NEW Cu²⁺ COORDINATION POLYHEDRA IN THE CRYSTAL STRUCTURE OF BURNSITE, KCdCu₇O₂(SeO₃)₂Cl₉

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Abstract

The crystal structure of burnsite, KCdCu₇O₂(SeO₃)₂Cl₉, a new mineral species from fumaroles of the Tolbachik volcano, Kamchatka, Russia, has been solved by direct methods and refined on the basis of F^2 for all unique data collected with monochromatic MoK α X-radiation and a CCD (charge-coupled device) detector to an agreement factor (*R*1) of 1.9%, calculated using the 460 unique observed reflections ($|F_0| = 4\sigma_F$). The mineral is hexagonal, space group $P6_3/mmc$, a 8.7805(8), c 15.521(2) Å, V 1036.3(2) Å³, Z = 2. There are two non-equivalent Cu²⁺ cations in the structure of burnsite. The Cu(1) cation is coordinated by three atoms of O and three of Cl; this is the first example of this type of mixed-ligand Cu²⁺ ϕ_6 (ϕ : O²⁻, Cl⁻) octahedron in a mineral. The octahedron is (2 + 4)-distorted owing to the Jahn–Teller effect, in contrast to the vast majority of Cu²⁺ oxysalt minerals that contain (4 + 2)- distorted Cu²⁺ octahedra. The Cu(2) cation is coordinated by two O and three Cl anions arranged at the vertices of a trigonal bipyramid; this is the first occurrence of a mixed-ligand Cu²⁺O₂Cl₃ trigonal bipyramid in a mineral. The structure contains one symmetrically independent K cation that is coordinated by six Cl anions in a trigonal prismatic arrangement. There is one unique Se⁴⁺ cation that is strongly bonded to three O anions on one side of the cation, owing to the presence of a s^2 lone-electron pair. The structure is described both in terms of cation-centered polyhedra and oxocentered OCu₄ tetrahedra. The structure of burnsite is closely related to the structures of other natural copper oxide chloride selenites discovered in the Tolbachik fumaroles (chloromenite, georgbokiite, ilinskite).

Keywords: burnsite, crystal structure, copper selenites, oxocentered tetrahedra.

Sommaire

Nous avons résolu la structure cristalline de la burnsite. KCdCu₇O₂(SeO₃)₂Cl₀, nouvelle espèce minérale provenant des fumerolles du volcan Tolbachik, péninsule de Kamchatka, en Russie, par méthodes directes, et nous l'avons affiné en utilisant les facteurs F^2 des réflexions uniques prélevées avec rayonnement monochromatique MoK α et un détecteur CCD à couplage de charges, jusqu'à un facteur de concordance (R1) de 1.9%, calculé pour les 460 réflexions uniques observées ($|F_0| = 4\sigma_F$). Il s'agit d'un minéral hexagonal, groupe spatial $P6_3/mmc$, a 8.7805(8), c 15.521(2) Å, V 1036.3(2) Å³, Z = 2. La structure contient deux cations non équivalents de Cu^{2+} . Le cation Cu(1) est coordonné à trois atomes de O et trois de Cl; c'est en fait le premier exemple de ce type de ligand mixte $Cu^{2+}\phi_6(\phi; O^{2-}, Cl^{-})$ octaédrique dans un minéral. L'octaèdre montre une distortion (2+4) due à l'effet Jahn-Teller, contrairement à la grande majorité d'oxysels de Cu²⁺, minéraux qui contiennent des octaèdres à Cu²⁺ montrant la distortion (4 + 2). Le cation Cu(2) est coordonné à deux atomes de O et trois anions de Cl aux coins d'une bipyramide trigonale. C'est le premier exemple d'un ligand mixte $Cu^{2+}O_2Cl_3$ en bipyramide trigonale dans un minéral. Le cation unique Cd^{2+} est symétriquement indépendant et coordonné à six anions de Cl situés aux coins d'un octaèdre régulier. La structure contient un cation K symétriquement indépendant, coordonné par six anions Cl dans un agencement prismatique trigonal. Il y a un cation Se⁴⁺ unique avec fortes liaisons à trois atomes O d'un côté du cation, à cause de la présence d'une paire isolée d'électrons s^2 . Nous décrivons la structure à la fois en termes de polyèdres entourant un cation et de tétraèdres OCu₄ ayant un atome d'oxygène au centre. La structure de la burnsite montre des ressemblances étroites avec la structure d'autres sélénites chlorurés naturels de cuivre découverts dans les fumerolles à Tolbachik (chloroménite, georgbokiite, ilinskite).

