Artroeite, PbAlF₃(OH)₂, a new mineral from the Grand Reef mine, Graham County, Arizona: Description and crystal structure

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ABSTRACT

Artroeite, PbAlF₃(OH)₂, space group $P\overline{1}$, a=6.270(2), b=6.821(3), c=5.057(2) Å, $\alpha=90.68(2)$, $\beta=107.69(2)$, $\gamma=104.46(2)^{\circ}$, V=198.6(2) Å³, Z=2, is a new mineral from the Grand Reef mine, near Klondyke, Graham County, Arizona. It occurs as colorless bladed crystals associated with quartz, fluorite, galena, anglesite, and an as yet undescribed mineral of composition PbCa₂Al(F,OH)₉. Artroeite has a Mohs hardness of about 2.5, a measured density of 5.36(2) g/cm³, and a calculated density of 5.47 g/cm³. It exhibits a perfect {100} cleavage and a good {010} cleavage. Optically it is biaxial (-) with $\alpha=1.629(1)$, $\beta=1.682(2)$, and $\gamma=1.691(2)$. Dispersion is strong, $r>\nu$. The six strongest powder diffraction lines are [d(I,hkl)] 4.42₁₀₀($\overline{101}$), 3.221₄₀(101), 2.595₇₀($\overline{121}$,021), 2.190₆₅(201,012,030), 2.030₅₀(022), 2.015₄₀($\overline{230}$) Å.

The structure was solved by direct methods and refined to R = 0.022 using X-ray diffractometer data (1096 independent reflections). In the structure, edge-sharing dimers of AlF₃(OH)₃ octahedra link together by bonds to Pb atoms to form approximately close-packed layers parallel to (101). The layers are linked to one another by one Pb-F bond and two H bonds per formula unit. Pb is coordinated to 6 F and 3 O atoms. Three F ligands are associated with the same Al octahedral face and correspond to much longer bonds. It is probable that the lone pair electrons are situated on that side of the Pb atom. The structure is compared with those of acuminite and tikhonenkovite.

INTRODUCTION

Kampf et al. (1989) described the new minerals grand-reefite, Pb₂SO₄F₂, pseudograndreefite, Pb₆SO₄F₁₀, laure-lite, Pb(F,Cl,OH)₂, and aravaipaite, Pb₃AlF₉·H₂O, from a single vug found in the oxidized zone of the Grand Reef mine, an epithermal Cu-Pb-Ag deposit in the Aravaipa mining district of Graham County, Arizona. Acidic supergene sulfate-rich solutions are thought to have dissociated Pb and F from layers of earlier-formed galena and fluorite that envelop the vug. The layers of galena and fluorite apparently also served as a barrier preventing Cu²⁺ from entering the vug.

A specimen containing a similar vug, collected in the same portion of the mine by Michael Shannon in 1981 and set aside as an unknown, was provided to the authors in 1992 by David Shannon. Although it did not contain any of the previously described minerals, it yielded two more new Pb-F species, one of which is described herein, including a determination of its crystal structure. The other new mineral is still under study.

The new mineral is named artroeite in honor of Arthur Roe (1912–1993) of Tucson, Arizona. Art Roe was an avid collector of microscopic minerals (micromounts) and a prolific author and lecturer to the amateur mineralogi-

cal community. He provided volunteer curatorial assistance to the mineral collections of the U.S. National Museum of Natural History, the Arizona-Sonora Desert Museum, and the University of Arizona. Art Roe served as professor and as chairman in the Chemistry Department at the University of North Carolina (1939–1959) and was the first director of the Office of International Science at the National Science Foundation. The mineral and name were approved by the Commission on New Minerals and Mineral Names, International Mineralogical Association. The holotype specimen is in the collection of the Natural History Museum of Los Angeles County (Cat. no. 39338).

