# Klebelsbergite, Sb<sub>4</sub>O<sub>4</sub>(OH)<sub>2</sub>SO<sub>4</sub>: redefinition and synthesis

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#### Abstract

Klebelsbergite is verified as a valid mineral species, in museum specimens from the type locality, Baia Sprie (Felsöbánya), Rumania.

Klebelsbergite occurs as pale yellow to orange-yellow crystals radially grown on stibnite. Wet-chemical analysis of synthetic klebelsbergite, supported by electron microprobe analyses of natural material, gives Sb 71.78, S 4.65, H<sub>2</sub>O 2.69, O 20.88% (by difference). The empirical cell contents are Sb<sub>4.07</sub>O<sub>5.00</sub>(SO<sub>4</sub>)  $\cdot$  1.03H<sub>2</sub>O with an ideal formula Sb<sub>4</sub>O<sub>4</sub>(OH)<sub>2</sub>SO<sub>4</sub>. The crystals are tabular or acicular elongated [001], flattened {010}. The mineral is orthorhombic, *Pcmb* or *Pc*2<sub>1</sub>*b*, *a* = 11.279(2), *b* = 14.909(3), *c* = 5.7648(6)A; *Z* = 4. D(meas) = 4.62(6), D(calc) = 4.69g/cm<sup>3</sup>. The strongest lines of the X-ray powder pattern are 6.22(s)(120), 3.892(s)(211), 3.545(s)(221), 3.150(s)(301), 3.131(vs)(231), 2.830(s)(012), 2.435(s)(132), 1.805(s)(213). Klebelsbergite has vitreous luster, yellow streak, and no distinct cleavage. The mineral is optically biaxial with large 2*V<sub>x</sub>* about 70°,  $\beta \approx 1.95$ , and is not pleochroic. VHN = 200 (10g load).

Single crystals of klebelsbergite were synthesized by boiling 2 g of  $Sb_2O_3$  with 600 ml of 1 N H<sub>2</sub>SO<sub>4</sub> for four days. Synthetic crystals are tabular, elongated [001], flattened {010}. The space group, cell parameters, and powder-diffraction data are essentially identical to those of natural klebelsbergite. The crystals decompose by dehydration at 273°C.

## Introduction

Klebelsbergite was first described by Zsivny in 1929 from Felsöbánya, Hungary (now Baia Sprie, Rumania). He reported that the mineral was essentially a basic antimony sulfate, containing a little  $H_2O$ , Fe, Mg, Na, K and traces of Bi and  $P_2O_5$  by microchemical tests. He also showed that it was monoclinic with a  $\beta$  angle of 91.8°, biaxial negative, X = b,  $Y:c = 1.8^{\circ}$ , n > 1.74. In the following 50 years, several reports concerning this mineral have appeared (Manilici et al., 1965; Radulescu and Dimitrescu, 1966; Antonovici and Vanghelie, 1972); however, none of them gave a complete description of this rare mineral. Antonovici and Vanghelie (1972), in the only detailed study of this mineral since Zsivny, examined museum specimens of klebelsbergite. They reported that the mineral contained 61.86 percent antimony by X-ray fluorescence analysis and more than one percent sulfur by spectrographic analysis.

They also presented an X-ray powder diffraction chart with the three strongest lines given as 6.19(100), 3.43(80), and 2.47A(90).

Our work gives for the first time the chemical formula, crystal data, and physical and chemical properties of klebelsbergite. This study has been done using museum specimens from the Smithsonian Institution labelled as klebelsbergite from the type locality, Baia Sprie, Rumania; NMNH #137949, R6535, B12631, B12640, B12647, and B16076.

A short version of this manuscript was submitted to the IMA Commission on New Minerals and Mineral Names. It was determined that a vote of the commission was not necessary (Max H. Hey, written communication).