(Traduit par la Rédaction)

Mots-clés: burnsite, structure cristalline, sélénites de cuivre, tétraèdres oxo-centrés.

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INTRODUCTION

Burnsite, $KCdCu_7O_2(SeO_3)_2Cl_9$, is a new potassium cadmium copper oxide chloride selenite mineral species that was recently described by Krivovichev *et al.* (2002) from the fumaroles of the Tolbachik volcano, Kamchatka, Russia. It forms aggregates of dark red transparent isometric grains up to 0.1 mm in maximum dimension. The crystal structure, which provides the first mineral examples of two new Cu²⁺ coordination polyhedra, is reported herein.

EXPERIMENTAL

X-ray diffraction

A single crystal of burnsite was selected by optical examination and mounted on a Bruker PLATFORM three-circle goniometer equipped with a 1K SMART CCD (charge-coupled device) detector and a crystal-to-detector distance of 5 cm. More than a hemisphere of three-dimensional intensity data was collected using monochromatic Mo $K\alpha$ X-radiation and frame widths of 0.3° in ω . The unit-cell dimensions (Table 1) were refined from 2697 reflections using least-squares techniques. The data were integrated and corrected for Lorentz, polarization and background effects using the

TABLE 1. CRYSTALLOGRAPHIC DATA FOR BURNSITE

- (8)	0 7005/0	(1)	10.00100			
a (Å)	8.7805(8)	c (Å)	15.521(2)			
V (Å')	1036.3(2)	Space group	P6 ₃ /mmc			
F_{acc}	1110					
Crystal size (mm)	0.09 × 0,08 × 0,04	Radiation	ΜοΚα			
Total reflections	6211	Unique reflections	532			
Unique $F_p \ge 4\sigma_p$	460					
Ri	0,019	wR2	0.046			
\$	0.99					
Formula	KCdCu ₇ O ₂ (SeO ₄) ₂ Cl ₂					
Z	2	μ (cm ⁻¹)	129.4			
D_{one} (g/cm ³)	3.85	,				
$R1 = \Sigma \left(F_{a} - F_{b} \right)$	/Σ⊮J×100					
	$F_{c}^{2})^{2}/\Sigma w(F_{c}^{2})^{2}]^{n}$					
$w = 1/[\sigma^2(F_0^2) + ($	$(0.0314 \times P)^2$, $P = [max(F_{12})^2$	$(9) + 2 \times F_r^2 / 3$				
$S = [\Sigma w(F_d - F_d)^2 / (m - n)]^{1/2}$, for 532 observations and 36 parameters						

Bruker program SAINT. A semi-empirical correction for absorption was done using 2003 intense reflections with the crystal modeled as an ellipsoid; the R_{int} was lowered from 8.36 to 3.97%. Additional information pertinent to the data collection is given in Table 1.

Structure solution and refinement

Scattering curves for neutral atoms, together with anomalous-dispersion corrections, were taken from International Tables for X-Ray Crystallography, Vol. IV (Ibers & Hamilton 1974). The Bruker SHELXTL Version 5 system of programs was used for the determination and refinement of the structure. It was solved by direct methods and refined on the basis of F^2 for all unique data. The final refinement included all atomic positional parameters not restricted by symmetry, with an allowance for anisotropic displacement of all atoms, and a weighting scheme of the structure factors. The refinement converged to an agreement index (R1) of 1.9%, calculated for the 460 unique observed reflections $(F_{0} > 4\sigma_{F})$. Final atomic parameters and selected interatomic distances are presented in Tables 2 and 3. respectively. Observed and calculated structure-factors are available from the Depository of Unpublished Data, CISTI, National Research Council, Ottawa, Ontario K1A 0S2, Canada.

