

ferroselite sample from the Virgin mine. Thanks are also due to John Moscal of the Phillips Electronic Instrument Co. and C. M. Taylor of the Materials Analysis Co., who performed the electron microprobe analyses, and to Professor E. N. Cameron, who kindly lent the germanium standard.

## REFERENCES

- BUR'YANOVA, E. Z., AND KOMKOV, A. I. (1955) *Dokl. Akad. Nauk. SSSR*, **105**, 812-813.  
COLEMAN, R. G. (1959) *Geochim. Cosmochim. Acta*, **16**, 296-301.  
GRANGER, H. C. (1963) *N. Mex. Bur. Mines Miner. Res. Mem.*, **15**, 21-37.  
——— (1966) *U.S. Geol. Surv. Prof. Pap.*, **550-C**, C133-C137.  
VLASOV, K. A. (ed.) (1966) *Geochemistry and mineralogy of rare elements and genetic types of their deposits, Vol. 2, Mineralogy of rare elements. Acad. Sci. USSR*, [Trans. Israel Program for Scientific Translations], p. 812-813.

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## NEW DATA ON SARMIENTITE

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Sarmientite,  $\text{Fe}_2(\text{AsO}_4)(\text{SO}_4)\text{OH} \cdot 5\text{H}_2\text{O}$ , was discovered by V. Angelelli in the locality of La Alcaparrosa, Department Barreal, San Juan province, Argentina, and described by Angelelli and Gordon (1941) who performed the optical and morphological study.

While studying some arsenates and sulfates, a specimen of sarmientite from the original locality was investigated at our laboratory; the present note records the results.

On optical examination the sample appeared to be homogeneous and free of impurities formed by very small crystals up to 20  $\mu\text{m}$  long.

Before the wet analysis was undertaken, a spectrochemical analysis revealed, in addition to the major elements: Ca, Cu (0.1%); Al, Mn (0.01-0.1%) and Si, Mg (0.001%).

The quantitative chemical analysis of the specimen now studied agrees with that given by Angelelli and Gordon (Table 1).

In the electron microscope, crystals showing a habit similar to that described by Angelelli and Gordon are seen (Fig. 1). After long exposure to the electron beam, the crystals alter visibly and crack.

## X-RAY DIFFRACTION DATA

After a very long search, three crystals permitting single crystal work were found. Rotation, Weissenberg and precession methods were employed to determine the dimensions of the unit cell. These values, slightly modified after data derived from X-ray powder patterns, are

TABLE 1. CHEMICAL ANALYSES OF SARMIENTITE  
 Wt. %

	Analysis in Angelelli and Gordon	Present study
As <sub>2</sub> O <sub>5</sub>	22.68	23.00
SO <sub>3</sub>	18.28	19.10
Fe <sub>2</sub> O <sub>3</sub>	36.57	33.63
FeO	n.d.	1.43
CaO	0.27	n.d.
H <sub>2</sub> O	22.86	22.60
Total	100.66	99.76

$a=6.55$ ,  $b=18.55$ ,  $c=9.70$  Å,  $\beta=97^{\circ}39'$ , monoclinic  $P2_1/c$ ,  $a:b:c=0.353:1:0.529$ , cell volume  $1168.08$  Å<sup>3</sup>,  $Z=4$ . The Weissenberg photographs were taken with  $c$  the rotation axis; precession photographs were obtained with  $c$  and  $a$  the precessing axes. Exposures times were 200–300 hours (Cu/Ni).

The spacings and intensities given in Table 2 were derived from diffractometer charts. The calculated  $d$  values were obtained and indexed by means of a Bull G.E. 625 computer.

Extinctions in the single crystal photographs lead to space group

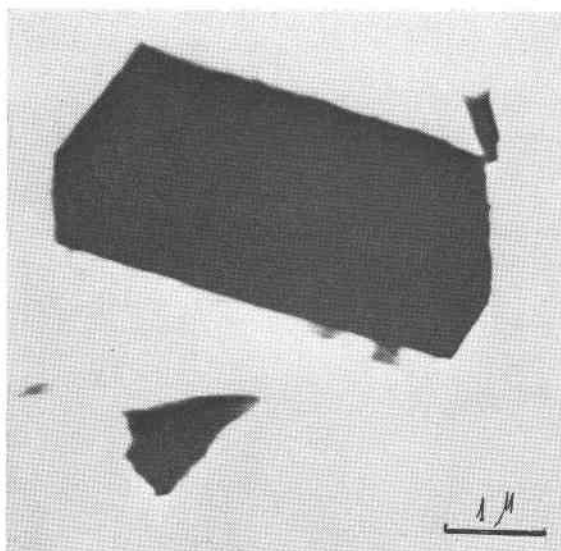


FIG. 1. Electron micrograph of a sarmientite crystal.