DESCRIPTION

Occurrence and paragenesis

Artroeite is found on a single specimen in a 15×5 mm quartz-lined vug in association with anglesite and another new mineral of composition PbCa₂Al(F,OH)₉. The second new mineral occurs as colorless tabular crystals, as a massive cavity filling, and as inclusions up to $10 \mu m$ in length in artroeite crystals. An intergrowth of fluorite and galena completely envelopes the vug. The galena is partially altered to anglesite. Linarite and mus-

TABLE 1. Descriptive characteristics of artroeite

	(Chemical data			
PbO Al ₂ O ₃ F H ₂ O O ≡ F Total	67.5 (67.0–68.0)* 15.6 (15.5–15.7)* 16.1 (15.8–16.5)* 6.0 6.8 98.4				
		Cell data**			
Tricline, I PbAIF ₃ (O V = 198.0	$H)_2; Z = 2$	$a = 6.270(2) \text{ Å}$ $\alpha = 90.68(2)^{\circ}$ $b = 6.821(3) \text{ Å}$ $\beta = 107.69(2)^{\circ}$ $c = 5.057(2) \text{ Å}$ $\gamma = 104.46(2)^{\circ}$			
		Morphology			
Maximum Forms: {	n crystal size: 1×0 100}, {010}, {001},	and flattened on $\{10\overline{1}\}$.7 \times 0.04 mm $\{10\overline{1}\}$, $\{011\}$, $\{11\overline{1}\}$, $\{121\}$ ation on [010] with C.P. $\{100\}$			
	Phy	ysical properties			
Color: co Streak: v	vhite itreous	cleavage: (100) perfect; (010) good tenacity: brittle fracture: choncoidal density: 5.36(2) _{meas} †; 5.43 _{calc} g/cm³			
Luster: vi Hardness					
	Or	otical properties			

- ** Four-circle data, 25 reflections ($2\theta = 18-22^{\circ}$).
- † Berman balance, toluene at 25 °C.

‡ Spindle stage.

covite are only present outside of the fluorite-galena envelope. Artroeite probably resulted from the reaction of acidic supergene solutions with galena and fluorite and was the last mineral to form in this assemblage.

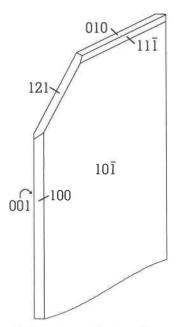


Fig. 1. Orthographic projection of an artroeite crystal.

TABLE 2. X-ray powder diffraction data for artroeite

l _{obs}	d _{obs}	d _{calc} *	I _{calc}	hki	lobs	d _{obs}	d _{calc} *	I _{calc}	hki
						- 000		9	212
15	6.57	6.575	13	010			2.033	24	022
10	5.74	5.759	12	100	50	2.030	2.032		131
30	5.04	5.048	32	110			2.022	8 12	230
15	4.79	4.796	16	001 101			(2.022	25	122
100	4.42	4.442	100 50	011	40	2.015	2.013	9	112
35	4.05	4.063	37	111			1.934	11	211
20	3.88	3.898			30	1.932	,	10	220
20	3.84	3.854	46	110 011			1.927	8	320
20	3.70	3.710	24		15	1.889	∫1.891	5	130
5	3.480	3.496 (3.288	5 29	111 020	10	1.826	1.888	6	112
35	3.267	3.251	40	120	10	1.020	1.810	13	311
40	3,221	3.218		101	25	1.808	1.803	7	302
40	3.170	3.173	50 6	111	10	1.746	1.748	6	222
<5 30	2.940	2.940	31	210	<5	1.717	1.746	5	322
30	2.940	2.940	19	201	20	1.707	1.705	10	140
30	2.880	2.879	29	200	10	1.670	1.670	8	311
<5	2.674	2.675	3	111	10	1.070	1.639	4	132
< 5	2.074	(2.599	30	021	10	1.639	1.635	7	203
70	2.595	2.599	24	121			(1.609	5	202
		(2.515	14	102	15	1.605	1.602	5	041
30	2.511	2.506	22	121	10	1.005	1.599	6	003
		(2.474	12	211			(1.596	5	231
20	2.454	2.450	22	221	10	1.590	1.592	4	122
		2.412	9	210	10	1.550	1.589	3	013
20	2.404	2.398	13	002			(1.513	4	123
20	2.334	2.336	16	112	10	1.515	1.512	4	041
10	2.245	2.242	7	211			1.501	4	241
10	2.243	(2.192	18	030	10	1.497	1.495	5	331
		2.187	9	012			1.456	4	341
65	2.190	2.183	15	212	15	1.456	1.452	4	232
		2.182	6	201	5	1,428	1.426	4	431
<5	2.082	2.089	8	121	5	1.420	(1.393	3	142
< 5	2.002	(2.071	15	311			1.389	4	123
25	2.068	2.066	10	221	25	1.388	1.387	4	023
20	2.000	2.064	7	131			1.386	5	231
		(2.004		101			(1,300		201

