*The Canadian Mineralogist* Vol. 42, pp. 1037-1045 (2004)

# THE CRYSTAL STRUCTURE OF CALCIUM CATAPLEIITE

# STEFANO MERLINO<sup>§</sup>, MARCO PASERO AND MARCO BELLEZZA

Department of Earth Sciences, University of Pisa, via S. Maria 53, I-56126 Pisa, Italy

### DMITRY YU. PUSHCHAROVSKY, ELENA R. GOBETCHIA, NATALIA V. ZUBKOVA AND IGOR V. PEKOV

Faculty of Geology, Moscow State University, Vorobievy Gory, RU-119992 Moscow, Russia

### Abstract

Calcium catapleiite, CaZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, the calcium-dominant analogue of catapleiite, Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, firstly described as hexagonal, space group *P*6<sub>3</sub>/*mmc*, with *a* 7.32, *c* 10.15 Å, is actually orthorhombic, space group *Pbnn*, *a* 7.378(1), *b* 12.779(1), *c* 10.096(1) Å, *V* 951.89 Å<sup>3</sup> and *Z* = 4. The crystals display three-fold twinning ("Drilling") about the **c** axis, with the three individuals rotated 120° one from the other. The crystal structure of calcium catapleiite from the type locality, the Burpala massif, northern Baikal Region, Siberia, Russia, has been determined by direct methods from single-crystal X-ray-diffraction data and refined to a final *R* = 0.0528 (*wR*<sup>2</sup> = 0.1604) for 2167 reflections with *F*<sub>0</sub> > 6 $\sigma$ (*F*<sub>0</sub>). It is characterized by a mixed framework of regular [ZrO<sub>6</sub>] octahedra and three-membered silicate rings [Si<sub>3</sub>O<sub>9</sub>], very similar to the framework of catapleiite. The H<sub>2</sub>O molecules also are similarly placed in the two minerals. The Ca<sup>2+</sup> cations in calcium catapleiite are distributed over two distinct sites, with 0.8 and 0.2 occupancies, which contrasts with the fully disordered distribution of the Na<sup>+</sup> cations in catapleiite.

Keywords: calcium catapleiite, crystal structure, ring silicate, syenitic pegmatite, Burpala massif, Russia.

### Sommaire

La calcium catapléiite, CaZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, l'analogue à dominance de calcium de la catapléiite, Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, a tout d'abord été décrite comme phase hexagonale, groupe spatial *P*6<sub>3</sub>/*mmc*, avec *a* 7.32, *c* 10.15 Å, mais elle est en fait orthorhombique, groupe spatial *Pbnn*, *a* 7.378(1), *b* 12.779(1), *c* 10.096(1) Å, V 951.89 Å<sup>3</sup> et *Z* = 4. Les cristaux font preuve de macles d'ordre trois (en allemand, "Drilling") autour de l'axe **c**, les trois individus montrant une rotation de 120° l'un par rapport à l'autre. La structure cristalline de la calcium catapléiite provenant de la localité-type, le massif de Burpala, dans la partie nord de la région de Baïkal, en Sibérie, Russie, a été déterminée par méthodes directes avec données obtenues en diffraction X sur monocristaux, et affinée jusqu'à un résidu *R* de 0.0528 (*wR*<sup>2</sup> = 0.1604) en utilisant 2167 réflexions pour lesquelles *F*<sub>0</sub> > 6σ(*F*<sub>0</sub>). La structure contient la même trame mixte d'octaèdres réguliers de [ZrO<sub>6</sub>] et d'anneaux de tétraèdres à trois membres [Si<sub>3</sub>O<sub>9</sub>] que la catapléiite. Les molécules de H<sub>2</sub>O sont aussi placées de façon semblable dans les deux minéraux. Les atomes de Ca<sup>2+</sup> dans la calcium catapléiite sont d'accupation 0.8 et 0.2, ce qui diffère nettement de la distribution complètement désordonnée des ions de sodium dans la catapléiite.

#### (Traduit par la Rédaction)

Mots-clés: calcium catapléiite, structure cristalline, silicate à anneaux, syénite pegmatitique, massif de Burpala, Russie.

## INTRODUCTION

Calcium catapleiite, CaZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, is a rare mineral found in hydrothermal assemblages associated with alkaline complexes. It was first discovered in syenitic pegmatites of the Burpala massif, northern Baikal Region, Siberia, Russia (Portnov 1964), and later found in the following alkaline complexes: Golden Horn in Okanogan County, Washington, U.S.A., Strange Lake in Quebec–Labrador and Ice River, British Columbia, Canada, Norra Karr, Sweden (*cf.* Gaines *et al.* 1997), Cnoc Rhaonastil, Islay, Scotland, U.K. (Preston *et al.* 1998), and the Afrikanda complex, Kola Peninsula, Russia (Chakhmouradian & Zaitsev 2002). In this article, we provide definitive information on the crystal structure of calcium catapleiite.

<sup>§</sup> *E-mail address*: merlino@dst.unipi.it

At the type locality, calcium catapleiite is found in aegirine–microcline pegmatite veins cross-cutting syenite along the northern and western margin zone of the Burpala massif. It occurs in the axial parts of veins (usually in cavities) as light yellow to cream-colored grains and lamellar crystals a few centimeters in size. Associated minerals include pyrophanite, pyrochlore, leucophanite, Ti-rich låvenite, loparite-(Ce), kupletskite, REE-bearing apatite, and Ca-enriched seidozerite. Some crystals of calcium catapleiite are partially replaced by aegirine and Ca-rich seidozerite (Portnov 1964).

