CHARLES PALACHE AND HARRY BERMAN, Harvard University, Cambridge, Mass.*

ABSTRACT

Boulangerite, a lead sulfantimonide, has been studied on new material, the first which has proved suitable for detailed crystallographic examination. It is monoclinic, prismatic, 2/m, pseudo-orthorhombic. Elements a:b:c=0.9158:1:0.3456; $\beta=100^{\circ}39\frac{1}{2}'$. Some 60 crystal forms were observed. Lattice constants (Berry): $a_0=21.14$ Å, $b_0=23.46$ Å, c_0 =8.07 Å, $\beta=100^{\circ}48'$. New analyses by Gonyer on Washington boulangerite confirmed Shannon's earlier analysis and Berry's cell content of Pb₄₀Sb₃₂S₃₈.

CRYSTALLOGRAPHY

Boulangerite is one of the many fibrous sulfantimonides of lead concerning the crystal form of which little of a definite nature has hitherto been known. Its slender striated prisms generally show no terminal faces. Sjögren (1897) measured crystals from Sala, Sweden, which he regarded as orthorhombic, and on the basis of a single domal face he was able to establish an axial ratio. Shannon (1921) found a single pyramid face on a crystal from Stevens County, Washington, and also inferred orthorhombic symmetry but obtained an axial ratio in very poor agreement with that of Sjögren.

New studies on boulangerite from three localities have been successful in revealing its true crystallographic nature. The mineral proves to be monoclinic with an astonishing wealth of forms. The results of these studies are presented in the following pages.

The first measurable crystals of boulangerite were found in small vugs of a massive, coarsely fibrous material from Rocker Gulch Placer Claim near Deerlodge, Montana (Specimen No. 92671). This material was supplied by Ward's Natural Science Establishment and bore the name geocronite. In addition to the dominant boulangerite, there is coarsely granular sphalerite present, crystals of pyrite, and, in a single cavity, crystals of bournonite, identified by crystal measurement, and showing the forms:—

c 001	e 210	$\Sigma 031$	z 201	p 223	0 213
a 100	к 013	x 102	φ 113	y 111	v 211
m 110	n 011	o 101	u 112	g 221	

The boulangerite crystals are needles of dark iron-gray color, less than a half millimeter in diameter and from a few to ten or more millimeters

* Contribution from the Department of Mineralogy and Petrography, Harvard University, No. 253.

in length. Since they project from the walls of vugs, they are all singly terminated. The needles are very brittle, the slightest pressure causing them to break into thin leaf-like flakes along a single pinacoidal cleavage, which proved to be parallel to $a\{100\}$. The prisms are deeply striated parallel to their length so that they gave generally a continuous chain of weak or colored signals on the goniometer, but the best crystals gave a consistent series of prism forms. The numerous terminal faces are so minute and so irregular as to give under the binoculars no definite clue of symmetrical distribution; but their weak signals gave position angles which when projected yielded a definite pattern, repeated, although with great variation in the forms present, on successive crystals. By plotting the measurements of each crystal on transparent paper and superimposing these projections, with rotation about the projection center, parallelism was obtained and the form series was developed.

The projection showed apparent orthorhombic symmetry as regards spacing of poles. However, the radial zones to prism faces did not intersect in the projection center but rather at a point about ten degrees off center—a point not represented by a crystal face on any of the crystals at first measured. This fixed the symmetry as probably monoclinic, a conclusion strengthened by each succeeding measurement and subsequently proved by x-ray study as reported on a later page.

Identical crystallographic characters including the wealth of forms were later found on boulangerite from the Gold Hunter Mine, Mullan, Idaho, described by Shannon (1918) as "mullanite"; and on crystals from Stevens County, Washington, also analyzed by Shannon (1925).

Boulangerite is, then, monoclinic but with pronounced pseudosymmetry both orthorhombic and tetragonal, as shown in the gnomonic projection, Fig. 1, which presents several unusual features. Although there is an unusually rich form series, one primary zone [010], parallel to the symmetry plane, is almost missing since both {100} and {001} are seldom found and only one orthodome, {101}, is at all common. The axial zone [001] is of course strong, the crystals being strongly prismatic in that direction and showing four well-established prism forms. The clinodome zone [100] is also well marked. A strange feature of this projection is the absence of poles in the central area. {001}, {011} and { $\overline{2}01$ } are among the rarest forms and the unit pyramid {111} has been found on but two crystals. If these be omitted, there are but three forms which have a slope angle (ρ) less than 40°.