Note: 114.6-mm Gandolfi camera, CuKα (Ni-filtered) radiation, observed intensities visually estimated, calculated intensities obtained from structure data.

* Calculated from the cell parameters given in Table 1.

Mineral data

A drawing of a single crystal of artroeite is provided as Figure 1, and the descriptive characteristics are provided in Table 1. X-ray powder diffraction data are given in Table 2.

CRYSTAL STRUCTURE DETERMINATION

Most details of the data collection and refinement are given in Table 3. Data were corrected for Lorentz and

TABLE 3. Details of data collection and refinement for artroeite

Diffractometer: Huber	$\mu = 429.85 \text{ cm}^{-1}$
Radiation/monochromator:	Absorption correction by
Mo <i>Kα</i> /Gr	Gaussian integration
$2\theta_{\text{max}} = 60^{\circ}$	Face (distance in mm):
Scan type: $\theta/2\theta$	001(0.040), 100(0.055),
Scan width: 3.0°	101(0.015), 101(0.015),
Scan duration: 3°/min	010(0.030), 010(0.030),
Standards: 3/100 ref.	121(0.040), 101(0.055)
Total reflections: 2310	$R_{\rm abs}^* = 0.100, R_{\rm w abs} = 0.119$
Unique reflections: 1155	$R_{\rm iso}^{**} = 0.033, R_{\rm wiso} = 0.043$
$R_{\text{merge}} = 0.033$	$R = 0.022, R_w = 0.031$
No. data $[F_0 > 6\sigma_F]$: = 1096	-

Prior to absorption correction.

^{**} Based upon isotropic displacement parameters.

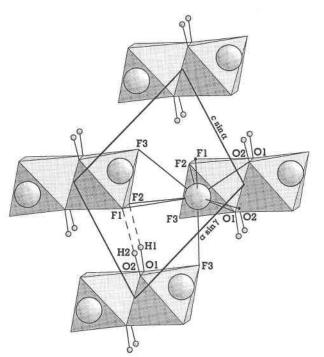


Fig. 2. The atomic structure of artroeite projected down [010] showing close-packed layers viewed on edge. Large circles are Pb atoms; small circles are H atoms; black dots indicate anions in octahedra above those shown.

polarization factors. Computations were carried out with the crystallographic computing package from the University of California, Los Angeles. Scattering factors for neutral atoms and anomalous dispersion coefficients were taken from *International Tables for X-ray Crystallography* (Ibers and Hamilton, 1974). The structure was solved using direct methods.