TABLE 2  
 X-RAY POWDER DATA FOR SARMIENTITE

<i>hkl</i>	<i>d</i> (obs) <sup>a</sup> Å	<i>d</i> (calc) <sup>b</sup> Å	<i>I</i>	<i>hkl</i>	<i>d</i> (obs) <sup>a</sup> Å	<i>d</i> (calc) <sup>b</sup> Å	<i>I</i>
020	9.29	9.275	100	052	2.93	2.937	7
011	8.53	8.535	22	211		2.921	
021	6.68	6.675	9	202		2.873	
110	6.13	6.127	35	230	2.87	2.874	16
120	5.31	5.318	15	142		2.863	
031	5.21	5.201	27	212		2.839	
12 $\bar{1}$	4.87	4.885	38	033	2.84	2.845	10
040	4.64	4.637	90	23 $\bar{1}$		2.847	
012		4.653		160	2.792	2.791	4 <sup>b</sup>
130	4.47	4.477	18	22 $\bar{2}$	2.734	2.745	10
121		4.453	80	13 $\bar{3}$	2.726	2.727	18
022	4.26	4.268	80	16 $\bar{1}$		2.723	
11 $\bar{2}$	4.02	4.036	32	231	2.668	2.669	18
140	3.76	3.774	9	240		2.659	
12 $\bar{2}$		3.777		062	2.600	48	
11 $\bar{2}$	3.56	3.570	10	152	2.60		2.598
141	3.43	3.424	40	23 $\bar{2}$	2.526	2.606	6
13 $\bar{2}$		3.437		202		2.538	
042	3.33	3.337	35	212	2.492	2.514	7
200	3.24	3.246	9	241		2.494	
150	3.22	3.221	8	133	2.436	2.500	9
20 $\bar{1}$		3.207		24 $\bar{2}$		2.442	
132	3.13	3.136	20	250	2.443		
15 $\bar{1}$		3.117		213	2.428		
220	3.06	3.064	70	25 $\bar{1}$	2.426		

<sup>a</sup> Values obtained using  $\text{CoK}\alpha = 1.7889 \text{ \AA}$ , Fe filter. Diffractometer.

<sup>b</sup> The calculated *d* values were obtained by means of a Bull G.E. 625 computer.

$P2_1/c$ . The morphology (Angelelli and Gordon, 1941) suggests that the crystals belong to class  $2/m$ .

The calculated density is 2.58 for a unit cell containing four formula units of  $\text{Fe}_2(\text{AsO}_4)(\text{SO}_4)\text{OH}\cdot 5\text{H}_2\text{O}$ . The density measured by Angelelli and Gordon is 2.58.

#### INFRARED SPECTRUM

The infrared absorption spectrum shown in Figure 2 was obtained in a Perkin Elmer model 337 grating spectrophotometer. A sample of 1.5 mg was mixed and ground with 210 mg of KBr.

Hydroxyl and water bands appear in the regions near 3.0 and 6.1  $\mu\text{m}$  respectively.

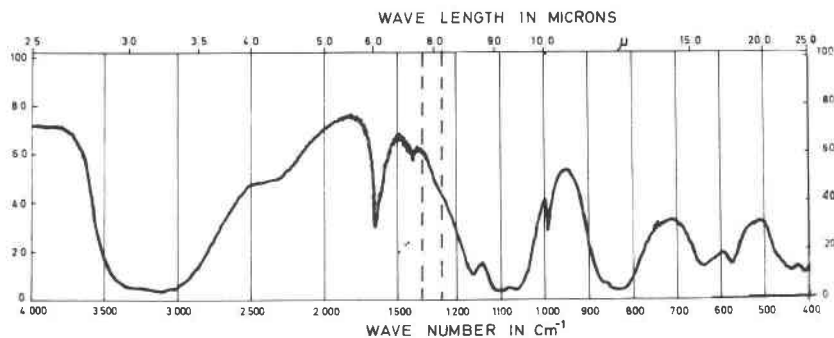


FIG. 2. Infrared absorption spectrum of sarmientite.

The following bands correspond to the vibrations of the sulfate ion, according to Adler and Kerr (1965); three intense bands in the 8- to 10- $\mu\text{m}$  region which correspond to the  $\tau_3$  vibrations, a weak band slightly above 10  $\mu\text{m}$  which represents the  $\nu_1$  vibration; two bands in the 15- to 18- $\mu\text{m}$  region are attributed to the  $\nu_4$  mode, and a band at 22  $\mu\text{m}$ , to the  $\nu_2$  mode. The bands in the 11- to 13- $\mu$  region are attributed to vibrations of the arsenate ions.

