Kangnas farm near Aggeneys, Bushmanland, South Africa is described. The mineral formed by reaction between iron oxides and hydroxides and a phosphate, probably apatite. It was identified by its characteristic X-ray diffraction pattern, but microprobe analyses presented here indicate that the chemical compositions are extremely variable.

KEYWORDS: dufrenite, EMP analyses, Proterozoic metamorphites, iron-formation.

Introduction

DUFRENITE is an iron phosphate with an ideal formula of $CaFe_6(PO_4)_4(OH)_6.2H_2O$ (Moore, 1970). The pure mineral is scarce in nature and, although first described in 1833 (Nriagu, 1984), it occurs at only a few localities. In the literature it is described as a secondary phosphate at Rock Run, Cherokee County, Alabama; at Wheal Phoenix, Cornwall, England; and at deposits in Saxony, Hesse, Westphalia and Thuringia, Germany (Nriagu, 1984).

The occurrence of dufrenite at Kangnas was first mentioned by Van der Westhuizen *et al.* (1986). This paper reports results of the first investigation of this rare mineral at the former locality and presents data on its range of chemical composition, as well as its mineralogy.

Occurrence and geological setting

Dufrenite occurs as a secondary iron phosphate in iron-formation on Kangnas farm, which is situated approximately 60 km east of the town of Springbok in Bushmanland (Fig. 1).

The Kangnas, K'oumoesk'naap (Fig. 2) and Ou-Eendop Formation (not shown in Fig. 2) form part of the Khurisberg Subgroup (SACS, 1980) which in turn is part of the Bushmanland Group (Strydom *et al.*, 1987). The Kangnas Formation represents the structurally lowermost unit and SWA/NAMIBIA Qrange River Aggeneys •Pofadder Kangnas Springbok R.S.A. ATLANTIC OCEAN •Vanrhynsdorp 50 0 100 km Cape Town

FIG. 1. Locality map of the Kangnas area in the north western Cape Province, South Africa.

overlies refoliated augen gneiss. The latter probably belongs to the Klein Namaqualand Suite (Strydom, 1985). The Kangnas Formation is approximately 70–180 m thick and can be subdivided into a basal brown quartzite unit (70–100 m thick), consisting of several second-order upward

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FIG. 2. Geological map of Kangnas and environs.

coarsening cycles which are composed of quartzsillimanite-biotite schist, and thin quartzite with interbedded schist layers. The upper iron-formation unit of the Kangnas Formation consists of quartz-sillimanite-biotite schist with interbedded white to brown quartzites and iron-formation beds and lenses. The quartzites have a massive appearance, but are laminated near the contact with the overlying K'oumoesk'naap Formation. The contact between the Kangnas and K'oumoesk'naap Formations is gradational from the laminated quartzites of the former to the medium to thickly bedded quartzites of the latter. The Khurisberg Subgroup was subjected to several episodes of complex and intense deformation, resulting in isoclinal folds. This produced a duplication of the stratigraphy at some localities (Strydom, 1985).

Petrographical relations

Dufrenite is a secondary mineral which predominantly occurs as small isolated, rounded grains in cracks in garnet, juxtaposed to garnet or adjacent to goethite in iron-formation (Fig. 3). Most of these grains display botryoidal features with



FIG. 3. Photomicrograph of dufrenite-bearing sample from Kangnas: Dufrenite (d) adjacent to garnet (gn), goethite (go), and quartz (q).

a radial fibrous structure. The grains range from 0.01 to 0.25 mm in diameter with a mean of 0.10mm. Stringers of dufrenite also exist, but are less common than the isolated grains. This highly pleochroic mineral is green, bluish-green to yellowish-green in colour, but the variation in colour is heterogeneous.

The iron-formation in which the dufrenite occurs consists of garnet-hematite quartzite. The garnet varies in composition from predominantly spessartine to spessartine-rich almandine. Goethite is omnipresent and results from the breakdown of hematite.

Mineral chemistry

Electron microprobe analyses of dufrenite and associated minerals were carried out with a CAMECA CAMEBAX electron microprobe using a 15 kV accelerating voltage, 40 nA beam current and $1 \mu m$ beam diameter. Mineral standards were used in conjunction with ZAF-correction techniques.

