MÖSSBAUER SPECTROSCOPY OF THE Ag-Au CHALCOGENIDES PETZITE, FISCHESSERITE AND UYTENBOGAARDTITE

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ABSTRACT

Mössbauer spectra of the 77.3 keV γ rays of ¹⁹⁷Au were measured at 4.2 K for natural and synthetic Ag₃AuTe₂ (petzite), synthetic Ag₃AuSe₂ (fischesserite) and synthetic Ag₃AuS₂ (uytenbogaardtite). All compounds studied exhibit large gradients in electric field at the gold nuclei, notably the largest so far found in any gold mineral. The isomer shifts and electric quadrupole interactions, and in particular the similarity of these parameters for Ag₃AuS₂ with those of Au₂S, suggest that the gold in the Ag₃AuX₂ compounds should be considered as monovalent.

Keywords: refractory Au minerals, invisible Au, structurally bound Au, Ag-Au chalcogenides, petzite, fischesserite, uytenbogaardtite, ¹⁹⁷Au Mössbauer spectroscopy, isomer shift, electric quadrupole interaction, Hollinger mine, Timmins, Hemlo deposit, Ontario.

SOMMAIRE

Nous avons étudié les spectres de Mössbauer des rayons γ (77.3 keV) de l'isotope ¹⁹⁷Au générés à 4.2 K et mesurés sur des échantillons naturels et synthétiques de Ag₃AuTe₂ (petzite), et des échantillons synthétiques de Ag₃AuSe₂ (fischesserite) et Ag₃AuS₂ (uytenbogaardtite). Tous ces composés font preuve d'un gradient intense dans le champ électrique autour du nucléus des atomes d'or, et en fait le plus intense qui ait été découvert dans une espèce aurifère. D'après les déplacements isomères et les interactions électriques quadrupolaires, et en particulier la similarité de ces paramètres dans le Ag₃AuS₂ avec ceux de Au₂S, l'or dans ces composés Ag₃AuX₂ serait monovalent.

(Traduit par la Rédaction)

Mots-clés: minéraux réfractaires de l'or, or invisible, or incorporé dans le réseau, chalcogénures de Ag-Au, petzite, fischesserite, uytenbogaardtite, spectroscopie Mössbauer ¹⁹⁷Au, déplacement isomère, interaction électrique quadrupolaire, mine de Hollinger, Timmins, gisement de Hemlo, Ontario.

INTRODUCTION

Gold occurs in nature as well-defined minerals (native gold, *i.e.*, metallic gold or gold alloys with silver, and compounds, mainly with Te, Bi or Sb), and as a dilute impurity within sulfides (mainly arsenopyrite and pyrite) (Wilson 1982, Harris 1990). Because gold locked up within sulfides commonly is refractory to conventional techniques of extraction (cyanidization), a better understanding of the mineralogical characteristics of gold and gold-bearing minerals is of great practical significance. Mössbauer spectroscopy can be used to characterize the nature of gold in such materials (Wagner *et al.* 1986, 1988, 1989, Friedl *et al.* 1991). In this context, it is of interest to understand the Mössbauer spectra of gold in well-defined minerals of gold. As part of a systematic survey of gold minerals by ¹⁹⁷Au Mössbauer spectroscopy, in the present work we report on a study of petzite, fischesserite and uytenbogaardtite. These minerals are silver-gold chalcogenides with the general formula Ag_3AuX_2 , where X stands for Te, Se, and S, respectively (Prince 1988a, b, c).