The lattice of boulangerite is pseudo-orthorhombic because x_0' is almost exactly half of p_0' , another case like that of brochantite recently described by Palache (1939). It is pseudotetragonal because p_0' and q_0' are so nearly equal and the ϕ of {110} is about 48°.



FIG. 1. Gnomonic projection of boulangerite. O indicates most common forms.

The combinations of forms found on eight crystals from Montana, five from Idaho, and two from Washington are shown in Table 1. From this table it appears that the most frequently occurring forms on Montana crystals are: $a, q, m, n, f, g, D, \delta, G, \phi, R, \rho, X$, and Λ ; on Idaho crystals c, m, f, x, α, E , and ϕ ; on Washington crystals f, x, A, α, G, ϕ , and θ . Of these forms only two, $f\{031\}$ and $\phi\{\overline{131}\}$ are among the commonest for all three localities. Several, however, are common on two localities. It is clear that many forms occur at all the localities and that a larger number are confined to one. This last fact may be due to the better and more numerous crystals measured from Montana. At least ten forms not listed were found once only on the Montana crystals.

Lo	cality				Mon	tana						Idaho	0		Washi	ington
Cr	ystal	1	2	3	4	5	6	7	8	1	2	3	4	5	1	2
с	001										x		x	x		
b	010	х	х		x		X									
a	100	х	x	х			x	х		x						
q	130	x		х	х	х		х								
P	120										х					
m	110	x	х	х	х	x	x	х	х	X	x	X			x	
п	210		х	х	х	x	х	х	х	x	х				x	
d	011										X		x			
е	021								x	х					х	
f	031	x		Х	X	X			х	х	x	x			x	х
g	041			х	x	х		x	х							
j	051		X		х	x			X							
0	061	x							х							
x	101							х	х	x			х	х	x	х
y	201												х	х	x	
p	111										x					
A	221			x				x				x			x	х
В	331	x		х	x	х										
D	441	x			x	х		х	х							
C	551		X				х									х
α	$\overline{2}21$								x	x	x	x	х	x	x	x
β	331		x		х			х	х							
δ	4 41	1	x		х	х		х	х							
E	121								x	х	х	x	х		x	
F	131			х		Х			х							
G	141	x	х	x	х	x	x	x	х						x	х
L	151						х									
e	T21								x		x	х				
ϕ	T 31		x		х	х			х	x	х	X			x	х
Y	T 41		x						x							
λ	T51				x	x				0						
M	211								х	x		х			h	
N	231			х	x	x		x							1.0	
R	251	x	X		х	х	x		х							
μ	211								x		х				X	
ν	231				х				x							
θ	241			x	x			x	x						x	x
ρ	251	x		x	x	x			x						100	
S	311		x					x	х						1	
Ţ	321	x	x	x												х

TABLE 1. COMBINATIONS OF FORMS ON BOULANGERITE

555

Locality				Mor	ntana	L					Idaho)		Wash	ington
Crystal	1	2	3	4	5	6	7	8	1	2	3	4	5	1	2
U 341	x	x			x	x						10000			
V 361				x	x			x							
σ 311				x			х	x					x	1	
τ 321								x	x		x				
ω 351		х			x			x							
W 411	х						x	x							
X 421	х		х	x	x			x							
г 411				x			x								
Δ 421				x			x	x							
Y 531	x	х				x									
X 541		x												x	
Θ 631				х			x							1.12	
Ω 641				x	X			x						x	
Λ 651		x		x	x			x	x						