DESCRIPTION OF THE CATION POLYHEDRA

The coordination environments of the various cations in the structure of burnsite are illustrated in Figure 1. The coordination polyhedra about the two symmetrically distinct Cu^{2+} cations are most unusual. The Cu(1)site is coordinated by three O and three Cl anions arranged at the vertices of a distorted octahedron. Burns & Hawthorne (1995b) examined the geometries of mixed-ligand $Cu^{2+}\varphi_6$ (φ : O^{2-} , OH^- , H_2O , CI^-) octahedra, with an emphasis on the stereochemical details of observed octahedra and prediction of the polyhedron geometries using Hartree–Fock calculations. The $Cu(1)\varphi_6$ octahedron in burnsite is the first example in a mineral of a mixed-ligand $Cu^{2+}\varphi_6$ octahedron that con-

TABLE 2. ATOMIC COORDINATES AND DISPLACEMENT PARAMETERS FOR BURNSITE

	х	y	I	*U _{eq}	$U_{\gamma\gamma}$	U_{22}	U_{M}	U_{23}	U_{13}	U_{12}
Cd	0	0	υ	0.0205(2)	0.0212(2)	0.0212(2)	0.0192(3)	0	0	0.0106(1)
Se	-2/3	-14	0.09730(4)	0.0131(2)	0.0147(2)	0.0147(2)	0.0099(3)	0	0	0.0073(1)
Cu(1)	-0.45264(3)	-0.54736(3)	0.09057(3)	0.0157(2)	0.0200(2)	0.0200(2)	0.0137(2)	0.0013(1)	-0.0013(1)	0.0150(2)
Cu(2)	-1/3	-245	14	0.0134(2)	0.0169(3)	0.0169(3)	0.0063(5)	0	0	0.0085(2)
ĸ	0	0	14	0.0609(9)	0.080(1)	0.080(1)	0.023(1)	0	0	0.0398(7)
CI(1)	-0.13697(8)	-0.2739(1)	0.10178(7)	0.0402(3)	0.0571(6)	0.0221(5)	0.0299(6)	0.0041(5)	0.0020(2)	0.0110(3)
CI(2)	-0.4945(1)	-0.5055(1)	14	0.0352(4)	0.0433(7)	0.0433(7)	0.0370(8)	0	0	0.0350(8)
O(1)	-1/3	-2/3	0.1267(2)	0.0112(8)	0.012(1)	0.012(1)	0.010(2)	0	0	0.0058(6)
<u>O(2)</u>	-0.5696(2)	-0.1392(3)	0.0447(2)	0.0210(6)	0.026(1)	0.011(1)	0.021(1)	-0.001(1)	-0.0005(5)	0.0053(6)

* $U_{eq} = (1/3) \Sigma \Sigma U_{ij} a_i^* a_j a_j^*$

tains three Cl anions; such an octahedron occurs in the synthetic compound Cu(OH)Cl (Effenberger 1981). Examples of Cu ϕ_6 octahedra with one, two and four Cl are listed in Burns & Hawthorne (1995b).

The distribution of bond lengths in the $Cu(1)\phi_6$ octahedron is worthy of comment. The expected Cu^{2+} –O

TABLE 3. SELECTED INTERATOMIC DISTANCES (Å) IN THE STRUCTURE OF BURNSITE

Cd-Cl(1),a,b,c.d,c	2.614(1) 6x	K-CI(1),i,d,k.l,e	3.104(1) 6x
Se-O(2),f.g	1.687(3) 3x	O(1)-Cu(1),h,j O(1)-Cu(2)	1.899(1) 3x 1.914(4)
Cu(1)-O(1)	1.899(1)	<o(1)-cu></o(1)-cu>	1.90
Cu(1) O(2)g	1.916(3)		
Cu(1)-O(2)e	2.126(3)	Cu(1)-Cu(1)h.j	3.1426(8) 2x
Cu(1)-Cl(2)	2.5554(6)	Cu(1)h-Cu(1)j	3.1426(8)
Cu(1)/Cl(1),h	2.612(1) 2x	Cu(2)-Cu(1).h.j	3.0685(5) 3x
		<cu-cu></cu-cu>	3.106
Cu(2)-O(1),i	1.914(4) 2x		
Cu(2)-Cl(2).h.j	2.452(2) 3x		

$$\begin{split} a &= y_2 \mathbf{1}, x_2 y, z_3 \mathbf{b} = (x + y_1 + x_2), z_3 \mathbf{c} = (x_1 y_1 + x_2 + z_3) \mathbf{d} = (x_1 y_2 + z_2 + 1/2) \mathbf{c} = (y_2 \mathbf{1}, x_2 y_2 + z_3 + z_3) \mathbf{d} = (y_1 \mathbf{1}, x_2 y_2 + z_3 + z_3) \mathbf{d} = (y_1 \mathbf{1}, x_2 y_2 + z_3 + z_3) \mathbf{d} = (y_1 \mathbf{1}, x_2 y_1 + z_3) \mathbf{d} = (y_1 \mathbf{1}, x_2 y_1 + z_3) \mathbf{d} = (y_1 \mathbf{1}, x_2 y_1 + z_3) \mathbf{d} = (y_1 \mathbf{1}, y_2 \mathbf{1}, z_$$