#### **Physical properties**

Klebelsbergite occurs as aggregates of very small platy or acicular crystals (Fig. 1), maximum size about  $0.8 \times 0.08 \times 0.02$  mm, with a pale yellow to orange-yellow color. The streak is white to yellow and the luster is vitreous. The mineral is brittle and has no distinct cleavage. The density determined by Ber-

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Fig. 1. SEM photograph of natural klebelsbergite from Baia Sprie, Rumania (225×).

man balance is 4.62(6) g/cm<sup>3</sup> (average of 6 measurements using more than 10mg of sample each time). The Vickers hardness number (VHN) is 200 with a 10g load.

Klebelsbergite is optically biaxial, negative, with 2V about 70°. Because of its high reactivity with index-of-refraction liquids of high index, the refractive indices could not be measured; however, quick observation before the crystals dissolved showed that  $\beta$  is near 1.95. Birefringence  $\beta - \gamma$  is about 0.035 (by retardation). The mean index of refraction, 1.95, was calculated from the Gladstone-Dale relationship using the ideal chemical formula, Sb<sub>4</sub>O<sub>6</sub>H<sub>2</sub>SO<sub>4</sub>, calculated density of 4.64 g/cm<sup>3</sup>, and the constants given by Mandarino (1976), and this supports the above observation that  $\beta$  is about 1.95. The optical orientation is X = b, Y = c and Z = a. The mineral is not pleochroic.

Klebelsbergite is weakly magnetic when it contains iron as an impurity. This feature enabled us to separate klebelsbergite from other minerals with the Franz isodynamic separator. The mineral is not luminescent in either long- or short-wave ultraviolet radiation.

# Crystallography

In the original description by Zsivny (1929), klebelsbergite was described as monoclinic; however, the true symmetry has been revealed to be orthorhombic by our X-ray diffraction studies of single crystals. Weissenberg and precession photographs show that the space group is either *Pcmb* or *Pc2*<sub>1</sub>*b*. The lattice constants are a = 11.279(2), b = 14.909(3), and c = 5.7648(6)A, calculated from the powder-diffraction data using the least-squares refinement program of Appleman and Evans (1973). The powder-diffraction pattern of klebelsbergite was obtained with a Guinier-Hägg camera using CuK $\alpha_1$  radiation and BaF<sub>2</sub> internal standard; the data are given in Table 1. Indexing of ambiguous reflections was aided by intensity measurements from single-crystal photographs.

Crystals of klebelsbergite are generally flattened to tabular on {010} and elongated in the [001] direction. Well-developed crystals were occasionally observed; however, the small crystal size and especially the thickness of less than 20  $\mu$ m prevented us from carrying out detailed measurements of the interfacial angles. Under the optical microscope, the apical angle, indicated by an arrow in Figure 2, was measured to be about 124°. This value is found to be close to the supplemental angle of (101)/( $\overline{101}$ ) = 54.14° calculated from the lattice constants. A section normal to the *c* axis shows a flattened hexagon; therefore the platy crystal can be described as having a major {010} pinacoid and a minor probably {120} prism.

As is seen in Figure 2, klebelsbergite crystals tend to grow as parallel blades elongated along the c axis, stacked in the *b*-axis direction in lamellar fashion. Precession photographs parallel to the (010) plane revealed very slight distortion among the stacked crystals; *i.e.* the (010) plane rotates slightly around the *b* axis. As a result, the crystals show a monoclinic appearance, and this is considered to be the reason klebelsbergite was described as monoclinic in the original description (Zsivny, 1929). Zsivny reported a beta angle of 91.8° with optical orientation Y:c = 1.8°, and this is well explained as the distortion around the *b* axis.

### **Synthesis**

Single crystals of klebelsbergite (maximum size about  $1 \times 0.1 \times 0.03$  mm) have been synthesized by boiling 2 g of Sb<sub>2</sub>O<sub>3</sub> with 600 ml of 1 N H<sub>2</sub>SO<sub>4</sub> solution for four days.