Due to the inability of the electron microprobe to analyse for H_2O and the non-availability of sufficient amounts of material for a modified Penfield determination (Kolthoff *et al.*, 1969), the (F,OH) was accepted to be 8 and the crystal water calculated accordingly to sum to 100% total. All analyses were recalculated using the above norms and are presented in Table 1. Structural formulae were computed on the basis of 24(O,OH,F) and the charge balances calculated. All data from literature were recalculated on the basis of 24(O,OH,F) to conform with the data from Kangnas.

From Tables 1 and 2 it is evident that there is a large variation in the chemical composition of the dufrenite from Kangnas, especially in the Fe₂O₃, Al₂O₃ and P₂O₅ concentrations. Al₂O₃, Na₂O, P₂O₅ and F are prominent constituents in the samples from Kangnas, while data from the literature show that Fe₂O₃, MnO and CaO are more abundant elsewhere, and with the exception of a small number of samples, they are almost devoid of Al₂O₃. It is clear that both Al₂O₃ and F may not always have been analysed for (see Table 1).

The correlation matrix (Table 3) which was generated from the data contained in Table 1 shows that Fe₂O₃ has a good correlation with both P_2O_5 and Al_2O_3 . As both these values are negative, it may be attributed to the breakdown of goethite by Al- and phosphatic solutions and the consequent formation of dufrenite. The relationship between these elements is also illustrated by Fig. 4awhere Al, P, and Fe replace one another in the structure. The replacement of Fe^{2+} by Mn and of Fe^{3+} by Al is also depicted in Fig. 4b. The amount of replacement of Fe²⁺ by Mn is very small in the case of the Kangnas samples, whereas the replacement of Fe^{3+} by Al is very low for samples from literature. When the minerals associated with dufrenite from Kangnas are considered, it is evident that the hematite contains negligible amounts of MnO, whereas the garnet is MnO-rich. End-member compositions of the garnet were calculated by the method of Rickwood (1966), while those of hematite were calculated by the method of Carmichael (1967).

Although a number of the analysed samples (Table 1 no. 6 and 13) show very high Al_2O_3 values there is, however, a wide variation in the Al₂O₃ content of the samples from Kangnas, ranging from 3.98 to 12.10 wt.% Al₂O₃ (Table 2). However, only two dufrenite samples from literature contain in excess of 1 wt.% Al₂O₃ (nos. 26 and 34). The presence of the high Al_2O_3 in some samples makes it possible to refer to these as aluminian dufrenite. This is, as far as known, the highest concentration of Al₂O₃ ever recorded for dufrenite; Hey (1955) reported that dufrenite sometimes contains appreciable aluminium, although he omitted any values. According to Moore (1984) replacement of Ca by Na is possible. This can also be observed in the samples from Kangnas (Table 1).

According to Frondel (1949), several basic iron phosphates have been confused with dufrenite and he redefined dufrenite as a monoclinic basic phosphate of ferric and ferrous iron (Fe²⁺ Fe³⁺₄(PO₄)₃(OH)₅.2H₂O). Palache *et al.* (1951), however, found that small amounts of Ca can substitute for Fe in the structure and the formula would then change and so becomes CaFe₄(PO₄)₃(OH)₅.2H₂O. Moore (1970) regards the formula for dufrenite to be $CaFe_2Fe_{10}(OH)$ -12(H₂O)₄(PO₄)₈ while Nriagu (1984) presents it as $CaFe_6(PO_4)_4(OH)_6.2H_2O$.

X-ray data

All the samples from Kangnas were subjected to close scrutiny by X-ray diffraction using Debye-Scherrer cameras (114.6 mm. diameter with Nifiltered Cu-K α radiation). The samples were prepared using the small-sample method of Hiemstra (1956). The *d*-values were measured and then compared with the data generated from the atomic coordinates of Moore (1970) by the Lazy Pul-

TABLE 1 CHEMICAL ANALYSES OF DIFFENITE AND OTHER PHOSPHATES

verix computer programme (Yvon *et al.*, 1977). These *d*-values also compare favourably with the data in literature (Table 4). The *d*-values from the Kangnas samples were subjected to the least-squares refinement programme of Appleman *et al.* (1972) for determination of cell parameters accepting a C2/c space group. These are also presented in Table 4 and clearly reflect the different phosphatic phases.

Conclusion

Dufrenite represents a secondary phosphatic alteration product of iron oxides and hydroxides.

