Hydrogen and the melting of silicates¹

ROBERT W. LUTH,² A. L. BOETTCHER

Institute of Geophysics and Planetary Physics and Department of Earth and Space Sciences University of California, Los Angeles, Los Angeles, California 90024

ABSTRACT

The solidi of albite, diopside, and quartz in the presence of H_2O-H_2 vapors have been determined from 5 to 30 kbar for vapors buffered by the assemblages iron-wüstite- H_2O (IW), hematite-magnetite- H_2O (HM), and pure H_2O . As the activity of H_2 in the vapor increases, the solidi in these systems increase systematically in temperature. The difference between the solidus of silicate + H_2O and that of silicate + H_2O + H_2 (HM) in all three systems is larger than could result from simple dilution of the vapor by H_2 , implying that H_2 dissolves in the silicate liquid and inhibits the solution of H_2O . Comparison of these results with data for the analogous $H_2O + CO_2$ systems suggests that H_2 is more soluble than CO_2 in these liquids under similar conditions of pressure and temperature. Variations in f_{H_2} and f_{O_2} in experimental apparatus in different laboratories can explain many apparent inconsistencies in experimental results. These systems also serve as models for melting and metasomatism under conditions of f_{O_2} and f_{H_2} consistent with values postulated for the Earth.

INTRODUCTION

The influence of volatile components, particularly H_2O , on phase relations in silicate systems has been intensively investigated for a number of years. The emphasis on the effect of H_2O stems from petrologic and experimental grounds. In magmas, H_2O is the dominant volatile, as revealed by analyses of volcanic gases and glasses. In the laboratory, H_2O lowers melting temperatures as much as hundreds of degrees, and it is an efficacious flux, greatly enhancing reaction rates and the attainment of equilibrium.

Nevertheless, H₂O-saturated conditions appear only under exceptional conditions in magmas, prompting petrologists to study silicate-H₂O systems at H₂O-undersaturated conditions. To achieve H2O-undersaturated conditions in the laboratory, the hydrous vapor is diluted with another component, commonly CO₂, which is the second-most-abundant volatile component in volcanic gases (Gerlach and Nordlie, 1975; Gerlach, 1982). However, Hill and Boettcher (1970) determined that CO₂ dissolves in some silicate liquids at pressures above ~ 15 kbar based on the shapes and relative positions of the basalt-H₂O and basalt-H₂O-CO₂ solidi. Subsequently, there have been numerous studies of $H_2O + CO_2$ -bearing systems to assess the effect of these volatiles on melting in the mantle (Mysen and Boettcher, 1975; Brey and Green, 1977; Eggler, 1978; Wyllie, 1979), to determine thermodynamic properties of H₂O-CO₂ fluids at high pressures (Eggler and Kadik, 1979; Bohlen et al., 1982, 1983), and

0003-004X/86/0304-0264\$02.00

to determine the mechanisms of solution of the volatiles and their effect on the structural and rheological properties of silicate liquids (Burnham, 1979; Mysen and Virgo, 1980; Stolper, 1982; Boettcher, 1984).

Our work focuses on a mixed-volatile system that has not been considered in any detail to date: the H_2O-H_2 system. This system is important in both geologic and experimental situations, because of the dissociation reaction

$$H_2O[v] = H_2[v] + \frac{1}{2}O_2[v].$$

In the presence of an assemblage that buffers the oxygen fugacity (f_{o_2}) at a fixed value, the ratio of the fugacity of $H_2O(f_{H_2O})$ to the fugacity of $H_2(f_{H_2})$ is fixed at a given pressure and temperature. With the additional constraint that the bulk composition of the vapor is in the O-H system, the composition of the vapor is invariant. Silicate components are negligible, typically composing <1 mol% of the vapor. In natural assemblages, there may be an oxidation-reduction reaction such as the following:

$$2Fe_2O_3[1] = 4FeO[1] + O_2,$$

where Fe_2O_3 and FeO represent ferric and ferrous components in the silicate liquid (Kennedy, 1948; Sack et al., 1980). Other assemblages, which lack a silicate liquid but contain Fe, may also buffer the f_{O_2} (Greenwood, 1975).

In the laboratory, the buffering reaction may be the oxidation of the walls of the pressure vessel by the gas pressure medium, or degassing and dehydration of the furnace assembly in solid-media high-pressure apparatus. Alternatively, f_{O_2} may be controlled by gas mixtures of $CO_2 + H_2$ or $CO_2 + CO$, or by oxygen buffers such as hematite + magnetite + H_2O (HM) or iron + wüstite + H_2O (IW). Even a vapor of nominally pure H_2O contains

¹ Institute of Geophysics and Planetary Physics Contribution No. 2628.

² Present address: Geophysical Laboratory, 2801 Upton Street, N.W., Washington, D.C. 20008.

a finite amount of H_2 and O_2 because of the dissociation reaction. In this case, the composition of this vapor is fixed at a given pressure and temperature by the dissociation reaction of H_2O , the restriction to a binary O-H vapor, and the restriction that the bulk composition remains H_2O . Oxygen buffers, from the most oxidizing one that is commonly used (HM) to the most reducing one (IW), all have lower f_{O_2} and higher f_{H_2} than does pure H_2O .

Previous work on the effect of H_2 on the phase relations of silicates is sparse. Nakamura (1974) studied the system SiO₂-H₂O-H₂ at 15 kbar and f_{H_2} buffered by the talc furnace assembly (~HM) and by the IW buffer. When buffered by the furnace assembly at low f_{H_2} , the system was supercritical, with a single fluid phase rather than discrete liquids and vapors, consistent with previous work on SiO₂-H₂O that indicated a critical end-point at ~10 kbar (Kennedy et al., 1962). In contrast, when the system was buffered by IW at higher f_{H_2} , the liquid and vapor were distinguishable, indicating that H_2 increased the critical point to >15 kbar.

Faile et al. (1967) and Faile and Roy (1970, 1971) studied the solubility of H₂ in SiO₂ glasses quenched from 800°C and $P_{\rm H_2} < 3$ kbar. Even at these low temperatures and pressures and without H₂O, they produced glasses with up to ~8 mol% H₂. They did not study the effect of H₂ on the phase relations of SiO₂.

To examine the effect of high $f_{\rm H_2}$ on silicate phase relations, we have determined vapor-saturated solidi in the systems SiO₂-H₂O-H₂, NaAlSi₃O₈-H₂O-H₂, and Ca-MgSi₂O₆-H₂O-H₂ at pressures between 5 and 30 kbar. We controlled $f_{\rm H_2}$ by using a wide range of buffers. Because the composition of the vapor in equilibrium with the buffer assemblages changes with pressure and temperature, the mole fraction of H₂O in the vapor ($X_{\rm H_2O}$) changes along the vapor-saturated solidi, which complicates comparisons with the analogous systems containing H₂O + CO₂, with fixed $X_{\rm H_2O}$ along a given solidus. However, we can use this sytem in solid-media high-pressure apparatus and thereby study these reactions at higher pressures than attainable in a gas-pressurized apparatus.

EXPERIMENTAL TECHNIQUES

Starting materials

The diopside was synthesized from MgO and CaO (both Puratronic grade, from Johnson Matthey Chemicals) and natural quartz. The CaO was fired at 1000°C for 24 h prior to weighing. The MgO was fired at 1000°C for 24 h, then 1500°C for 4 h, followed by 16 h at 1000°C prior to weighing. The quartz is that used in the study of SiO₂-H₂O-CO₂ by Boettcher (1984); it was ground, boiled in HNO₃, rinsed thoroughly in distilled H₂O, and fired at 1000°C for 24 h before weighing. The oxides were weighed in stoichiometric proportions, mixed thoroughly, and fused at 1450°C for 8 h, quenched, reground, and remelted several times until a homogeneous glass was obtained. This glass was crystallized at 1-atm pressure, 1300°C for 12 h, then recrystallized hydrothermally at 5 kbar, 1100°C for 24 h. This resulted in wellcrystallized, moderately fine grained (10–20 μ m) diopside.

The albite is a natural albite from the Franciscan Formation of California that contains 0.03 wt% CaO, 0.04 wt% K_2O , 0.01

wt% Fe₂O₃, and 0.003 wt% MgO. For most experiments, the albite was in the low structural state. This material is ideal for melting reactions in NaAlSi₃O₈-H₂O-H₂, because above \sim 3 kbar, low albite is the stable polymorph at temperatures near the albite + H₂O solidus (Goldsmith and Jenkins, 1984). For some experiments, the albite was converted to the high structural state by recrystallization at 20 kbar, 1200°C for 24 h. In contrast to the rapid disordering of low albite, the ordering of high albite takes months at the conditions of interest, and therefore high albite persists metastably for the duration of our experiments.

Approximately 5 mg of the appropriate starting material, together with deionized, distilled H₂O, were sealed into 1.6-mm diameter Pt capsules. Masses were measured to the nearest microgram on a Mettler M3 microbalance and were reproducible to $\pm 3 \mu g$. The inner capsule containing silicate + H₂O was sealed inside a 3.5-mm diameter Pt capsule together with the buffer assemblage, either iron-wüstite-H₂O (IW), wüstite-magnetite-H₂O (WM), nickel-nickel oxide-H₂O (NNO), or hematite-magnetite-H₂O (HM). The buffer assemblage fixes the f_{H_2} in the buffer capsule; diffusion of H₂ through the wall of the inner capsule buffers the f_{H_2} inside the sample capsule, in turn fixing the f_{O_2} and f_{H_2O} in the sample capsule.