Petzite (Ag_3AuTe_2) is the widely occurring natural silver-gold telluride known since 1845. At temperatures below 320°C, it has a cubic structure (Frueh 1959, Cabri 1965, Chamid et al. 1978), comparable with that of garnet. The space group is $I4_132$. The rare mineral fischesserite (Ag₃AuSe₂) was first described in 1971 (Johan et al. 1971); below 270°C, it has a cubic I432 crystal structure. similar to that of petzite. The very rarely occurring mineral uytenbogaardtite (Ag₃AuS₂) is an analog of petzite and fischesserite. Messien et al. (1966) reported the low-temperature form of Ag₃AuS₂ (stable below 185°C) to be cubic, but Graf (1968), Smit et al. (1970) and Barton et al. (1978) later demonstrated that the symmetry of the roomtemperature modification is actually tetragonal, $P4_{1}22$ or $P4_{1}$. The observed reflections show that the space group is primitive and that the only possible elements of non-unit translational symmetry are a 4_1 or a 4_2 axis. Despite these differences, the crystal structures of the three minerals are closely related. The packing of anions is approximately body-centered: gold is coordinated by two, silver by four chalcogen atoms distributed in a tetrahedron. Gold atoms are located on the 1/8, 1/8, 1/8, etc., positions in the cubic (or almost cubic) unit-cell, which contains eight gold atoms. Each gold atom has two close Te, Se or S neighbors ("dumbbell") located along the body diagonal, and six Ag atoms as next-nearest neighbors. Because of this linear two-fold coordination, the environment of gold in these minerals is very different from that, for instance, of the gold-silver ditellurides, in which there is a six-fold coordination of Te atoms to the gold. Also, the distance between Au and its nearest neighbors, 2.61 Å in petzite (Chamid et al. 1978), is markedly shorter than any bond length in the gold ditellurides (Tunell & Pauling 1952, Pertlik 1984a, b, c).

The information about the chemical state of gold obtained by Mössbauer spectroscopy stems mainly from the electric quadrupole interaction and from the isomer shift of the Mössbauer pattern. The former splits the Mössbauer line of ¹⁹⁷Au into a doublet whose separation, the quadrupole splitting (QS), depends on both the symmetry and the degree of covalency of the bonds between the gold atoms and their ligands. The isomer shift (IS) is a measure of the electron density at the gold nuclei. Both the electron density at the gold nuclei and the isomer

shift increase with the valence state of gold and with the covalency of the bonds between the gold and its ligands. For the linear, two-fold-coordinated Au⁺ compounds as well as for the planar four-fold-coordinated Au³⁺ compounds, distinctive approximately linear correlations between the electric quadrupole splitting and the isomer shift have been found (Bartunik et al. 1970, Faltens & Shirley 1970). Using these correlations, one can determine the valence state of gold in such compounds from the observed values of the isomer shift and the quadrupole splitting. For bonding situations that deviate from the two-fold linear or four-fold planar coordination, however, a distinction between Au⁺ and Au³⁺ on the basis of Mössbauer spectroscopy may be difficult, the more so since one normally cannot determine the sign of the electric quadrupole interaction, which is expected to be negative for linear Au⁺ and positive for planar Au³⁺ compounds (Parish 1982, 1984).

EXPERIMENTAL

Mössbauer spectra of petzite have been measured for two natural and three synthetic specimens. One of the natural specimens is a 5-mm-thick slab of rock from the Golden Spectre property, Hemlo, Ontario. The specimen contains small inclusions of petzite, intergrown with hessite (Ag₂Te) and chalcopyrite (CuFeS₂), highly disseminated in quartz and K-feldspar. It also contains a large amount of native gold. The second natural specimen consists of single-crystal pieces with a total volume of about 10 mm³, which were extracted from ore of the Hollinger mine, Timmins, Ontario. This specimen was obtained from the Royal Ontario Museum, where it is registered as M13740. It contains a very small amount of visible native gold. Unfortunately, we could not obtain natural specimens containing fischesserite and uytenbogaardtite in quantities measurable by Mössbauer spectroscopy.

A synthetic specimen of petzite was obtained from the Royal Ontario Museum, where it is registered as R395. The specimens produced in the course of this work were made by fusing the elements in the required stoichiometric amounts in evacuated and sealed quartz tubes. In the case of Ag_3AuTe_2 and Ag_3AuSe_2 , the charges were melted at 1000°C for 12 hours. After grinding and sealing again, the specimens were annealed at 900°C for 24 hours and then cooled from 900°C to 700°C for 6 days. Two specimens of Ag₃AuS₂ were prepared by heating to 1200°C for 12 hours. The first of them was prepared with the exact stoichiometric proportions of the elements, whereas in the preparation of the second one, excess sulfur was included. The temperature of the synthesis seems