Sample	No. 1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18
SiO2 Al2C3 Fe2O3* MnO CaO Na2O K2O P2O5 F H2O+ Total F×O	00.30 04.21 49.23 00.21 00.24 01.38 00.02 31.87 02.54 11.07 101.07	00.70 05.02 54.84 00.44 00.29 01.41 00.02 23.66 02.05 12.43 100.86 -00.86	00.34 05.13 48.43 00.40 00.33 00.95 00.02 30.81 01.19 12.48 100.80	00.25 04.46 48.52 00.62 00.34 01.24 00.02 32.28 02.78 10.66 101.17 -01.17	01.19 04.63 65.86 00.72 00.23 00.42 00.01 13.09 00.92 13.32 100.39	00.82 12.10 39.01 00.84 01.50 01.56 00.40 32.64 02.71 09.56 101.14 -01.14	00.40 07.81 42.22 00.10 00.44 00.35 00.39 33.18 01.68 14.14 100.71 -00.71	01.16 09.07 42.80 00.26 01.37 00.07 32.16 02.64 11.32 101.11 -01.11	00.34 08.27 44.01 00.31 00.33 01.07 00.35 31.52 03.51 11.77 101.48 -01.48	01.69 03.98 48.09 00.00 00.46 02.30 00.11 32.90 00.00 10.47 100.00 -00.00	00.32 08.55 42.35 00.25 01.02 00.02 32.50 03.32 101.40 -01.40	00.40 07.45 43.39 00.27 00.24 01.01 00.03 32.67 02.88 12.87 101.21 -01.21	00.16 11.99 38.55 00.50 00.04 00.00 00.02 32.78 03.22 14.10 101.36 -01.35	00.33 07.40 43.74 00.35 00.23 01.93 00.06 32.27 02.40 12.30 101.01 -01.01	00.00 00.00 56.20 06.76 00.00 00.00 27.84 00.00 09.20 100.00 100.00	00.00 00.00 63.45 00.00 00.00 00.00 27.72 00.00 08.83 100.00	00.00 00.87 53.15 00.00 01.68 00.00 00.00 31.10 00.00 13.20 100.00	00.00 00.00 55.63 00.00 01.50 00.00 30.26 00.00 12.61 100.00
Total	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00	100.00
						STRUC	TURAL F	ORMULAE	TO THE	BASIS	OF 22(0	,он, ғ)						
Si Al Fe3+ Mn Ca Na K P	00.05 00.75 05.58 00.03 00.04 00.04 00.00 04.06	00.11 00.95 06.64 00.06 00.05 00.44 00.00 03.22	00.05 00.92 05.55 00.05 00.05 00.28 00.00 03.97	00.04 00.79 05.47 00.08 00.05 00.36 00.00 04.09	00.21 00.96 08.70 00.11 00.04 00.14 00.00 01.95	00.12 02.03 04.19 00.10 00.23 00.43 00.07 03.94	00.06 01.38 04.75 00.01 00.07 00.10 00.07 04.20	00.17 01.57 04.72 00.03 00.04 06.39 00.01 03.99	00.05 01.46 04.96 00.04 00.05 00.31 00.07 04.00	00.25 00.68 05.28 00.00 00.07 00.65 00.02 04.06	00.05 01.51 04.76 00.03 00.04 00.30 00.00 04.11	00.06 01.31 04.89 00.03 00.04 00.29 00.01 04.14	00.02 02.10 04.30 00.06 00.01 00.00 00.00 04.11	00.05 01.30 04.92 00.04 00.04 00.56 00.01 04.08	00.00 00.00 06.60 00.89 00.00 00.00 00.00 03.68	00.00 00.00 07.33 00.00 00.00 00.00 00.00 03.60	00.00 00.16 06.19 00.00 00.28 60.00 00.00 04.08	00.00 00.00 06.52 00.00 00.25 00.00 00.00 03.99
F OH	01.66 06.34	01.43 06.57	01.27 06.73	01.81 06.19	00.70 07.30	01.68 06.32	01.09 06.91	01.68 06.32	02.29 05.71	00.00 08.00	02.16 05.84	01.88 06.12	02.08 05.92	01.56 06.44	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00
SUM	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00
н20	02.39	03.39	02.98	02.23	04.15	01.39	03.60	02.37	03.02	01.09	03.47	03.37	04.01	02.91	00.79	00.52	02.82	02.55