All atoms are in general positions, and all were assigned full occupancy. F and O were easily discerned on the basis of distances from Al, isotropic displacement factors, and bond valence sums. There was no conclusive evidence for assigning partial O occupancy to any F site, as was suggested by the chemical analysis. H atoms were located from analysis of difference-Fourier maps. The scale factor and the positional and anisotropic displacement parameters of all atoms except H were varied in the final cycle of refinement. Final positional and displacement

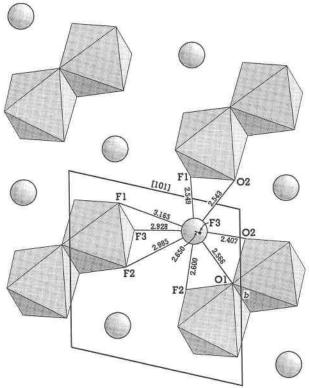


Fig. 3. An approximately close-packed layer in artroeite oriented parallel to (101) showing bond distances (Å) to one Pb. Large circles are Pb atoms; black dot indicates F3 in layer above that shown.

parameters are given in Table 4, selected interatomic distances and angles are given in Table 5, and observed and calculated structure factors are listed in Table 6.1

DISCUSSION OF THE STRUCTURE

Structure description

The artroeite structure is depicted in Figures 2 and 3. It contains AIF₃(OH)₃ octahedra that share an O1-O1'

TABLE 4. Positional and displacement parameters for artroeite

	×	y	Z	U _{eq} *	U ₁₁	U ₂₂	U_{33}	U_{12}	U ₁₃	U_{23}
Pb	0.30975(4)	0.19142(3)	0.20352(5)	0.00960(5)	0.00866(8)	0.00607(6)	0.0140(1)	0.00235(5)	0.00339(7)	0.00156(5)
Al	0.8110(3)	0.3273(3)	0.8364(4)	0.0077(3)	0.0085(5)	0.0047(4)	0.0100(7)	0.0016(3)	0.0031(5)	0.0012(4)
F1	0.7950(8)	0.1378(6)	0.5751(9)	0.0130(8)	0.017(1)	0.0069(8)	0.015(2)	0.0030(8)	0.009(1)	0.0003(9)
F2	0.7407(7)	0.4891(6)	0.5627(8)	0.0104(8)	0.012(1)	0.0081(8)	0.011(2)	0.0043(8)	0.002(1)	0.0028(9)
F3	0.5137(8)	0.2308(7)	0.8113(10)	0.0142(8)	0.009(1)	0.0105(9)	0.023(2)	0.0001(8)	0.006(1)	0.002(1)
01	0.8629(8)	0.5514(7)	0.0981(10)	0.0094(8)	0.008(1)	0.0073(9)	0.013(2)	0.0027(8)	0.005(1)	0.001(1)
02	0.8981(8)	0.1432(7)	0.0988(9)	0.0082(8)	0.010(1)	0.0076(9)	0.007(2)	0.0030(8)	0.003(1)	0.0021(9)
H1	0.807	0.509	0.253	0.02	()		` ,	()	. ,	• • •
H2	0.862	0.122	0.255	0.02						

^{*} $U_{eq} = (U_{11} + U_{22} + U_{33})/3.$

¹ A copy of Table 6 may be ordered as Document AM-95-573 from the Business Office, Mineralogical Society of America, 1130 Seventeenth Street NW, Suite 330, Washington, DC 20036, U.S.A. Please remit \$5.00 in advance for the microfiche.

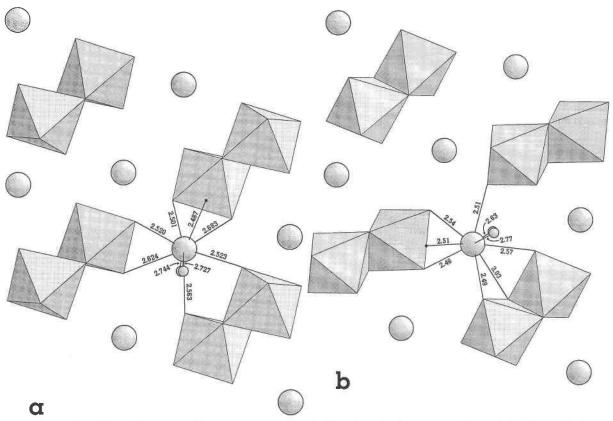


Fig. 4. Puckered layers in (a) acuminite parallel to (010) and (b) tikhonenkovite perpendicular to [100]. Large circles are Sr atoms; small circles are O atoms of H₂O; black dots indicate anions in octahedra above those shown. Bond distances (Å) to one Sr are shown.