TABLE 1—Continued

TABLE 2a. MEASURED ANGLES ON BOULANGERITE FROM MONTANA

		No,	Meas,	Aver.		Rar	ıge	Calcu	ulated
6			φ	ρ	φ	a 18559 -	ρ	φ	ρ
Ъ	010	4	0°27′	90°00′	- 2°38'-+	2°35′		0°00′	90°00′
a	100	5	88 44	90 00	89 04 -	94 48		90 00	90 00
q	130	6	20 20	90 00	19 39 -	21 09	-	20 19 ¹ / ₂	90 00
т	110	22	47 55 ¹ / ₂	90 00	47 15 -	48 42		48 01	90 00
n	210	18	65 39	90 00	65 00 -	66 15		$65 \ 46\frac{1}{2}$	90 00
е	021	1	15 14	35 53			-	15 14	35 37
f	031	5	10 25	$46 \ 34\frac{1}{2}$	9 21 -	10 55	46°25′-46°43′	10 171	46 30
g	041	6	7 58	54 091	7 31 -	8 27	54 14 -55 00	7 45	54 22
j	051	5	6 13	60 06	5 59 -	6 31	59 49 -60 30	6 13	$60 \ 05\frac{1}{2}$
0	061	2	$5\ 07\frac{1}{2}$	64 21	4 46 -	5 29	64 18 -64 24	5 11	$64\ 20\frac{1}{2}$
x	101	2	$90\ 04\frac{1}{2}$	29 52	89 28 -	90 41	29 50 - 29 54	90 00	$29 \ 46\frac{1}{2}$
A	221	2	53 39 ¹ / ₂	50 25	52 58 - 3	54 21	49 45 -57 05	$54 \ 08\frac{1}{2}$	49 43
В	331	5	52 18	59 25	52 02 -	52 26	59 17 -59 34	52 16 ¹ / ₂	59 27
D	441	6	51 18	65 39	51 05 -	51 27	65 32 -65 55	51 161	65 39
С	551	2	50 08	$70 \ 01\frac{1}{2}$	49 50 -	50 26	70 00 -70 03	50 39 ¹ / ₂	$69 51\frac{1}{2}$
α	221	1	-39 37	41 56	_			-39 59 ¹ / ₃	42 03 1
β	331	5	$-42\ 37\frac{1}{2}$	55 00	42 12 -	42 52	54 40 -55 22	$-4254\frac{1}{2}$	54 451
δ	441	5	-44 21	62 59	43 34 -	45 52	62 40 -63 30	$-44\ 16\frac{1}{2}$	62 37

		No.	Meas.	Aver.		Ran	ige	Calcul	ated
				0	φ		ρ	φ	ρ
E	121	1	39 43	42 10			<u></u>	39 37	41 54
F	1.31	3	29 09	49 51	29 05 -	29 11	49 22 -50 20	$28 53\frac{1}{2}$	$49 \ 49\frac{1}{2}$
G	141	10	22 29	56 16	21 16 -	23 16	56 00 -57 04	22 29	56 14 ¹ / ₂
L	151	1	18 14	61 16				18 19 ¹ / ₂	61 13
e	1 21	1	-15 19	35 45				-15 49	$35 \ 41\frac{1}{2}$
φ	T 31	5	-1052	46 08	10 09 -	11 23	45 34 -46 33	$-10\ 41\frac{1}{2}$	46 32
γ	141	2	- 8 20	$54\ 37\frac{1}{2}$	8 11 -	8 28	54 30 -54 45	$-803\frac{1}{2}$	54 23 ¹ / ₂
λ	151	3	- 6 02	59 57	5 39 -	6 29	59 47 -60 05	- 6 28	60 06
Μ	211	1	69 41	45 37				$70\ 07\frac{1}{2}$	$45\ 28\frac{1}{2}$
N	231	4	42 30	54 40	41 43 -	42 51	54 25 -55 00	42 41	54 40
R	251	8	29 06	$63\ 14\frac{1}{2}$	28 42 -	29 30	63 00 -63 37	$28\ 57\frac{1}{2}$	$63\ 08\frac{1}{2}$
μ	211	1	-59 53	33 15	-			-59 12	34 01
ν	231	3	-27 50	50 03	27 01 -	28 46	49 38 -50 30	-29 13	$49 54\frac{1}{2}$
ϕ	$\overline{2}41$	5	$-22\ 35\frac{1}{2}$	56 26	22 19 -	23 04	56 09 -56 45	-2245	$56\ 17\frac{1}{2}$
ρ	251	6	$-18\ 32\frac{1}{2}$	61 14	18 10 -	19 05	60 55 -61 33	-18 33	61 15
S	311	3	75 30	54 25	75 12 -	75 44	54 20 -54 32	$75 \ 32\frac{1}{2}$	54 09
T	321	3	62 45	56 30	62 41 -	62 50	56 16 -56 54	62 43	56 27
U	341	5	44 05	62 28	43 45 -	44 20	62 00 -62 52	$44\ 06\frac{1}{2}$	62 33
V	361	3	32 54	67 58	32 49 -	33 00	67 12 -68 35	32 521	67 57
σ	311	4	$-69\ 55\frac{1}{2}$	45 28	-69 31 -	70 10	45 19 -45 33	-70 102	45 402
τ	321	2	-54 10	49 46	-54 09 -	54 11	49 43 -49 49	-54 21	49 52
ω	351	3	-28 56	63 29	-28 43 -	29 08	63 12 -64 00	-29 09	$63\ 11\frac{1}{2}$
W	411	3	79 09	60 19	78 25 -	80 01	60 00 -60 42	78 40	00 222
X	421	8	68 17	61 33	67 49 -	68 39	61 16 -61 50	68 09 <u>5</u> .	01 422
Г	411	2	-75 35	54 31	-75 32 -	75 38	54 22 -54 40	-75 37	54 17 ¹ / ₂
Δ	421	4	$-62\ 35$	56 25	- 62 13 -	63 09	56 12 -56 32	-6251	50 34
Y	531	5	64 00	66 50	- 63 26 -	64 42	66 3667 12	63 481	00 302
Ζ	541	1	56 25	68 13		6		56 45	68 22
Θ	631	3	-63 46	67 04	- 63 40 -	63 53	66 53 -67 20	$-63\ 53\frac{1}{2}$	67 00
Ω	641	3	-56 36	68 21	- 56 19 -	56 47	68 15 -68 27	$-56\ 50\frac{1}{2}$	68 25