The crystals obtained are colorless and transparent, having the same luster as natural crystals. The crystal habit is also similar to that of natural crystals: tabular on  $\{010\}$ , elongated along [001], with the dominant form  $\{010\}$  (Fig. 3). The crystals often developed radially with a parallel growth along [001] as in natural crystals. The VHN is 130 with 10 g

# Table 1. X-ray powder-diffraction data for klebelsbergite

Table 1. (continued)

Natural Klebelsbergite				Synthetic klebelsbergite				
hie1	d calc	d obs	I*	hkl	d calc	d obs	I	
020	7.45	7.47	w	020	7.44	7.45	w	
20	6.22	6.22	S	120	6.21	6.21	s	
.11	4.85	4.85	ms	111	4.856	5.L30 4.854	ms	
21	4.228	4.231	mw	121	4.228	4.228	m	
11	3.892	3.892	s	211	3,893	3.892	vs	
40	3.727	3.731	vw	040	3.722	3.720	w	
	0.544			131	3.569	3.568	W	
20	3.340	3.343	S	140	3,546	3,545	S	
01	2.140	3,150		201	3.355	2.1/0	ш	
31	3.149	3.131	S	231	3.130	3.140	S	
40	3.110	3.110	vw	240	3.106	3.105	W	
11	3,081	3.083	ms	311	3.081	3.081	s	
21	2 901	2.902	шw	321	2.901	2.901	m	
002	2.882	2.882	w	002	2.885	2.886	w	
00	2.830	2.830	s	012	2.833	2.833	s	
41	2.737	2.736	m	241	2.735	2.735	m	
122	2.688	2.688	៣៩	022	2,690	2.691	ms	
20	2.637	2,639	W	420	2.636	2.637	mw	
22	2.615	2.618	VW	122	2_617	2.617	vw	
				151	2.576	2.577	vvi	
12	2.529	2.531	vw	212	2.531	2,531	w	
11	2 :4 37	2.497	tu	411	2.497	2.490	ms	
22	2.485	2.485	W	060	2.482	2.481	m	
32	4.993	2.435	5	222	2 438	2.428	S	
				421	2.398	2.396	m	
51	2.397	2.398	mw	251	2.396			
04.2	2.280	2.280	mw	042	2.280	2,281	m	
40	2.249	22.248	w	440	2.247	2.247	mw	
.61	2 237	2.239	V.W	161	2.235	2.235	w	
				351	2.100	2.164	W	
20	2.159	2.159		520	2.158	2.158	w	
				242	2 114	2 115	1757	
41	2.095	2,094	mw	441	2.094	2.094	ms	
132	2.078	2.078	w	332	2,078	2.079	m	
50	0.000	0.020		052	2.072	2.072	VVI	
.52	2.038	2.038	AAA	152	2.038	2.038	VW	
02	2.016	2.015	m	402	2.016	2.017	ms	
+12	1.997	1,998	m	412	1.998	1.998	ms	
361	1.951	1,950	2002	361	1.965	1,965	w ms	
				422	1.946	1.946	s	
252	1.945	1.945	ms	252	1 945			
51	1.930	1.931	m	451	1,929	1.928	5	
0.2	1 004	1 00/		540	1.929			
271	1.894	1 883	mw	271	1.896	1,896	ms	
162	1.882	1.005	til w	600	1.879	1.001+1	\$	
000	1.880	1.880	w					
13	1.879							
60	1.864	1.861	m	460	863	1.863	Ş	
				102	1.050	1.057	001	
				123	1.838	1.837	VW	
520	1.823	1.823	mW	620	1.822	1.822	ms	
			1999	352	1.815	1.815	vw	
13	1.806	1.805	н	213	1.807	1.807	s	
511	1.775	1.774	m	611	1.7743	1.7735	ms	
42	1.773	1 7/0		442	1.7730	1 3/04		
23	1.767	1.768	VW	223	1.7684	1.7688	W	
512	1.764	1.764		512	1.7644	1.7040	w	
				371	1.7627	1.7627	W	
				551	1.7165	1.7170	vw	
072	1.713	1.711	шw	072	1.7121	1.7124	m	
13	1.700	1.700	mw	313	1.7011	1.7012	m	
43	1.689	1 690	vvw	143	1.6896	1.6895	w	
	1,002	r = 00Z	11	640	1.6777	1.6774	m Vu	
32	1 473	1 673		530	1 4720	1 4707	-	
132	1.0/3	1.0/3	m	452	1.6697	1.6696	DR VV-V	
272	1.639	1.640	vvw	272	1.6383	1.6384	VW	
43	1.635	1.635	vw	243	1.6354	1.6349	W	
				471	1.6289	1.6291	w	
				333	1.6187	1.6193	vv	
5.2		1		542	1.6035	1.6032	VW	
133	1.599	1,598	VVW	153	1.5995	1.5994	W	
100	1.575	1.575	140	602	1.5748	1.5746	шw	
/20			and the second sec		and the second sec			