Most of the pure-H₂O experiments were run without an outer capsule. Previous experiments, using samples of different buffer assemblages, established that the $f_{\rm H_2}$ of the furnace assembly at conditions of interest to this study was lower than that of the HM buffer, as expected because of the lack of hydrous parts in the furnace assembly (Boettcher et al., 1981). Therefore, influx of H₂ into the capsule was negligible, and thus these experiments were in the silicate + H₂O subsystem. To insure that there were no thermal effects because of no outer capsule, we loaded quartz into an outer capsule, together with the sample capsule, mimicking the thermal situation of the other experiments where the double-capsule technique was used. The results of this experiment were consistent with those using a single capsule.

Apparatus and run procedure

The experiments were run in piston-cylinder apparatus, using 2.54-cm diameter furnace assemblies composed of NaCl, pyrex, graphite, BN, and MgO (Boettcher et al., 1981). Temperatures were monitored using Pt-Pt₉₀Rh₁₀ thermocouples encased in mullite ceramic, with the tip of the thermocouple in contact with the capsule. We used the hot piston-in technique by bringing the pressure to about 5% below the final value, then increasing temperature and finally pressure to the desired values. The experiments were controlled with a precision of $\pm 2^{\circ}$ C and ± 0.05 kbar. At the end of the run, the experiment was quenched by shutting off the power, resulting in a temperature drop of 500°C or more within 20 s, and to room temperature in less than 3 min.

Following the run, the buffer capsule was weighed, punctured, heated, and then reweighed to confirm that H_2O was present. Optical microscopy or X-ray diffraction confirmed the presence of the required solid phases of the buffer. The sample capsule was cleaned, weighed, punctured, heated at 112°C, and then reweighed. A weight loss at this stage confirmed that H_2O remained in the capsule throughout the experiment. In typical hypersolidus experiments, the weight loss was 75–90% of the amount of H_2O originally loaded in the capsule, indicating that some H_2O remained in the quenched liquid. However, the weight losses for capsules from hyperliquidus experiments indicate that >50% of the volatiles exsolved from the liquid during the quenching of the experiment. After the sample capsule was opened, the products were examined in immersion oils using a petrographic microscope.

Run products

The nature of the run products differed somewhat from system to system. In the NaAlSi₃O₈-H₂O-H₂ and SiO₂-H₂O-H₂ systems, quenched liquids were always bubble-rich glasses that had nearly completely degassed during the quench. In the CaMgSi₂O₆-H₂O-H₂ systems, liquids quenched to crystals, either large, single crystals with many vapor bubbles, glassy inclusions, and undulatory or patchy extinction, or radial sheaves of crystals with interstitial glass. In either case, the quenched liquid was clearly distinguishable from the well-crystalized, usually smaller crystals of primary diopside.

In all three systems, vapor quenched to "fish-roe"—small, isotropic, glass spheres of lower refractive index than quenched liquid. In the SiO₂-H₂O-H₂ system, vapor also quenched to large, thin plates of bubble-free glass, which often appeared to have formed from coalesced fish-roe spheres. These plates of vapor could be distinguished from quenched liquid by their lack of bubbles and lower index of refraction (n < 1.50). At pressures above the critical points in SiO₂-H₂O-H₂, liquid and vapor were indistinguishable; there was a single fluid phase that became increasingly abundant at the expense of the crystalline quartz as the temperature increased. In the CaMgSi₂O₆-H₂O-H₂ system, at pressures above ~25 kbar, vapor commonly quenched as large, thin plates or acicular pieces of isotropic, or nearly isotropic, material of low refractive index (n < 1.50).

To ensure equilibrium, runs were reversed at several pressures in all three systems. These were done by first subjecting a sample to conditions of pressure, temperature, and run duration, known from our other experiments to be sufficient to melt, or partially melt, the charge. The temperature was then lowered to below the solidus determined by previous runs. Run products in a successful reversal contained no quenched liquid. Data for intracrystalline diffusion of hydrogen through Pt (Ebisuzaki et al., 1968) indicate that equilibrium with respect to f_{H_2} would be established between the buffer and sample capsules in <5 s even at the lowest temperatures studied.

Calculation of vapor composition

The composition of the H₂O-H₂ vapor was calculated using the expressions for f_{02} , for the various buffers from Huebner (1971), the free energy of dissociation of H₂O from Robie et al. (1979), and the modified Redlich-Kwong (MRK) equation of state of Holloway (1977). When buffered by an O_2 buffer, the $X_{O_2}^{v}$ was always orders of magnitude less than the $X_{H_2}^{\nu}$, and the vapors were considered to be binary H2O-H2 mixtures. Because of the many factors of varying precision used in the calculation, there are large uncertainties in the $X_{H_2}^{\nu}$. The f_{O_2} of the buffer was calculated from experimental data at atmospheric pressure and the molar volumes of the solid phases; these extrapolations have been experimentally constrained only at 2 kbar (Chou, 1978). For example, changing log (f_{O_2}) by 1 changes $X_{H_2}^{\nu}$ for a HM-buffered vapor by a factor of 3. Uncertainties in the free energy of dissociation for H₂O lead to uncertainties in the $f_{\rm H_2}/f_{\rm H_2O}$ ratio and thereby in the $X_{H_2O}^{\nu}$. An error in the free energy of dissociation of H₂O of 10% may change X_{H}^{ν} , for a HM-buffered vapor by a factor of 15. The MRK equation of state, used to convert $f_{\rm H_2}$ to $X_{H_{2}}$, has potentially large uncertainties for two reasons. First, there are no fugacity data for pure H₂ or H₂-H₂O mixtures at pressures higher than 3 kbar (Shaw, 1963; Shaw and Wones, 1964; Presnall, 1969) or for H₂O above 10 kbar (Burnham et al., 1969) against which to calibrate the MRK equation. Second, the MRK equation uses a simple mixing rule that may not be realistic for H₂O-H₂ vapors at the conditions of our experiments. In this regard, it should be noted that de Santis et al. (1974), the authors who published the modification of the Redlich-Kwong equation of state used by Holloway (1977), specifically stated that because of the simplifications used in the formulation of this equation of state, it would not provide highly accurate results, and furthermore, they recommended that the equation be used in the region 25–700°C and at pressures up to 1500 bars.

RESULTS

Experimental results are presented in Tables 1 through 3 and Figures 1 through 4. There are two features that the systems studied share. First, increasing $f_{\rm H_2}$ at a given pressure systematically increases the temperature of the vapor-saturated solidus, with a large difference between the pure-H₂O and HM-buffered solidi, disproportionate to the increase in the calculated $X_{\rm H_2}^v$. Second, the amount of quenched vapor observed in the run products decreases as $X_{\rm H_2}$ increases, indicating that the solubility of silicates in the vapor decreases with decreasing $X_{\rm H_2O}^v$, similar to results for systems containing H₂O + CO₂ (Shettel, 1973; Eggler, 1975; Boettcher, 1984). For simplicity, individual features of the three systems will be presented in turn.

CaMgSi₂O₆-H₂O-H₂

There is a large ($\sim 80^{\circ}$) temperature difference between the pure-H₂O solidus and that buffered by HM at ~ 10 kbar, but the difference decreases with increasing pressure (Fig. 1). For the IW-buffered solidus, there is a pronounced temperature minimum between 15 and 24 kbar. Below 15 kbar, the IW-buffered solidus is $\sim 100^{\circ}$ C higher than the solidus for pure H₂O, whereas above 15 kbar, the difference between the two solidi increases to $\sim 130^{\circ}$ C at 27.5 kbar.

NaAlSi₃O₈-H₂O-H₂

In contrast to the results for the CaMgSi₂O₆-H₂O-H₂ system, there are no abrupt changes in slope for the solidus of NaAlSi₃O₈-H₂O-H₂ (IW), although we examined this system only to ~17 kbar because of the reaction of albite to jadeite + quartz (Fig. 2). The difference in temperature between the pure-H₂O and the IW-buffered solidi remains relatively constant at ~50°C.

Following the study of Goldsmith and Jenkins (1984) on the effect of the structural state of albite on vaporpresent melting, we determined the metastable vapor-saturated solidus of high albite + H_2O (Fig. 3), which is at lower temperatures than the stable solidus of low albite + H_2O . This temperature difference agrees with the data of Goldsmith and Jenkins (1984), and it is in accord with the requirement that the metastable polymorph melts at a lower temperature than does the stable polymorph. Comparison of the X-ray diffractograms of the albite before and after the experiments showed no ordering during the experiments.