# **REVIEW OF PREVIOUS INVESTIGATIONS**

The crystal structure of catapleiite Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O was described by Brunowsky (1936) in the space group P6<sub>3</sub>/mmc, a 7.40, c 10.05 Å. As reported by Vlasov (1966), both optically uniaxial and biaxial crystals exist, and different crystallographic systems (monoclinic, hexagonal, orthorhombic) have been given by various authors, even before the development of the X-ray studies. Through measurements on catapleiite crystals at ordinary and elevated temperatures, Brögger (1890) established the existence of two modifications: monoclinic (pseudohexagonal) at ordinary temperature and hexagonal at 140°C (Vlasov 1966). The structure of the monoclinic modification has been determined and refined by Chao et al. (1973) in the space group I2/c, a 12.779, b 7.419, c 20.157 Å, β 90.41°, and by Ilyushin et al. (1981) in the space group B2/b, with a 23.917(2), b 20.148(2), c 7.432(1) Å, γ 147.46(1)°. The setting of Chao et al. (1973), metrically almost orthorhombic and pseudo-hexagonal ( $a \approx b\sqrt{3}$ ), may be obtained from that of Ilyushin et al. (1981) through the transformation matrix [110/001/010]. The monoclinic structure, as described by Ilyushin et al. (1981), and the hexagonal one, as described by Brunowsky (1936) are essentially the same: they contain a mixed framework built up of [ZrO<sub>6</sub>] octahedra and [Si<sub>3</sub>O<sub>9</sub>] rings of tetrahedra; however, the sites occupied by Na atoms and H2O molecules are exchanged in the monoclinic model with respect to the hexagonal one.

The synthetic analogue of catapleiite was obtained in the hydrothermal system  $ZrO_2$ -SiO\_2-Na<sub>2</sub>O-H<sub>2</sub>O (P = 700 bars, 300° < T < 700°C) by Baussy *et al.* (1974), who indicated a hexagonal cell, space group *P*6<sub>3</sub>/*mmc*, with *a* 7.388(2), *c* 10.068(3) Å. In an investigation of the thermal behavior of the synthetic analogue of catapleiite (Ilyushin & Dem'yanets 1988), the reversible phase-transition from the monoclinic modification to the hexagonal one was established to occur at T = 155°C (the reverse transformation occurs at 145°C).

Calcium catapleiite, CaZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, was considered to be an end member of a solid-solution series with catapleiite, Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O, and consequently was originally described as presenting "the crystal lattice of catapleiite" (Portnov 1964), which, at the time, according to the indication of Brunowsky (1936), was assumed to be hexagonal, space group *P6<sub>3</sub>/mmc*. However, the subsequent structural investigations carried out on catapleiite with modern X-ray methods (Chao *et al.* 1973, Ilyushin *et al.* 1981) did not confirm that indication and suggested that the hexagonal symmetry of catapleiite as claimed could be not correct (Piatenko *et al.* 1999), probably the result of an involuntary omission of a set of very weak reflections and of pervasive polysynthetic twinning [a careful study of the multiple twinning in catapleiite was made by Chen & Chao (1974)].

For these reasons, the crystallographic data normally reported for calcium catapleiite, indicating that it is hexagonal, space group  $P6_3/mmc$ , a 7.32, c 10.15 Å (cell parameters calculated by analogy with catapleiite in ICDD 13–371), are not reliable. This lack of reliable data concerning the metrics and the symmetry of calcium catapleiite, as well as our interest in deriving its actual structural arrangement and defining the difference between it and that of catapleiite, stimulated us to carry out a characterization by XRD using some fragments kindly provided to us by N.V. Vladykin, taken from Portnov's holotype specimen.

### EXPERIMENTAL

Chemical analyses were carried out by means of a Camebax MBX electron microprobe equipped with a LINK 10000 energy-dispersion X-ray spectrometer, operating at 10 kV and 10 nA. The electron beam was rastered over an area of  $10 \times 10 \,\mu\text{m}^2$  to minimize damage to the sample. We used the following standards: albite (Na), orthoclase (K), amphibole (Ca, Si), and wadeite (Zr). The results obtained (Table 1) are very close, at least for the main constituents, to the data published by Portnov (1964). Results of our analysis correspond to the following empirical formula, calculated on

TABLE 1. CHEMICAL COMPOSITION OF CALCIUM CATAPLEIITE FROM THE BURPALA MASSIF, RUSSIA

	1	2
Na <sub>2</sub> O wt.%	0.32	0.38 (0.15-0.44)
K <sub>2</sub> Õ	0.10	0.05 (0.00-0.08)
CaO	13.82	13.59 (13.32-13.74)
REE <sub>2</sub> O <sub>3</sub>	0.28	b.d.l.
Fe <sub>2</sub> O <sub>3</sub>	0.36	b.d.l.
Al <sub>2</sub> O <sub>3</sub>	0.60	b.d.l.
SiO,	44.49	44.99 (44.51-45.39)
TiO,	0.06	b.d.1.
ZrO <sub>2</sub>	31.00	31.03 (30.58-31.33)
$H_{2}O^{+}$	9.15	9.01*
H <sub>2</sub> O <sup>-</sup>	0.18	-
Total	100.36	99.05

NOTE: 1: wet-chemical data (Portnov 1964), 2: our electron-microprobe data for the sample under study (average values from four analyses); b.d.l.: below the detection limit; \*: calculated so as to give 2  $H_2O$  molecules per formula unit.

the basis of 11 O atoms:  $(Ca_{0.97}Na_{0.05})_{\Sigma 1.02}Zr_{1.01}Si_{2.99}$  O<sub>9</sub>•2H<sub>2</sub>O.