and Cu²⁺-Cl bond lengths from sums of effective ionic radii are 2.09 and 2.40 Å, respectively [Cu²⁺ = 0.73 Å, $O^{2-} = 1.36 \text{ Å}$ (Shannon 1976); $Cl^{-} = 1.67 \text{ Å}$ (Whittaker & Muntus 1970)]. However, as noted by Burns & Hawthorne (1995b), mixed-ligand $Cu^{2+}\phi_6$ octahedra invariably involve bond-length distortion owing to the Jahn-Teller effect (Jahn & Teller 1937). Although a holosymmetric coordination about the Cu2+ cation is impossible because the ligands are mixed, a pseudoholosymmetric octahedral geometry would involve a near-degenerate electronic state, and either an elongation [(4+2) distortion] or a compression [(2+4) distortion] of the ligand geometry will likely occur, removing the near-degenerate electronic state and resulting in a net stabilization of the octahedron. The Cu(1) ϕ_6 octahedron involves two compressed Cu2+-O bond lengths of 1.899(1) and 1.916(3) Å, with the corresponding ligands in a *trans* arrangement. There is also one Cu²⁺-O bond-length of 2.126(3) Å, which is longer than the expected distance for an undistorted octahedron. The three Cu(1)-Cl bond lengths range from 2.5554(6) to 2.612(1) Å, and are thus longer than the expected distance for an undistorted octahedron. The four elongate $Cu^{2+}-\phi$ bonds extend to ligands in the equatorial posi-

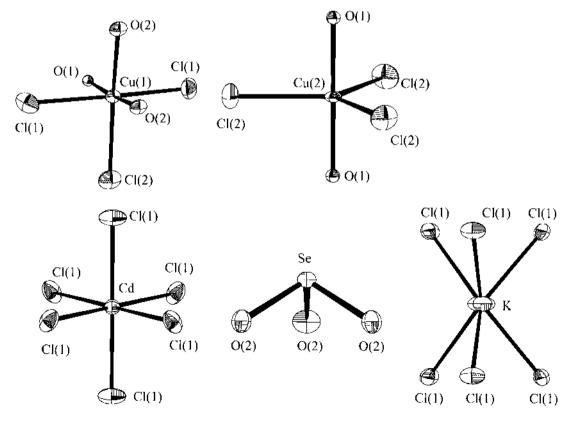


FIG. 1. Cation coordination polyhedra in the structure of burnsite.

tions of the octahedron, with the compressed bonds extending to the two apical ligands of the octahedron. The $Cu(1)\phi_6$ octahedron may therefore be described as (2 + 4)-distorted, which is extremely unusual in Cu²⁺ oxysalt minerals (Burns & Hawthorne 1996).

Burns & Hawthorne (1996) proposed that some Cu²⁺ octahedra in minerals may involve a dynamic Jahn-Teller distortion, in which case the structure determined by X-ray diffraction reveals coordination polyhedra with time-averaged geometries. Given the rarity of (2 + 4)distorted Cu²⁺ octahedra in minerals, is it possible that the Cu(1) ϕ_6 octahedron in burnsite involves a dynamic distortion? If this is the case, the dynamic interchange of elongation directions of (4 + 2)-distorted octahedra may result in an apparent (2 + 4)-distorted octahedron. Dynamic Jahn-Teller distortion of an octahedron involves movement of the ligands approximately parallel to the bond lengths, and thus should be reflected by anisotropic displacement ellipsoids that are elongate (sub)parallel to the bond direction. Inspection of the displacement ellipsoids for the ligands of the Cu(1) octahedron (Fig. 1) shows that they are somewhat elongate (sub)perpendicular to the bond direction, which is inconsistent with a dynamic distortion of the octahedron. The Cu(1) ϕ_6 octahedron thus appears to be a persuasive example of a (2 + 4)-distorted octahedron.