L1.7	d aala	d obc	T.*	bk1	d calc	d obs	T
hki	d calc	d obs	1.	IKI	d care		-
462	1.565	1.5655	m	462	1.5649	1.5644	ms
082	1.565	1 5520	223	082	1.5539	1 5534	W.F.
923	1.553	1.3329	m	253	1.5533	1.3334	ms
182	1.550	1.549	w	711	1.5432	1,5431	VW
711	1.5435	1.5435	VVW	622	1.5407	1.5403	VW
622	1,5406	1.5401	VVW	651	1.0323	1.3321	new
291	1.5322	1.5318	13	291	1.5306	1.5306	m
433	1.5126	1 5132	W	433	1.5133	1,5132	កា ឃ
103	1.3003	1.5007	V W	481	1.4998	1.4997	vw
				353	1.4844	1.4841	W
1.10.0	1.4781	1.4777	vvw	1.10.0	1.4762	1.4764	w
472	1.4640	1.4630	vw	472	1.4634	1.4633	w
	1 1051	1 / 255	945	642	1.4503	1.4502	W
523	1,43.54	1.4377	w	014	1.4359	1,4550	a.
				741	1.4321	1.4318	vw
192	1 4248	1.4238	TING	114	1.4244	1.4242	ms
114	1.4238		1000	192	1.4237		
024	1.4150	1.4149	w	024	1.4163	1.4163	m
800	1.4099	1_4095	VW	800	1.4095	1.4091	mW
124	1.4093	1,4040	VW	124	1,4053	1.4053	mw
712	1.4002	1.4005	vw	712	1.4003	1.4003	m
3.10.0	1-3859	1,3859	mΨ	2.10.1	1.3968	1.3969	W
	x = 3037			652	1.3921	1.3919	mω
034	1.3842	1.3839	м	034	1.3853	1.3850	m
				820	1.3849	1 2822	101.1
				751	1.3760	1.3755	VW
(7)	1 2401	1 2697	0.00	671	1 3682	1.3681	1.7
671	1.3684	1,3000	VVW.	482	1.3676	1.5001	
811	1.3638	1.3636	inw	811	1.3635	1.3631	ms
				234	1 . 34 53	1.3456	mwi
- and the second s		1 7/01		044	1.3452	1. 36.00	
314	1.3403	1.3401	vw.	392	1 2200	1, 2207	
	1 22/2	1 22/0	1213	613	1 3232	1 3225	w
680	1.3293	1.3240	VW	680	1.3225	115225	
000				840	1.3182	1.3180	vw
761	1.3162	1_3158	VVW	761	1.3156	1.3155	W
334	1.2989	1.2988	VW	334	1.2999	1.2990	mw
0(1	1 2055	1 2054		271	1.2951	1 2951	
841 473	1.2855	1 2733	VVW	752	1.2718	1.2001	w
424	1.2647	1.2646	m	802	1.2665	1.2656	TR.
254	1.2645			672	1.2657		
0.0.0	1 0/07	1 2/93		900	1.2529	1.2531	W
822	1,2486	1.2463	VW	022	1,2400	1.2401	w
851	1 2445	1,2446	W	851	1.2441	1.2438	mw
0.12.0	1.2424	1.2421	VW	0.12.0	1.2428	1.2405	W
				713	1.2308	1.2306	vw
				354	1.2272	1.2274	vw
				832	1.2272		
				0.11.2	1.2255	1.2251	mw
				901	1,2252		
				911	1.2203	1.2205	vw
				264	1 + 2178	1,2179	W
				444	1.2140	1.2139	mw
				592	1.2108	1,2110	100
				021	1 2082	1 2080	
				573	1.2057	1.2057	VW
4.10.2	1,1987	1.1985	w	842	1.1990	1.1991	mw
				772	1.1732	1.1733	vw
				3.12.1	1.1545	1.1547	W
				2.10.3	1.1526	1.1526	VVV
				583	1.1505	1.1508	VW
				604	1.1444	1,1442	W
				464	1.1406	1.1406	w
				374	1.1379	1.1381	vw
				5.11.1	1.1378	52	
				673	1.1364	1.1363	vw
				951	1,1324	1.1345	W WW
				10,0.0	1.1276	1.1277	W
				554	1,1252	1,1253	υw
				4.11.2	1.1239	1.1239	w
				1,13.1	1.1179	1.1181	W
				833	1.1151	1,1096	mw mw
				000	1,1005	1.1000	10W
				235	1.1024	1.1027	VV6
				1-1-2	1.00711	A + U 7 / 4	14
				564	1.0914	1.0918	W