A crystal fragment (dimensions  $0.2 \times 0.1 \times 0.1$  mm<sup>3</sup>) was mounted on a Nonius Kappa CCD diffractometer, equipped with a MoK $\alpha$  X-ray tube ( $\lambda = 0.71073$  Å) operating at 55 kV and 30 mA. The intensity data were collected with a crystal-to-detector distance of 25 mm. As a first step, 20  $\phi$ -rotation images with framewidths of 1° were collected for the initial orientation matrix and cell setting. A total of 334 frames, using frame-widths of 2° and exposure of 60 s, were collected, corresponding to 93.8% of the reciprocal space. The reflections were collected up to a resolution of 0.50 Å ( $\theta_{max} = 45.29^\circ$ ) in *ca*. 6.5 hours.

On the basis of the structural information available for catapleiite and related structures, we carefully examined the frames for a possible doubling of the c parameter. Unexpectedly, whereas the c parameter was found to be *ca*. 10 Å, we found the *a* parameter to be doubled (ca. 14.8 instead of ca. 7.4 Å). The extraction of the data was therefore carried out assuming a hexagonal cell ( $a \approx 14.8$ ,  $c \approx 10.0$  Å; it will be denoted H2 in the following). The three-dimensional data were reduced using the DENZO-SMN program from Nonius, which is based on the HKL package (Otwinowski & Minor 1997). A semi-empirical absorption correction was accomplished in DENZO-SMN using the redundancy of data measurements. A total of 71220 reflections were extracted from the collected images. The value of  $R_{int}$ , computed on squared structural amplitudes by merging equivalent reflections in the P63/mmc space group after absorption correction, was 6.3%. A set of 30128 independent reflections was obtained.

At this stage, a careful inspection of the intensities of the reflections in the hexagonal cell revealed that all reflections with both h and k even were, on average, about ten times more intense than the other reflections. thus pointing to a subcell H1 with a halved a axis. The whole diffraction-pattern, presented in Figure 1, has been interpreted as resulting from a triple twin (in German, "Drilling") of an orthorhombic cell O, with  $a \approx$ 7.4 Å,  $b \approx \sqrt{3}a$ ,  $c \approx 10.1$  Å, with rotation axis [001]<sub>+120°</sub>. This type of reticular merohedral twinning may also be described as a multiple twin with twin axes  $[110]_{180^\circ}$  and  $[310]_{180^\circ}$ ; the latter two twin laws were already observed, together with four others, by Chen & Chao (1973) in crystals of catapleiite. The relationships among the various reciprocal lattices are illustrated in Figure 1. The reflections in the orthorhombic setting are of two types, namely those that correspond to the reciprocal lattice points of the hexagonal H1 cell (subcell reflections, with  $h_0 + k_0 = 2n$ , indicated with large circles in Fig. 1), and those that are not present in the H1 diffraction pattern (characteristic reflections, with  $h_{\rm O} + k_{\rm O} = 2n + 1$ , indicated with small circles in Fig. 1). The subcell reflections are common to the three orthorhombic individuals, whereas the characteristic reflections are not. To test the "Drilling" hypothesis, we selected a few rather strong *characteristic* reflections  $(n = 10, \text{ with } 2\theta \text{ angles ranging between } 15 \text{ and } 40^\circ)$  and measured their intensities in the three individuals; the ratios of the intensities was found to be constant, indicating that the volume proportions of the three individuals were 50, 40 and 10%. Of course, the structure determination and refinement were carried out by selecting the largest individual, for which the *hkl* indices were adequately converted from the hexagonal H2 setting; moreover, a different scale-factor was assigned to the common *subcell* reflections, to place them on the same scale as the *characteristic* non-superimposed reflections.

The orthorhombic unit-cell has a 7.378(1), b 12.779(1), c 10.096(1) Å, space-group symmetry Pbnn, determined on the basis of the systematic absences. The non-standard space-group was adopted to keep the correspondence with the H1 and H2 settings.

Eventually, a set of 21775 reflections was obtained in the orthorhombic unit-cell. The set was reduced to 4002 independent squared amplitudes ( $R_{int} = 9.81\%$ ). Only those reflections with  $F_o > 6\sigma(F_o)$  were considered as observed, and used in the least-square refinement. A higher-than-usual cutoff was chosen to overcome minor inaccuracies due to the treatment of twinned data. The structure of calcium catapleiite was solved through direct methods using the SIR–92 program (Altomare *et al.* 1994) and refined with the SHELXL–97 software package (Sheldrick 1997).

In total, the positions of one Zr, two Ca, two Si, and six O atoms, one of which belonging to a  $H_2O$  molecule, were determined and refined. In the last cycles of leastsquare refinement (the final index of reliability *R*1 was 0.053 for 2167 observed reflections), the positions of the H atoms belonging to the  $H_2O$  molecule were found in the difference-Fourier map, and refined with loose metrical constraints on the relevant O–H distances through the DFIX option of SHELX–97. Further details concerning the collection of the single-crystal X-ray data and the structure refinement are given in Table 2.