edge to form $[Al_2F_6(OH)_4]^{4-}$ dimers. The structure can be visualized as consisting of approximately close-packed layers oriented parallel to (10 $\overline{1}$), composed of Pb atoms and octahedral dimers that are linked by Pb-F and Pb-O bonds. The layers are joined in the third dimension by one Pb-F bond and two H bonds per formula unit.

Following the scheme of Hawthorne (1984), artroeite is classified as a simple aluminofluoride based on a finite octahedral cluster, the aforementioned edge-sharing dimer. Artroeite is most closely related to the SrAlF₄(OH)·H₂O dimorphs, acuminite (Andersen and Ploug-Sørensen, 1991) and tikhonenkovite (Pudovkina and Pyatenko, 1967). In all three of these structures, isolated edge-sharing dimers are linked by bonds to large ninefold-coordinated divalent cations. The layers parallel to (101) in artroeite (Fig. 3), parallel to (010) in acuminite (Fig. 4a), and

TABLE 5. Selected bond lengths (Å) and angles (°) for artroeite

Pb-O2	2.407(5)	Al-F3	1.779(5)	O1-H1	0.972
-02	2.543(5)	-F2	1.799(4)	H1···F2	1.736
-F1	2.549(4)	-F1	1.800(4)	O1-F2	2.691(6)
-01	2.566(5)	-02	1.892(5)	01-H1-F2	166.9
-F2	2.600(4)	-01	1.907(5)		
-F3	2.650(5)	-01	1.927(5)	O2-H2	0.890
-F3	2.928(5)	(AI-F)	1.793	H2···F1	1.800
-F2	2.985(4)	(Al-O)	1.909	O2-F1	2.676(6)
-F1	3.165(5)			O2-H2-F1	167.6

perpendicular to [100] in tikhonenkovite (Fig. 4b) are remarkably similar in projection; however, in acuminite and tikhonenkovite the layers are markedly puckered.

Al octahedron

In artroeite the three F atoms and three O atoms in each AlF₃(OH)₃ octahedron occupy opposite faces. The Al-F bonds range from 1.779 to 1.800 Å, and the Al-O bonds range from 1.892 to 1.927 Å (Table 5). The octahedral angles range from 79.9 to 97.1°. The Al octahedron is somewhat less regular than those in acuminite and tikhonenkovite in that the 1.779-Å Al-F bond is shorter than any in either of the latter structures and the 1.927 Å Al-O bond is longer. The Al octahedra in acuminite and tikhonenkovite only have two O ligands, but in all three structures the vertices of the shared edges are occupied by O atoms, and the corresponding O-Al-O angles are the smallest (79.9° in artroeite, 79.8° in acuminite, and 78.5° in tikhonenkovite), the result of the Al³⁺-Al³⁺ cation repulsion effect.

Pb coordination and the lone pair effect

In artroeite Pb is coordinated by 6 F and 3 O atoms. The Pb-O bonds and three of the Pb-F bonds are within normal limits, ranging from 2.407 to 2.566 Å and 2.549 to 2.650 Å, respectively. The other three Pb-F bonds are

much longer, ranging from 2.928 to 3.165 Å (Table 5). The longer Pb-F bonds are all on the same side of the Pb atom, and the F atoms participating in these bonds are all contained within the same Al octahedral face. The lopsided coordination of Pb is ascribed to the lone pair effect.