n	20.00	06.00	03.00	08.00	03.00	07.00	10.00	10.00	10.00	13.00	10.00	05.00	05.00	05.00				
SAMPLE	No. 19	20	21	22	23	24	25	26	27	28	29	30	31	32	33	34	35	
Si02 Al203 Fe203* MnO CaO Na20 K20 F205 F H20+ Total F=0 Total	00.00 00.00 58.30 00.00 00.00 00.00 31.80 09.96 100.00 -00.00 100.00	00.00 00.00 61.55 00.00 00.00 00.00 31.82 00.00 06.63 100.00 -00.00	00.00 00.21 57.67 00.40 01.12 00.00 00.00 31.76 00.00 08.84 100.00 -00.00 100.00	00.00 00.29 57.62 00.24 00.00 00.00 01.00 10.19 100.00 100.00 100.00	$\begin{array}{c} 00.00\\ 00.00\\ 56.94\\ 02.24\\ 00.00\\ 00.00\\ 32.86\\ 00.00\\ 07.96\\ 100.00\\ -00.00\\ 100.00 \end{array}$	00.00 00.00 57.97 02.84 00.00 00.00 00.00 30.43 00.00 08.76 100.00 -00.00 100.00	00.00 00.00 51.54 05.51 00.99 00.00 00.00 31.87 00.00 10.09 100.00 -00.00 100.00	00.00 04.44 49.42 00.64 00.00 00.00 31.26 00.00 14.04 100.00 -00.00 100.00	$\begin{array}{c} 00.00\\ 00.00\\ 64.74\\ 00.00\\ 00.00\\ 00.00\\ 28.39\\ 00.00\\ 06.87\\ 100.00\\ -00.00\\ 100.00 \end{array}$	00.00 00.00 47.53 05.71 00.00 00.00 32.40 00.00 13.91 100.00 -00.00 100.00	00.00 02.02 03.00 00.00 00.00 27.71 00.00 10.27 100.00 100.00	00.00 00.00 59.14 02.09 00.00 00.00 25.20 00.00 13.57 100.00 100.00	00.00 80.03 60.00 00.00 00.00 00.00 06.25 00.00 13.72 100.00 100.00	00.00 00.00 66.49 00.00 00.00 20.11 00.00 13.40 100.00 100.00	00.00 79.86 00.00 00.00 00.00 05.66 00.00 14.48 100.00 100.00	00.00 01.75 73.58 00.26 00.55 00.00 13.69 00.00 10.17 100.00 10.00 100.00	02,64 00.00 52,73 01.50 02.84 00.00 15.90 00.00 15.90 024.39 100.00 100.00	
						STRUC	TURAL F	ORMULAE	TO THE	BASIS	OF 22(0	,OH,F)						
Si Al Fe3+ Mn Ca Na K P	00.00 00.00 06.59 00.00 00.00 00.00 00.00 00.00 04.05	00.00 06.77 00.00 00.00 00.00 00.00 00.00 00.00	00.00 00.04 06.47 00.05 00.18 00.00 00.00 00.00 00.00	00.00 06.53 00.03 00.00 00.00 00.00 00.00 00.00	00.00 06.31 00.28 00.00 00.00 00.00 00.00 00.00	00.00 00.00 00.36 00.00 00.00 00.00 00.00 00.00	00.00 05.91 00.71 00.16 00.00 00.00 04.11	00.00 00.80 05.69 00.03 00.10 00.00 00.00 00.00 00.00	00.00 07.32 00.00 00.00 00.00 00.00 00.00 00.00 00.00	00.00 05.56 00.06 00.95 00.00 00.00 04.26	00.00 07.26 00.00 00.00 00.00 00.00 00.00 00.00 00.00	00.00 00.00 07.30 00.29 00.00 00.88 00.88 00.88 00.88	00.00 00.00 10.47 00.00 00.00 00.00 00.00 00.00 00.92	00.00 00.00 08.51 00.00 00.00 00.00 00.00 00.00 00.00	00.00 00.00 11.77 00.00 00.00 00.00 00.00 00.00 00.94	00.00 00.36 09.55 00.04 00.10 00.00 00.00 00.00 00.00	00.51 00.00 07.72 00.25 00.59 00.00 00.00 00.00 02.62	
F OH	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	00.00 08.00	
SUM	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	08.00	
H20	00.90	00.77	00.39	01.12	00.09	00.42	01.12	03.17	-0.56	03.21	01.32	03.43	3.95	03.60	05.46	01.85	11.83	