The atom coordinates and site occupancies, displacement parameters, selected interatomic distances, and bond-valence sums (Brese & O'Keeffe 1991, Ferraris & Ivaldi 1988) are reported in Tables 3a, 3b, 4, and 5, respectively. A table of observed and calculated structure-factors is available from the Depository of Unpublished Data, CISTI, National Research Council, Ottawa, Ontario K1A 0S2, Canada. The crystal structure of calcium catapleiite is shown in Figure 2.

Finally, for comparative purposes, the structure of calcium catapleiite was also refined in the hexagonal H1 subcell [space group  $P6_3/mmc$ , a 7.378(1), c 10.096(1) Å], using only those reflections common to the three twin individuals. In this case, the final *R* factor was found to be 4.68% for 704 unique reflections. Despite its slightly higher index of reliability, the orthorhombic model is definitely preferable, as it takes into account those reflections omitted in the refinement of



FIG. 1. Drawing of the reciprocal lattices of the hexagonal cell with a = 7.4 Å (H1, in green), the hexagonal cell with a = 14.8 Å (H2, in blue), and the orthorhombic cell (O, in red) of the "Drilling".

TABLE 2.	INFORMATION ON DATA COLLECTION AND REFINEMENT
	FOR CALCIUM CATAPLEIITE

Chemical formula	CaZrSi <sub>3</sub> O <sub>9</sub> •2H <sub>2</sub> O
Unit-formula weight	393.59
Radiation and wavelength (Å)	ΜοΚα; 0.71069
$\mu$ (mm <sup>-1</sup> )	2.12
Diffractometer, monochrometer	Nonius CCD, graphite
Space group	Pbnn
Z	4
Unit-cell dimensions (Å)	a 7.378(1), b 12.779(1), c 10.096(1)
Unit-cell volume (Å3)	951.89
F(000)	768.0
$D_{calc}$ (g/cm <sup>3</sup> )	2.746
Crystal size (mm <sup>3</sup> )	$0.2 \times 0.1 \times 0.1$
Index ranges	$-12 \leq h \leq 14, -14 \leq k \leq 14, -19 \leq l \leq 20$
Standard reflection	Every 100 reflections
No. collected reflections	21775
No. reflections with $I > 3\sigma(I)$	2167, $R_{int} = 0.0981$
Refinement method	full-matrix least-squares on F2
Weights	$1/[\sigma^2(F_0^2) + (0.0376P)^2 + (8.10P]],$
	$P = [max(F_p)^2 + 2(F_c)^2]/3$
$R(F), wR(F^2), \text{GooF}$	0.0528, 0.1601, 1.025
Number of refined parameters	91
Largest residues (e/A3)	1.54, -5.51

Numbers in parentheses are estimated standard deviations referred to the last digit.

the hexagonal model, and properly describes the Ca distribution in calcium catapleiite, as will be discussed in the next section and is illustrated in Figure 3.

### DISCUSSION

The topology of calcium catapleiite corresponds to that already described for catapleiite (Ilyushin *et al.* 1981). The latter authors refined the structure of catapleiite in a non-standard space group (B2/b) and unconventional monoclinic setting with a very large

obtuse angle  $\gamma = 147^{\circ}$ . As stated above, their unit cell can be transformed into a nearly metrical orthorhombic cell that matches our unit cell, apart from the doubled *c* parameter.

In common with catapleiite, the most specific feature of calcium catapleiite is a mixed framework of regular [ZrO<sub>6</sub>] octahedra and three-membered [Si<sub>3</sub>O<sub>9</sub>] silicate rings. This kind of mixed framework is also displayed in the isotypic pairs of Ti- and Zr-minerals and analogous synthetic compounds: wadeite,  $K_2Zr[Si_3O_9]$ (Henshaw 1955), the synthetic Ti-dominant analogue of wadeite (Choisnet *et al.* 1973), bazirite BaZr[Si<sub>3</sub>O<sub>9</sub>] (Hawthorne 1987), benitoite BaTi[Si<sub>3</sub>O<sub>9</sub>] (Fischer 1969), pabstite Ba(Sn,Ti)[Si<sub>3</sub>O<sub>9</sub>] (Hawthorne 1987), and the synthetic Ca-dominant analogue of pabstite CaSn[Si<sub>3</sub>O<sub>9</sub>] (benitoite structure-type; Nekrasov 1973).

The three-membered  $[Si_3O_9]$  rings in calcium catapleiite are built up of two crystallographically independent  $[SiO_4]$  tetrahedra, one lying on the twofold axis parallel to **b**, the other on a general position. The calcium cations and the H<sub>2</sub>O molecule are located in tunnels formed by the framework of tetrahedra and octahedra and running parallel to [100], [110] and [10]. The relevant bond-distances are presented in Table 4. The H<sub>2</sub>O molecule O(6) is connected through weak hydrogen bonds to O(1) (-x + 3/2, y, -z + 3/2) [O(6)...O(1) distance 3.07 Å] and possibly to O(2) (-x + 1, y + 1/2, z) and O(5) (-x + 1, -y + 1, -z + 2) with a bifurcated hydrogen bond [O(6)...O(2) and O(6)...O(5) distances 3.18 and 3.20 Å, respectively].