TABLE 1. UNIT-CELL DATA OBTAINED FOR THE AG3AUX2 (X=Te, Se, S) COMPOUNDS STUDIED IN THE PRESENT WORK*

COMPOUND	SPACE GROUP	a (Å)	c (Å)	V (Å ³)	REFERENCE	
AggAuTe ₂		10.3850(4)	•	1120	Chamid <i>et al.</i> (1978)	
	cubic	10.338(4)	-	1105	this work, R395	
AggAuSez	1432	9.967(3)		990	Johan et al. (1971)	
	cubic	9.974(1)	·	992	this work, synth. 1000°C	
AggAuSg	P 4122 or P 41	9.68	9.81	919	Barton et al. (1978)*	
	tetragonal	9.71(1)	9.80(4)	925	this work, synth. 500°C	

*The references refer to the recent crystallographic work on the respective minerals. *The cell edges of one of two specimens (Tambang Sawah) are quoted; experimental errors were not given in the reference.

to have been too high, as in both cases a significant amount of sulfur did not react. Thus, the third specimen of Ag₃AuS₂ was prepared at 500°C for 100 hours and then cooled to room temperature during 20 hours. This procedure also was used to prepare a sample of petzite. The chemical composition of the specimens was determined by electronmicroprobe analysis, which showed significant fractions of native gold in petzite and in the uytenbogaardtite synthesized at high temperatures. The indexing of the X-ray powder-diffraction photographs yielded good agreement of the unitcell parameters with the published structural data (Table 1). The results of electron-microprobe analysis of the native gold inclusions present in the specimens are in fair agreement with the calculated compositions of the gold-silver alloys derived from the Mössbauer spectra.

The 18-h-half-life ¹⁹⁷Pt sources for the ¹⁹⁷Au Mössbauer measurements were prepared by irradiation of about 200 mg of enriched ¹⁹⁶Pt metal in a neutron flux of 2×10^{13} n/s·cm² for one day. The Mössbauer γ -rays were detected with a planar intrinsic Ge detector, which yielded count rates of up to 2×10^5 /s in the window of the single-channel analyzer set on the 77.3 keV line. The absorbers of synthetic specimens were fine powders distributed uniformly over an area of 2 cm². The absorber thickness was between about 60 and 200 mg/cm². The gold content of the petzite-bearing slab of rock was not determined, but was sufficient for measurement of a Mössbauer spectrum. All measurements were performed with both the source and the absorber cooled to 4.2 K in a liquid He bath cryostat. The spectra were least-squares-fitted with superpositions of Lorentzian lines. All line positions and isomer shifts are given with respect to the source, *i.e.*, to ¹⁹⁷Au in Pt metal. In order to convert them to shifts relative to metallic gold, 1.23 mm/s must be added to the given values.

RESULTS

The ¹⁹⁷Au Mössbauer spectra of some of the specimens are shown in Figure 1, and the relevant Mössbauer parameters are compiled in Table 2.



FIG. 1. ¹⁹⁷Au Mössbauer spectra of petzite (Ag₃AuTe₂), fischesserite (Ag₃AuSe₂) and uytenbogaardtite (Ag₃AuS₂).

TABLE 2. MÖSSBAUER RESULTS FOR THE INVESTIGATED GROUP OF MINERALS*

NESORBER	IS (mm/s)	QS (mm/s)	W (mm/s)	RI(%)	REMARKS
Petzite Ag ₃ AuTe ₂					
l'immins, Ont. (M13740)	+1.19(2) -0.68(16)	4.56(5)	1.85(4) 1.68(40)	92(2) 8(2)	Au* Au ₇₆ Ag ₂₄
Hemio, Ont.	+1.14(4) -0.74(2)	4.80(7)	1.79(8) 2.20(5)	31(1) 69(1)	Au+ Au78A922
Synthetic (R395)	+1.18(4) -0.75(9)	4.58(7) -	1.90(5) 2.01(20)	75(2) 25(2)	Au≁ Au79A921
Synthetic (1000°C/12 h)	+1.12(1) -0.59(1)	4.80(4) -	1.68(3) 2.30(2)	20(1) 80(1)	Au⁺ Au ₇₂ Ag ₂₈
Synthetic (500°C/100 h)	+1.12(3) -0.81(4)	4.52(6) -	2.14(4) 2.33(8)	56(1) 44(1)	Au+ Au ₈₀ Ag ₂₀
Fischesserite Ag ₃ AuSe ₂					
Synthetic (1000ºC/12 h)	+1.20(1)	4.96(3)	1.94(2)	100	Au+
Uytenbogaardtite Ag ₃ Au	 S ₂				
Synthetic	+1.57(2)	5.38(4)	1.94(4)	79(1)	Au+
(1200°C/12 h)	-0.63(6)		2.05(17)	21(1)	Au74Ag2
Synthetic	+1.60(2)	5.42(5)	1.96(9)	64(1)	Au+
(1200°C/12 h)	-0.54(3)	•	2.58(8)	36(1)	Au ₇₀ Ag ₃
Synthetic (500°C/100 h)	+1.59(1)	5.36(3)	1.96(2)	100	Au+

*16 is the isomer shift relative to the source of Au in Pt metal, QS is the electric quadrupole splitting, W the FWHM line width, and Ri the relative intensity of the respective components. Figures in brackets indicate uncertainty in last digits proported.