The Cu(2) cation is coordinated by five ligands in a symmetric trigonal bipyramidal arrangement. The coordination polyhedron contains two O anions located in the apical positions, and three of Cl in the equatorial positions. Although Cu²⁺ occurs in trigonal prismatic coordination in several minerals (Burns & Hawthorne 1995a), burnsite is the first mineral to contain a mixed-ligand Cu²⁺O₂Cl₃ trigonal bipyramid.

There is a single symmetrically distinct Cd^{2+} cation that is coordinated by six Cl(1) anions arranged at the

vertices of an octahedron. The K cation is coordinated by six Cl(1) anions in a trigonal prismatic arrangement. The structure contains one symmetrically distinct Se⁴⁺ cation, which is coordinated by three atoms of O that are arranged at the vertices of a trigonal pyramid. Note that all of the O atoms are located on the same side of the Se⁴⁺ cation, owing to the presence of a s^2 lone-electron pair on the other side of the cation.

DESCRIPTION OF THE STRUCTURE

The structure of burnsite is complex in that it involves five different types of cation coordination polyhedra. Here we provide two alternative descriptions of the structure, one based upon cation-centered polyhedra, and the other involving anion-centered tetrahedra. We consider each of these approaches to be of merit; both emphasize important structural features and facilitate comparison with related structures.

Description of the structure using cation-centered polyhedra

The structure of burnsite is shown projected along [001] in Figure 2a, and onto (100) in Figure 2b, with cation polyhedra shown in each case. From these projections, it is apparent that burnsite possesses a complex framework structure. Description of the structure is facilitated by considering first the Cu(1) ϕ_6 octahedra. As illustrated in Figure 3a, Cu(1) ϕ_6 octahedra occur as trimers of octahedra formed by the sharing of edges, with all three octahedra sharing the O(1) vertex. Each trimer is linked to three identical trimers by the sharing of O(2)–O(2) octahedron edges, resulting in layers of octahedra that are parallel to (001) that involve linkages only by the sharing of octahedron edges. The

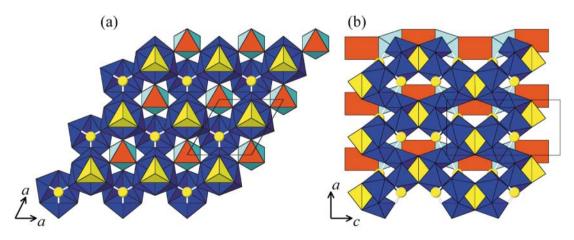


FIG. 2. Polyhedral representations of cation-centered polyhedra in the structure of burnsite. Legend: $Cu(1)\phi_6$ octahedra: dark blue, $Cu(2)O_2Cl_3$ triangular bipyramids: yellow, $CdCl_6$ octahedra: light blue, KCl_6 trigonal prisms: red, and Se atoms: yellow balls.

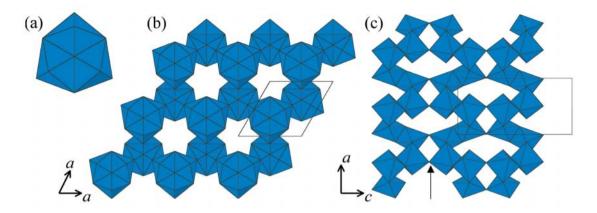


FIG. 3. The framework of $Cu(1)\phi_6$ octahedra in the structure of burnsite. (a) The trimer of edge-sharing $Cu(1)\phi_6$ octahedra, (b) the layer of edge-sharing octahedra parallel to (001), (c) linkage of layers of edge-sharing octahedra to form a framework. The Cl(2) vertex, which is shared between adjacent layers, is indicated by the arrow. The unit cell is outlined in each case.

layers of Cu(1) ϕ_6 octahedra are stacked along [001], and are connected by sharing the Cl(2) vertex (shown by an arrow in Fig. 3c), resulting in a framework of Cu(1) ϕ_6 octahedra that share vertices and edges.

The Cu(2)O₂Cl₃ trigonal bipyramid is located between two translationally equivalent trimers of Cu(1) φ_6 octahedra, which belong to adjacent layers of edge-sharing Cu(1) φ_6 octahedra (Figs. 2a, b). The Cl(2) vertices shared between adjacent layers of Cu(1) φ_6 octahedra constitute the equatorial ligands of the trigonal bipyramid, with the apical ligands corresponding to O(1), which are shared between the three Cu(1) φ_6 octahedra of each trimer. The SeO₃ trigonal pyramids are linked to the trimers of Cu(1) φ_6 octahedra by sharing the O(2) vertices, but on the side opposite the Cu(2)O₂Cl₃ trigonal bipyramid. The SeO₃ trigonal pyramid is located toward the center of individual layers of Cu(1) φ_6 octahedra (Fig. 2b).