\* Relative intensities estimated visually ( s, strong; m, medium; w, weak; v, very; b, broad. )



Fig. 2. SEM photograph of natural klebelsbergite from Baia Sprie, Rumania (400×).

load, as compared to 200 measured on natural crystals. The powder-diffraction data obtained with a Guinier-Hägg camera using  $CuK\alpha_1$  radiation and CaF<sub>2</sub> standard are given in Table 1. The data agree very well with those for natural klebelsbergite. The crystal symmetry and the space group are the same as the natural mineral, i.e. orthorhombic Pcmb or  $Pc2_1b$ . Cell dimensions refined from the powder data are a = 11.2763(7), b = 14.890(1) and c =5.7707(4)A, in excellent agreement with natural klebelsbergite; the difference is almost within one standard deviation. The measured density of the synthetic material is 4.61(4) g/cm<sup>3</sup> which agrees well with the calculated density 4.69 g/cm<sup>3</sup> and that measured on natural crystals, 4.62(6) g/cm<sup>3</sup>. Thus the synthetic crystals have been identified as klebelsbergite.

## Chemistry

Wave-length scans using an ARL-SEMQ electron microprobe disclosed that klebelsbergite contains major Sb and S and minor Fe. No other elements ( $Z \ge 9$ ) were detected. The presence of sulfate ion was confirmed by a spot test with BaCl<sub>2</sub> solution.

As very little material was available, wet-chemical analyses were carried out only on synthetic klebelsbergite. Sb was determined by KBrO<sub>3</sub> titration, and  $SO_4^{2-}$  gravimetrically with BaCl<sub>2</sub> solution (Hillebrand and Lundell, 1953). Water was determined by the Penfield method and the sample was heated to 300°C. The result is given in Table 2, which yields the empirical formula  $Sb_{4,07}O_{5,00}SO_4 \cdot 1.03H_2O$  (on the basis of  $SO_4 = 1$ ), and ideally  $Sb_4O_6H_2SO_4$ , which can be expressed as either  $Sb_4O_4(OH)_2SO_4$  or  $Sb_4O_5SO_4H_2O$ . On the basis of IR spectroscopy,  $Sb_4O_4(OH)_2SO_4$  is favored (see below); however,  $Sb_4O_5SO_4 \cdot H_2O$  cannot be ruled out completely.<sup>2</sup>