SiO₂-H₂O-H₂

As in the other systems, increasing $f_{\rm H_2}$ at a given pressure increases the temperature of the vapor-saturated solidus of quartz + vapor, and again, there is a measurable dif-

Table 1. Experimental results for CaMgSi₂O₆-H₂O-H₂

Table 1-Continued

	D	т	H=O						_				
Run ∉	kbar	°C	loaded wt%	x _{H2}	Duration h	Results	Run #	p kbar	°C	H2O loaded wt%	x _{H2}	Duration h	Results
Pure H ₂ 0							NNO Buft	fer					
1012 1016 1008 1021 994 1007 118 6 R	7.5 7.5 7.5 7.5 7.5 7.5 7.5 7.5	1170 1180 1190 1200 1210 1190	17.1 13.8 18.6 15.0 15.8 19.3 20.5	3.0x10-7 3.4x10-7 3.8x10-7 3.8x10-7 4.3x10-7 4.8x10-7 3.8x10-7 3.8x10-7	33383393	Di+V Di+V Di+L+V Di+L+V Di+L+V Di+L+V Di+L+V (see #1008)	1027 984 972 973 976 980	7.5 10 10 10 10 10	1260 1210 1230 1240 1250 1260	17.5 15.1 20.1 13.3 16.0 18.0	2.4x10-3 1.5x10-3 1.6x10-3 1.6x10-3 1.6x10-3 1.6x10-3 1.6x10-3	333324	Di+V Di+V Di+V Di+V Di+V Di+V
1005	10	1170	15.5	5.6x10-8	3	Di+V Di+V	1030 1026 1031	20 20 20	1220 1230 1240	17.9 19.0 19.8	3.1x10-4 3.2x10-4 3.3x10-4	3 3 3	Di+V Di+V Di+(L)+V
992 990 987 985	10 10 10 10	1170 1180 1190 1190 1200	17.5 19.3 42.1 18.3	5.6x10-8 6.4x10-8 7.4x10-8 7.4x10-8 8.5x10-8	2 11 12 12 12	D1+V Di+V Di+L+V Di+L+V Di+L+V	1032 1028 WM_Buffe	27.5 27.5	1190 1200	16.8 19.5	8.6x10-5 8.9x10-5	3	Di+V Di+L+V
982 967 966* 915 1189R	10 10 10 10 10	1210 1220 1230 1240 1190 1175	14.7 16.7 18.3 17.8 21.7	9.9x10-8 1.1x10-7 1.3x10-7 1.5x10-7 7.4x10-8 6.0x10-8	a a N a a a	Di+L+V Di+L+V Di+L+V Di+L+V (see # 987) Di+V	680 684 643 647 655 646	7.5 7.5 10 10 10	1270 1280 1250 1260 1260	18.9 21.0 19.3 17.9 18.3	1.0x10-3 2.9x10-3 1.0x10-3 2.9x10-3 2.9x10-3 1.6x10-3	2224	Di+V Di+L+V Di+V Di+V Di+V
1024 1023 1003	20 20 20	1170 1180 1190	17.5 14.7 20.4	2.5x10-17 5.0x10-17 9.7x10-17	333	Di+V Di+L+V Di+L+V	653 649	10	1270	20.3 21.3 20.4	1.6x10-2 1.6x10-2 1.6x10-2	2 4 2	Di+(L)+V Di+L+V Di+L+V
1023 1004 1018	20 20 20	1200 1210 1220	15.6 17.8 17.8	1.8x10-16 3.6x10-16 7.0x10-16	3333	Di+L+V Di+L+V Di+L+V Di+L+V	673 672	20 20	1240 1250	25.1 26.1	1.0x10-3 2.8x10-3	2 2	Di+V Di+L+V
1015 1213	27.5	1170 1180	19.3 18.2	< 10-20	3 0,5	Di+V Di+V	IW Buffe 762	r 7.5	1280	19.2	0.42	0.5	Di+V
HM Buffer 695	7.5	1200	26.6	< 10-20	15	Di+L+V Di+V	767 769 770 1188r	7.5 7.5 7.5 7.5	1290 1300 1310 1310	20.1 28.3 21.9 18.3	0.42 0.42 0.41 0.41	0.5 0.5 0.5 0.5	Di+V Di+V Di+L+V (see # 770)
696 691 702 682 718 676 677 1187R	7.5 7.5 7.5 7.5 7.5 7.5 7.5 7.5 7.5	1220 1240 1250 1260 1260 1270 1280 1270 1255	25.8 23.0 24.4 23.5 30.1 25.8 19.3 19.9	1.1x10-5 1.1x10-5 1.1x10-5 1.1x10-5 1.1x10-5 1.1x10-5 1.2x10-5 1.2x10-5 1.1x10-5 1.1x10-5	15 15 15 2 18 2 2 3 3	Di+V Di+V Di+V Di+V Di+V Di+V Di+L+V Di+L+V (see # 676) Di+V	687 738 772 746 771 745 742 1191R	10 10 10 10 10 10 10 10	1295 1270 1290 1300 1300 1310 1330 1300 1300	18.3 27.5 22.3 23.5 24.7 30.3 17.8 20.2 18.6	0.42 0.39 0.39 0.38 0.38 0.38 0.38 0.38 0.38	0.5 2 0.5 2 0.5 2 2 0.5	Di+L+V Di+V Di+V Di+L+V Di+L+V L+V Di+L+V Di+L+V (see # 771)
644 725 735 920 642 921 639 922 726 7368	10 10 10 10 10 10 10 10 10	1220 1230 1240 1250 1250 1260 1260 1260 1270 1280	23.6 22.2 14.3 23.8 17.1 19.7 18.1 17.7 22.3	7.5×10-6 7.6×10-6 7.8×10-6 7.9×10-6 7.9×10-6 8.0×10-6 8.0×10-6 8.1×10-6 8.2×10-6	2 23 18 2 2 2 2 2 18 2	Di+V Di+V Di+V Di+V Di+V Di+V Di+(L)+V Di+(L)+V L+V (see # 922)	760 773 774 783 777 780 779 778	20 20 20 20 20 20 20 20 20	1285 1240 1250 1270 1280 1290 1290 1300 1310	23.5 15.0 16.9 27.3 18.9 24.7 22.5 13.4	0.39 0.28 0.28 0.28 0.28 0.28 0.28 0.28 0.28	0.5 2 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0.5	Di+V Di+V Di+V Di+L+V Di+L+V Di+L+V Di+L+V Di+L+V Di+L+V
021	10	1250	21.2	7.9x10-6	2	Di+V	1210 1212 1211	24 24 24	1280 1290	21.2 18.1	0.24	0.5	Di+V Di+V
657 661 669 659	20 20 20 20 20	1240 1250 1250 1260	22.4 24.8 35.9 30.0	2.0x10-6 2.0x10-6 2.0x10-6 2.0x10-6 2.1x10-6	2 2 14 2	Di+(L)+V Di+L+V Di+L+V Di+L+V Di+L+V	1203 1193 1195	27.5 27.5 27.5	1200 1240 1270	21.7 17.1 18.8	0.23 0.23 0.23	0.5	Di+L+V Di+V Di+V Di+V
824 826 823 1214 822 793	27.5 27.5 27.5 27.5 27.5 27.5 27.5	1180 1190 1200 1210 1220 1230	20.9 27.3 31.6 18.5 30.9 27.6	5.9x10-7 6.1x10-7 6.3x10-7 6.5x10-7 6.7x10-7 6.9x10-7	0.5 0.5 0.5 0.5 0.5 0.5	Di+V Di+V Di+V Di+V Di+L+V Di+L+V	1200 1204 1206 1201 1207R	27.5 27.5 27.5 27.5 27.5 27.5 27.5	1290 1300 1310 1320 1320 1295	20.1 20.3 19.3 19.7 20.2	0.23 0.23 0.23 0.23 0.23 0.23 0.23 0.23	0,5 0,5 0,5 0,5 0,5 0,5	Di+V Di+V Di+L+V Di+L+V (see #1201) Di+V
802 1215R	27.5 27.5 27.5	1240 1220 1205	31.3 22.3	7.1x10-7 6.9x10-7 6.4x10-7	0.5 0.5 0.5	D1+L+V (see # 822) Di+V	1216 1209 1208	30 30 30	1300 1310 1330	16.3 22.4 22.0	0.21 0.21 0.22	0.5 0.5 0.5	Di+V Di+L+V Di+L+V

Abbreviations: Di-diopside, L=liquid, V=vapor, R=successful reversal, r=attempted reversal, ()=trace amount, *=run in outer capsule filled with mechanical buffer of SiO₂. $\chi^{}_{H^2}$ calculated as discussed in text.

ference between the pure H_2O and the HM-buffered curves. The three solidi terminate in critical points, where liquid and vapor become indistinguishable. In the SiO₂-H₂O system, this point (K₂ in Fig. 4) is a critical end-point (Kennedy et al., 1962). However, for the SiO₂-H₂O-H₂ compositions, K_A and K_B are critical points, not critical endpoints, as explained by Boettcher (1984). The outrageous shape of the vapor-saturated solidus of SiO₂-H₂O-H₂ (IW) has been tightly constrained by reversal experiments at pressures bracketing each abrupt change in slope.