 $-50\ 28$ 70 $00\frac{1}{2}$ $\left|-50\ 10\ -\ 50\ 41$ 69 45 -70 18 $\right|$ -50 $45\frac{1}{2}$ 69 $53\frac{1}{2}$

TABLE 2a—Continued

651

4

Λ

		No.	Meas.	aver.	Ra	nge	Calco	ulated
			φ	ρ	φ	ρ	φ	ρ
С	001	3	88°58′	12°01′	88°58′-90°00′	10°42′-13°20′	90°00′	10°391/
a	100	1	88 40	90 00	100		90 00	90 00
Ρ	120	2	29 05	90 00	28 24 - 29 46		29 $03\frac{1}{2}$	90 00
т	110	8	48 23	90 00	47 41 -49 03		48 01	90 00
п	210	3	66 11	90 00	65 03 -67 00	—	65 461	90 00
d	011	3	29 37	21 25	28 20 - 30 48	21 13 -21 40	28 34 ¹ / ₂	21 29
е	021	Î	16 27	41 42		-	15 14	35 47
f	031	6	10 33	46 32	9 44 -11 29	46 06 -46 54	10 171	46 30
x	101	3	90 24	29 46	90 00 -91 00	29 37 -29 53	90 00	29 $46\frac{1}{2}$
у	201	2	-90 23	30 05	90 23 -90 23		-90 00	$30\ 06\frac{1}{2}$
Þ	111	1	60 44	33 37		-	58 52	$33\ 45\frac{1}{2}$
A	221	1	54 14	49 14			$54 \ 08\frac{1}{2}$	49 43
α	221	5	-39 57	41 32	- 38 36 - 41 23	41 23 -42 10	$-39\ 59\frac{1}{2}$	$42\ 03\frac{1}{2}$
\boldsymbol{E}	121	5	39 56	41 39	39 42 -40 14	41 06 -42 00	39 37	41 54
e	T21	4	-15 22	35 38	-15 03 -16 00	35 34 -35 55	-15 49	$35 \ 41\frac{1}{2}$
φ	<u>1</u> 31	4	-10 38	46 02	-10 15 -11 19	45 08 -46 00	$-10 41\frac{1}{2}$	46 32
M	211	2	69 43	45 48	69 21 -70 05	45 30 -46 06	$70\ 07\frac{1}{2}$	45 28 1
μ	211	1	-56 14	34 48	(amp)		-59 12	34 01
σ	311	Ĩ	-69 12	45 10	_		$-70\ 16\frac{1}{2}$	$45 40\frac{1}{2}$
τ	321	3	-54 19	49 34	53 31 -55 06	49 10- 49 52	-54 21	49 52
Λ	651	1	-51 17	69 50	— ,		$-50\ 45\frac{1}{2}$	$69\ 53\frac{1}{2}$

TABLE 2b. MEASUREMENTS ON BOULANGERITE FROM IDAHO.