1-14, dufrenite from Kangnas; 15-20, dufrenite, Table 3, Frondel, 1949; 21-24, rockbridgeite, Table 9, ibid.; 25-28, unidentified dufrenite-like minerals, Table 11, ibid.; 29-35, altered dufrenite and rockbridgeite, Table 2, ibid. FeO from literature converted to Fe203 to correspond to format of Kangnas samples from present study.

	Data fr	om Kangnas	CHANDARD	CMATT DOM	TADOROT
		MENN	DEUTATIO	J VALUE	VALUE
N-VA	NIADLE	PIEAD	DEVIATION	A VALUE	VALUE
NO.	Name				
1	Si	.6000	. 4541	.1600	1.6900
2	A1	7.1479	2.7219	3,9800	12,1000
3	Fe3	46.5029	7.0937	38,5500	65.8600
4	Mn	. 3764	. 2327	.0000	.8400
5	Ca	. 3700	. 3407	.0400	1.5000
6	Na	1.1436	.6110	.0000	2.3000
7	ĸ	.1100	.1492	.0100	.4000
8	P	30.3093	5.4974	13.0900	33.1800
ğ.	-	2.3257	.9629	.0000	3.5100
	Literat	ure			
1	Si	.1257	. 5761	.0000	2.6400
2	Al	.3600	1.0232	.0000	4.4400
3	Fe3	60.2648	8.8376	47.5300	80.0300
4	Mn	1.0710	1.9000	.0000	6.7600
5	Ca	.7157	1.3772	.0000	5.7100
6	Na	.0000	.0000	.0000	.0000
7	ĸ	.0000	.0000	.0000	.0000
8	P	25.9852	8.5273	5.6600	32.8600
9	F	.0000	.0000	.0000	.0000
-	-				

Table 2. Statistical parameters of the chemistry of dufrenite and related phosphates.

Table 3. Comparison of correlation matrices of Kangnas dufrenite with data from literature.

	Kangnas	Literature
Si02 - F	- 0.7438	0.0000
Al2O3 - Fe2O3	- 0.7356	- 0.1789
Fe203 - P205	- 0.9006	- 0.8051

It occurs mostly as botryoidal fibres and masses, commonly associated with goethite and less often with magnetite and/or hematite. The mineral displays a large colour variation from green to yellow.

A comparison of the unit cell dimensions and consequently the cell volume indicates a large variation in these parameters. At present insufficient data are available to bring unit cell dimensions and chemistry into relation with each other.

From Table 1 it is evident that 'pure dufrenite' very seldom occurs in nature. Dufrenite could better be regarded as a general term encompassing iron phosphates displaying the characteristic dufrenite X-ray powder pattern, as the chemical composition is too variable to absolutely define its chemical ranges. From the data presented (Table 1) is seems that Ca substitution as proposed by Moore (1970) and Nriagu (1984) is minor even in samples reported as "pure dufrenite". The present authors have shown that substitution by other elements, viz. Al, Ca, Na, Si and Mn, also takes place.

TABLE 4. D-VALUES AND UNIT CELL PARAMETERS OF DUFRENITE FROM VARIOUS LOCALITIES.

	11	2	3	4
	12.30	12.00		
	6.90	6.78	6.92	
	6.54	6.41	6.46	6.46
	6.10	5.99	6.11	
	5.59		5.54	
	5.05	5.00	5.06	5.00
	4.84	4.80	4.85	
	4.40	4.35	4.40	4.40
	4.15	4.11	4.16	4.20
	4.04		4.05	
	3.79	3.76	3.79	3.80
	3.67		3.68	
		3.63	3.63	
	3.54	3.51		
	3.42	3.39	3.44	3.39
	3.24	3.21	3.24	3.20
	3.17	3.15	3.18	3.16
	3.01	2.99	3.007	2,990
	2.88	2.86		2.86
	2.81	2.795		
	2.64	2.627		
•	24 60	25 94	25 70	25 00
ĥ	5 14	5 126	4 88	5 18
00	10 07	13 78	13 71	14 02
0	100925/	111012/	1099467	111934/
β	100 25	111 12	103 40	111 34

 First 20 reflections; dufrenite from Hirschberg, Germany (JCPDS, 1980).

 First 18 reflections; dufrenite from Cornwall, England (JCPDS, 1980).

- Dufrenite from Kangnas (Meyer, 1986) sample LKG-009.
- 4. Dufrenite from Kangnas, this study, DKAG-16.



FIG. 4a. Ternary relationship between Al, Fe and P in dufrenite from Kangnas (squares) and other phosphates from literature (triangles)—Table 1. FIG. 4b. Similar diagram of Al, Fe and Mn for dufrenite from Kangnas (squares) and phosphates from literature (triangles)—Table 1.

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