DISCUSSION

The solidus of silicate- H_2O should be univariant, yet the products of our hypersolidus experiments contain solid + liquid + vapor. This behavior may be caused by the incongruent melting of the silicates, thereby causing a divariant melting interval. However, the same behavior is found in the SiO₂- H_2O system, where there can be no such

Table 2. Experimental results for NaAlSi₃O₈-H₂O-H₂

Table 2-Continued

Run #	P kbar	°C	H2O loaded wt%	x _{H2}	Duration n	Results
Pure H ₂ 0,	low albit	e				
842 810 799 843 849R 1132R	ភូមិសូសួសួ	720 730 750 775 775 720 730 720	26.0 24.0 25.0 20.5 20.3 16.5	8.3x10-10 1.1x10-9 1.6x10-9 2.8x10-9 2.8x10-9 8.3x10-10 1.1x10-9 8.3x10-10	8 24 5 17 8 15.5	Ab+V Ab+L+V Ab+L+V (see # 843) Ab+V (see # 810) Ab+V
850 851 856 1025 1039 853 1131R	10 10 10 10 10 10 10	680 690 700 700 700 700 700 690	17.3 16.9 20.7 44.5 19.4 18.5 19.2	1.6x10-13 1.6x10-13 2.6x10-13 4.0x10-13 4.0x10-13 4.0x10-13 4.0x10-13 4.0x10-13 2.6x10-13	8 15 24 8 20 8 15	Ab+V Ab+V Ab+V Ab+L+V Ab+L+V Ab+L+V Ab+L+V (see # 853) Ab+V
836 839 809 844R	15 15 15 15 15	630 640 650 650 640	25.4 22.2 31.0 21.4	4.4x10-30 4.1x10-29 3.3x10-28 3.3x10-28 4.1x10-29	7.5 8 6.5 13	Ab+V Ab+V Ab+L+V (see # 809) Ab+V
Pure H20	, high albi	ite				
873 894 893	5 5 5	720 730 740	16.0 21.0 17.8	8.3x10-10 1.1x10-9 1.3x10-9	8 8 14	Ab+V Ab+V Ab+L+V
872 867 868 895 866 874 881 882	10 10 10 10 10 10 10	660 670 680 690 690 700 710	18.2 35.7 20.3 27.5 15.9 20.6 19.7	6.1x10-14 1.0x10-13 1.6x10-13 1.6x10-13 2.5x10-13 2.5x10-13 4.0x10-13 6.1x10-13	8 15 25 9 14 19 19	Ab+V Ab+V Ab+L+V Ab+L+V Ab+L+V Ab+L+V Ab+L+V Ab+L+V (Ab)+L+V
891 883	15 15	630 640	17.5	4.4x10-30 4.1x10-29	8	Ab+V Ab+L+V
HM Buffe	r, low alb	ite				
785 816 808 3045 819 817 801 1133R	ភតាតាលាតាតាត	730 740 750 760 760 770 770 750	22.7 16.2 26.5 20.0 21.5 23.1 31.7 19.5	6.4x10-6 6.6x10-6 6.6x10-6 6.7x10-6 6.9x10-6 6.9x10-6 7.1x10-6 7.1x10-6 6.7x10-6 6.7x10-6	24 22 24 9 7,5 24 7,5 14	Ab+V Ab+V Ab+V Ab+V Ab+V Ab+(L)+V Ab+(L)+V (see # 801) Ab+V
784 857 3026	10 10 10	680 700 700	24.8 23.0 20.0	1.9x10-6 2.0x10-6 2.0x10-6	24 8 24	Аь+V Аь+V Аь+V

Run #	P kbar	T °C	H2O loaded wt%	x _{H2}	Duration h	Results	
855 858 854 1151R	10 10 10 10	710 710 730 710 690	21.8 19.6 23.7 21.1	2.1x10-6 2.1x10-6 2.3x10-6 2.1x10-6 2.1x10-6 2.0x10-6	5 9 15 8 14	Ab+L+V Ab+L+V Ab+L+V (see # 858) Ab+V	
800 805 818 820 807	15 15 15 15 15	660 680 690 700 700	27.8 29.3 23.2 23.0 29.1	5.7x10-7 6.4x10-7 6.7x10-7 7.0x10-7 7.0x10-7	24 24 24 7.5 24	Ab+V Ab+V Ab+L+V L+V L+V	
1064 1052 1063 1051	16 17 17 17	680 660 670 680	19.6 18.4 20.5 18.5	4.6x10-7 3.7x10-7 3.9x10-7 4.1x10-7	8 8 8	Ab+L+V Ab+V Ab+L+V Ab+L+V	
NNO Buffe	er, low all	oite					
1011 1010 1006	5 5 5	740 750 760	15.7 17.2 19.5	1.6x10-3 1.6x10-3 1.7x10-3	8 8 8	Ab+V Ab+L+V Ab+L+V	
995 998 996	10 10 10	700 710 720	16.9 18.3 18.9	4.2x10-4 4.4x10-4 4.6x10-4	8 8 8	Ab+V Ab+L+V Ab+L+V	
997 1000 999	15 15 15	680 690 700	17.2 19.1 19.4	1.1x10-4 1.2x10-4 1.3x10-4	8 8 8	Ab+V Ab+L+V Ab+L+V	
IW Buffer	r, low alb	ite					
845 846 835 831 1140R	555555	750 760 770 790 770 760	19.2 17.5 23.5 25.8 18.5	0.53 0.53 0.53 0.53 0.53 0.53 0.53	6 6 6 7,5 6	Ab+V Ab+V Ab+L+V Ab+L+V (see # 835) Ab+V	
1128 1129 1138 1139	10 10 10 10	720 730 740 750	19.1 17.4 16.2 18.9	0.45 0.44 0.44 0.43	8 8 10	Ab+V Ab+V Ab+L+V Ab+L+V	
828 834 829	15 15 15	700 710 720	25.8 28.1 23.3	0.36 0.36 0.36	7.5 7.5 7.5	Ab+V Ab+L+V Ab+L+V	
1066 1068	17 17 17	680 690 700	19.6 16.2 20.0	0.32 0.33 0.33	8 8 8	Ab+V Ab+L+V Ab+L+V	

incongruency, so that this cannot be a general explanation. Alternatively, there may be insufficient H_2O in the capsule to allow the formation of 100% H_2O -saturated liquid. If this were the case, however, there would be no quenched vapor in the run products, contrary to what is observed. Finally, and most probably, we have allowed insufficient time for all of the solid to melt. However, determination of the solidus does not require complete melting, and our reversal experiments verify that our curves represent equilibrium.

The increase in temperature of the vapor-saturated solidi with increasing $X_{H_2}^{\nu}$ may largely result from a simple lowering of the $a_{H_{2}0}^{\nu}$, with H₂ acting as as an inert diluent in the vapor. However, several factors indicate that H₂ is soluble in silicate liquids. First, increasing the $X_{H_2}^{\nu}$ from <10⁻¹⁰ (calculated for pure H₂O) to ~10⁻⁵ (calculated for HM-buffered vapor) increases the temperature of the vapor-saturated solidi at a given pressure disproportionately to the calculated decrease in $X_{H_{2}O}^{\nu}$. Second, vapor-saturated solidus temperatures at a given $a_{H_{2}O}^{\nu}$ are lower in the presence of H_2O-H_2 vapors than in the presence of a vapor of either H_2O-CO_2 or H_2O-X , where X is an inert, insoluble diluent, suggesting that H_2 is soluble in these liquids, probably more so than is CO_2 . Each of these factors will be discussed in turn.

The difference between the solidus for pure H_2O and that buffered by HM may be caused by the reaction of H_2 with one of the phases. Trace amounts (<0.1 wt%) of H_2O , either as OH groups or as H_2O molecules, have been found in garnets (Aines and Rossman, 1984a), olivine (Beran and Putnis, 1983), and quartz (Aines and Rossman, 1984b; Aines et al., 1984), whereas there is no evidence for molecular H_2 in these minerals. Because H_2O dissolved in the solid stabilizes the solid + vapor assemblage relative to liquid, dehydration of the solid resulting from lowered a'_{H_2O} stabilizes the liquid and lowers the temperature of the vapor-saturated solidus. However, this effect would decrease solidus temperatures with increasing X'_{H_2O} , opposite to what is observed.

