

PHOSINAITE-(Ce) FROM MONT SAINT-HILAIRE, QUEBEC: NEW DATA AND STRUCTURE REFINEMENT

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ABSTRACT

Phosinaite-(Ce), ideally $\text{Na}_{13}\text{Ca}_2\text{REE}[\text{Si}_4\text{O}_{12}](\text{PO}_4)_4$, has been found in sodalite syenite xenoliths at the Poudrette Quarry, Mont Saint-Hilaire, Quebec. The mineral occurs as colorless to pale brown anhedral, elongate grains up to 2 mm set in the groundmass, and euhedral prismatic crystals up to 2×1 mm in vugs in sodalite. Crystals commonly are coated in a brown film of an amorphous Na-Ca-K silicate, previously believed to be a hydrocarbon. Important associated phases include microcline, analcime, albite, villiaumite, astrophyllite, thermonatrite, vitusite, revdite and vuonnemite. Average results of electron-microprobe analyses ($n = 4$) gave Na_2O 31.68, CaO 8.92, MnO 0.89, La_2O_3 4.18, Ce_2O_3 6.52, Nd_2O_3 0.99, Pr_2O_3 0.26, ThO_2 1.84, Al_2O_3 0.11, SiO_2 18.36 and P_2O_5 23.69 wt.%, corresponding to $(\text{Na}_{12.72}\text{Ca}_{0.14})_{\Sigma 12.86}(\text{Ca}_{1.84}\text{Mn}_{0.16})_{\Sigma 22.00}(\text{Ce}_{0.49}\text{La}_{0.32}\text{Th}_{0.09}\text{Nd}_{0.07}\text{Pr}_{0.02})_{\Sigma 0.99}[(\text{Si}_{3.77}\text{P}_{0.15}\text{Al}_{0.04})_{\Sigma 3.96}\text{O}_{12}] (\text{PO}_4)_4$, based on 28 oxygen atoms. Phosinaite-(Ce) from Mont Saint-Hilaire is found to have a distinctive concentration of Th compared to that from the Khibina and Lovozero massifs, Russia. Single-crystal study showed it to be orthorhombic, $P22_1$, with a 12.297(2), b 14.660(3), c 7.245(1) Å, V 1306.1(1) Å³ and $Z = 2$. The observed powder-pattern is consistent with that previously published; the reflection reported at 6.92 Å was not observed in our material. The crystal structure was refined using single-crystal X-ray diffraction data to $R = 3.1\%$ and $wR^2 = 7.99\%$. In addition to confirmation of the basic features of the structure originally proposed for the mineral, site-occupancy refinement for one octahedrally coordinated site gave $(\text{Na}_{0.66}\text{Ca}_{0.34})$. A high degree of order of Si and P was established on the basis of results from electron-microprobe analyses, site-occupancy refinement, mean T-O bond distances and calculated bond-valence sums, suggesting that the mineral be classified as a silicophosphate.

Keywords: phosinaite-(Ce), new occurrence, silicophosphate, Mont Saint-Hilaire, Quebec, X-ray data, chemical composition, crystal structure.

SOMMAIRE

Nous avons trouvé la phosinaite-(Ce), de composition idéale $\text{Na}_{13}\text{Ca}_2\text{TR}[\text{Si}_4\text{O}_{12}](\text{PO}_4)_4$ (TR : terres rares) dans des xénolithes de syénite à sodalite dans la carrière Poudrette, mont Saint-Hilaire, Québec. Il s'agit d'un minéral incolore à brun pâle, se présentant en grains xénomorphes, allongés, jusqu'à 2 mm en longueur dans la pâte, et en grains idiomorphes atteignant 2×1 mm dans des cavités dans la sodalite. Les cristaux sont en général recouverts d'une pellicule brune d'un silicate amorphe de Na-Ca-K, que l'on avait auparavant considéré comme un hydrocarbure. Sont associés à la phosinaite-(Ce) microcline, analcime, albite, villiaumite, astrophyllite, thermonatrite, vitusite, revdite et vuonnemite. Les analyses ($n = 4$) à la microsonde électronique ont donné Na_2O 31.68, CaO 8.92, MnO 0.89, La_2O_3 4.18, Ce_2O_3 6.52, Nd_2O_3 0.99, Pr_2O_3 0.26, ThO_2 1.84, Al_2O_3 0.11, SiO_2 18.36 et P_2O_5 23.69% (en poids), ce qui, sur une base de 28 atomes d'oxygène, correspond à $(\text{Na}_{12.72}\text{Ca}_{0.14})_{\Sigma 12.86}(\text{Ca}_{1.84}\text{Mn}_{0.16})_{\Sigma 22.00}(\text{Ce}_{0.49}\text{La}_{0.32}\text{Th}_{0.09}\text{Nd}_{0.07}\text{Pr}_{0.02})_{\Sigma 0.99}[(\text{Si}_{3.77}\text{P}_{0.15}\text{Al}_{0.04})_{\Sigma 3.96}\text{O}_{12}] (\text{PO}_4)_4$. La phosinaite-(Ce) du mont Saint-Hilaire possède une concentration distinctive en Th que les échantillons des massifs de Khibina et de Lovozero (Russie) ne montrent pas. Une étude sur cristal unique montre qu'elle est orthorhombique, $P22_1$, avec a 12.297(2), b 14.660(3), c 7.245(1) Å, V 1306.1(1) Å³ et $Z = 2$. Le spectre de diffraction X obtenu sur poudre concorde avec celui qui est déjà dans la littérature, exception faite de la réflexion à $d = 6.92\text{\AA}$, qui n'a pas été observée dans nos échantillons. Nous avons affiné la structure sur cristal unique jusqu'à un résidu R de 3.1% ($wR = 7.99\%$). En plus de confirmer les grandes lignes de la structure déjà connues, nos résultats ont permis d'affiner l'occupation d'un des sites octaédriques: $(\text{Na}_{0.66}\text{Ca}_{0.34})$. Une distribution essentiellement ordonnée des atomes Si et P, établie par nos résultats d'analyses à la microsonde électronique, nos affinements de l'occupation des sites, les longueurs moyennes des liaisons T-O, et les valences de liaison, font penser qu'il s'agit d'un silicophosphate.

(Traduit par la Rédaction)

Mots-clés: phosinaite-(Ce), nouvel exemple, silicophosphate, mont Saint-Hilaire, Québec, données de diffraction X, composition chimique, structure cristalline.