TABLE 3. ANGLE TABLE. BOULANGERITE

Monoclinic, prismatic $\frac{2}{m}$

 $\begin{array}{l} a:b:c=0.9158:1:0.3456; \ \beta = 100^{\circ}39\frac{1}{2}'\\ p_{0}:q_{0}:r_{0}=0.3774:0.3396:1; \ \mu = 79^{\circ}20\frac{1}{2}'\\ \end{array}$

		$r_2: p_2: q_2 = 2.944$	3:1.1111:1;	$p_0' = 0.3840$	$q_0' = 0.3456$	$x_0' = 0.183$	82
F	forms	ϕ	ρ	ϕ_2	$\rho_2 = B$	С	Α
С	001	90°00′	10°39 <u>1</u> ′	79°201/2'	90 00'		79°201/2'
b	010	0 00	90 00		0 00	90°00′	90°00′
a	100	90 00	90 00	0 00	90 00	$79\ 20\frac{1}{2}$	—
**r	140	15 31 ¹ / ₂	90 00	0 00	15 31 ¹ / ₂	87 52	74 281
*q	130	$20\ 19\frac{1}{2}$	90 00	0 00	$20\ 19\frac{1}{2}$	86 19	69 40 ¹ / ₂
*P	120	$29\ 03\frac{1}{2}$	90 00	0 00	$29\ 03\frac{1}{2}$	84 51	$60\ 56\frac{1}{2}$