Alternatively, if the calculated values of $X_{H_2}^{\nu}$ for the HM

Table 3. Experimental results for SiO₂-H₂O-H₂

Table 3-Continued

Run #	p kbar	T °C	H20 loaded wt%	x ^v _{H2}	Duration h	Results	Run #	P kbar	T °C	H20 Noaded wt%	x _{H2}	Duration h	Results
Pure H ₂ 0							1156R	10.5 10.5	1110 1120	18.5	0.41 0.41	2	L+V (see #1135)
1086 1042 1159R	5555	1050 1060 1060 1050	19.9 19.9 18.2	2.5x10-7 2.8x10-7 2.8x10-7 2.8x10-7 2.5x10-7	2 2 2 2	Qz+V Qz+L+V (see #1042) Qz+V	1104 1109 1102	10,5 12 12 12	1100 1090 1100 1110	14.3 20.6 17.2	0.41 0.39 0.39 0.39	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	Qz+V Qz+V Qz+V Oz+L+V
1037 1040	7 7	1050 1060	17.4 18.0	8.4x10-8 9.7x10-8	22	Qz+V Qz+L+V	1093 1090	12 12	1120 1130	17.3 21.7	0.39 0.39	22	Qz+L+V Qz+L+V
1038 1034 1035 1033 1058R HM Buffer	000000	1015 1030 1040 1050 1050 1050	15.7 17.4 17.0 12.0 15.7	1.1x10-8 1.4x10-8 1.7x10-8 2.0x10-8 2.0x10-8 1.7x10-8	2 2 2 2 2 2 2 2 2	Qz+V Qz+V Qz+V Qz+L+V (see #1033) Qz+V	1148 1088 1085 1082 1081 1136r 1147R	15 15 15 15 15 15 15 15	1100 1110 1120 1140 1160 1110 1100 1110 1090	18.1 18.7 19.9 19.2 19.8 18.7 16.3	0,36 0,36 0,36 0,36 0,36 0,36 0,36 0,36	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	0z+V (Qz)+L+V (Qz)+L+V L+V (see #1088) L+V (see #1088) Qz+V
1092 1097 1099 1106 1113 1107	5555	1060 1080 1090 1110 1120 1130	17.4 18.5 17.4 19.7 19.2 17.0	1.2x10-5 1.3x10-5 1.3x10-5 1.3x10-5 1.3x10-5 1.3x10-5 1.4x10-5	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	Qz+V Qz+V Qz+V Qz+V Qz+L+V Qz+L+V	1119 1120 1118 1144R	17.5 17.5 17.5 17.5 17.5	1050 1060 1070 1060 1050	20.3 18.9 18.9 20.2	0.33 0.33 0.33 0.33 0.33	2222222	Qz+V Qz+L+V Qz+L+V (see #1120) Qz+V
1161r 1166R	5 5 5 5	1120 1110 1120 1120	21.3 16.8	1.3x10-5 1.3x10-5 1.3x10-5 1.3x10-5 1.3x10-5	2224	(see #1113) Qz+L+V (see #1113) Qz+V	1095 1096 1091 1145r	20 20 20 20	1050 1060 1070 1060	20.4 17.1 20.6 22.0	0.30 0.30 0.30 0.30	2222	Qz+V Qz+L+V Qz+L+V (see #1096)
1116 1098 1100	7777	1050 1070 1070	18.2 18.8 19.1	8.8x10-6 9.1x10-6 9.1x10-6	224	Qz+V Qz+V Qz+V	1155R	20 20 20	1060 1040	19,1	0.30 0.30 0.30	222	(see #1096) Qz+V
1101 1110 1105	7 7 7	1080 1090 1100	22.2 19.2 19.5	9.3x10-6 9.5x10-6 9.6x10-6	222	Qz+V Qz+V Qz+L+V	1103 1124 1125	25 25 25	1000 1020 1030	18.8 19.5 19.9	0.24 0.24 0.24	2 2 2 2	Qz+V Qz+V Qz+V
1123 1127 1117	9 9 9	1060 1070 1080	18.3 23.3 20.8	6.6x10-6 6.7x10-6 6.8x10-6	2 2 2	Qz+V Qz+V Qz+L+V	1108 1112 1114 1111	25 25 25 25	1040 1060 1070 1080	20.1 21.5 18.1 20.9	0.24 0.25 0.25 0.25	2 2 2 2 2 2 2	Qz+L+V Qz+L+V Qz+L+V Qz+L+V
IW Buffer 1087 1089 1060 1061 1059	55777	1160 1170 1070 1090	21.4 16.5 16.7 21.3	0.48 0.48 0.46 0.46	2 2 2 2	Qz+V Qz+L+V Qz+V Qz+V	1176 1177 1174 1173 1179R	27.5 27.5 27.5 27.5 27.5 27.5 27.5	1020 1030 1040 1060 1040 1020	21.5 18.8 20.2 20.9 18.6	0.22 0.22 0.22 0.22 0.22 0.22 0.22	~~~~~	Qz+V Qz+V Qz+L+V Qz+L+V (see #1174) Qz+V
1073 1074 1075 1072 1071 1170	777777	1130 1140 1150 1160 1180 1200	21.3 18.0 18.8 20.2 19.6 14.9 19.7	0.46 0.45 0.45 0.45 0.45 0.45 0.45	22222	Qz+V Qz+V Qz+V Qz+V Qz+L+V Qz+L+V	1162 1163 1164 1168 1169	30 30 30 30 30	1020 1040 1060 1080 1100	17.8 20.4 17.4 20.1 20.1	0.20 0.20 0.20 0.20 0.20 0.20	22222	Qz+F Qz+F Qz+F Qz+F Qz+F
1076 1080 1079 1141R	99999	1150 1160 1170 1160 1150	20.9 19.9 18.4 14.2	0.43 0.43 0.43 0.43 0.43 0.43	2 2 2 2 2 2	Qz+V Qz+L+V (Qz)+L+V (see #1080) Qz+V	1180 1182 1181 1183	35 35 35 35	1100 1120 1140 1160	20.5 18.9 22.4 17.3	0.17 0.17 0.17 0.17 0.17	2 1 1	Ct+F Ct+F Ct+F Ct+F
1122 1135 1121 1146r	10.5 10.5 10.5 10.5	1110 1120 1130 1120	19.9 21.2 21.7 21.4	0.41 0.41 0.41 0.41	2 2 2	Qz+V L+V L+V (see #1135)	Abbreviat fluid, R= XY ₂ calcu	ions: Qz= successful lated as d	quartz, Ci reversal, iscussed i	t≃coesite, L , r=attempte n text,	=liquid, V=\ 1 reversal,	/apor, F=super ()=trace amo	critical unt,

buffer are systematically too low because of the uncertainties discussed previously in the section on Calculation of Vapor Composition, the difference between pure-H₂O and HM-buffered solidi may simply result from the difference in $X_{H_2}^{\nu}$ with H₂ acting as an inert diluent. However, this requires that the combined uncertainties in the calculation of the vapor composition would change $a_{H_2}^{\nu}$ by a factor of 10⁺⁴, which appears to be unrealistic.

Of great interest is the possibility that the difference between the vapor-saturated solidus for pure H_2O and that buffered by HM results from solubility of H_2 in the liquids. For example, H_2 may associate with H_2O in the liquid, increasing the molar volume of the H_2O-H_2 complex in the liquid, decreasing a'_{H_2O} and stabilizing the solid + vapor assemblage to higher pressures and temperatures. H_2 could also lower the a'_{H_2O} by forming bonds with ions in the liquid or by occupying sites as molecular H_2 that otherwise would be filled by H_2O . The distortion of the liquid caused by H_2 molecules occupying sites that are appropriate for H_2O molecules would increase the free energy of the liquid and thereby increase melting temperatures. One possible reaction is to form Si-H bonds in the liquid. Infrared evidence for Si-H bonds was obtained by Faile and Roy (1970) in SiO₂ glass containing 2 mol% H_2 that had been subjected to gamma-radiation and a neutron flux. They postulated that the irradiation provided the energy to break H-H and Si-O-Si bonds to form Si-O-H and Si-H bonds. Perhaps the higher temperatures in our study provided the necessary activation energy to form Si-H bonds.

Would formation of Si-H bonds strengthen or weaken neighboring bonds? If Si-H bonds strengthen neighboring Si-O bonds, it would be more difficult to locally break a Si-O-Si bridge to form Si-O-H groups. Alternatively, if



Fig. 1. Pressure-temperature projection of the vapor-saturated solidi of diopside + vapor for various values of the activity of hydrogen in the vapor $(a_{H_2}^u)$. Closed symbols represent diopside + liquid + vapor; open symbols are diospide + vapor. Values of $a_{H_2}^u$ are buffered by pure H₂O, hematite-magnetite-H₂O (HM), nickel-nickel oxide-H₂O (NNO), wüstite-magnetite-H₂O (WM), and iron-wüstite-H₂ (IW). Where determined, the NNObuffered solidus coincides with the HM-buffered solidus.

Si-H bonds weaken the Si-O bonds of those Si cations, those bonds would preferentially react with H₂O to form Si-O-H groups, effectively ordering the Si-O-H groups in the liquid. The resulting decrease in entropy would increase the free energy of the liquid and increase melting temperatures. Either way, disruption of the framework of the silicate liquid could lead to less favorable conditions for Si-O-H bonding than in the absence of H₂, thereby decreasing $a'_{\rm H_2O}$ and increasing solidus temperatures.

One possible explanation of the difference between the pure- H_2O and the HM-buffered solidi of CaMgSi₂O₆- H_2O - H_2 and of NaAlSi₃O₈- H_2O - H_2 is the differing extent of incongruent solubility of silicate components in the vapor as $X_{H_2}^*$ increases. With extensive incongruent solubility of the crystalline phase in the vapor, the residual solid departs from NaAlSi₃O₈ or CaMgSi₂O₆ composition and melts at a different temperature. However, the same effect is seen in SiO₂- H_2O - H_2 , where there is only one silicate component and no possibility of incongruent solution of silicates in the vapor.

Comparison of the effects of H_2 on the vapor-saturated solidi with those of an inert diluent and with those of CO_2 provide insight into the solubility of H_2 in these silicate liquids. We calculated the phase relations for silicate- H_2O -Xusing thermochemical data, activity coefficients for the components in the vapor provided by the MRK equation,



Fig. 2. Pressure-temperature projection of the vapor-saturated solidi of albite + vapor for various values of $a_{H_2}^*$. Closed symbols represent albite + liquid + vapor; open symbols are albite + vapor. Values of $a_{H_2}^*$ are buffered by the same assemblages as in Fig. 1. Where determined, the NNO-buffered solidus coincides with the HM-buffered solidus. The albite = jadeite + quartz reaction is from Boettcher and Wyllie (1968).

and the assumption that X was completely insoluble in the liquids. The data for silicate-H₂O-CO₂ were taken from the literature.