Quantitative analyses of natural klebelsbergite were carried out with the electron microprobe using an operating voltage of 15kV and a sample current of 0.15 microamps. Valentinite, barite, and olivine were used as standards for Sb, S, and Fe, respectively. The results are given in Table 2. Concentration of iron varies from 0.0 to 0.5 percent and is thought to be present substituting for antimony as explained below. The rather large discrepancy of sulfur content between the two analyses, wet-chemical and microprobe, is considered to be due to inadequate data correction of the microprobe data. The ideal chemical formula  $Sb_4O_6H_2SO_4$  with Z = 4 is suitable for the space groups Pcmb or  $Pc2_1b$ . The theoretical density for the ideal formula is  $4.69 \text{ g/cm}^3$ , which agrees well with the measured density 4.62(6) g/cm<sup>3</sup>. Klebelsbergite is readily soluble in HCl(1:1), and NaOH(10 percent), slowly soluble in tartaric acid (20

<sup>&</sup>lt;sup>2</sup> After this paper was accepted for publication, we received a preprint from Drs. S. Menchetti and C. Sabelli of the Università di Firenze describing their recent determination of the klebelsbergite structure (Menchetti and Sabelli, in preparation). They found that the correct formula is indeed  $Sb_4O_4(OH)_2SO_4$ , and that the space group is  $Pc2_1b$ .



Fig. 3. SEM photograph of synthetic klebelsbergite (60×).

percent), and very slowly soluble in  $HNO_3(1:1)$  and  $H_2SO_4(1:1)$ .

The color of klebelsbergite is more or less yellow; however, the depth of the yellow color varies from specimen to specimen. To elucidate the origin of the color, microprobe analyses were made on various specimens with different depths of color. The result indicates that there is a clear positive correlation between the iron concentration and the depth of the color. Iron-free synthetic klebelsbergite is colorless. Therefore, the origin of the yellow color is ascribed to the iron impurity. To establish the valence of iron, spot tests were made using 0.1 N potassium ferri- and ferro-cyanide solutions. HCl(1:3), HNO<sub>3</sub>(1:1) and tartaric acid (20 percent) were used to dissolve the crystals. The tests clearly indicate the presence of both trivalent and divalent iron; however, trivalent iron is much more abundant.

### **Thermal treatments**

## DTA/TGA

DTA/TGA spectra were taken under He atmosphere on a 100 mg sample of synthetic klebelsbergite. Heating rate was 4°C/min, and the temperature ranged from room temperature to 940°C. The spectra are characterized by an endothermic peak at 273°C with a weight loss of 2.4 percent, which is attributed to a dehydration reaction. The value corresponds well with the result of quantitative analyses of water (see Table 2). After the thermal analysis, the specimen was X-rayed and was found to be Sb<sub>2</sub>O<sub>4</sub>.

## **Decomposition**

Decomposition experiments were carried out in an evacuated Pyrex glass tube (15 cm in length; 5 mm ID). A 100 mg sample of synthetic klebelsbergite was placed at the end of a sealed glass tube, then the other end was closed. The tube was placed in a horizontal furnace; the end of the tube containing the sample was located inside the furnace along with a thermocouple, while the other end (about 4 cm) was kept outside of the furnace. On heating to 300°C, a condensation of vapor was observed at the cool end. The specimen was heated for 1 hour, after which the dehydration product and the condensed liquid were tested.

The product kept the original shape; however, it became translucent with a slightly brownish color. The precession photograph of a "single crystal" which had been heated in the same way as above gave diffuse spots which approach a powder pattern,

Table 2. Chemical analysis of klebelsbergite

	1	2	3	4	5	6
Sb	71.78	0.5896	4.066	4	71.51	70.8
S	4.65	0.1450	1	1	4.71	3.2
Fe	2022 C					0.5
H20	2,69	0.1493	1.030	1	2.64	
⊃*̃	20.88	1.305	9.000	9	21.14	
Total.	100.00				100.00	

2. Molecular quotient

Mole number basis S=1 (or SO<sub>4</sub>=1)