TABLE 3-Continued

Fo	orms	φ	ρ	ϕ_2	$\rho_2 = B$	С	A
*m	110	48 01	90 00	0 00	48 01	82 06	41 59
*n	210	$65 \ 46\frac{1}{2}$	90 00	0 00	$65 \ 46\frac{1}{2}$	$80\ 17\frac{1}{2}$	$24\ 13\frac{1}{2}$
**l	310	73 18	90 00	0 00	73 18	79 48	16 42
**k	410	77 19	90 00	0 00	77 19	79 36 1	12 41
**i	510	79 48	90 00	0 00	79 48	$79 \ 30\frac{1}{2}$	10 12
**h	710	$82 \ 40\frac{1}{2}$	90 00	0 00	$82 \ 40\frac{1}{2}$	79 26	7 19 1
d	011	$28 \ 34\frac{1}{2}$	21 29	$79\ 20\frac{1}{2}$	$71 \ 14\frac{1}{2}$	$18 \ 45\frac{1}{2}$	79 55
е	021	15 14	35 37	$79\ 20\frac{1}{2}$	$55 48\frac{1}{2}$	$34\ 11\frac{1}{2}$	81 12
f	031	$10\ 17\frac{1}{2}$	46 30	$79\ 20\frac{1}{2}$	44 28	45 32	$82\ 33\frac{1}{2}$
g	041	7 45	54 22	79 $20\frac{1}{2}$	$36\ 21\frac{1}{2}$	53 38 ¹ / ₂	$83 \ 42\frac{1}{2}$
j	051	6 13	$60\ 05\frac{1}{2}$	$79\ 20\frac{1}{2}$	$30\ 29\frac{1}{2}$	59 30 ¹ / ₂	84 37
0	061	5 11	$64\ 20\frac{1}{2}$	$79\ 20\frac{1}{2}$	$26\ 08\frac{1}{2}$	$63 51\frac{1}{2}$	85 191
**u	102	90 00	20 49	69 11	90 00	$10 \ 09\frac{1}{2}$	69 11
x	101	90 00	$29 \ 46\frac{1}{2}$	$60\ 13\frac{1}{2}$	90 00	19 07	$60\ 13\frac{1}{2}$
У	$\overline{2}01$	-90 00	$30\ 06\frac{1}{2}$	$120\ 06\frac{1}{2}$	90 00	40 46	$120\ 06\frac{1}{2}$
Þ	111	58 52	$33 \ 45\frac{1}{2}$	$60\ 13\frac{1}{2}$	73 18	$25 \ 10^{\frac{1}{2}}$	$61 \ 35\frac{1}{2}$
A	221	$54\ 08\frac{1}{2}$	49 43	46 17	$63\ 27\frac{1}{2}$	41 26	$51 \ 48\frac{1}{2}$
В	331	$52\ 16\frac{1}{2}$	59 27	$36\ 43\frac{1}{2}$	58 12	51 17	47 04
D	441	$51 \ 16\frac{1}{2}$	65 39	30 07	55 15 ¹ / ₂	57 33	44 42
C	551	$50\ 39\frac{1}{2}$	$69\ 51\frac{1}{2}$	$25\ 22\frac{1}{2}$	$53\ 28\frac{1}{2}$	61 47	$43\ 26\frac{1}{2}$
α	221	$-39\ 59\frac{1}{2}$	$42\ 03\frac{1}{2}$	$120\ 06\frac{1}{2}$	$59\ 07\frac{1}{2}$	49 27	115 30
β	331	$-42\ 54\frac{1}{2}$	$54 \ 45\frac{1}{2}$	$133\ 56\frac{1}{2}$	53 15 ¹ / ₂	$62\ 20\frac{1}{2}$	123 47
δ	441	$-44\ 16^{\frac{1}{2}}$	62 37	$143\ 25\frac{1}{2}$	$50\ 31\frac{1}{2}$	70 17	$128 \ 18\frac{1}{2}$
E	121	39 37	41 54	$60\ 13\frac{1}{2}$	59 02 ¹ / ₂	35 53	$64 \ 47\frac{1}{2}$
F	131	28 53 ¹ / ₂	$49 \ 49\frac{1}{2}$	$60\ 13\frac{1}{2}$	48 01	45 23	68 20
G	141	22 29	$56\ 14\frac{1}{2}$	$60\ 13\frac{1}{2}$	$39\ 48\frac{1}{2}$	$52 \ 46\frac{1}{2}$	$71\ 27\frac{1}{2}$
L	151	$18 \ 19\frac{1}{2}$	61 13	$60\ 13\frac{1}{2}$	$33 \ 41\frac{1}{2}$	$58\ 23\frac{1}{2}$	74 00 ¹ / ₂
e	121	-15 49	$35 \ 41\frac{1}{2}$	$101 \ 04\frac{1}{2}$	55 51	$39 \ 45\frac{1}{2}$	99 09
ϕ	T 31	$-10 \ 41\frac{1}{2}$	46 32	$101 \ 04\frac{1}{2}$	44 30	49 22 1	$97 \ 44\frac{1}{2}$
γ	141	$-803\frac{1}{2}$	$54\ 23\frac{1}{2}$	$101 \ 04\frac{1}{2}$	$36\ 23\frac{1}{2}$	56 53	96 33
λ	151	- 6 28	60 06	$101 \ 04\frac{1}{2}$	30 31 ¹ / ₂	61 51	95 36
M	211	$70\ 07\frac{1}{2}$	$45\ 28\frac{1}{2}$	46 17	$75 58\frac{1}{2}$	36 03	$47 53\frac{1}{2}$
N	231	42 41	54 40	46 17	53 09	47 52 ¹ / ₂	$56\ 25\frac{1}{2}$
R	251	$28 57\frac{1}{2}$	$63 \ 08\frac{1}{2}$	46 17	38 41	$58\ 24\frac{1}{2}$	$64\ 24\frac{1}{2}$
μ	$\overline{2}11$	-59 12	34 01	$120\ 06\frac{1}{2}$	$73\ 21\frac{1}{2}$	$43\ 28\frac{1}{2}$	$118\ 43\frac{1}{2}$
ν	$\overline{2}31$	-29 13	$49 54\frac{1}{2}$	$120\ 06\frac{1}{2}$	$48\ 06\frac{1}{2}$	$55 \ 40\frac{1}{2}$	111 55늘