Petzite

The spectra of the five different specimens of petzite exhibit two well-separated absorption lines. The different intensities of the two resonant lines can be attributed to the presence of different amounts of native gold (electrum) in the different specimens. The right-hand peak is always at virtually the same position, + 3.48 \pm 0.03 mm/s, and has nearly the natural linewidth. The left-hand peak is always more intense and slightly broadened, which indicates that it is composed of the left peak of a quadrupole doublet attributable to petzite and a peak due to impurities of a gold-silver alloy. The left-hand peak of petzite should be at the same place in all cases, whereas the position of the alloy peak may vary with the Au/Ag ratio, since Au-Ag alloys have a more positive isomer shift (Huray et al. 1976) than pure gold with its resonance line at -1.23 mm/s. These considerations leave only the disposition of lines as shown in Figure 1 for the consistent fitting of all of the petzite spectra. The Timmins specimen (M13740) yielded the spectrum with the weakest line for the gold-silver alloy. We therefore consider the Mössbauer parameters of petzite obtained for this sample (IS = +1.19mm/s, QS = 4.56 mm/s) as the most reliable that can presently be given.

The relative intensities (RI) of the petzite and

gold-silver alloy components given in Table 2 refer to the areas of the respective components in the Mössbauer spectra. The relative amounts of petzite and gold-silver alloy may be different from these values because the Lamb-Mössbauer f-factors of the two components may differ somewhat. In the course of this work, no efforts were made to measure the ratio of these f-factors.

The composition of the gold-silver alloys observed in the spectra can be estimated using an empirical linear relationship between the content of silver, x, in $Au_{100-x}Ag_x$ alloys and the isomer shift IS of the Mössbauer line (Huray *et al.* 1976), which can be written as

 $x[at.\%] = 44 \bullet (IS[mm/s] + 1.23).$

The compositions given in Table 2 have been derived from this relation and the observed isomer shifts. Note that the composition of the metallic phase in both the petzite and the uytenbogaardtite samples is always much richer in gold than the nominal Au/Ag ratio in these specimens (AuAg₃). This seems to be a systematic deviation, the alloy phase being always about Au₃Ag whatever the total amount of gold present in the metallic form. The excess silver is presumably bound in a silver telluride such as hessite (Ag₂Te), which is invisible to Mössbauer spectroscopy.

Fischesserite

The ¹⁹⁷Au spectrum of synthetic Ag_3AuSe_2 exhibits a symmetrical quadrupole doublet and shows no trace of a gold-silver alloy. The Mössbauer parameters of this doublet are very close to those of petzite (Table 2).

Uytenbogaardtite

The Mössbauer spectrum of the synthetic Ag_3AuS_2 sample prepared at 500°C is symmetrical and shows no traces of impurities, whereas the specimens prepared at 1200°C contain some gold-silver alloy. Fits in analogy to those for the petzite spectra yield the results given in Table 2.

DISCUSSION

The common feature of the Mössbauer spectra of the three compounds investigated is a rather large quadrupole splitting (QS \approx 4.5-5.4 mm/s). Large magnitudes of the quadrupole splitting, which is a measure of the gradient in the electric field at the Au nuclei, are in accordance with the linear two-fold coordination of the gold atoms expected from the similarity of the cell edges in all three compounds, although a determination of the atomic position so far has been made only for petzite. The similarity of the Mössbauer parameters



FIG. 2. Plot of the magnitude of the electric quadrupole splitting *versus* the isomer shift (relative to gold in a platinum matrix) for Au^+ and Au^{3+} compounds, as well as the gold minerals studied in the present work.