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INTRODUCTION

During the course of a mineralogical investigation of sodalite syenite xenoliths found in the Poudrette Quarry, Mont Saint-Hilaire, Québec, an unidentified mineral was encountered and temporarily designated UK64 (Chao *et al.* 1990). It was studied by optical, chemical and crystal-structure analyses, and subsequently found to be the rare species phosinaite-(Ce), which had previously been recorded only from the Khibina and Lovozero massifs of Russia (Kapustin *et al.* 1974). The occurrence of new material has therefore provided an opportunity to compare the physicochemical characteristics of phosinaite-(Ce) from all three localities and to investigate the extent of Si and P order in the structure.

OCCURRENCE AND PROPERTIES

The Canadian phosinate-(Ce) is a late-stage phase that occurs as anhedral, elongate grains up to 2 mm set in the groundmass, as well as euhedral, prismatic crystals up to 2×1 mm in small vugs and fractures in colorless sodalite. Although more than 72 minerals have been found in these xenoliths (Chao *et al.* 1991), the most important phases associated with phosinaite-(Ce) include microcline, analcime, albite, villiaumite,

astrophyllite, thermonatrite, vitusite, revdite and vuonnemite.

The euhedral crystals of phosinaite-(Ce) are elongate on [001] and are commonly bounded by the forms {100}, {010} and {001}. The mineral is colorless to pale brown and commonly is coated with a brown film previously believed to be a hydrocarbon. However, a powder X-ray-diffraction pattern and results of an energy-dispersion analysis indicate that this film is actually an amorphous Ca-Na-K silicate. Phosinaite-(Ce) is quite brittle, with perfect {001}, distinct {110} and good {010} cleavages. The measured density was found to be 2.97 g/cm^3 , as determined by flotation in bromoform diluted with acetone.

Optically, the Mont Saint-Hilaire material is non-pleochroic with a very low birefringence. The observed interference figures were found to be quite poor, and no measurements of $2V$ or dispersion could be made. The indices of refraction are: $\alpha = \beta = 1.568(1)$ and $\gamma = 1.572(1)$, as determined in Na light ($\lambda = 589.3 \text{ nm}$), and the optical orientation is $X = b$, $Y = c$, $Z = a$, with a variable sign of elongation, all of which are consistent with those observed in the Russian material. A summary of the optical and physical properties of phosinaite-(Ce) from the three localities is provided in Table 1.

CHEMICAL COMPOSITION

TABLE 1. PHYSICAL PROPERTIES OF PHOSINAITE-(Ce)

Habit	1 Anhedral aggregates	2 Columnar crystals up to 5×1 mm	3 Prismatic crystals up to 2×1 mm, elongate, anhedral grains up to 2 mm
Colour	Colorless to pale rose	colorless, pale rose to brownish rose	colorless to pale brown
Hardness (Mohs')	3.5	3.5	4
Forms Observed	None	{110} predominant, also {100}, {010}, {001}	{100}, {010}, {001}
Cleavage	{100} perfect, {010}, {110} imperfect	As in 1.	{001} perfect, {110} distinct {010} good
D_{min} (g/cm ³)	2.62	3.00	2.97(1)
D_{calc} (g/cm ³)	3.17*	3.21**	3.09
Optics			
α	1.567	1.570	1.568(1)
β	1.569	1.572	1.568(1)
γ	1.570	1.573	1.572(1)
$2V_{\text{max}}^{(\pm)}$	(+68°)	(+70°)	(+)*
$2V_{\text{obs}}^{(\pm)}$	70°	70°	0
Elongation	(\pm)	(\pm)	(\pm)
Orientation	$X = b$ $Y = c$ $Z = a$	$X = b$ $Y = c$ $Z = a$	$X = b$ $Y = c$ $Z = a$

* Calculated in this study for the formula:

(Na_{2.7}K_{0.6})_{23.0}(Ca_{0.71}Cr_{0.28})_{20.97}Si_{1.10}P_{0.59}O_{7.10}

** Calculated in this study for the formula:

(Na_{2.04}K_{0.60})_{23.07}(Ca_{0.31}Cr_{0.28}Mn_{0.07})_{20.68}Si_{1.30}P_{0.95}O_{7.33}

† - $2V$ could not be determined.

‡ - Calculated in this study.

1. Khibina massif, Kapustin *et al.* (1974).

2. Lovozero massif, " " "

3. Mont Saint-Hilaire, this study.

Electron-microprobe analyses were carried out on the Mont Saint-Hilaire phosinaite-(Ce) using a Cambridge Microscan MK5 instrument. The operating conditions were 15 kV, 30 nA (estimated), with the beam rastered over an area of $10 \times 10 \mu\text{m}^2$. Wavelength-dispersion data were collected using the following standards: albite (Na), Kakanui hornblende (Al, Si, Ca), Durango fluorapatite (P), tephroite (Mn), REE-bearing glasses (La, Ce, Pr, Nd, Sm) and synthetic ThO₂ (Th); the data were reduced using a modified version of EMPADR VII (Rucklidge & Gasparini 1969). In addition to the elements listed above, S, Zr, U and Fe were sought but not detected.

Results of these analyses indicate that the material is chemically homogeneous. The average of these analyses (Table 2) gives the empirical formula [based on 28 atoms of oxygen as determined from crystal-structure analysis]: (Na_{12.72}Ca_{0.14})_{21.86}(Ca_{1.84}Mn_{0.16})_{22.00}(Ce_{0.49}La_{0.32}Th_{0.09}Nd_{0.07}Pr_{0.02})_{20.99}[Si_{3.77}P_{0.15}Al_{0.04}]_{23.96}O₁₂]PO₄ with $Z = 2$ or, ideally, Na₁₃Ca₂REE[Si₄O₁₂]PO₄.

