

## A stable carbon isotopic record of climatic change from a tropical mountain ecosystem in Colombia

A. Boom

Hugo de Vries-laboratory, Dept of Palynology and Paleo/Actuology, University of Amsterdam, Kruislaan 318, 1098 SM Amsterdam, Netherlands Centre for Geo-ecological Research (ICG)

Netherlands Institute for Sea Research, (NIOZ), Dept of Marine Biogeochemistry and Toxicology, P.O. Box 59, 1790 AB Den Burg, The Netherlands

FOM Institute for Atomic and Molecular Physics, Kruislaan 407, 1098 SJ Amsterdam, the Netherlands

S. Schouten

J. S. Sinninghe Damsté

J. W. de Leeuw

Netherlands Institute for Sea Research, (NIOZ), Dept of Marine Biogeochemistry and Toxicology, P.O. Box 59, 1790 AB Den Burg, The Netherlands

J. J. Boon

FOM Institute for Atomic and Molecular Physics, Kruislaan 407, 1098 SJ Amsterdam, the Netherlands

H. Hooghiemstra

Hugo de Vries-laboratory, Dept of Palynology and Paleo/Actuology, University of Amsterdam, Kruislaan 318, 1098 SM Amsterdam, Netherlands Centre for Geo-ecological Research (ICG)

The altitudinal position of the highplain of Bogotá at 2550 m elevation renders its unique archive in the lake sediments sensitive to past climatic change. At present the upper forest line lies at 3200 meters, but during glacial time, it descended to around 2000–1800 m. Above the forest line lies the paramo, a vegetation belt dominated by grasses and shrubs. Palynological studies of a 568-m long core, taken in the centre of this large drained lake, resulted in a pollen record showing changes in climate and vegetation during the last 3.2 million years.

Here we present organic geochemical data from this core Funza II that give new insights into the regional climatic history. Bulk analyses of the total organic carbon (TOC) and irmGC-MS analyses of the lipids, show significant isotopic excursions. The excursions of the *n*-alkanes, derived from leaf-wax lipids follow similar trends with  $\delta^{13}\text{C}$  values ranging from  $-30$  to  $-12\text{‰}$  (e.g. Fig. 1). These isotopic excursions of terrestrial signals can be interpreted as shifts in terrestrial plant composition from  $\text{C}_3$  to  $\text{C}_4$  dominated ecosystems.  $\text{C}_4$ -Ecosystems mainly include grasses that are highly adapted to aridity or severely  $\text{CO}_2$ -stressed environments, while  $\text{C}_3$  ecosystems usually consist of trees and other grasses that are less adapted to environmental stress. At present,  $\text{C}_4$  ecosystems are mainly found

in tropical savannahs or water stressed environments. The grass paramo that is found above the forest line of today is completely dominated by  $\text{C}_3$  grasses. This implies that previous reconstructions of the vegetation within glacial times need to be revised. The paramo as we know it today, did not, descended to lower altitudes. More likely is that the  $\text{C}_4$  savannah now occurring at altitudes below that of the Andean forest, triggered by the glacial low  $p\text{CO}_2$  situation, actually started to colonise higher altitudes and out competed the  $\text{C}_3$  grasses of the paramo.

The aquatic ecosystem also responded to climatic changes by changing fractionation effects towards  $^{13}\text{C}$ . Octahydrobotryococcene and mono cyclic botryococcenes, biomarkers for *Botryococcus braunii* algae have extremely heavy  $\delta^{13}\text{C}$  values up to  $+5\text{‰}$ . Such high values can only be explained when bicarbonate is assimilated without any significant fractionation by the algae. Possibly, at some stages dissolved  $\text{CO}_2$  was not available to the algae while bicarbonate still was. Isotopically heavy mono cyclic botryococcenes have also been found. The timing of these events fall in glacial intervals, an explanation for this phenomena would be that the lake had started to dry out and experienced increasing alkaline conditions possibly due decreased rainfalls during glacial times. This combined with the low

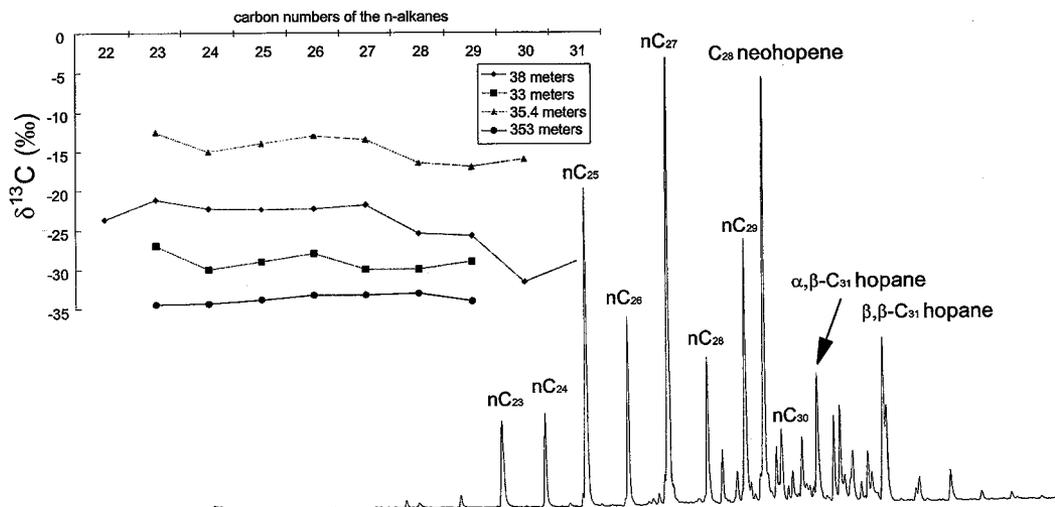


FIG. 1. Chromatogram from the apolar fraction of the 38 m depth sample, with predominant *n*-alkanes. The inset shows their  $\delta^{13}\text{C}$  values (diamonds) and those of *n*-alkanes present in other samples. Samples of 35.4 and 353 meters depth represent end members, where terrestrial alkanes are mainly derived from C<sub>4</sub> and C<sub>3</sub> plants, respectively.

atmospheric  $p\text{CO}_2$ , would make dissolved  $\text{CO}_2$  virtually unavailable for the algae. The terrestrial ecosystem would respond to this increasing aridity by favouring even more C<sub>4</sub> grasses on the land. The shallow parts of the lake could at some stages be colonised by C<sub>4</sub> sedges.

Thus the stable carbon isotope record of this tropical mountain ecosystem shows changes that can be related to climatic change and suggests an increased aridity during glaciations and reflects

decreases in atmospheric  $p\text{CO}_2$ . It also shows that C<sub>4</sub> grasses are able to compete with C<sub>3</sub> grasses at lower temperatures while under low  $p\text{CO}_2$ . It may be possible to make estimations of the palaeotemperature when evaluating the  $\delta^{13}\text{C}$  values of the *n*-alkanes. These temperature estimations will result in less drastic glacial cooling of this tropical area than previously thought and will correlate better with palaeo-oceanographical derived temperatures.