It is worthy of note that a complex of tetrahedra in the form of a three-membered ring is quite rare in minerals. It almost exclusively occurs in the structures of cyclosilicates, as in walstromite BaCa<sub>2</sub>[Si<sub>3</sub>O<sub>9</sub>] (Dent THE CRYSTAL STRUCTURE OF CALCIUM CATAPLEIITE



FIG. 2. Projection of the crystal structure of calcium catapleiite on (001), drawn with the program ATOMS (Dowty 1995).



FIG. 3. Positions of the calcium cations in calcium catapleiite, with an indication of the corresponding occupancies, around the columns of [ZrO<sub>6</sub>] octahedra running along **c**. (a) Actual distribution in calcium catapleiite, (b) ordered distribution in calcium catapleiite, and (c) fully disordered distribution of calcium cations, corresponding to the actual distribution of sodium cations in catapleiite (Ilyushin *et al.* 1981) and to the distribution found in the refinement carried out in the hexagonal H1 subcell of calcium catapleiite.

1041

Atom	x	у	Z	Wyckoff multiplicity	SOF
Zr	0.7442(1)	0.2500	0.0000	0.5	1.0
Si(1)	0.4545(1)	0.3467(1)	0.7488(1)	1	1.0
Si(2)	0.2500	0.5528(1)	0.7500	0.5	1.0
Ca(1)	0.5000	0.5000	0.0000	0.5	0.776(4)
Ca(2)	0.7485(7)	0.7500	0.0000	0.5	0.224(4)
O(1)	0.4285(4)	0.4746(2)	0.7374(3)	1	1.0
O(2)	0.2500	0.2949(4)	0.7500	0.5	1.0
O(3)	0.2803(5)	0.6148(3)	0.8839(3)	1	1.0
O(4)	0.5602(6)	0.3030(3)	0.6250(3)	1	1.0
O(5)	0.5434(6)	0.3293(5)	0.8904(3)	1	1.0
0(6)	0.7630(6)	0.5827(4)	0.9143(5)	1	1.0
H(1)	0.885(6)	0.555(8)	0.892(7)	1	1.0
H(2)	0.768(9)	0.581(7)	0.013(1)	1	1.0

# TABLE 3a. FRACTIONAL COORDINATES OF ATOMS, WYCKOFF SITE-MULTIPLICITIES AND OCCUPANCY COEFFICIENTS FOR CALCIUM CATAPLEIITE

Numbers in parentheses are estimated standard deviations referred to the last digit. SOF: site-occupancy factor.

TABLE 3b. DISPLACEMENT PARAMETERS FOR CALCIUM CATAPLEIITE

Atom	U <sub>11</sub>	U <sub>22</sub>	U <sub>33</sub>	U <sub>23</sub>	U <sub>13</sub>	U <sub>12</sub>	$U_{\text{eq}}\!/\!U_{\text{iso}}$
Zr	0.0098(2)	0.0113(2)	0.0066(2)	0.0001(1)	0	0	0.0092(1)
Si(1)	0.0119(3)	0.0154(4)	0.0087(3)	-0.0007(4)	0.0002(4)	0.0024(3)	0.0119(2)
Si(2)	0.0150(5)	0.0108(5)	0.0086(4)	0	-0.0018(6)	0	0.0115(2)
Ca(1)	0.0188(6)	0.0134(5)	0.0147(5)	-0.0025(5)	-0.0046(5)	0.0043(4)	0.0157(3)
Ca(2)	0.0106(14)	0.0213(18)	0.0116(14)	0.0044(14)	0	0	0.0145(8)
O(Ì)	0.0160(11)	0.0147(10)	0.0229(15)	0.0036(11)	0.0021(11)	0.0010(9)	0.0179(5)
O(2)	0.0121(14)	0.0181(16)	0.0354(23)	0	-0.0017(19)	0	0.0219(8)
0(3)	0.0259(16)	0.0266(15)	0.0159(12)	-0.0106(12)	0.0039(11)	-0.0016(12)	0.0228(7)
O(4)	0.0393(21)	0.0317(19)	0.0165(14)	-0.0035(14)	-0.0098(14)	-0.0107(16)	0.0292(8)
0(5)	0.0290(16)	0.0297(16)	0.0115(11)	-0.0039(11)	0.0066(12)	0.0148(14)	0.0234(7)
0(6)	0.0300(21)	0.0390(22)	0.0455(24)	-0.0024(20)	-0.0090(19)	0.0069(19)	0.0382(9)
нà		. ,		. ,			0.09(3)*
H(2)							0.04(2)*

Numbers in parentheses are estimated standard deviations referred to the last digit. \*  $U_{iso}$  values.

Glasser & Glasser 1968), margarosanite PbCa<sub>2</sub>[Si<sub>3</sub>O<sub>9</sub>] (Freed & Peacor 1969) and the high-temperature polymorph (T > 1125°C) of CaSiO<sub>3</sub> ( $\alpha$ -CaSiO<sub>3</sub>, pseudowollastonite; Yang & Prewitt 1999), in  $\alpha$ '-CaSiO<sub>3</sub>, which forms at high pressure (Ilyukhin *et al.* 1979), and in synthetic compounds of the benitoite structure-type: BaTiGe<sub>3</sub>O<sub>9</sub> (Robbins 1960); TlTiGe<sub>3</sub>O<sub>9</sub>, CsBX<sub>3</sub>O<sub>9</sub> (*B* = Ti, Sb; *X* = Si, Ge) (Choisnet *et al.* 1973), in metafluorberillates, with the general formula  $Me^+Me^{2+}$  (BeF<sub>3</sub>)<sub>3</sub> ( $Me^+$  = Rb, NH<sub>4</sub>, Tl;  $Me^{2+}$  = Mg, Ni, Co, Zn) (Favre *et al.* 1969), in NaBaNd[Si<sub>3</sub>O<sub>9</sub>] (Malinovskii *et al.* 1984) and in (NH<sub>4</sub>)<sub>2</sub>ZrGe<sub>3</sub>O<sub>9</sub> (Liu *et al.* 2003), with the wadeite structure-type. Only very rarely do silicate three-membered rings occur in more condensed silicate anions, as in the sheet silicate zussmanite (Lopes-Vieira & Zussman 1969). They never occur in framework silicates (zeolites and zeolite-like materials).