There is good agreement between the major-element content of phosinaite-(Ce) from all three localities and that calculated for the ideal composition (Table 2). In terms of minor elements, the Mont Saint-Hilaire material exhibits a distinctly high Th concentration. The Russian phosinaite-(Ce) contains water in the range of 2 to 6 wt.% (Table 2). However, it is now considered to

TABLE 2. CHEMICAL COMPOSITION OF PHOSINAITE-(Ce)

	1	2	3	4	5
Na ₂ O wt.%	31.68	28.05	28.78	29.76	33.48
K ₂ O	-	0.72	0.47	0.50	-
CaO	8.92	12.10	5.39	8.90	9.32
MnO	0.89	-	1.90	0.16	-
Fe ₂ O ₃	-	-	0.49	0.45	-
REE ₂ O ₃	11.95	13.22	13.80	11.67	13.64
ThO ₂	1.84	-	-	-	-
SiO ₂	18.36	18.56	23.78	20.00	19.97
P ₂ O ₅	23.69	21.45	20.50	25.50	23.59
H ₂ O	-	5.87	4.41	2.38	-
TOTAL	97.33	99.97	99.52	99.32	100

- This study; Mont St. Hilaire, Rouville Co., Québec; SiO₂ includes Al₂O₃ 0.15 wt.%; REE₂O₃ includes La₂O₃ 4.18, Ce₂O₃ 6.52, Nd₂O₃ 0.99 and Pr₂O₃ 0.26 wt. %.
- Kapustin *et al.* (1974); average of two analyses; Khibina, Russia; Spectrographic analysis gave La 16.8, Ce 53.9, Pr 4.4, Nd 19.5, Sm 3.0, Eu 0.5, Gd 0.7, Tb 0.3, Er 0.3, Y 0.6.
- Kapustin *et al.* (1974); Mt. Karnasurt, Lovozer, Russia; Spectrographic analysis as above; Analysis also gave Li₂O 0.24 wt. %.
- Krutik *et al.* (1981); Khibina Massif, Russia; Analysis also gave MgO 0.28 and SrO 0.30 wt. % respectively.
- Composition calculated for: Na₁₃Ca₂Ce[Si₄O₁₂](PO₄)₄.

be nonessential on the basis of results of a thermogravimetric analysis (Krutik *et al.* 1981) and refinement of the crystal structure, reported below. The water is most likely adsorbed, possibly resulting from supergene alteration, as commonly occurs in minerals from peralkaline rocks (Krutik *et al.* 1981).

X-RAY CRYSTALLOGRAPHY AND STRUCTURE REFINEMENT

Single-crystal X-ray precession photographs of phosinaite-(Ce) from Mont Saint-Hilaire confirmed that the mineral is orthorhombic and belongs to the unique space-group *P*2₁2₁ (#18). X-ray powder-diffraction data were collected using Ni-filtered CuK α radiation ($\lambda = 1.5418 \text{ \AA}$) and a 114.6-mm Gandolfi camera. The unit-cell parameters refined from these data using the program CELREF (Appleman & Evans 1973) are: *a* 12.297(2), *b* 14.660(3), *c* 7.245(1) \AA , with a *V* of 1306.1(1) \AA^3 and *Z* = 2, in good agreement with those determined from previous studies (Table 3). The indexed powder-patterns of phosinaite-(Ce) from Mont Saint-Hilaire, the Russian localities and that calculated from the refined crystal-structure are presented in Table 4. Agreement between estimated and calculated intensities is good, with minor discrepancies owing to errors in the visual estimates and the recording geometry of the Gandolfi camera. The 6.92 \AA reflection originally reported by Kapustin *et al.* (1974) is both unindexable on the cell adopted for

TABLE 3. CELL DIMENSIONS OF PHOSINAITE-(Ce)

Space Group	<i>P</i> 2 ₁ 2 ₁ or <i>P</i> 2 ₂ 2 ₁	1	2	3
		i	ii	<i>P</i> 2 ₁ 2 ₁
<i>a</i> (\AA)	12.23(4)	12.24(4)	7.234(3)	12.297(2)
<i>b</i>	14.62(4)	14.59(4)	14.670(4)	14.660(3)
<i>c</i>	7.21(4)	7.20(4)	12.231(4)	7.245(1)
<i>V</i> (\AA^3)	1289(1)	1286(1)	1298(1)	1306.1(1)
<i>Z</i>	2	2	2	2

1. Kapustin *et al.* (1974); i) Khibina massif, Russia;

ii) Lovozero massif, Russia

2. Krutik *et al.* (1981); Khibina massif, Russia

3. This study; Mont St. Hilaire, Québec; dimensions refined from the X-ray powder-diffraction pattern.

phosinaite-(Ce) and inconsistent with the observed diffraction-symmetry. A 12-hour powder pattern of Lovozero phosinaite-(Ce) was made in the context of this study, but failed to reveal any indication of this enigmatic reflection. Therefore, it is believed to be

TABLE 4. X-RAY-DIFFRACTION DATA FOR PHOSINAITE-(Ce)

	1		2		3	
	Lovozero	Mont Saint-Hilaire	Mont Saint-Hilaire	Calculated	Calculated	hkl
	<i>I</i> _{ext} ¹	<i>d</i> _{max} ¹	<i>I</i> _{ext} ²	<i>d</i> _{max} ²	<i>I</i> _{calc} ³	<i>d</i> _{calc} ³
55	7.44	70	9.420	75	9.421	110
50	6.92	20	7.338	24	7.330	020
9	6.39	10	6.531	28	6.495	011
		40	6.265	39	6.296	120
				43	6.242	101
15	5.72	20	5.740	72	5.743	111
				14	5.670	210
12	4.71	25	4.718	16	4.752	121
				52	4.711	220
		10	4.541	26	4.688	201
15	4.11	5	4.080	16	4.541	130
		60	3.949	13	4.100	300
				66	4.051	031
20	3.94			7	3.949	221
				17	3.948	310
			15	14	3.848	131
				9	3.826	230
30	3.62			10	3.665	040
40	3.51			0	3.622	002
		20	3.468	27	3.512	140
18	3.37	25	3.384	15	3.466	311
				16	3.383	231
				11	3.381	112
				26	3.208	321
20	3.13			7	3.160	141
			30	4	3.148	240
			40	4	3.140	330
				59	3.140	122
18	3.01			3	3.053	212
7	2.89	15	2.899	10	2.910	032
				13	2.887	241
7	2.82			0	2.852	150
			10	20	2.835	420
100	2.74		2.819	9	2.830	401
				100	2.732	340
				10	2.718	051
		100	2.718	86	2.714	302
				14	2.646	250
			70	2.577	87	2.576
						042

TABLE 4. X-RAY-DIFFRACTION DATA FOR PHOSINAITE-(Ce) — Cont'd.