CHARLES PALACHE AND HARRY BERMAN

Fe	orms	ϕ	ρ	ϕ_2	$\rho_2 = B$	С	A
θ	$\overline{2}41$	-2245	56 $17\frac{1}{2}$	$120\ 06\frac{1}{2}$	39 54	60 56	108 46
ρ	$\overline{2}51$	-18 33	61 15	$120\ 06\frac{1}{2}$	33 47	$65\ 05\frac{1}{2}$	$106\ 11\frac{1}{2}$
S	311	75 32 ¹ / ₂	54 09	$36\ 43\frac{1}{2}$	$78 \ 49\frac{1}{2}$	43 53	$38\ 17\frac{1}{2}$
T	321	62 43	56 27	$36\ 43\frac{1}{2}$	$67 \ 32\frac{1}{2}$	47 09	$42\ 12\frac{1}{2}$
U	341	$44\ 06\frac{1}{2}$	62 33	$36\ 43\frac{1}{2}$	50 25	55 26 1	51 51
V	361	$32 \ 52\frac{1}{2}$	67 57	$36\ 43\frac{1}{2}$	38 53	62 29	$59\ 47\frac{1}{2}$
σ	311	$-70\ 16\frac{1}{2}$	$45 \ 40\frac{1}{2}$	$133 \ 56\frac{1}{2}$	$76\ 01\frac{1}{2}$	55 48	132 20
au	321	-5421	49 52	$133 \ 56\frac{1}{2}$	$63\ 32\frac{1}{2}$	58 46	$128 \ 24\frac{1}{2}$
ω	351	-2909	$63\ 11\frac{1}{2}$	$133 \ 56\frac{1}{2}$	$38\ 47\frac{1}{2}$	$68\ 43\frac{1}{2}$	115 46
W	411	78 40	$60\ 22\frac{1}{2}$	30 07	80 10	49 57	31 32
X	421	$68 \ 09\frac{1}{2}$	$61 \ 42\frac{1}{2}$	30 07	$70\ 52\frac{1}{2}$	51 54	35 11
г	411	-75 37	54 17불	$143 \ 25\frac{1}{2}$	68 22	$64 \ 39\frac{1}{2}$	141 52
Δ	421	-62 51	56 34	$143\ 25\frac{1}{2}$	67 37	66 10	137 57
\boldsymbol{Y}	531	$63 \ 48\frac{1}{2}$	$66\ 56\frac{1}{2}$	25 22 ¹ / ₂	$66\ 02\frac{1}{2}$	57 29	$34\ 20\frac{1}{2}$
Ζ	541	56 45	68 22	$25\ 22\frac{1}{2}$	$59\ 21\frac{1}{2}$	59 35 ¹ / ₂	38 59
Θ	631	$-63\ 53\frac{1}{2}$	67 00	154 42	$66\ 06\frac{1}{2}$	$76\ 38\frac{1}{2}$	145 45
Ω	$\overline{6}41$	$-56\ 50\frac{1}{2}$	68 25	154 42	59 $25\frac{1}{2}$	77 26	141 07
Λ	651	$-5045\frac{1}{2}$	69 531	154 42	53 331	78 16	136 391

TABLE 3—Continued

* Forms also found by Sjögren.

** Forms found only by Sjögren.

Since the crystals are all so minute that the faces are practically invisible even under the binoculars, it was not deemed advisable to attempt to figure them, as drawings would only be conventional representations at best. Crystal eight from Montana with faces of 33 forms is by far the most complex one seen.

For the reason given in the last paragraph no attempt has been made here to characterize the individual forms and their relative importance must be judged solely by frequency of occurrence.

Table 2a contains a summary of the angular measurements obtained from the eight Montana crystals and Table 2b similar figures for those from Idaho. The Washington crystals were of far inferior quality and are not recorded. If there appears to be a somewhat wide range in the observed angles of individual forms, it must be remembered that all of these measurements were obtained from crystal facets so minute that often the signal was but the faintest spot of light. It was impossible to distinguish the faces on most of the crystals sufficiently clearly to recognize their symmetry relations. Only after projection could the orientation of the crystal be determined, and it was a matter of astonishment to find

how closely these weak reflections repeated on each new projection the pattern of the common gnomonic lattice. The Montana crystals gave, on the whole, the better reflections, but both sets of measurements were used to calculate elements with the following concordant results:

Boulangerite, Montana	$a:b:c=0.9158:1:0.3456; \beta=100^{\circ}39\frac{1}{2}'$
Boulangerite, Idaho	$a:b:c=0.9252:1:0.3437; \beta=100^{\circ}52'$

The first set of elements, based on measurements of 90 faces of 17 forms on 8 crystals, was accepted as best established and was used in the calculation of an angle table, Table 3.