The relatively low stability of three-membered silicate rings reflects the large deviations of the Si-O-Si angles ( $\approx 130^{\circ}$ ) from 147°, which is typical of a ring of tetrahedra with low energy (Chakoumakos et al. 1981). The structure of calcium catapleiite does not represent an exception in this sense: the values of the bond angles Si(1)-O(2)-Si(1)(-x + 1/2, y, -z + 3/2) and Si(1)-O(1)-O(1)-O(1)-O(1)-O(1)-O(1)Si(2) are 132.59° and 133.19°, respectively, similar to values found in the other silicates with three-membered rings. As the Si-O-Si angles narrow, the higher repulsion between the Si atoms and the decreasing s-character of the bridging oxygen (Obr) orbitals (decreasing  $\pi$ -bond strengths) are factors leading to longer Si-O<sub>br</sub> bonds (Gibbs et al. 1972) and lower stability of the ring. According to Chakoumakos et al. (1981), the s-component in the hybrid orbitals of the bridging oxygen can be estimated simply on the basis of the equation d(Si- $O_{br}$ ) =  $-3.5 \times 10^{-3} \delta + 1.78$ , where  $\delta$  is the percent fractional s-character. This s-component in calcium catapleiite is 36.1%, which is significantly lower than in the silicates with four- or six-membered silicate rings (close to 47%) (Pushcharovsky 1986).

The main difference between the structures of catapleiite and calcium catapleiite lies in the occupancy

TABLE 4. SELECTED INTERATOMIC DISTANCES (Å) IN THE STRUCTURE OF CALCIUM CATAPLEIITE

Ca(1)	- O(6)	$2.373(5) \times 2$	Si(1) - O(4)	1.576(4)
	- O(5)	2.467(3) × 2	- O(5)	1.589(3)
	- O(3)	$2.481(4) \times 2$	-O(1)	1.648(3)
	- O(1)	$2.723(3) \times 2$	- O(2)	1.649(2)
<ca(1< td=""><td>1)-0&gt;</td><td>2.511</td><td><si(1) o="" –=""></si(1)></td><td>1.616</td></ca(1<>	1)-0>	2.511	<si(1) o="" –=""></si(1)>	1.616
Ca(2)	– O(6)	$2.310(5) \times 2$	Si(2) - O(3)	$1.582(3) \times 2$
	-0(2)	$2.588(1) \times 2$	-O(1)	$1.658(3) \times 2$
	-O(5)	2.625(6) × 2	<si(2) -="" o=""></si(2)>	1.620
	- 0(4)	$2.710(6) \times 2$	.,	
<ca(2< td=""><td>2) – O&gt;</td><td>2.561</td><td></td><td></td></ca(2<>	2) – O>	2.561		
Zr	- O(4)	$2.033(4) \times 2$		
	- 0(3)	$2.096(3) \times 2$		
	- O(5)	$2.109(3) \times 2$		
$\langle 7r - 0 \rangle$		2 079		

of the larger, [8]-coordinated cationic sites. In catapleiite, Na is statistically distributed among five different sites (four with multiplicity 4 and one with multiplicity 8), all of which present a occupancy. The investigations of the thermal behavior of catapleiite (Ilyushin & Dem'yanets 1988) indicate that there is no variation in the distribution of Na among these sites at higher T; "the transition to the more symmetrical hexagonal phase leads to symmetrization of the framework and to equalization of the bond lengths" for the various Na atoms; thus the transition (at 155°C) is of the displacement kind and "the distances through which the Na atoms move are much smaller than the interatomic distances" (Ilyushin & Dem'yanets 1988). In calcium catapleiite, cations Ca(1) and Ca(2) are distributed between two sites, both with multiplicity 4, with largely different occupancies, 0.8 and 0.2, respectively, as schematically represented in Figure 3a. This figure shows the positions of the calcium cations, with an indication of the corresponding occupancies, around the columns of [ZrO<sub>6</sub>] octahedra running along c. The actual distribution is compared to a perfectly ordered distribution [Fig. 3b: occupancies 1.0 and 0.0 at Ca(1) and Ca(2) sites, respectively] and with a distribution (Fig. 3c) corresponding to that in catapleiite, with occupancy instead of owing to the different cationic charge. In the case of occupancy  $\frac{1}{2}$  in both Ca(1) and Ca(2) sites, the orthorhombic symmetry of calcium catapleiite would be preserved. We may presume that in the case of calcium catapleiite, a phase transformation of the order-disorder type will take place at high temperature, made possible by Ca diffusion along the tunnels of the framework; however, we cannot reliably foresee which would be the final step in the transformation, namely whether it will correspond to the equalization of the occupancies in the Ca(1) and Ca(2) sites with preservation of the orthorhombic symmetry, or to that described in Figure 3c. with transition to a hexagonal form (space group  $P6_3$ / *mmc*) with unit-cell H1. It would be very interesting to carry out a high-temperature investigation to monitor the transformation path and to define the transition tem-