	1 Lovozero	2		3		<i>hkl</i>
		<i>I</i> _{est} ¹	<i>d</i> _{meas} ¹	<i>I</i> _{est} ²	<i>d</i> _{meas} ²	
9	2.504	<5	2.486	12	2.546	222
				2	2.522	142
				4	2.485	251
				10	2.449	431
				14	2.396	160
				10	2.390	350
				5	2.346	402
				9	2.339	113
				1	2.329	501
				1	2.315	061
11	2.317	10	2.315	13	2.314	412
				3	2.279	052
				5	2.274	161
				2	2.271	260
				15	2.232	20
				15	2.233	422
				15	2.197	530
				10	2.178	22
				10	2.138	13
				4	2.065	252
18	2.038	20	2.050	23	2.050	600
				10	2.037	11
				3	1.999	162
				6	1.990	143
				10	1.988	12
				10	1.965	1.966
				9	1.916	243
				1	1.914	615
				30	1.913	5
				1	1.913	460
17	1.832	35	1.835	1	1.912	271
				10	1.883	1
				5	1.883	550
				6	1.843	413
				6	1.843	153
				20	1.832	080
				0	1.812	180
				20	1.811	23
				15	1.800	5
				5	1.768	1.770
15	1.804	5	1.811	5	1.768	433
				20	1.691	12
				15	1.674	0
				10	1.663	1.663
				5	1.656	0
				5	1.635	1
				5	1.618	0
				3	1.618	1.625
				5	1.604	533
				5	1.576	324
10	1.685	20	1.604	15	1.604	190
				5	1.576	642
				10	1.561	1.560
				5	1.545	8
				15	1.520	1.519
				15	1.507	11
				Plus others	1.510	750

- Data from Kapustin *et al.* (1974) for a specimen from Mt. Karnasut, Lovozero massif, Russia; The 6.92 Å line is not indexable on the cell reported by the authors.
- Data from this study for a specimen from Mont Saint-Hilaire; 114.6 mm Gandolfi camera, CuK α radiation ($\lambda = 1.5418 \text{ \AA}$); *I*_{est} visually estimated; *d*_{calc} determined from the cell refined from the powder data: *a* 12.297, *b* 14.660, *c* 7.245 Å.
- I*_{calc} determined using the program DISPOW (Gabe *et al.* 1989).

either erroneous or perhaps caused by impurities trapped in some of the Lovozero material.

X-ray intensity data were collected on an automated four-circle diffractometer using an optically clean, prismatic crystal mounted so as to rotate about *c*. The cell dimensions given in Table 5 were derived by a

TABLE 5. MISCELLANEOUS DATA FOR THE REFINEMENT OF THE CRYSTAL STRUCTURE OF PHOSINAITE-(Ce)

Space Group	<i>P</i> 2 ₁ 2 ₁ (#18)	Diffractometer	Enraf-Nomius CAD4
<i>a</i> (Å)	12.270(1)*	Radiation	MoK α (50kV, 20 mA)
<i>b</i>	14.648(2)*	Monochromator	Graphite
<i>c</i>	7.231(1)*	Crystal Shape	Equant, Prismatic
<i>V</i> (Å ³)	1299.6(1)	Crystal Size	0.25 × 0.06 × 0.04 mm
<i>Z</i>	2	μ (MoK α)	4.76 mm ⁻¹
Chemical Formula	Na _{1.1} Ca _{0.9} REE(Si ₂ O ₅)(PO ₄) ₄		
Intensity Data Collection	6-2θ Scanning Mode		
2θ Limit	55°		
Intensity Standards	Three every 10,000 seconds		
Orientation Standards	Three every 400 reflections		
Number of Unique Reflections	2158		
Number of Observed Reflections	1868		
Criterion for Significance	<i>F</i> >6σ(<i>F</i>)		
Final <i>R</i> for all Observed Reflections	3.1%		
Final <i>wR</i> ² for all Observed Reflections	7.99%		

* - refined from four-circle diffractometer data.

least-squares method (Busing 1970) using 24 automatically aligned reflections in the range 5° < 2θ < 42° permuted four ways ($\pm h$ at ± 20). A total of 2158 reflections were collected to 2θ = 55° (a hemisphere of data), of which 1868 were classified as observed [*F* > 6σ(*F*)]. Pertinent information regarding data collection is presented in Table 5.

Data measurement and reduction (Lorentz, polarization, background, scaling) were done using the NRC-VAX package of computer programs (Gabe *et al.* 1989). No absorption correction was applied owing to the low linear absorption-factor of the mineral and the small size of the crystal used.

The crystal structure of phosinaite-(Ce) was determined independently of that by Krutik *et al.* (1981) using the SHELXTL PC package of programs (Sheldrick 1990) and subsequently refined using SHELXL-93 (Sheldrick, in prep.). Scattering curves for neutral atoms and anomalous dispersion coefficients were taken from Cromer & Mann (1968) and Cromer & Liberman (1970), respectively. A sharpened Patterson synthesis and subsequent difference-Fourier maps were used to identify all of the cation and anion positions. Refinement of this model with isotropic thermal factors gave an *R* of 5.2%. Site-occupancy refinement revealed that only those sites assigned to rare-earth elements (REE), and assumed to be Ce, and one Na atom were significantly less than ideal, suggesting deficiencies in the scattering powers at these sites. These results are consistent with those of Krutik *et al.* (1981), who assigned a (Na,Ca) occupancy to the Na site and a refined partial occupancy of Ce_{0.67} to the REE site (the remainder of which was attributed to a partial vacancy). Refinement of the site occupancy for the (Na,Ca) site in this study converged to Na 0.66(1) and Ca 0.34(1). This gives a valence of 1.34 *vu*, in good agreement with the value of 1.204 *vu* obtained from bond-valence calculations (Table 6). Refinement of the REE site, which lies on a

TABLE 6. EMPIRICAL BOND-VALENCES ($v.u.$) IN PHOSINAITE-(Ce)⁺

	Na(1)	Na(2)	Na(3)	Na(4)	Na(5)	Na(6)	Na(7)	Na(8)	(Na,Ca)	Ca	REE	Si(1)	Si(2)	P(1)	P(2)	ΣV
O(1)	0.137						0.200	0.073 ^{x2}		0.433					1.235	2.078
O(2)	0.185	0.278		0.098 ^{x2}					0.230						1.195	1.986
O(3)			0.146 ^{x2}	0.109 ^{x2}		0.153 ^{x2}	0.278	0.139 ^{x2}							1.218	2.047
O(4)							0.207		0.087	0.157, 0.337		1.129				1.917
O(5)	0.165		0.062 ^{x2}				0.251		0.177					1.238		1.893
O(6)		0.144	0.121 ^{x2}	0.124 ^{x2}	0.124 ^{x2}			0.189 ^{x2}	0.184	0.062	0.426	0.272 ^{x2}		1.120		1.905
O(7)														1.199		2.060
O(8)		0.199								0.263		0.339 ^{x2}				2.000
O(9)	0.044	0.121			0.194 ^{x2}				0.120		0.431 ^{x2}	1.067				1.977
O(10)							0.185	0.083 ^{x2}		0.217, 0.382			1.126			1.993
O(11)						0.107 ^{x2}		0.118 ^{x2}				0.917	0.907			2.049
O(12)	0.120									0.090		0.950	1.011			2.109
O(13)								0.262			0.570 ^{x2}		1.154			1.989
O(14)	0.298	0.201, 0.037			0.074 ^{x2}									1.293		1.903
ΣV	0.949	0.980	0.658	0.662	0.784	0.898	1.305	0.826	1.204	2.042	3.224	4.063	4.198	4.850	4.847	