Vapor-saturated solidi of CaMgSi₂O₆-H₂O and Ca- $MgSi_2O_6-H_2O-X$ (Fig. 5) were calculated using the model of Eggler and Burnham (1984). To calculate the solidus, this model required the temperature of the solidus of CaMgSi₂O₆-H₂O at one pressure, as determined by Eggler and Rosenhauer (1978). If this temperature is revised downward to agree with our data on the solidus of CaMgSi₂O₆-H₂O, the calculated solidus becomes lower in temperature than the experimental solidus, with the difference between the two increasing from ~20°C at 10 kbar to ~100°C at 27.5 kbar. Without a better thermodynamic model for the phase relations in CaMgSi₂O₆-H₂O, it is impossible to say if H₂ is dissolving in the liquid. However, comparing the vapor-saturated solidus of CaMg-Si₂O₆-H₂O-CO₂ and that of CaMgSi₂O₆-H₂O-H₂ (IW) (Fig. 6) suggests that below ~ 15 kbar, H₂ and CO₂ have similar effects on the melting temperatures at $X_{\rm H_{2O}}^{\nu} \approx 0.70$. Above this pressure, the solidus of CaMgSi₂O₆-H₂O-H₂ (IW) increases in temperature with increasing pressure, whereas the solidus of CaMgSi₂O₆-H₂O-CO₂ continues to decrease in temperature with increasing pressure. This difference may result from an increase in the solubility of CO2 or



Fig. 3. Pressure-temperature projection of the vapor-saturated solidi of high albite + vapor and low albite + vapor in the system NaAlSi₃O₈-H₂O. Closed symbols represent albite + liquid + vapor; open symbols are albite + vapor.

 H_2O or both in the CaMgSi₂O₆- H_2O -CO₂ liquid with increasing pressure, or a decrease in the solubility of H_2 or H_2O or both in the CaMgSi₂O₆- H_2O - H_2 liquid with increasing pressure. The proposed pressure-induced structural changes in these liquids cannot be defined without more data on the solubility of the volatile components; this will be hard to acquire because these hydrous liquids do not quench to glass, but rather to an intergrowth of quench crystals and glass, complicating analysis of the run products. Experiments at temperatures well above the liquidus quench to glass, but temperature-induced structural degradation of the liquids renders the results less useful (Boettcher et al., 1982).

The temperature of the vapor-saturated solidus of NaAlSi₃O₈-H₂O-H₂ (IW) is consistently lower at all pressures than the solidus of NaAlSi₃O₈-H₂O-CO₂ (Fig. 7) or that of NaAlSi₃O₈-H₂O-X (Fig. 8), calculated using the data of Burnham (1979). The difference between the experimental and calculated solidi (Fig. 8) is ~175°C, much larger than the difference between the calculated and experimental solidi for NaAlSi₃O₈-H₂O (~30°C), suggesting that H₂ as well as H₂O dissolves in the liquid. Further, the difference between the solidus of NaAlSi₃O₈-H₂O-H₂ and that of NaAlSi₃O₈-H₂O-CO₂ (Fig. 7) suggests that the solubility of H₂ is higher in the liquid than is that of CO₂, or that the calculated values for $X'_{\rm H₂O}$ for the IW buffer are too low. Data on H₂O-CO₂ mixtures



Fig. 4. Pressure-temperature projection of the vapor-saturated solidi of quartz + vapor for various values of $a_{H_2}^z$. Closed symbols represent quartz + liquid + vapor; open symbols are quartz + vapor. Values of $a_{H_2}^z$ are buffered by pure H₂O, hematitemagnetite-H₂O (HM) and iron-wüstite-H₂O (IW). The α - β quartz transition is from Cohen and Klement (1967), and Bohlen and Boettcher (1982) determined the quartz = coesite transition to 1000°C, here extrapolated to higher temperatures. K₂ is a critical end-point in the system SiO₂-H₂O. K_A and K_B are critical points for SiO₂-H₂O-H₂.

suggest that activity coefficients in these vapors are approximately unity (Bohlen et al., 1982, 1983; Kerrick and Jacobs, 1981; Boettcher, 1984) and that they may be modeled, to a first approximation, by a variant of the MRK equation of state (Holloway, 1977; Kerrick and Jacobs, 1981). The data available on H₂O-H₂ mixtures (Shaw, 1963) are consistent with activity coefficients for both components on the order of unity. As discussed previously, the calculated $X_{H_{2O}}^{\nu}$ may be incorrect, but they would have to be too small, instead of too large, as would be required for the HM-buffered vapors. If these calculated values for $X_{H_{2O}}^{\nu}$ are correct, then the difference in the solidi must reflect the higher solubility of H₂ relative to CO₂ in the liquid.

The experimentally determined solidus of $SiO_2-H_2O-H_2$ (IW) lies consistently at lower temperature for a given pressure than does the corresponding solidus of $SiO_2-H_2O-CO_2$ (Fig. 9) or that of SiO_2-H_2O-X (Fig. 10), modified from Boettcher (1984), for the same $a_{H_2O}^y$. The difference in temperature increases with increasing pressure, imply-



Fig. 5. Pressure-temperature projection of experimentally determined and calculated vapor-saturated solidi of diopside + vapor. Note the difference between the calculated and experimental CaMgSi₂O₆-H₂O solidi. The solidus of CaMgSi₂O₆-H₂O-H₂ (IW) is shown with the values of aH₂O along the curve. The CaMgSi₂O₆-H₂O-X solidus was calculated assuming an inert component in the vapor and the same $a_{\rm H_2}^{\nu}$ as in the experimental curve at any given pressure.

ing a concomitant increase in the solubility of H₂. Compared to the NaAlSi₃O₈-containing systems, the SiO₂-containing systems show a greater difference in temperature between the solidus for SiO₂-H₂O-H₂ (IW) and the equivalent SiO₂-H₂O-CO₂ solidus, which may be a real difference in the solubility of H₂ or an artifact of comparing such different systems in this fashion. Nevertheless, the explanation offered above for the NaAlSi₃O₈-bearing systems, that H₂ is more soluble than is CO₂ at low $X_{H_2O}^{\nu}$, may also be appropriate for SiO₂-H₂O-H₂ and SiO₂-H₂O-CO₂.

The abrupt changes in slope of the vapor-saturated solidus of $SiO_2-H_2O-H_2$ (IW) may be explained by pressureinduced transformations in either the liquid or the vapor. However, we favor a mechanism involving a decrease in the free energy of the liquid caused by structural changes that increase the solubility of H_2O and/or H_2 . In any case, the abruptness of the changes in slope of the solidus suggest that there must be important structural changes in the liquid with increasing pressure, making the calculation of melting equilibria, even in this simple case, hazardous unless complemented by tightly constrained experimental determinations.

The difference in behavior of H_2 and CO_2 may also be seen in their respective effects on the critical end-point in SiO₂-H₂O. Dilution of the vapor phase with ~30 mol%



Fig. 6. Pressure-temperature projection of the solidus of $CaMgSi_2O_6-H_2O-H_2$ (IW) and that of $CaMgSi_2O_6-H_2O-CO_2$ with $X_{H_{2O}}^{*}$ fixed at 0.60. Values of $a_{H_{2O}}^{*}$ along the $CaMgSi_2O_6-H_2O-H_2$ (IW) solidus are shown by the numbers along the curve. The solidus of $CaMgSi_2O_6-H_2O-CO_2$ is from Eggler and Rosenhauer (1978).

 H_2 shifts the critical point from ~10 kbar to ~29 kbar, whereas only 5 mol% CO₂ increases the pressure of the critical point to ~26 kbar. This difference suggests that the liquid and vapor in SiO₂-H₂O change character less when diluted by H₂ than when diluted by CO₂. The solubility of both H₂O and H₂ in the liquid must be higher than the solubility of H₂O and CO₂ under the same conditions, and the solubility of SiO₂ in the vapor must be greater in a H₂O-H₂ vapor than in a H₂O-CO₂ vapor.

Our results reveal that experimental apparatus must be calibrated to determine their intrinsic $f_{\rm H_2}$. Such differences may explain why, for example, the vapor-saturated solidus for CaMgSi₂O₆-H₂O determined by Eggler and Rosenhauer (1978) in an internally heated pressure vessel lies between our solidus for pure-H₂O and that for HMbuffered vapor below 10-kbar pressure. Their internally heated pressure vessel had an argon pressure medium and an intrinsic f_0 , near the NNO buffer. If the argon contained some H_2O , the f_{H_2} would be lower than that for hydrous vapor in equilibrium with NNO, possibly lower than the $f_{\rm H_2}$ for vapors in equilibrium with HM, which would be consistent with our results. At higher pressures, Eggler and Rosenhauer (1978) used a solid-media high-pressure apparatus with intrinsic f_{O_2} slightly lower than the HM buffer (Eggler et al., 1974), and their solidus above 10 kbar is consistent with the vapor having an $f_{\rm H_2}$ between that of vapors buffered by HM and IW based on our results.