#### THERMAL PROPERTIES

The weight loss of sarmientite on heating was determined by means of a thermogravimetric balance. A 204 mg sample, was heated from room temperature to 980°C, at 5°C/min. The resulting curve is shown in Figure 3. The loss in weight between 200 and 685°C is 22.95 percent;

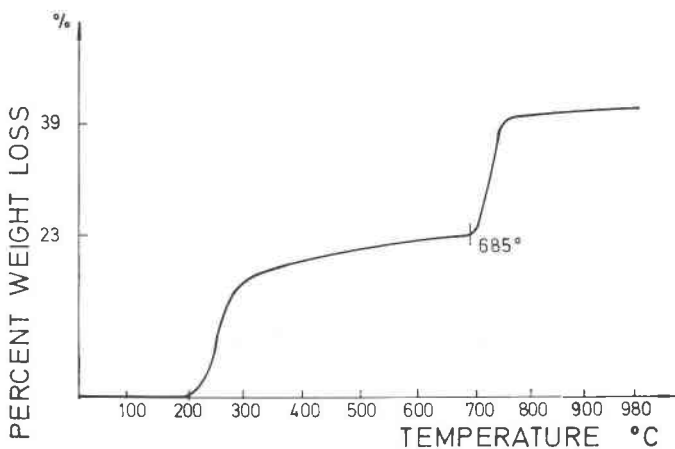


FIG. 3. Thermogravimetric curve of sarmientite.

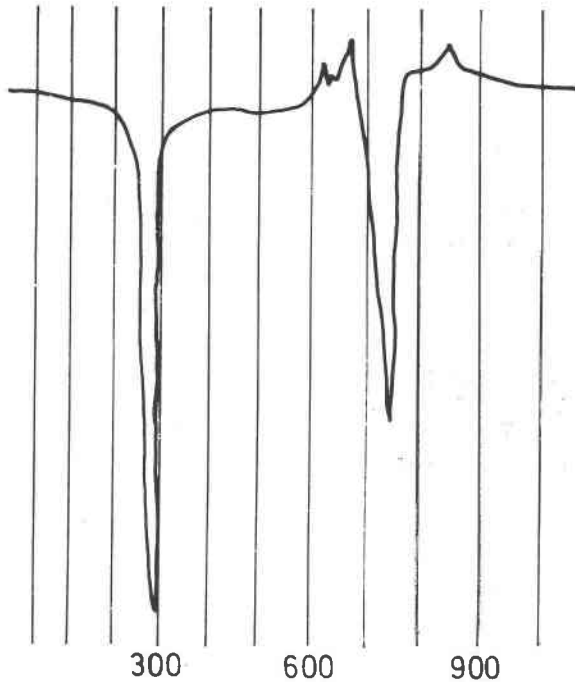


FIG. 4. Differential thermal analysis curve of sarmientite, temperature in °C.

most of this loss occurs between 200 and 300°C, and represents the water content; it agrees closely with the water determination in the chemical analyses (Table 1). A loss of weight of 16.4 percent occurred between 685 and 980°C; it agrees reasonably well with the  $\text{SO}_3$  determination in the chemical analyses.

A differential thermal analysis of sarmientite yielded the curve of Figure 4 in which are shown major endothermic peaks at 200–300°C (loss of water) and at 700–800°C (loss of  $\text{SO}_3$ ) and small exothermic peaks at 600–700°C and 850°C which probably represent recrystallization processes.

When heated at 300°C for one hour sarmientite yields a buff colored product, amorphous to X-rays. A sample held at 650°C for two hours lost all the  $\text{SO}_3$ ; a reddish-brown, crystalline material, remained. The spacings in Å and the relative intensities of the strongest lines, as obtained from a diffractometer chart are: 3.04 (100), 3.13 (50), 3.26 (50), 2.99 (30), 2.49 (30), 3.68 (30), 3.53 (30). The line at 3.68 Å belongs to hematite; the spacings of the remaining lines correspond to those of strong lines in the angelellite pattern (Weber, 1959).

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## REFERENCES

- ADLER, H. H. AND P. F. KERR (1965) Variations in infrared spectra, molecular symmetry and site symmetry of sulfate minerals. *Amer. Mineral.*, **50**, 132-147.
- ANGELELLI, V. AND S. G. GORDON (1941) Sarmientite, a new mineral from Argentine. *Acad. Sci. Phil., Not. Nat.* **92**.
- WEBER, K. (1959) Eine Kristallographische Untersuchung des Angelellits. *Neues Jahrb. Mineral., Monatsh.*, **7**, 152-58.

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## JOHACHIDOLITE, A REVISED CHEMICAL FORMULA

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Iwase and Saito (1942) described a new mineral which was discovered in nepheline veins in limestone at Johachido, Kankyohokudo, Korea, which they named after the locality. The impossible formula given,  $H_6Na_2Ca_3Al_4F_3B_6O_{20}$ , has appeared in several books dealing with mineralogy or the chemistry of boron compounds. Only the analysis appears in Dana (1951). A balanced formula,  $Na_2Ca_3Al_4B_6O_{14}(OH)_3F_3$ , is given in Hey (1962), and Strunz (1957) offers the formula  $Ca_3Na_2Al_4H_4[(F, OH)BO_3]_6$ , but neither represent the analysis well. A new calculation seemed necessary.

Since impurities due to nepheline,  $NaAlSiO_4$ , and apatite,  $Ca_5(PO_4)_3F$ , probably make some contribution<sup>1</sup>, column 2 presents the amended analysis. The atomic proportions were calculated on a basis of total anions=45 (column 3), this being the smallest number whereby an approximately integral number of cations in each group could be ac-

<sup>1</sup> Private communication from M. H. Hey.