Valence for the (Na,Ca) site was calculated using the weighted average (Na_{0.66}Ca_{0.34}); + calculated using the constants of Brese & O'Keeffe (1991).

special position and has an ideal site-occupancy factor of 0.5, converged to Ce 0.42(1). Attempts to refine the site incorporating the scattering curve of Th were not successful. The site was subsequently assumed to fully occupied by Ce, in accordance with results from the electron-microprobe analyses. This gives a calculated valence of 3 $v.u.$, close to the value of 3.224 $v.u.$ determined from bond-valence summations and much more reasonable than the calculated value of 2.0 $v.u.$

given by Krutik *et al.* (1981). Calculated bond-valence sums for all the oxygen atoms present gave values all close to 2 (Table 6), indicating the absence of hydroxyl and water.

A further refinement of this model using anisotropic displacement factors and a weighting scheme converged to $R = 3.1\%$ and $wR^2 = 7.99\%$, a significant improvement over the value of $R = 7.2\%$ reported by Krutik *et al.* (1981). No improvement in the final

TABLE 7. POSITIONAL AND THERMAL PARAMETERS FOR PHOSINAITE-(Ce)

	x	y	z	U_{11}	U_{22}	U_{33}	U_{23}	U_{13}	U_{12}	U_{eq}
Si(1)	0.3220(1)	0.8891(1)	0.3098(2)	204(8)	115(7)	191(8)	-8(6)	-36(7)	14(6)	170(3)
Si(2)	0.3646(1)	0.9111(1)	0.7449(2)	136(6)	118(6)	152(7)	1(6)	5(7)	9(6)	135(3)
P(1)	0.9736(2)	0.1255(1)	0.2295(3)	138(7)	198(7)	229(9)	-37(7)	-1(6)	-21(5)	189(4)
P(2)	0.3064(1)	0.6183(1)	0.7658(2)	124(6)	129(6)	139(7)	1(6)	-4(6)	-14(5)	131(3)
Na(1)	0.1829(3)	0.7588(2)	0.5440(5)	376(16)	370(16)	410(22)	80(14)	18(14)	129(13)	386(8)
Na(2)	0.0030(3)	0.3491(3)	0.2237(5)	306(18)	637(23)	371(20)	1(19)	11(16)	28(16)	439(9)
Na(3)	0.1714(3)	0.5	0.5	198(17)	566(24)	339(20)	0	0	0	358(10)
Na(4)	0.1604(4)	0.5	0	278(21)	390(20)	393(20)	0	0	0	353(9)
Na(5)	0.8772(3)	0.5	0	331(21)	492(23)	280(19)	0	0	0	368(9)
Na(6)	0.4504(3)	0.5	0	258(7)	295(17)	289(19)	0	0	0	281(8)
Na(7)	0.3307(2)	0.6398(2)	0.2635(4)	181(13)	143(10)	228(13)	-1(10)	21(10)	-23(9)	184(6)
Na(8)	0.4561(3)	0.5	0.5	306(19)	394(20)	247(19)	14(22)	0	0	316(9)
(Na,Ca)	0.1902(2)	0.7662(1)	0.0240(3)	481(13)	203(9)	222(13)	-15(8)	7(10)	-20(8)	302(8)
Ca	0.4971(1)	0.7500(1)	0.5210(3)	181(5)	166(4)	224(8)	-32(5)	-3(5)	-45(3)	191(3)
REE	0.13955(5)	0	0	120(2)	130(2)	118(2)	0	0	0	142(1)
O(1)	0.6492(4)	0.8390(3)	0.0896(6)	275(23)	305(22)	186(21)	-37(18)	-11(21)	-49(25)	255(10)
O(2)	0.8124(4)	0.1480(3)	0.7076(7)	176(20)	258(21)	392(29)	14(22)	55(21)	-42(19)	275(10)
O(3)	0.6873(4)	0.0144(2)	0.7513(6)	318(20)	100(20)	284(22)	21(17)	-45(20)	-7(15)	239(9)
O(4)	0.3741(4)	0.7976(3)	0.2360(8)	304(22)	132(17)	445(28)	-40(20)	104(24)	20(17)	294(10)
O(5)	0.1442(4)	0.3424(4)	0.7200(8)	168(20)	481(29)	433(29)	-16(26)	-6(26)	78(23)	361(12)
O(6)	0.9812(4)	0.9808(3)	0.7585(9)	418(27)	322(28)	534(36)	86(26)	-122(29)	-8(22)	424(14)
O(7)	0.3720(4)	0.3522(3)	0.0636(6)	361(25)	268(21)	167(18)	-29(17)	39(20)	132(21)	266(10)
O(8)	0.0335(4)	0.1476(4)	0.0487(7)	362(26)	335(25)	343(31)	90(22)	74(22)	86(23)	346(12)
O(9)	0.1963(3)	0.9039(3)	0.2614(7)	154(18)	277(21)	322(25)	23(21)	7(20)	20(17)	251(10)
O(10)	0.4728(3)	0.8576(3)	0.7912(8)	149(19)	207(19)	506(34)	55(23)	-7(22)	39(17)	287(11)
O(11)	0.3962(3)	0.9789(2)	0.2514(7)	216(19)	132(19)	392(27)	-1(18)	62(20)	-6(14)	247(10)
O(12)	0.6704(7)	0.6152(4)	0.0359(7)	1335(68)	338(28)	214(28)	-10(22)	271(34)	-91(37)	629(22)
O(13)	0.2610(4)	0.1058(3)	0.1330(9)	366(28)	255(24)	582(36)	94(26)	-258(28)	-36(22)	401(14)
O(14)	0.9712(5)	0.3363(5)	0.8987(9)	424(35)	895(53)	476(39)	348(39)	53(30)	-192(37)	598(20)

Note: Anisotropic temperature factors have the form: $\exp. -2\pi^2(h^2a^2U_{11} + k^2b^2U_{22} \dots + 2hka^2b^2U_{12})$, all U values are in $\text{Å}^2 \times 10^4$; estimated standard deviations in parentheses.

