Blue amphibole from Precambrian metabasalts, Savage River, Tasmania

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Abstract

A blue sodic amphibole occurs in mafic greenschists associated with a Precambrian iron deposit at Savage River, Tasmania. Its composition falls between magnesioriebeckite and actinolite. Crystallization of the amphibole probably reflects P, T conditions intermediate between glaucophane-lawsonite schist and greenschist facies, and demonstrates partial closing of the miscibility gap between sodic and calcic amphiboles at these conditions.

A blue sodic amphibole occurs in Precambrian mafic schists and metabasalts in the Savage River area of northwest Tasmania. The rocks are interlayered with banded magnetite-pyrite-silicate deposits and are situated in a multiply-deformed metamorphic belt (Spiller 1974). The overall metamorphic grade is greenschist facies, but the presence of blue sodic amphibole suggests transition to the higher-pressure 'blueschist' facies (Brown 1974, Katagas, 1974). This note presents further data on the occurrence and composition of blue amphibole intermediate between actinolite and riebeckite.

The mineral assemblage found with the blue amphibole consists of:

1. actinolite, albite, epidote, chlorite, titanite, magnetite, pyrite, and quartz in metabasalt 7281;
2. actinolite, albite, epidote, green mica, chlorite, titanite, magnetite, pyrite, quartz, and tourmaline in metabasalt 7463;
3. actinolite, albite, talc, green mica, calcite, quartz, magnetite, titanite, pyrite, and tourmaline in mafic schist 7475.

Textural relations are variable and the rocks range from actinolite-dominated, strongly-foliated mafic schists to more equigranular metabasalts. The minerals listed do not constitute equilibrium assemblages, and blue amphibole grains are rimmed by actinolite in all rocks. In one metabasalt (7281) the blue amphibole occurs as bladed crystals reaching 0.5 cm in size, rimmed by epidote, actinolite, and albite.

The optical properties for blue amphibole examined in this work are as follows: \( \alpha = \) colorless to pale brown; \( \beta = \) blue; \( \gamma = \) lavender blue; \( 2V = 0-40^\circ \), optically negative, birefringence low to moderate. These properties are similar to those of glaucophane-crossite, and Spry (1964) in describing similar rocks from the Pieman River area, west Tasmania, identified the mineral as glaucophane.

Analyses of the blue amphibole indicate the most sodic compositions are magnesioriebeckite with a significant actinolite component (>18% molecular proportion) in solid solution (Table I). Complete solid solution between the magnesioriebeckite and actinolite end-members is not demonstrated (Fig. 1) for the limited data presented here, but the analyses do indicate compositions plotting in the proposed compositional gap (cf. Katagas 1974). The separation of the compositions into two distinct fields in Fig. 1 points to a compositional gap smaller than that proposed by either Coleman and Papike (1968) or by Iwasaki (1963), and it is likely that the extent of this gap is sensitive to temperature \( T \) (Klein, 1969) or to \( T \) and pressure \( P \) (Brown, 1974).

The composition of the Savage River blue amphibole is similar to blue sodic amphibole with significant actinolite (>18% molecular proportions) in solid solution described by Katagas (1974) and Brown (1974) from metamorphic terrains containing greenschist facies assemblages. This composition contrasts with the typical blue sodic amphibole with low actinolite (<18% molecular proportions) in solid solution occurring in mafic rocks of the glaucophane-lawsonite schist facies (Black, 1973; Boucquet, 1974; Coleman and Papike, 1968; and Himmelberg and Papike, 1969).

This work shows that analysis of blue amphibole is essential to distinguish between optically similar am-
Fig. 1. Compositions of blue amphiboles from Savage River, Tasmania, compared with compositions of other amphiboles intermediate between tremolite/actinolite and magnesio-riebeckite/riebeckite. Dashed lines denote miscibility gap between calcic and sodic amphiboles proposed by Coleman and Papike (1968) while dotted lines refer to the miscibility gap proposed by Iwasaki (1963). The lower group of solid symbols correspond to amphiboles with pleochroic schemes similar to that listed in the text, while the upper group have pleochroic schemes with a greenish-blue character.

Fig. 2. Interpretation (modified from Brown, 1974) of miscibility between actinolite and Na-amphibole as indicated by data from the Savage River amphiboles, compared with Otago, Shuksan, and Franciscan examples.

The identification of blue amphibole alone is insufficient to infer the high-pressure, low-temperature conditions of the glaucophane-lawsonite-schist facies (Turner, 1968). However, the narrowing of the miscibility gap between sodic and calcic amphiboles may provide some indication of the $P, T$ conditions of formation of blue amphiboles of this type, and Brown (1974) suggests that the coexistence of blue sodic amphibole with albite and iron oxide (magnetite) allows the maximum content of the sodic end-member in the sodic–calcic amphibole solid solution series for a given $T/P$ ratio. Brown (1974) and Katagas (1974) predict that with increasing $T/P$ the apparent miscibility gap between the sodic and calcic amphiboles is reduced or closed. The appropriate $T/P$ condition proposed is intermediate to the glaucophane–lawsonite–schist facies and the greenschist facies, within the range of 5–10 kilobars and 300–400°C. The oxygen fugacity indicated by the blue amphibole-bearing assemblages is at least higher than the quartz–mag-
Thus it is inferred that the magnetite–albite–magnesiorebeckite-bearing rocks of the Savage River area reflect crystallization at \( P, T \) conditions intermediate between the greenschist and glaucophane-lawsonite–schist facies, with subsequent adjustment of the mineral assemblage, corresponding to an increase in \( T \) or fall in \( P \), or both. The end result is a typical greenschist facies mineralogy with relict cores of blue amphibole.

**Acknowledgments**

Savage River Mines provided assistance in the field and Macquarie University research funds supported the analytical work. The Research School of Earth Sciences, Australian National University, allowed use of the TPD probe. Drs. R. H. Flood, S. E. Shaw and R. H. Vernon critically read the manuscript and offered helpful comments.

**References**


Manuscript received, January 19, 1976; accepted for publication, May 5, 1976.