Figures 3 and 4 clearly depict the Pb and Sr coordinations in artroeite, acuminite, and tikhonenkovite. The Sr coordinations in acuminite and tikhonenkovite differ from the Pb coordination in artroeite in three important ways: (1) each Sr bonds to two H₂O groups in lieu of bonding to two octahedral ligands; (2) the Sr coordination polyhedra do not share faces with Al octahedra; and (3) the Sr coordinations are not lopsided, the one long Sr-O bond in tikhonenkovite notwithstanding.

Moore et al. (1993) showed that, when the bonds between the lone pair cation and the anion are disregarded, the relative electrostatic undersaturations of the anions surrounding the lone pair cation predict whether the lone pair effect will be manifested in Pb- ϕ distances and determine in which direction the lone pair cation will shift, relative to the center of its anion coordination polyhedron. The analysis of the Pb coordination sphere in artroeite by this approach is complicated by the fact that four of the five distinct anions form two bonds, each of a different length, to Pb. Each of the three F atoms forms one normal-length and one long bond to Pb. Nevertheless, when the bond valences of the anions, listed in Table 7, are considered, after deducing the bond valence contribution by Pb, it is seen that O2, by far the most undersaturated anion, forms the two shortest bonds to Pb. Furthermore, it is seen in Figure 2 that both Pb-O2 bonds are roughly on the same side of the Pb atom and the three long Pb-F bonds are on approximately the opposite side of the Pb atom. The Pb can, therefore, be visualized as having shifted off the center position of its coordination polyhedron in the direction of O2 in response to the high degree of undersaturation of that anion. The lone pair electrons then are localized on the opposite side of the Pb atom in the direction of the long Pb-F bonds.

Although a variety of factors, e.g., temperature, solution chemistry, and activity of H₂O, may serve to determine whether the artroeite, acuminite, or tikhonenkovite structure is preferred, the presence of lone pair electrons on the large cation may be crucial to the stability of the artroeite structure.

TABLE 7. Bond valences (vu) in artroeite

	F1	F2	F3	01	02	Total
Pb	{ 0.25 0.05	0.21 0.08	0.19 0.09	0.29	{ 0.45 } 0.31 }	1.92
Al	0.50	0.50	0.53	0.50	0.52	3.02
H1		0.17		0.83		1
H2	0.17				0.83	1
Total	0.97	0.96	0.81	2.09	2.11	

Note: based upon Brese and O'Keeffe (1991), except for H bonds, which are based on Brown and Altermatt (1985) with O-H 0.95 Å.

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REFERENCES CITED

Andersen, E.K., and Ploug-Sørensen, G. (1991) The structure of acuminite, a strontium aluminum fluoride mineral. Zeitschrift für Kristallographie, 194, 221–227.

Brese, N.E., and O'Keeffe, M. (1991) Bond-valence parameters for solids. Acta Crystallographica, B47, 192–197.

Brown, I.D., and Altermatt, D. (1985) Bond-valence parameters obtained from a systematic analysis of the inorganic structure database. Acta Crystallographica, B41, 244–247.

Hawthorne, F.C. (1984) The crystal structure of stenonite and the classification of the aluminofluoride minerals. Canadian Mineralogist, 22, 245–251.

Ibers, J.A., and Hamilton, W.C., Eds. (1974) International tables for X-ray crystallography, vol. IV, 366 p. Kynoch, Birmingham, U.K.

Kampf, A.R., Dunn, P.J., and Foord, E.E. (1989) Grandreefite, pseudo-grandreefite, laurelite, and aravaipaite: Four new minerals from the Grand Reef mine, Graham County, Arizona. American Mineralogist, 74, 927-933.

Moore, P.B., Davis, A.M., Van Derveer, D.G., and Sen Gupta, P.K. (1993) Joesmithite, a plumbous amphibole revisited and comments on bond valences. Mineralogy and Petrology, 48, 97–113.

Pudovkina, Z.V., and Pyatenko, Yu.A. (1967) Crystal structure of tikhonenkovite, Sr₂[Al₂F₈(OH)₂]·2H₂O. Doklady Akademii Nauk SSSR, Earth Sciences Section, 174, 117–120.

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