TABLE 8. SELECTED INTERATOMIC DISTANCES (\AA) AND BOND ANGLES ($^\circ$) FOR PHOSINAITE-(Ce)

$\text{Si}(\text{O}_2)_4$ tetrahedron	Mean	109.3	Mean	104.0	
$\text{Si}(2) - \text{O}(13)$	1.571(5)	$\text{O}(13) - \text{Si}(2)$	-1.10(1)	119.0(3)	
- $\text{O}(10)$	1.580(5)	- $\text{O}(12)$	105.9(4)	- $\text{O}(10)$	2.716(7)
- $\text{O}(12)^*$	1.620(5)	- $\text{O}(11)$	109.5(3)	- $\text{O}(11)$	2.547(8)
- $\text{O}(11)^*$	1.650(4)	- $\text{O}(12)$	107.7(4)	$\text{O}(10)$	2.639(7)
Mean	1.608	- $\text{O}(11)$	106.4(2)	- $\text{O}(11)$	2.594(6)
		$\text{O}(12) - \text{Si}(2)$	- $\text{O}(11)$	107.9(3)	
		$\text{O}(12)$	- $\text{O}(12)$	- $\text{O}(11)$	2.653(7)
		Mean	109.4	Mean	2.622

	Mean	SD	Mean	SD		
P(1)O ₄ tetrahedron						
P(1) - O(14)	1.509(6)	O(14) - P(1)	- O(5)	110.4(4)	O(14) - O(5)	2.492(9)
- O(5)	1.525(5)	- O(8)	113.5(3)	- O(8)	2.547(8)	
- O(8)	1.537(5)	- O(6)	107.4(4)	- O(6)	2.475(8)	
- O(6)	1.562(5)	O(5) - P(1)	- O(8)	110.6(3)	O(5) - O(8)	2.517(7)
Mean	1.533	O(6)	111.5(3)	O(6)	2.552(7)	
		O(8) - P(1)	- O(6)	103.2(3)	O(8) - O(6)	2.430(7)
		Mean	109.4	Mean	2.502	

P(2) ₀ tetrahedron	P(2)	O(1) - P(2)	O(3)	108.9(3)	O(1)	O(3)	2.484(6)
- O(1)	1.523(4)	O(1) - P(2)	- O(3)	108.9(3)	O(1)	O(3)	2.484(6)
- O(3)	1.531(4)		- O(7)	111.7(3)	- O(7)	- O(7)	2.533(6)
- O(7)	1.537(5)		- O(2)	109.2(3)	- O(2)	- O(2)	2.495(6)
- O(2)	1.538(5)	O(3) - P(2)	- O(7)	108.6(3)	O(3)	- O(7)	2.491(6)
Mean	1.532		- O(2)	109.8(2)	- O(2)	- O(2)	2.511(6)
		O(7) - P(2)	- O(2)	108.6(3)	O(7)	- O(2)	2.496(7)
				Mean	109.7	Mean	2.502

$\text{Na}^+(\text{I})\text{O}_2$	octahedron								
Na(1)	- O(14)	2.248(7)	O(14) - Na(1)	- O(2)	102.8(3)	O(1)	- O(2)	2.495(6)	
	- O(2)	2.425(6)		- O(5)	106.5(2)		- O(5)	3.389(7)	
	- O(5)	2.467(7)		- O(12)	103.4(3)		- O(9)	4.684(6)	
	- O(1)	2.535(6)		- O(9)	79.1(2)		- O(12)	3.314(7)	
	- O(12)	2.583(7)	O(2) - Na(1)	- O(5)	100.2(2)	O(2)	- O(5)	3.753(8)	
	- O(9)	2.525(6)		- O(1)	60.4(2)		- O(12)	4.308(8)	
Mean	2.536			- O(12)	118.7(2)		- O(9)	3.653(9)	
			O(5) - Na(1)	- O(1)	85.3(2)	O(5)	- O(9)	3.669(7)	
				- O(9)	84.6(2)		- O(14)	3.780(8)	
			O(1) - Na(1)	- O(12)	80.7(2)	O(9)	- O(12)	2.593(8)	
				- O(9)	116.8(2)		- O(14)	3.359(8)	
			O(12) - Na(1)	- O(5)	95.3(2)	O(12)	- O(14)	3.727(1)	

	$O(2,2) - Na(2)$	$O(2,2) - O(2)$	$3.53(2)$	$O(12) - O(14)$	$3.19(11)$
$Na(2)O_2$ octahedron					
$Na(2)$ - $O(2)$	2.273(6)	$O(2) - Na(2)$	$O(8) - 103.7(2)$	$O(2) - O(8)$	$3.674(8)$
- $O(4)$ (1)	2.394(8)	- $O(4)$ (6)	96.5(2)	- $O(4)$ (8)	3.484(8)
- $O(8)$ (8)	2.397(7)	- $O(6)$ (6)	84.8(2)	- $O(6)$ (6)	3.234(7)
- $O(6)$ (6)	2.516(6)	- $O(4)$ (4)*	86.2(2)	- $O(4)$ (4)*	3.653(9)
- $O(9)$ (5)	2.581(5)	$O(8) - Na(2)$	$O(5) - 84.0(2)$	$O(8) - O(5)$	3.209(8)
$O(4)$ (4)*	2.015(9)	- $O(9)$ (9)	77.3(2)	- $O(9)$ (9)	3.103(7)
Mean	2.529	- $O(14)$ (4)	68.9(2)	- $O(14)$ (4)*	3.059(8)

O(14) - Na ₂)	-O(6)	160.9(3)	O(14) - O(6)	3.786(10)	
	-O(9)	84.9(2)		-O(9)	3.359(8)
	-O(14)*	111.1(2)		-O(14)*	4.475(9)
O(6) - Na ₂)	O(9)	76.2(2)	O(6) - O(9)	3.143(7)	
O(9) - Na ₂)	-O(14)*	113.3(2)	O(9) - O(14)*	4.526(9)	

O(6) - Na(3)	- O(6)*	86.7(3)	O(6) -	O(6)*	3.544(13)
O(6)* - Na(3)	- O(5)	112.8(2)	x2	- O(5)	4.503(8)
O(6)* - Na(3)	- O(5)	56.1(2)	x2	O(6)* - O(5)	2.552(7)
Na(4)O ₆ octahedron					
Na(4) - O(6)	2.573(6)	x2	O(6) - Na(4)	- O(6)*	94.8(3)
- O(3)	2.620(5)	x2	- O(3)	89.2(2)	x2
- O(2)	2.661(5)	x2	- O(2)	114.0(2)	x2
Mean	2.618		O(6)* - Na(4)	- O(2)	76.3(2)
			- O(3)	- O(6)*	- O(2)
			- O(3)	- O(3)*	3.234(7)
			- O(3)	88.6(2)	O(3)* - O(3)
			- O(3)	- O(3)	3.666(9)