The KCl₆ trigonal prisms and CdCl₆ octahedra are linked to form columns that are parallel to [001] by the sharing of triangular faces (Fig. 2b). The columns extend through the framework of Cu(1) ϕ_6 octahedra, and are linked to the framework by the sharing of each of their ligands with the Cu(1) ϕ_6 octahedra (Figs. 2a, b).

Description of the structure using anion-centered tetrahedra

The two symmetrically distinct atoms of O in the structure of burnsite assume very different chemical roles. The O(2) atom is strongly bonded to a Se⁴⁺ cation, as well as to two Cu(1) cations. The O(1) atom is tetrahedrally coordinated by three Cu(1) and one Cu(2) atoms, thus forming an OCu₄ oxocentered tetrahedron. The OCu₄ tetrahedron is common in structures of Cu²⁺ oxysalts that also contain O atoms that are not part of

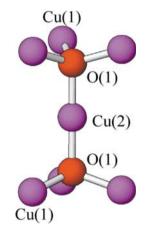


FIG. 4. The O₂Cu₇ finite clusters in the structure of burnsite.

the oxocentered tetrahedra (Effenberger 1985, Filatov *et al.* 1992, Starova *et al.* 1991, 1997, 1998, Gorskaya *et al.* 1992, Krivovichev *et al.* 1998a, 1999b). The mean <O(1)–Cu> bond length in burnsite is 1.90 Å, which is in accord with the typical <O–Cu> bond length in OCu₄ tetrahedra (Effenberger 1985, Filatov *et al.* 1992). The mean <Cu–Cu> distance in the OCu₄ tetrahedron in burnsite is 3.106 Å, in good agreement with the average value of 3.14 Å for such tetrahedra given by Krivovichev & Filatov (1999a).

In the structure of burnsite, two OCu_4 tetrahedra are linked by a common vertex, resulting in the O_2Cu_7 unit shown in Figure 4. The connectivity of this unit is analogous to the Si₂O₇ group in silicates, except that the cations and anions are inverted.

The bond valences associated with cation-anion bonds in the structure of burnsite may be grouped according to their approximate strengths (Table 4): Se⁴⁺-O bonds [1.40 valence units (vu)], Cu²⁺–O bonds [except the Cu(1)–O(2) bond; 0.53-0.55 vu] and other bonds (0.19-0.35 vu). The structure may be described on the basis of the polymerization of those polyhedra of higher bond-valence (Hawthorne 1983). The SeO₃ trigonal pyramid involves the strongest bonds, followed by the Cu²⁺–O bonds. The Cu²⁺–O bonds are of two types, those that are incident upon the O(1) atom (bonds within the OCu₄ tetrahedron), and those that link OCu₄ tetrahedra and SeO₃ trigonal pyramids. The OCu₄ tetrahedra and SeO₃ pyramids are linked to each other according to the "face-to-face" principle (Krivovichev et al. 1999b); their Cu-Cu-Cu and O-O-O triangular

TABLE 4. BOND-VALENCE* (ra) ANALYSIS FOR BURNSITE

Aiom	Cd	к	Cu(1)	Cu(2)	Se(1)	Σ
Cl(1)	0.35%	0.21**	0.19 ²⁴	0.2031		0.75
Cl(2) O(1)			0.22 ²⁻⁴ 0.55 ³⁻⁴	0.30 ³⁴ 0.53 ²⁴		0.7 4 2.18
O(2)			0.53, 0.30		1.4014	2.23
Σ	2.10	1.26	1.98	1.96	4.20	

*calculated using parameters from Brese & O'Kceffe (1991)

faces are parallel to each other and have the same orientation. The Cu²⁺–O and Se⁴⁺–O bonds thus form an extended three-dimensional $[O_2Cu_7](SeO_3)_2$ framework, which is illustrated in Figure 5a. The columns formed by the sharing of triangular faces of CdCl₆ octahedra and KCl₆ trigonal prisms (Fig. 5b) are located within large channels in the $[O_2Cu_7](SeO_3)_2$ framework (Fig. 6).