TABLE 5. BOND-VALENCE SUMS (v.u.) FOR CALCIUM CATAPLEIITE

	Zr	Ca(1)*	Ca(2)*	Si(1)	Si(2)	ΣH <sup>v</sup>	$\Sigma_{\mathbf{a}} v$
O(1)		0.101*21		0.935	0.912*21	+0.12	1.960
O(2)			0.042 <sup>×2→×2↓</sup>	0.937 <sup>×2</sup>		+0.05	2.008
O(3)	0.651×21	0.193×21			1.120*21		1.964
O(4)	0.771×21		0.030*21	1.139			1.940
O(5)	0.628*21	0.201 <sup>×21</sup>	0.038 <sup>×2↓</sup>	1.099		+0.05	2.016
O(6)		0.259 <sup>×21</sup>	0.086 <sup>×21</sup>			-0.22	0.125
$\Sigma_{c}v$	4.100	1.508	0.392	4.110	4.064		

\* The bond strength was calculated considering the partial occupancy of Ca sites resulting from the structural refinement.

perature and the structural arrangement of the high-temperature form.

### ACKNOWLEDGEMENTS

We are greatly indebted to Dr. N.V. Vladykin for providing a sample of calcium catapleiite. The facilities for the intensity-data collection were provided by Centro di Strutturistica Diffrattometrica, Dipartimento di Chimica, Università di Ferrara. G. Cruciani and S. Zanardi kindly assisted during the data collection. Reviews by P. Piilonen, A.M. McDonald and A. Lalonde, and the scientific and editorial indications of R.F. Martin, significantly improved the content and presentation of this work. This work was supported in part by the Russian Scientific Foundation (Grant 03-05-64054), by the Grant N1642.2003.5 and by the Russian Universities Program. SM and MP acknowledge the support by MIUR, through grants to the national projects "Structural complexity and mineral properties: microstructures, modularity, modulations" and "Crystal chemistry and alteration of ore minerals". DP acknowledges a grant for a research stay (Project "Crystal structures and topology of rare and new minerals", within the framework of the agreement for scientific and technological cooperation between Italy and Russia).

#### References

- ALTOMARE, A., CASCARANO, G., GIACOVAZZO, C., GAGLIARDI, A., BURLA, M.C., POLIDORI, G. & CAVALLI, M. (1994): SIR 92 – a program for automatic solution of crystal structures by direct methods. J. Appl. Crystallogr. 27, 435.
- BAUSSY, G., CARUBA, R., BAUMER, A. & TURCO, G. (1974): Minéralogie expérimentale dans le système ZrO<sub>2</sub>–SiO<sub>2</sub>– Na<sub>2</sub>O–H<sub>2</sub>O. Corrélations pétrogénétiques. *Bull. Minéral.* 97, 433-444.
- BRESE, N.E. & O'KEEFFE, M. (1991): Bond-valence parameters for solids. Acta Crystallogr. B47, 192-197.
- BRÖGGER, W. C. (1890): Die Mineralien der Syenitpegmatitgänge der südnorwegischen Augit- und Nephelinsyenite. Z. Krystallogr. Mineral. 16, 74-94.
- BRUNOWSKY, B. (1936): Die Struktur des Katapleits (Na<sub>2</sub>ZrSi<sub>3</sub> O<sub>9</sub>•2H<sub>2</sub>O). Acta Physicochim. URSS **5**(6), 863-892.
- CHAKHMOURADIAN, A.R. & ZAITSEV, A.N. (2002): Calcite amphibole – clinopyroxene rock from the Afrikanda Complex, Kola Peninsula, Russia; mineralogy and a possible link to carbonatites. III. Silicate minerals. *Can. Mineral.* 40, 1347-1374.
- CHAKOUMAKOS, B.C., HILL, R.J. & GIBBS, G.V. (1981): A molecular orbital study of rings in silicates and siloxanes. *Am. Mineral.* **66**, 1237-1249.