$\text{Na}(\text{S}_5)\text{O}_5$ octahedron		- O(2)*	111.5(2)*	O(2)	0.4369(7)
$\text{Na}(\text{S}_5)$	- O(9)*	- Na(4)	O(2)	56.1(2)*	O(3)*
- O(6)	2.573(7)	- O(9)	135.9(3)	O(9)	O(9)*
- O(14)	2.763(8)	- O(6)	78.2(2)	O(6)	4.460(10)
Mean	2.581	- O(14)*	80.7(2)*	O(14)	3.143(7)
		- O(6)	53.4(1)	O(9)*	O(9)*
		- O(6)	94.8(3)	O(6)	4.608(7)
		- O(14)	90.3(2)*	O(14)	3.789(13)
		- O(14)*	56.1(2)	O(14)	3.788(10)

	$O(14) - Na(5)$	$O(14)* - O(14)$	$S_{11}(12) x2$	$O(14) - O(14)$	$2.47(5)$	$5.02(16)$
$Na(6)_6$ octahedron						
$Na(6)$	$O(7)^*$	$2.416(5) x2$	$O(7)^* - Na(6)$	$O(3)$	$86.8(2) x2$	$O(7)^* - O(3)$
	$-O(3)$	$2.495(5) x2$	$O(11)$	$92.9(1) x2$	$O(11) - O(11)$	$3.657(7)$
	$-O(11)$	$2.627(5) x2$	$O(10)^*$	$121.6(1) x2$	$O(10)^* - O(11)$	$4.403(7)$
Mean	2.513					
			$O(7)^* - Na(6)$	$O(3)$	$60.9(1) x2$	$O(7)^* - O(3)$
			$O(3)$	$Na(6)$	$O(3)^* - O(3)$	$2.491(6)$
			$O(3)^*$	$O(3)^*$	$94.5(2) x2$	$O(3)^* - O(3)$
			$O(11)^*$	$O(11)$	$88.7(1) x2$	$O(11)^* - O(11)$
			$O(11)$	$Na(6)^*$	$88.2(2)$	$O(11) - O(11)$
						$3.582(6)$
						$3.656(6)$
						$3.656(1)$

TABLE 8.—*Cont'd.*

(Na,Ca)O ₄ polyhedron	O(13)	(-Na,Ca)	O(8)	76.6(2)	O(13) - O(8)	2.929(7)
(Na,Ca)-O(13)	2.360(5)	O(13) - (Na,Ca)	- O(8)	76.6(2)	O(13) - O(8)	2.929(7)
- O(8)	2.364(6)	- O(2)	- O(2)	104.0(2)	- O(2)	3.761(7)
- O(2)	2.412(5)	- O(9)	- O(9)	72.4(2)	- O(9)	2.970(8)
- O(5)	2.509(6)	- O(4)	- O(4)	80.4(2)	- O(4)	3.329(8)
- O(9)	2.653(5)	- O(7)	- O(7)	95.0(2)	- O(7)	3.896(7)
- O(4)	2.773(6)	O(8) - (Na,Ca)	- O(2)	102.8(2)	O(8) - O(2)	3.727(7)
- O(7)	2.901(6)	- O(5)	- O(5)	108.6(2)	- O(5)	3.958(8)
Mean	2.567	- O(9)	- O(9)	76.2(2)	- O(9)	3.104(7)
		O(2) - (Na,Ca)	O(5) - O(3)	93.2(2)	O(2) - O(3)	3.575(8)
		- O(4)	- O(4)	121.1(2)	- O(4)	4.517(8)

Ca_2O_7 , polyhedron				
Ca - O(1)	2.277(5)	O(1) - Ca	- O(10)*	95.5(2)
O(5) - $\langle \text{Na}, \text{Ca} \rangle$	O(9) 90.5(2)	O(5) - O(9)	2.497(7)	
O(4) -	83.1(2)	O(5) - O(4)	3.670(7)	
O(7) -	87.5(2)	O(7) - O(4)	3.507(7)	
O(9) - $\langle \text{Na}, \text{Ca} \rangle$	O(4) 59.4(2)	O(9) - O(4)	2.375(8)	
O(7) -	125.2(2)	O(7) - O(4)	4.933(7)	
O(4) - $\langle \text{Na}, \text{Ca} \rangle$	O(7) 65.9(2)	O(4) - O(7)	3.089(7)	

- O(7) -	2.283(7)	- O(4)*	102.8(2)	- O(4)*	3.599(7)
- O(10)*	2.332(5)	- O(10)	83.9(2)	- O(10)	3.074(7)
- O(4)*	2.328(5)	- O(4)	82.5(2)	- O(4)	3.262(7)
- O(10)	2.532(6)	- O(12)	79.5(2)	- O(12)	3.313(7)
- O(4)	2.655(6)	- O(10)*	83.2(2)	O(7)	- O(10)* 3.202(7)
- O(12)	2.385(7)	- O(4)*	84.1(2)	- O(4)*	3.089(7)
Mean	2.465	- O(10)	98.1(2)	- O(10)	3.480(7)
		- O(4)	91.1(2)	- O(4)	3.534(7)
		- O(12)	93.7(2)	- O(12)	3.377(10)
O(10)*- Ca	O(4)*	O(10)*- Ca	O(4)*	O(4)*	2.980(6)

	α_{eff}	α_{eff}	α_{eff}	α_{eff}
$O(4)^*$ - Ca	- O(4) 111.7(2)		- O(4) 4.292(8)	
	- O(12) 57.0(2)		- O(12) 2.585(8)	
$O(10)$ - Ca	- O(10) 99.7(2)	$O(4)^*$ - O(10) 3.555(8)		
	- O(12) 132.2(2)		- O(12) 4.748(9)	
$O(4)$ - Ca	- O(4) 73.2(2)	$O(10)$ - O(4) 2.980(5)		
	- O(12) 55.7(2)	$O(4)$ - O(12) 2.579(8)		

residual was noted when an extinction condition was included, and the final difference-map showed no maxima greater than $1\text{ e}^-/\text{\AA}^3$.