Bohlen et al. (1982) discussed discrepant data for vaporsaturated solidi in the system NaAlSi₃O₈-H₂O. Bohlen et



Fig. 7. Pressure-temperature projection of the solidus of NaAlSi₃O₈-H₂O-H₂ (IW) and the solidi of NaAlSi₃O₈-H₂O-CO₂ with various values of $X_{\rm H_2O}^*$. Values of $a_{\rm H_2O}^*$ along the NaAlSi₃O₈-H₂O-H₂ (IW) solidus are shown by the numbers along the curve. Each solidus in the system NaAlSi₃O₈-H₂O-CO₂ (Bohlen et al., 1982) has a fixed value of $X_{\rm H_2O}^*$ given by the number along the curve.

al. bracketed the solidus at 10 and 15 kbar at 670-680°C and 635-645°C, respectively. Boettcher and Wyllie (1969) had 10- and 15-kbar brackets at 680-690°C and 650-660°C, respectively. W. C. Luth et al. (1964) reported brackets of 10.3 kbar, 695°C and 10.0 kbar, 720°C. Morse (1970) bracketed the solidus at 5 kbar at 758 \pm 3°C in both cold-seal and internally heated apparatus. His results are dissonant with ours (Fig. 2) and with those obtained previously by us in piston-cylinder and gas-pressure apparatus (see Bohlen et al., 1982) and with the values of Tuttle and Bowen (1958) extrapolated from 4 kbar. The differences between these results are probably caused by differences in the $f_{\rm H_2}$ in the various apparatus. Bohlen et al. used an "anhydrous" solid-media pressure cell that would have the lowest $f_{\rm H_2}$ (Boettcher et al., 1981). Boettcher and Wyllie (1969) used a talc furnace assembly, with a higher $f_{\rm H_2}$, around NNO, resulting from the H₂O released during dehydration of the talc. The intrinsic f_0 , of both apparatus used by Morse was near the NNO buffer (Morse, pers. comm., 1984), and his results agree with our results for NaAlSi₃O₈-H₂O-H₂ (NNO) as well as with our results for NaAlSi₃O₈-H₂O-H₂ (HM).

Comparing our results for the system $SiO_2-H_2O-H_2$ with earlier work, our 9-kbar solidus for $SiO_2-H_2O-H_2$ (HM) is bracketed between the solidus of Kennedy et al. (1962) and that of Stewart (1967). Both Stewart and Kennedy et



Fig. 8. Pressure-temperature projection of experimentally determined and calculated vapor-saturated solidi of albite + vapor. The experimental curve is for NaAlSi₃O₈-H₂O-H₂ (IW), with the $a_{H_{2}O}^{*}$ indicated by the numbers along the solidus. The NaAlSi₃O₈-H₂O-X solidus is calculated assuming an inert diluent in the vapor and the same $a_{H_{2}O}^{*}$ as for the experimental curve at a given pressure.

al. used internally heated pressure vessels with argon pressure media, yet the solidus determined by Stewart is ~40°C lower than the solidus determined by Kennedy et al. Stewart used the internally heated pressure vessel at the Geophysical Laboratory in Washington, D.C., the same apparatus used by Eggler and Rosenhauer (1978) and Morse (1970), with a pressure medium of light-bulb-grade argon (Morse, pers. comm., 1984). The results of all of these studies suggest an $f_{\rm H_2}$ for this apparatus between that for a vapor of pure H₂O and that for a vapor buffered by HM. The higher results obtained by Kennedy et al. may be explained by a higher $f_{\rm H_2}$ in their apparatus generated by a higher $X_{\rm H_2O}^{\rm v}$ in the pressure medium. Conceivably, many inconsistent phase-equilibrium data may be reconciled when the $f_{\rm H_2}$ of the experimental apparatus is considered.

The effect of f_{H_2} in these melting experiments has many petrologic applications other than illustrating that f_{H_2} is an important factor in phase-equilibrium experiments. Recent measurements of intrinsic f_{O_2} in xenoliths from the upper mantle (Arculus and Delano, 1981) were interpreted as indicating that parts of the upper mantle have an f_{O_2} more reducing than IW. The presence of native iron in granulite xenoliths from a kimberlite is further evidence for conditions of low f_{O_2} in regions of the continental lower crust and upper mantle (Haggerty and Toft, 1985). Under these conditions, the melting of silicates in the presence of a hydrous vapor phase would occur at higher temper-

Fig. 9. Pressure-temperature projection of the solidus of SiO₂-H₂O-H₂ (IW) and the solidi of SiO₂-H₂O-CO₂ with various values of $X_{H_{2O}}^{\nu}$. Values of $a_{H_{2O}}^{\nu}$ along the SiO₂-H₂O-H₂ (IW) solidus are shown by the numbers along the curve. Each solidus in the system SiO₂-H₂O-CO₂ (Boettcher, 1984) has a fixed value of $X_{H_{2O}}^{\nu}$ given by the number along the curve.

atures than in the presence of pure H_2O , but at lower temperatures than in the presence of a H_2O-CO_2 vapor with the same $a_{H_2O}^v$. Further, under these reducing conditions, fluids in the upper mantle would be a mixture of H_2O and H_2 , carrying smaller amounts of dissolved silicates than would pure H_2O , thereby lessening the effect of metasomatizing fluids on changing the chemistry of the material with which they interact. Increasing $X_{H_2}^v$ may also affect partitioning of trace elements between solid and fluid phases, altering the trace-element characteristics of the metasomatized rocks.

The lower solubility of H_2 in silicate liquids relative to H_2O is also significant for models of degassing of the Earth and migration of H_2 in the mantle. H_2 is less likely to trigger melting than is H_2O , and, unlike H_2O , H_2 cannot react extensively with solid phases in the mantle, suggesting that H_2 will not become trapped in either a solid phase or a silicate liquid, and it may escape from the mantle (Welhan and Craig, 1979) or provide some control on the oxidation state of the mantle.

In the crust, most igneous rocks equilibrated at values of f_{0_2} near those of the QFM buffer. In the presence of a hydrous vapor buffered by QFM, melting would occur at temperatures as much as 80°C higher than the vaporsaturated solidus in the presence of pure H₂O. Furthermore, under such relatively reducing conditions in the crust, the low solubility of silicates in the vapor with relatively high values of $X_{H_2}^{\nu}$ will diminish the metasomatizing ability of the vapor. Zones of intense alteration

Fig. 10. Pressure-temperature projection of experimentally determined and calculated vapor-saturated solidi of quartz + vapor. The experimental curve is for SiO₂-H₂O-H₂ (IW), with the numbers along the solidus representing the $a_{H_{2O}}^{\nu}$. The SiO₂-H₂O-X solidus is calculated assuming an inert diluent in the vapor, with the same $a_{H_{2O}}^{\nu}$ as for the experimental curve at any given pressure.

probably occur only under very oxidizing conditions, and this is consonant with the mineralogy of many pegmatites, where fluid-phase metasomatism commonly is pervasive and extensive.

ACKNOWLEDGMENTS

Discussions with Ed Stolper and Bruce Bathurst have been most helpful. Reviews by Stolper and David Eggler are appreciated. Scott Boettcher, Steve Chambers, and Les Faus assisted with the conception and conductions of the experiments. We are indebted to J. M. Christie and R. Coleman for providing the quartz and albite. The National Science Foundation supported this research with grants EAR78-16413, EAR82-10695, and EAR83-06410 to Boettcher and with a Graduate Fellowship to Luth.

REFERENCES

- Aines, R.D., and Rossman, G.R. (1984a) The hydrous component in garnets: pyralspites. American Mineralogist, 69, 1116–1126.
 —— (1984b) Water in minerals? A peak in the infrared. Journal
- of Geophysical Research, 89, 4059–4071. Aines, R.D., Kirby, S.H., and Rossman, G.R. (1984) Hydrogen speciation in synthetic quartz. Physics and Chemistry of Minerals, 11, 204–212.
- Arculus, R.J., and Delano, J.W. (1981) Intrinsic oxygen fugacity measurements: techniques and results for spinels from upper mantle peridotites and megacryst assemblages. Geochimica et Cosmochimica Acta, 45, 899–913.
- Beran, A., and Putnis, A. (1983) A model of the OH positions in olivine, derived from infrared-spectroscopic investigations. Physics and Chemistry of Minerals, 9, 57–60.
- Boettcher, A.L. (1984) The system SiO₂-H₂O-CO₂: melting, solubility mechanisms of carbon, and liquid structure to high pressures. American Mineralogist, 69, 823–833.
- Boettcher, A.L., and Wyllie, P.J. (1968) Jadeite stability measured in the presence of silicate liquids in the system NaAlSiO₄-SiO₂-H₂O. Geochimcia et Cosmochimica Acta, 32, 999–1012.





(1969) Phase relationships in the system NaAlSiO₄-SiO₂- H_2O to 35 kilobars pressure. American Journal of Science, 267, 875–909.