Ideal mole number
Theoretical Sb/OcHoS

5. Theoretical  $Sb_4O_6H_2SO_4$ 6. Natural klebelsbergite by microprobe analysis

\* by difference

and this indicates that a single crystal changes into the polycrystalline state by dehydration. The X-ray powder diffraction pattern of the dehydration product has the following 8 strongest lines: 4.55(4), 3.35(1), 3.12(10), 2.80(1), 2.665(6), 1.926(3), 1.886(3), 1.615(3)A; the pattern as a whole shows a similarity to that of senarmontite, Sb<sub>2</sub>O<sub>3</sub>. The product dissolved in HCl solution (1:1) and gave a brown colloidal residue. Successive addition of BaCl<sub>2</sub> solution (10 percent) into the HCl solution caused an abundant white precipitate to form. This observation indicates that the sulfate ion remains in the structure after heating to 300°C. However, there is evidence of a partial decomposition of sulfate ion; the condensed liquid was strongly acidic, indicating that a small part of the sulfate ion decomposed on heating and the resultant  $SO_2$  gas dissolved in the liquid.

#### Infrared spectroscopy

IR spectra were measured by the conventional KBr method for the region from 4000 to 645 cm<sup>-1</sup>. Figure 4 shows the spectra of synthetic klebelsbergite (a), of synthetic klebelsbergite heated at 300°C for 1 hour (b), and of natural klebelsbergite (c). Since spectra of both natural and synthetic klebelsbergite have essentially the same features, we treat them together in the following discussion; the values in the parentheses are those for synthetic klebelsbergite. A sharp absorption of 3435 (3435) cm<sup>-1</sup> is attributed to the O-H stretching vibration, and it disappeared on heating to 300°C for 1 hour: spectrum (b). Judging from the absence of any sizable peak in the 1650 cm<sup>-1</sup> region, this absorption is more likely to be due to the stretching vibration of hydroxyl ion than to that of water of hydration; the O-H bending frequency usually occurs near 1650 cm<sup>-1</sup> in the cases of water of hydration and absorbed water, but not in the case of hydroxyl ion (Gadsden, 1975). Although there is a very weak broad absorption in the 1650



Fig. 4. Infrared spectra: (a) synthetic klebelsbergite; (b) synthetic klebelsbergite heated at 300°C for 1 hr.; (c) natural klebelsbergite from Baia Sprie, Rumania.

 $cm^{-1}$  region, it can be attributed to the absorbed water in the KBr, since spectrum (b) has a similar peak. The broad absorption in the 3450 cm<sup>-1</sup> region in spectrum (b) can be assigned to the same origin.

Absorptions at 1130 (1130), 1080 (1085), 1050 (1045) cm<sup>-1</sup> are here attributed to the  $v_1$  vibration of  $SO_4^{2-}$  ion, 955 (955) and 935 (930) cm<sup>-1</sup> are to  $\nu_1$ , 635 (635) and 600 (600) cm<sup>-1</sup> are to  $\nu_4$  (Gadsden, 1975; Miller and Wilkins, 1952; Miller et al., 1960). In the free tetrahedral  $SO_4^{2-}$  ion,  $v_1$  is the non-degenerate stretching vibration and is infrared inactive, and  $\nu_3$ and  $v_4$  are triply degenerate vibrations and are infrared active. Distortion of the tetrahedron causes  $v_1$ to be active, and removes the degeneracy of  $v_3$  and  $v_4$ . therefore the clear multiplet character of the above peaks in the  $v_3$  and  $v_4$  regions is explained by the presence of significant distortion of the SO<sub>4</sub><sup>2-</sup> tetrahedron in the klebelsbergite structure. The doublet character of  $v_1$  may possibly be explained by the presence of several kinds of S-O bonds. On heating, as seen in spectrum (b), the above splittings in the  $v_3$ and  $v_4$  regions become degenerate, and  $v_1$  absorption almost extinguished. This implies that the symmetry of  $SO_4^{2-}$  increases and becomes close to  $T_d$  symmetry. There is also evidence of a partial decomposition of  $SO_4^{2-}$  ion on heating (see "Thermal treatment"). This may be related to the apparent increase in  $SO_4^{2-}$  symmetry shown in the IR spectra.