Final positional and thermal parameters are presented in Table 7, and selected bond-lengths and angles, in Table 8. Observed and calculated structure-factors are available from the Depository of Unpublished Data, CISTI, National Research Council of Canada, Ottawa, Ontario K1A 0S2.

DESCRIPTION OF THE STRUCTURE

The structure of phosinaite-(Ce) may be best described as a series of mixed layers of tetrahedra and octahedra that are stacked along a (Fig. 1). Two different types of layers are present: a (SiP) layer with the composition $[\text{Na}(\text{SiO}_4)_2(\text{PO}_4)]$, and a (NaP) layer with the composition $[\text{Na}(\text{PO}_4)]$; there are two (SiP) layers on each side of a central (NaP) layer.

The (SiP) layer has SiO_4 and PO_4 tetrahedra that are cross-linked by $\text{Na}(7)\text{O}_6$ octahedra. Within the layer, two types of four-membered rings are present: Si_4O_{12} rings and mixed $\text{Na}(2)\text{O}_6\text{-PO}_4$ rings, which are both centered about a $\bar{2}$ axis and alternate along a . These rings form channels along a , the cores of which are occupied by REEO_8 and $\text{Na}(8)\text{O}_8$ polyhedra. The remaining NaO_6 octahedra and both CaO_7 and $(\text{Na,Ca})\text{O}_7$ polyhedra are positioned in the interstices between the rings. The (NaP) layer consists of PO_4 tetrahedra corner-linked to $\text{Na}(2)\text{O}_6$ octahedra, each PO_4 group sharing three of its corners with four $\text{Na}(2)\text{O}_6$ octahedra. Linking of the (SiP) and (NaP)

layers along a is accomplished among corner-sharing tetrahedra and octahedra, reinforced by edge- and face-sharing Na, Ca and REE polyhedra. Essential ordering of Si and P is indicated by refinement of their scattering powers [Si(1) 96(1)%, Si(2) 98(1)%, P(1) 99(1)%, and P(2) 99(1)%], as constrained by results of electron-microprobe analyses (Table 2), calculated bond-valence sums (Table 7) and mean T-O bond distances (Table 8), and is most likely driven by considerations of charge and polymerization (McDonald 1992). The fact that Si and P are well ordered in phosinaite-(Ce) suggests that it should be classified as a silicophosphate (McDonald 1992, McDonald *et al.* 1994) rather than a disordered silicate or phosphate. A more in-depth discussion of the crystal chemistry of silicophosphates will be presented in a forthcoming paper. Although phosinaite-(Ce) exhibits the strong layering typical of all silicophosphates, it is only the second one known to contain both SiO_4 and PO_4 tetrahedra in the same layer [the other being clinophosinaite, $\text{Na}_3\text{Ca}(\text{SiO}_3)(\text{PO}_4)$; McDonald 1992].

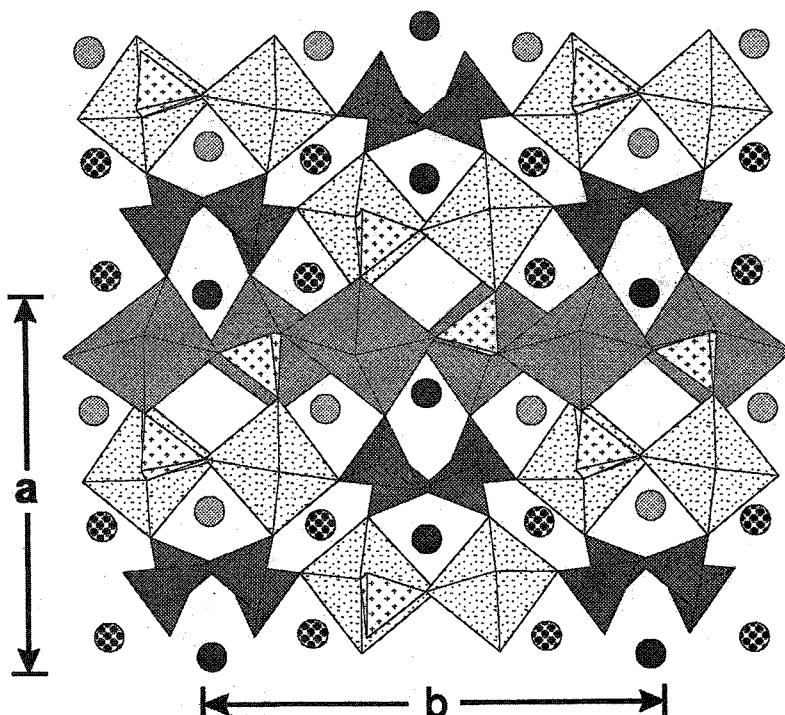


FIG. 1. The crystal structure of phosinaite-(Ce) projected down c . The (NaP) and (SiP) layers are stacked parallel to a . The SiO_4 tetrahedra are heavily shaded, the PO_4 tetrahedra are filled with crosses, the $\text{Na}(2)\text{O}_6$ octahedra are moderately shaded and the $\text{Na}(7)\text{O}_6$ octahedra are lightly stippled. Calcium ions are lightly stippled circles, Na ions filled with large dots, and REE ions heavily shaded.

The four-membered silicate rings present in phosinaite-(Ce) are composed of two independent SiO_4 tetrahedra whose apical directions alternately point up and down in a {UDUD} manner. They have point symmetry 2, deviating considerably from the maximum 4/mmm symmetry observed in the four-membered silicate rings in baotite [$\text{Ba}_4\text{Ti}_8\text{O}_{16}(\text{Si}_4\text{O}_{12})\text{Cl}$; Nekrasov *et al.* 1970], where the Si atoms are all coplanar. Silicate rings are, in general, energetically less favorable than comparable chain motifs, because of the repulsive forces acting on adjacent tetrahedra (Liebau 1985). As such, silicate rings will develop only if a stabilizing factor, capable of reducing the net repulsive forces acting on the tetrahedra, is present. Such factors typically include the presence of highly electronegative cations (*e.g.*, Pb, Cu, Sn), additional ligands [*e.g.*, H_2O , CO_3 , (OH)], and the absence of cations that tend to form rigid coordination polyhedra (*e.g.* Mg, Fe, Al, Li; Liebau 1985). Since none of these features can be called upon to explain the presence of the Si_4O_{12} rings in phosinaite, we instead postulate that the {UDUD} topology of these rings slightly increases the distance between adjacent Si atoms, thus reducing the associated repulsive forces and stabilizing the ring configuration.

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