- Boettcher, A.L., Windom, K.E., Bohlen, S.R., and Luth, R.W. (1981) Low friction, anhydrous, low- to high-temperature furnace sample assembly for piston-cylinder apparatus. Reviews of Scientific Instruments, 52, 1903–1904.
- Boettcher, A.L., Burnham, C. Wayne, Windom, K.E., and Bohlen, S.R. (1982) Liquids, glasses, and the melting of silicates to high pressures. Journal of Geology, 90, 127–138.
- Bohlen, S.R., and Boettcher, A.L. (1982) The quartz = coesite transformation: A precise determination and the effects of other components. Journal of Geophysical Research, 87, 7073–7078.
- Bohlen, S.R., Boettcher, A.L., and Wall, V.J. (1982) The system albite-H₂O-CO₂: A model for melting and activities of water at high pressures. American Mineralogist, 67, 451–462.
- Bohlen, S.R., Boettcher, A.L., Wall, V.J., and Clemens, J.D. (1983) Stability of phlogopite-quartz and sanidine-quartz: A model for melting in the lower crust. Contributions to Mineralogy and Petrology, 83, 270–277.
- Brey, G., and Green, D.H. (1977) Systematic study of liquidus phase relations in olivine melilitite + H₂O + CO₂ at high pressures and petrogenesis of an olivine melilitite magma. Contributions to Mineralogy and Petrology, 61, 141–162.
- Burnham, C. Wayne. (1979) The importance of volatile constituents. In H.S. Yoder, Jr., Ed., The evolution of the igneous rocks, 439-482. Princeton University Press, New Jersey.
- Burnham, C. Wayne, Holloway, J.R., and Davis, N.F. (1969) Thermodynamic properties of water to 1000°C and 10,000 bars. Geological Society of America Special Paper 132.
- Cohen, L.H., and Klement, W., Jr., (1967) High-low quartz inversion: Determination to 35 kilobars. Journal of Geophysical Research, 72, 4245–4251.
- Chou, I-M. (1978) Calibration of oxygen buffers at elevated P and T using the hydrogen fugacity sensor. American Mineralogist, 63, 690–703.
- de Santis, R., Breedveld, G.J.F., and Prausnitz, J.M. (1974) Thermodynamic properties of aqueous gas mixtures at advanced pressures. Industrial Engineering Chemistry, Process Design and Development, 13, 374–377.
- Ebisuzaki, Y., Kass, W.J., and O'Keefe, M. (1968) Solubility and diffusion of hydrogen and deuterium in platinium. Journal of Chemical Physics, 49, 3329–3332.
- Eggler, D.H. (1975) CO_2 as a volatile component of the mantle: the system Mg₂SiO₄-SiO₂-H₂O-CO₂. Physics and Chemistry of the Earth, 9, 869–881.
- ---- (1978) The effect of CO₂ upon partial melting of peridotite in the system Na₂O-CaO-Al₂O₃-MgO-SiO₂-CO₂ to 35 kb with an analysis of melting in a peridotite-H₂O-CO₂ system. American Journal of Science, 278, 305–343.
- Eggler, D.H., and Burnham, C. Wayne. (1984) Solution of H₂O in diopside melts: A thermodynamic model. Contributions to Mineralogy and Petrology, 85, 58–66.
- Eggler, D.H., and Kadik, A.A. (1979) The system NaAlSi₃O₈-H₂O-CO₂ to 20 kbar pressure: I. Compositional and thermodynamic relations of liquids and vapors coexisting with albite. American Mineralogist, 64, 1036–1048.
- Eggler, D.H., and Rosenhauer, M. (1978) Carbon dioxide in silicate melts: II. Solubilities of CO₂ and H₂O in CaMgSi₂O₆ (diopside) liquids and vapors at pressures to 40 kbar. American Journal of Science, 278, 64–94.
- Eggler, D.H., Mysen, B., and Hoering, T.C. (1974) Gas species in sealed capsules in solid media, high-pressure apparatus. Carnegie Institution of Washington Year Book 73, 228–232.
- Faile, S.P., and Roy, D.M. (1970) Mechanism of color center destruction in hydrogen impregnated radiation resistant glasses. Materials Research Bulletin, 5, 385–390.
- —— (1971) Dissolution of hydrogen in fused silica. American Ceramic Society Journal, 54, 533–534.

Faile, S.P., Schmidt, J.J., and Roy, D.M. (1967) Irradiation effects

in glasses: Suppression by synthesis under high-pressure hydrogen. Science, 156, 1593-1595.

- Gerlach, T.M. (1982) Interpretation of volcanic gas data from tholeiitic and alkaline mafic lavas. Bulletin Volcanologique, 45, 235-244.
- Gerlach, T.M., and Nordlie, B.E. (1975) The C-O-H-S gaseous system, Part I: Composition limits and trends in basaltic cases. American Journal of Science, 275, 353–376.
- Goldsmith, J.R., and Jenkins, D.M. (1984) The high-low albite relations revealed by hydrothermal melting and reversal of degree of order. (abs.) EOS (American Geophysical Union Transactions), 65, 307.
- Greenwood, H.J. (1975) Buffering of pore fluids by metamorphic reactions. American Journal of Science, 275, 573–593.
- Haggerty, S.E., and Toft, P.B. (1985) Native iron in the continental lower crust: Petrological and geophysical implications. Science, 229, 647–649.
- Hill, R.E.T., and Boettcher, A.L. (1970) Water in the Earth's mantle: melting curves of basalt-water and basalt-water-carbon dioxide. Science, 167, 980–982.
- Holloway, J.R. (1977) Fugacity and activity of molecular species in supercritical fluids. In D.G. Fraser, Ed., Thermodynamics in geology, 161–182. Reidel, Dordrecht, Holland.
- Huebner, J.S. (1971) Buffering techniques for hydrostatic systems at elevated pressures. In G.C. Ulmer, Ed., Research techniques for high pressure and high temperature, 123–177. Springer-Verlag, New York.
- Kennedy, G.C. (1948) Equilibrium between volatiles and iron oxides in rocks. American Journal of Science, 246, 529–549.
- Kennedy, G.C., Wasserburg, G.J., Heard, H.C., and Newton, R.C. (1962) The upper three-phase region in the system SiO₂-H₂O. American Journal of Science, 260, 501–521.
- Kerrick, D.M., and Jacobs, G.K. (1981) A modified Redlich-Kwong equation of state for H₂O, CO₂, and H₂O-CO₂ mixtures at elevated pressures and temperatures. American Journal of Science, 281, 735–767.
- Luth, W.C., Jahns, R.H., and Tuttle, O.F. (1964) The granite system at pressures of 4 to 10 kilobars. Journal of Geophysical Research, 69, 759–773.
- Morse, S.A. (1970) Alkali feldspars with water at 5 kb. Journal of Petrology, 11, 221-251.
- Mysen, B.O., and Boettcher, A.L. (1975) Melting of a hydrous mantle: I. Phase relations of natural peridotite at high pressures and temperatures with controlled activities of water, carbon dioxide, and hydrogen. Journal of Petrology, 16, 520–548.
- Mysen, B.O., and Virgo, D. (1980) Solubility mechanisms of carbon dioxide in silicate melts: A Raman spectroscopic study. American Mineralogist, 65, 885–899.
- Nakamura, Y. (1974) The system SiO₂-H₂O-H₂ at 15 kbar. Carnegie Institution of Washington Year Book 73, 259–263.
- Presnall, D.C. (1969) Pressure-volume-temperature measurements on hydrogen from 200°C to 600°C and up to 1800 atmospheres. Journal of Geophysical Research, 74, 6026–6033.
- Robie, R.A., Hemingway, B.S., and Fisher, J.D. (1979) Thermodynamic properties of minerals and related substances at 298.15 K (25°C) and one bar (10 pascals) pressure and at higher temperatures (revised 1979). U.S. Geological Survey Bulletin 1452.
- Sack, R.O., Carmichael, I.S.E., Rivers, M., and Ghiorso, M.S. (1980) Ferric-ferrous equilibria in natural silicate liquids at 1 bar. Contributions to Mineralogy and Petrology, 75, 369–377.
- Shaw, H.R. (1963) Hydrogen-water vapor mixtures: Control of hydrothermal atmospheres by hydrogen osmosis. Science, 139, 1220–1222.
- Shaw, H.R., and Wones, D.R. (1964) Fugacity coefficients for hydrogen gas between 0° and 1000°C, for pressures to 300 atm. American Journal of Science, 262, 918–929.
- Shettel, D.C. (1973) Solubility of quartz in H₂O-CO₂ fluids at 5 kb and 500°-900°C (abs.) EOS American Geophysical Union Transactions, 54, 480.
- Stewart, D.B. (1967) Four-phase curve in the system CaAl₂Si₂O₈-

 SiO_2 -H₂O between 1 and 10 kilobars. Schweizerische Mineralogische und Petrographische Mitteilungen, 47, 35–59.

- Stolper, E. (1982) The speciation of water in silicate melts. Geochimica et Cosmochimica Acta, 46, 2609-2620.
- Tuttle, O.F., and Bowen, N.L. (1958) Origin of granite in the light of experimental studies in the system NaAlSi₃O₈-KAlSi₃O₈-SiO₂-H₂O. Geological Society of America Memoir 74.
- Welhan, J.A., and Craig, H. (1979) Methane and hydrogen in East Pacific Rise hydrothermal fluids. Geophysical Research Letters, 6, 829-831.
- Wyllie, P.J. (1979) Magmas and volatile components. American Mineralogist, 64, 469–500.

MANUSCRIPT RECEIVED APRIL 8, 1985 MANUSCRIPT ACCEPTED NOVEMBER 4, 1985