In addition, spectrum (b) shows the appearance of a rather strong peak at 745 cm<sup>-1</sup>. This value reminds us that senarmontite (the low-temperature form of Sb<sub>2</sub>O<sub>3</sub>) has very strong absorption at 740 cm<sup>-1</sup> (Miller and Wilkins, 1952). X-ray powder-diffraction patterns of the heated sample show a possible structural similarity to senarmontite (see "Thermal treatment"). Combined with the above change in the absorption peaks of SO<sub>4</sub><sup>2-</sup> ion, this observation may indicate that the structure of klebelsbergite changes to an antimony oxide-type structure on heating.

## Association and paragenesis

Klebelsbergite is found as fan-shaped aggregates composed of many small crystals (Fig. 1) from Baia Sprie, Rumania. Associated minerals are stibnite, barite, and valentinite. We observed that klebelsbergite crystals almost always grew on stibnite crystals. Large euhedral crystals of barite (more than 1 cm) with stibnite inclusions are commonly observed. Surfaces of the stibnite crystals often show evidence of corrosion. It seems clear that the source of the antimony in klebelsbergite is stibnite. The mineral relationships suggest that stibnite crystallized first, followed by barite, and finally klebelsbergite.

A characteristic feature of the mineral assemblage is the intimate association of valentinite with klebelsbergite. Valentinite is found as small massive spherical crystals on the surface of stibnite crystals. The color is pale yellow to dark yellow, and the appearance resembles that of klebelsbergite except for the habit. Microprobe analyses of the valentinite crystals showed the presence of iron in unusually high concentrations. The concentration of iron was generally higher at the rim than at the core of the spherical crystals; in some cases iron concentrations of up to 5 weight percent were observed. The yellow color is likely due to the iron. Valentinite is very probably an alteration product of stibnite.

The above investigation clearly shows that klebelsbergite is a valid mineral species. It is the only naturally-occurring sulfate of antimony. Hereby the NMNH specimens are all designated as neotypes.

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## References

- Antonovici, S. and I. Vanghelie (1972) Klebelsbergitul, un mineral rar. Dări de seamă ale şedinţelor, 58, 13-19.
- Appleman, D. E. and H. T. Evans, Jr. (1973) Job 9214: Indexing and least-squares refinement of powder diffraction data. Natl. Tech. Inf. Serv., U. S. Dep. Commerce, Springfield, Virginia, Document PB-216 188.
- Gadsden, J. A. (1975) Infrared Spectra of Minerals and Related Inorganic Compounds. Butterworth, Cambridge, Massachusetts.
- Hillebrand, W. F. and G. E. F. Lundell (1953) Applied Inorganic Analysis, 2nd ed. Wiley, New York.
- Mandarino, J. A. (1976) The Gladstone-Dale relationship—Part I: derivation of new constants. *Can. Mineral.*, 14, 498-502.
- Manilici, V., D. Giuşcă and V. Stiopol (1965) Studium Zăcămîntului de la Baia Sprie. Mem. Com. Geol., Rep. Socialistă Romănia, 7, 1-95 (not seen; extracted from Mineral. Abstracts, 18, p. 246, 1967).
- Miller, F. A. and C. H. Wilkins (1952) Infrared spectra and characteristic frequencies of inorganic ions. Anal. Chem., 24, 1253– 1294.
- -----, G. L. Carlson, F. F. Bentley and W. H. Jones (1960) Infrared spectra of inorganic ions in the cesium bromide region. *Spectrochim. Acta, 16,* 135-235.
- Radulescu, D. and R. Dimitrescu (1966) Mineralogia topographica a Romaniei. Bucharest. Edit. Acad. Rep. Socialiste Romănia. (not seen; extracted from Mineral. Abstracts, 18, p. 66, 1967).
- Zsivny, V. (1929) Klebelsbergit, egy új ásvány Felsöbányáról. Mat. Természettud. Értesítő, 46, 19-24.

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