DISCUSSION

Burnsite, $KCdCu_7O_2(SeO_3)_2Cl_9$, was found in close association with other rare copper oxide selenite chlorides that have been described from the Tolbachik fumaroles: georgbokiite, $Cu_5O_2(SeO_3)_2Cl_2$, chloromenite, $Cu_9O_2(SeO_3)_4Cl_6$, ilinskite, $NaCu_5O_2(SeO_3)_2Cl_3$, and an unnamed selenite with the ideal formula $[Cu^+Cl_2]Pb$ $Cu_5O_2(SeO_3)_2Cl_3$. Each of these minerals involve structures based upon structural units composed of oxocentered tetrahedra. However, the coordination polyhedra about Cu^{2+} are variable in these structures, and include square planar arrangements, triangular bipyramids and octahedra.

The OCu₄ oxocentered tetrahedra combine in a variety of ways in the structures of Tolbachik minerals. In burnsite, two OCu₄ tetrahedra form O_2Cu_7 finite clusters, whereas in chloromenite, OCu₄ tetrahedra are

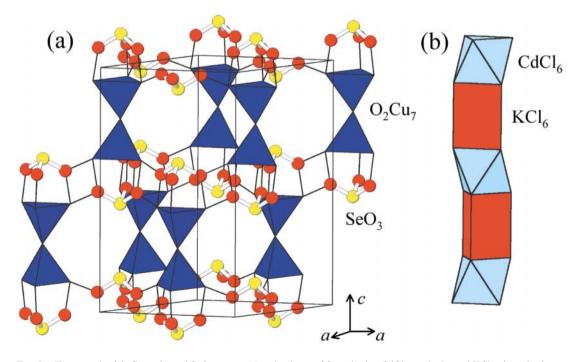


FIG. 5. Framework of O_2Cu_7 units and SeO_3 groups (a) and column of face-sharing CdCl₆ octahedra and KCl₆ trigonal prisms (b) in the structure of burnsite.

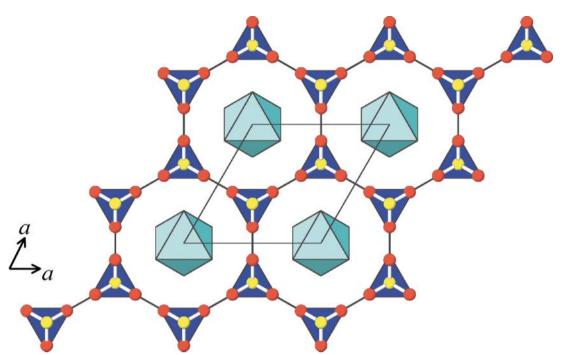


FIG. 6. Projection of the crystal structure of burnsite along the c axis showing the framework of O_2Cu_7 units and SeO_3 groups and the column of face-sharing CdCl₆ octahedra and KCl₆ trigonal prisms. The legend is the same as in Figure 5.

linked by corner-sharing into O₂Cu₆ chains (Krivovichev et al. 1998a). Identical chains with a formula O_2Cu_5Pb occur in the structure of $[Cu^+Cl_2]PbCu_5O_2$ (SeO₃)₂Cl₃ (unpublished results). In georgbokiite, OCu₄ tetrahedra are linked by sharing vertices and edges, resulting in O₂Cu₅ chains (Krivovichev et al. 1999a). The structure of ilinskite is based upon O₂Cu₅ sheets of corner-sharing OCu₄ tetrahedra. Note that with the sole exception of [Cu⁺Cl₂]PbCu₅O₂(SeO₃)₂Cl₃, this assemblage of minerals can be synthesized using chemical transport reactions (Galy et al. 1979, Shuvalov et al. 2000a, b, c, Bastide et al. 2000). In these experiments, crystals of the mineral analogues are formed directly from the gas phase. As noted by Filatov et al. (1992), the ability to synthesize these compounds by vapor transport provides evidence that oxocentered complexes may exist in gas phases, and they may play an important role in the transport of Cu and other metals involved in mineral-forming processes within fumaroles.

The description of structures of minerals and inorganic compounds in terms of anion-centered tetrahedra provides a useful tool for understanding occurrences of complex minerals such as those in the fumaroles of the Tolbachik volcano. For further references on anion-centered tetrahedra, see Krivovichev *et al.* (1998b, 2001), Krivovichev & Burns (2000a, b, 2001a, b, 2002), Li *et al.* (2000, 2001), Hawthorne *et al.* (2000) and Krivovichev & Filatov (2001).

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