Reference has been made earlier to the paper in which Sjögren first described crystals of boulangerite. His results have been correlated with our established position by the following transformation formula and they show excellent correspondence.

Transformation, Sjögren to Palache041/800/002Elements of Sjögren, calculated to new position: a:b:c=0.9315:1:0.3383; $\beta=100^{\circ}27\frac{1}{2}'$

Four of his prisms were found on our crystals. Five others, probably very weak forms, were not found by us but are included in the angle table as well as his single terminal form, which becomes the dome {102}.

The measurements of Shannon (1921) could not be satisfactorily correlated with our position.

X-RAY CRYSTALLOGRAPHY

Dr. Hurlbut reports as follows on his study of x-ray photographs made in 1939 on crystals of boulangerite from Montana. From measurements of a rotation photograph: $d_{001}=8.00$ Å. From Weissenberg zero and second-layer photographs: $d_{010}=23.16$ Å and $d_{100}=21.10$ Å. $\mu=79^{\circ}19'$ calculated from x-ray data. Hence, $a_0=21.47$, $b_0=23.16$, $c_0=8.00$.

 $\begin{array}{c} a_0:b_0:c_0=0.9166:1:0.3475; \ \beta=100^\circ41' \ \text{compared with} \\ a:b:c=0.9158:1:0.3456; \ \beta=100^\circ39\frac{1}{2}' \\ \text{Space group } P2_1/a \ \text{fixed by the conditions} \\ hkl \ \text{all present} \\ 0k0 \ \text{present only with } k \ \text{even} \\ h00 \ \text{present only with } h \ \text{even} \\ h0l \ \text{present only with } h \ \text{even} \end{array}$

This cell taken with the measured specific gravity, 5.98, gives $M_0 = 14468$.

These results agree very closely with the measurements published by Berry (1940), which were made on boulangerite from Mullan Co., Idaho, the so-called mullanite for which the crystal measurements are given above.

CHEMICAL COMPOSITION

Boulangerite from Rocky Gulch, Montana, was analyzed by Gonyer on a carefully selected sample prepared by Dr. Berman. The density was determined on the microbalance as 5.98 ± 0.02 , the mean of many measurements by several observers. Gonyer also analyzed the Stevens Co., Washington, boulangerite and confirmed Shannon's analysis. This second analysis is interpreted in the following table prepared by Dr. Berman.

TABLE 4. ANALYSIS OF BOULANGERITE FROM WASHINGTON

	1	2	3	4	5
Pb	55.42	55.28	55.91	39.03	40
Sb	25.69	25.40	25.69	30.53	32
S	18.89	18.19	18.40	83.03	88
Fe		.39			
Insol.		.62			
Total	100.00	99.88	100.00		

1. Pb₅Sb₄S₁₁.

2. Boulangerite from Cleveland Mine, Stevens Co., Washington, U.S.N.M. 95414. *Pb* and *S* average of two determinations. F. A. Gonyer, analyst.

3. Column 2 recalculated to 100%.

4. Atoms per unit cell with $M_0 = 14468$ and d = 5.98.

5. 8(Pb₅Sb₄S₁₁).

The experience of Mr. Gonyer in making this analysis proved that the complete separation of lead and antimony was not attained by some well-tried methods, leading to low results for lead. This observation suggests that possibly the numerous earlier analyses of boulangerite with low lead content may have been at fault through this determinative error. For instance, the so-called plumosite of English writers with a formula given as $Pb_2Sb_2S_5$ has physical characteristics differing in no way from boulangerite

REFERENCES

BERRY, L. G., Univ. of Toronto Studies, Geol. Series No. 44, 5 (1940).

PALACHE, C., Am. Mineral., 24, 463 (1939).

SHANNON, L. V., Am. Jour. Sci., 45, 66 (1918).

-----, Am. Jour. Sci., 1, 423 (1921).

——, Jour. Wash. Acad. Sci., 15, 195 (1925). Sjögren, H., Geol. För. Förh., 19, 153 (1897).