of all three compounds in the Ag_3AuX_2 series confirms that the coordination of gold is very similar in all three cases. On the basis of the known, approximately linear relation between IS and QS in covalent auric and aurous compounds (Bartunik et al. 1970, Parish 1982, 1984), of which a part is shown in Figure 2, one finds the investigated compounds to be between the Au^+ and the Au^{3+} regions. In both valence states, the IS and QS increase as the gold-ligand bonds become more covalent, and greater electron density is placed on the gold atom. An unambiguous distinction between Au^+ and Au^{3+} could be made if the sign of the gradient in the electric field could be determined, since the gradients are expected to be negative in the Au^+ and positive in the Au^{3+}

compounds (Parish 1984). In the present case, however, only the magnitudes of the quadrupole interaction can be determined from the Mössbauer spectra of powder absorbers. Even oriented singlecrystal specimens, which normally allow the sign of the quadrupole interaction to be determined (Prosser et al. 1975), would not be helpful in the present case, where the cubic (Ag₃AuTe₂ and $Ag_{1}AuSe_{2}$, or at least nearly cubic ($Ag_{3}AuS_{2}$) symmetry would preclude a net alignment of the tensors of the gradient in the electric field in the absorber, and hence the determination of the sign of the electric quadrupole interaction. However, the linear coordination of the gold between its two nearest tellurium neighbors in Ag₃AuTe₂ (Frueh 1959), together with its position in the plot of the

quadrupole splittings versus the isomer shifts (Fig. 2), strongly support the notion that gold in petzite is monovalent. From the similarity of the Mössbauer parameters of petzite, fischesserite and uytenbogaardtite, one then concludes that the same is true for the oxidation state of the gold in the latter two minerals. This view is supported by the similarity of Mössbauer parameters of Ag₃AuS₂ with those of the Au⁺ compound Au₂S (Faltens & Shirley 1970), which has the Cu₂O type structure $(O_{b}4, Pn3m)$ (Hirsch *et al.* 1966), in which the gold is linearly coordinated to two neighboring sulfur atoms at a distance of 2.17 Å. Figure 2 shows that Au₂S also falls between the Au⁺ and Au³⁺ regions on the QS versus IS plot and is remarkably close to Ag₃AuS₂, which indicates very similar bonding situations and gold-sulfur distances.

The Au-Te distances are 2.61 Å in petzite (Chamid et al. 1978). Because of the smaller size of the unit cell, the distances between Au and Se in fischesserite are expected to be smaller than the Au-Te distances in petzite. This may be the reason for the larger quadrupole splitting. The quadrupole splitting along the series $Ag_3AuTe_2 - Ag_3AuSe_2 -$ Ag₃AuS₂ increases with decreasing unit-cell volume, and hence with decreasing gold-chalcogen distance. It is tempting to attribute this to increasing covalency, with the Au-S bond being the most covalent. It is, however, difficult to decide to what extent the increasing covalency is attributable to the nature of the ligands and to what extent it is attributable to the decreasing bond-length of the gold-ligand bond.

The difference in isomer shifts between pure Au metal (IS = -1.23 mm/s) and the Ag₃AuX₂ (IS \approx +1.2-1.6 mm/s) shows that the density of electrons at the Au nucleus in the Ag_3AuX_2 compounds is larger than that of Au metal. The observed magnitude of the isomer shift in petzite (IS = +1.2mm/s) is only slightly smaller than in gold-silver ditellurides (IS $\approx + 1.4 - 1.7 \text{ mm/s}$) (Wagner *et al.* 1986, 1988). The explanation of the smaller values of the isomer shift, and thus smaller charge-transfer from tellurium to gold atoms, may rest in the fact that in the ditellurides, the gold atoms are bonded to six fairly close tellurium atoms, whereas in petzite the gold has only two close neighbors. In addition, whereas most of the other phases in the system Ag-Au-Te have metallic bonding, conductivity data suggest that covalent bonding is of considerable significance in petzite (Rucklidge & Stumpfl 1968).

CONCLUSIONS

The electric quadrupole interactions observed in Ag_3AuX_2 compounds are the largest so far found in any gold mineral. In particular, its large

quadrupole splitting allows petzite to be easily distinguished from the ditelluride minerals calaverite, sylvanite and krennerite. Notably, the Mössbauer analyses can be made on rather large pieces of rock without separation of the gold mineral, as long as the overall content of gold in the sample exceeds about 100 ppm.

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