- CHAO, G.Y, ROWLAND, J.R & CHEN, T.T. (1973): The crystal structure of catapleiite. *Geol. Soc. Am., Abstr. Programs* **5**(7), 572.
- CHEN, T.T. & CHAO, G.Y. (1973): Twinning in catapleiite. Geol. Soc. Am., Abstr. Programs 5(7), 573-573.
- CHOISNET, J., DESCHANVRES, A. & RAVEAU, B. (1973): Évolution structurale de nouveaux germanates et silicates de type wadéite et de structure apparentée. J. Solid State Chem. 7, 408-417.
- DENT GLASSER, L.S. & GLASSER, F.P. (1968): The crystal structure of walstromite. Am. Mineral. 53, 9-13.
- Dowry, E. (1995): Atoms 3.2. A Computer Program for Displaying Atomic Structures. Kingsport, Tennessee.
- FAVRE, C., FUR, Y. & ALÉONARD, S. (1969): Nouveaux fluorobéryllates de structure bénitoïte. Bull. Minéral. 92, 274-277.
- FERRARIS, G. & IVALDI, G. (1988): Bond valence vs. bond length in O...O hydrogen bonds. Acta Crystallogr. B44, 341-344.
- FISCHER, K. (1969): Verfeinerung der Kristallstruktur von Benitoit BaTi[Si<sub>3</sub>O<sub>9</sub>]. Z. Kristallogr. **129**, 222-243.
- FREED, R.L. & PEACOR, D.R. (1969): Determination and refinement of the crystal structure of margarosanite, PbCa<sub>2</sub>Si<sub>3</sub>O<sub>9</sub>. Z. Kristallogr. **128**, 213-218.
- GAINES, R.V., SKINNER, H.C.W., FOORD, E.E., MASON, B. & ROSENZWEIG, A. (1997): Dana's New Mineralogy (8<sup>th</sup> ed.). John Wiley & Sons, New York, N.Y.
- GIBBS G.V., HAMIL, M.M., LOUISNATHAN, S.J., BARTELL, L.S. & YOW, HSIUKANG (1972): Correlations between Si–O bond length, Si–O–Si angle and bond overlap populations calculated using extended Hückel molecular orbital theory. *Am. Mineral.* 57, 1578-1613.
- HAWTHORNE, F.C. (1987): The crystal chemistry of the benitoite group minerals and structural relations in (Si<sub>3</sub>O<sub>9</sub>) ring structures. *Neues Jahrb. Mineral.*, *Monatsh.*, 16-30.
- HENSHAW, D.E. (1955): The structure of wadeite. *Mineral. Mag.* **30**, 585-595.
- ILYUKHIN, V.V., KUZNETSOV V.A., LOBACHEV A.N. & BAKSHUTOV V.S. (1979): Calcium Hydrosilicates. Synthesis of Single Crystals and Crystal Chemistry. Nauka, Moscow, Russia (in Russ.).
- ILYUSHIN, G.D. & DEM'YANETS, L.N. (1988): Crystal-structural features of ion transport in new OD structures: catapleiite Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O and hilairite Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•3H<sub>2</sub>O. *Sov. Phys. Crystallogr.* **33**(3), 383-387.
- \_\_\_\_\_, VORONKOV, A.A., ILYUKHIN, V.V., NEVSKY, N.N. & BELOV, N.V. (1981): The crystal structure of the native monoclinic catapleiite Na<sub>2</sub>ZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O. *Dokl. Akad.*

Nauk SSSR 260, 623-627 (in Russ.) [(1981) Sov. Phys. Dokl. 26, 808].

- LIU, Z.-C., WENG, L.-H., CHEN, Z.-X. & ZHAO, D.-Y. (2003): (NH<sub>4</sub>)<sub>2</sub> ZrGe<sub>3</sub>O<sub>9</sub>: a new microporous zirconogermanate. *Acta Crystallogr.* C59, 29-30.
- LOPES-VIEIRA, A. & ZUSSMAN, J. (1969): The crystal structure of the mineral zussmanite. *Mineral. Mag.* **37**, 49-60.
- MALINOVSKII, Y.A., BONDAREVA, O.S. & BATURIN, S.V. (1984): The crystal structure of NaBaNd[Si<sub>3</sub>O<sub>9</sub>]. *Dokl. Akad. Nauk SSSR* **275**, 372-375 (in Russ.).
- NEKRASOV, I.YA. (1973): The hydrothermal synthesis of tin silicates. Dokl. Acad. Sci. USSR, Earth-Sci. Sect. 212(3), 126-130.
- OTWINOWSKY, Z. & MINOR, W. (1997): Processing of X-ray diffraction data collected in oscillation mode. *Methods Enzymol.* 276, 307–325.
- PIATENKO, YU.A., KUROVA, T.A., CHERNICOVA, N.M., PUDOVKINA, Z.V., BLINOV, V.A. & MAXIMOVA, N.V. (1999): Niobium, Tantalum and Zirconium in Minerals. Crystal Chemistry Guide. IMGRE, Moscow, Russia (in Russ.).
- PORTNOV, A.M. (1964): Calcium catapleiite, a new catapleiite variety. Dokl. Akad. Nauk SSSR 154(3), 607-609 (in Russ.).

- PRESTON, J., HOLE, M., BOUCH, J. & STILL, J. (1998): The occurrence of zirconian aegirine and calcic catapleiite (CaZrSi<sub>3</sub>O<sub>9</sub>•2H<sub>2</sub>O) within a nepheline syenite, British Tertiary igneous province. *Scot. J. Geol.* **34**, 173-180.
- PUSHCHAROVSKY, D.YU. (1986): Structural Mineralogy of Silicates and Their Synthetic Analogues. Izdat. Nedra, Moscow, Russia (in Russ.).
- ROBBINS, C.R. (1960): The compound BaTiGe<sub>3</sub>O<sub>9</sub>. J. Am. Ceram. Soc. **43**, 610.
- SHELDRICK, G.M. (1997): SHELX–97. Program for the Solution and Refinement of Crystal Structures. Siemens Energy and Automation, Madison, Wisconsin.
- VLASOV, K.A., ed. (1966): Geochemistry and Mineralogy of Rare Elements and Genetic Types of their Deposits. II. Mineralogy of Rare Elements. Israel Program of Scientific Translations, Jerusalem, Israel (367-370).
- YANG, HEXIONG & PREWITT, C.T. (1999): On the crystal structure of pseudowollastonite (CaSiO<sub>3</sub>). Am. Mineral. 84, 929-932.
- Received January 9, 2004, revised manuscript accepted June 20, 2004.