2*M*₁-phlogopite from Black Hills (South Australia): The first case of configurational polytype in micas

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ABSTRACT

The $2M_1$ -phlogopite from the potassic gabbronorite (Black Hill, Australia) has been studied by electron microprobe and single-crystal X-ray diffraction analyses. The crystal-chemical formula was $(K_{0.95}Na_{0.01})(Al_{0.15}Mg_{1.27}Fe_{1.16}^{1+}Fe_{0.04}^{3+}Ti_{0.38}^{4+})(Si_{2.85}Al_{1.15})O_{10.76}F_{0.11}Cl_{0.03}OH_{1.10}$. The structural analysis has shown that the crystal has the cell parameters a = 5.352(1), b = 9.268(1), c = 20.168(1) Å, and $\beta = 95.10(1)^{\circ}$ and exhibits symmetry lowering from the ideal space group C2/c to C1. An octahedral cation ordering pattern was revealed from the refined site-scattering powers. Specifically, using the scattering species Mg vs. Fe, it was found that the M1 site at z = 0 was occupied principally by Mg (~77%) and subordinately by Fe (~23%), whereas that at z = 0.5 was completely occupied by Fe; the M2 sites at z = 0 displayed ~88% Mg and ~12% Fe, whereas those at z = 0.5 were occupied by ~86% Fe and ~14% Mg. The analysis of geometrical features shows that the Ti uptake in the structure via the Ti-oxy mechanism induces structural distortions of different extents on the z = 0 and z = 0.5 layers, with stronger effects for the layer at z = 0. Minor chemical and structural differences, instead, affect the T sheets at z = 0 and z = 0.5.

Keywords: 2M₁-phlogopite, octahedral cation ordering, symmetry lowering, structure refinement

INTRODUCTION

Trioctahedral micas are principally 1*M* polytype and only subordinately 2*M*₁, 2*M*₂, and 3*T* (Bailey 1984), whereas dioctahedral micas are prevalently 2*M*₁ polytypes, although 3*T* and 1*M* structures have been found (Brigatti and Guggenheim 2002). For such a reason, polytype 2*M*₁ in trioctahedral micas are comparatively less studied than the 1*M* polytype. The 2*M*₁ polytype has ideally *C*2/*c* symmetry.

The crystal-chemical features of $2M_1$ -biotites with C2/c space group have been analyzed by several authors (for a review see Scordari et al. 2012 and references therein), whereas the cases of $2M_1$ trioctahedral micas with symmetry lowering have been less documented. Specifically, Takeda and Ross (1975) and Ohta et al. (1982) studied the coexistence of 1M and $2M_1$ polytypes in Ruiz Peak (New Mexico) oxy-biotite. They found that the unit layer of $2M_1$ polytype is characterized, relative to the unit layer of 1M polytype, by a shift along $\pm b$ of the upper and lower triads of octahedral oxygen atoms. As a consequence, the layer symmetry is better described in space group $C\overline{1}$. Other instances of $2M_1$ micas showing symmetry different from C2/c, namely Cc (Lin and Guggenheim 1983; Lahti and Saikkonen 1985; Rieder et al. 1996), C1 (Slade et al. 1987), C1 (Swanson and Bailey 1981), Am (Bujnowski et al. 2009) are characterized by remarkably rare composition and/or cation ordering.

Lin and Guggenheim (1983) refined in space group Cc the structure of a mica specimen with composition intermediate

between a bityite- $2M_1$ and a margarite- $2M_1$, i.e., a trioctahedral and a dioctahedral mica, respectively. The nearly complete tetrahedral ordering of Al, Be relative to Si and the pattern that these atoms take on resulted in the violation of the center of symmetry of the C2/c space group, leading to the final Cc symmetry. The authors also reported octahedral cation ordering (Li and vacancy partitioned at M1; Al at M2 and M3 octahedral sites, respectively) that, however, was not involved in the symmetry lowering. Lahti and Saikkonen (1985) document ordering of tetrahedral Si and Al+Be in bityite from pegmatites of the Eräjärvi area at Orivesi, Finland on the basis of the analysis of infrared absorption bands. By contrast, Rieder et al. (1996) analyzed zinnwaldite from mine Barbora, Krupka, Czech Republic, and found very low degree of tetrahedral ordering but remarkable octahedral ordering: Li and Fe²⁺ prefer Ml (M3) to M2, whereas Al does the opposite. Slade et al. (1987) refined the ephesite structure in space group C1. In this case, both tetrahedral and octahedral cation ordering was revealed. The presence of reflections violating the cglide plane and the occurrence of stacking disorder were also highlighted. Swanson and Bailey (1981) studying a lepidolite from Biskupice, Czech Republic, found the presence of some *h*0*l* reflection pointing to a small deviation from the monoclinic symmetry C2/c and concluded that the most likely space group of this mica was $C\overline{1}$ not on the basis of structure refinement but as a result of a second harmonic generation test.

In the present work, a $2M_1$ -phlogopite from potassic gabbronorite plutons of the Black Hills region, south Australia, petrologically characterized by Turner (1996) is considered.

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Coexistence of 1M and $2M_1$ polytype in the same host rock has been documented, but the investigation on 1M polytype was reported elsewhere (see Schingaro et al. 2005). In the present work, the characterization of the $2M_1$ polytype was accomplished by combination of chemical and structural analyses.

EXPERIMENTAL METHODS

The phlogopite analyzed here (labeled "BHG"), belongs to the Black Hills potassic gabbronorite dated 489 \pm 39 Ma, and formed under high-temperature (1200–1000 °C), low-pressure (~1 kbar), and moderate f_{0_2} conditions (Turner 1996). The crystal sample was investigated combining electron probe microanalysis (EPMA) and single-crystal X-ray diffraction (SCXRD).

Chemical composition was measured on the crystal embedded in epoxy resin and polished. A Cameca SX-50 electron microprobe operating at 15 kV accelerating voltage, 15 nA specimen beam current, and 10 μ m beam diameter was used. The analyses were carried out with wavelength-dispersive spectrometers for Na, K, Ba, F, Cl, Ni, Ti, and Mn, and with energy-dispersive spectrometers for Si, Al, Mg, and Fe. The following standards were employed: jadeite (Na), orthoclase (K), barite (Ba), apatite (F), sylvite (Cl), nickeline (Ni), rutile (Ti), rhodonite (Mn), wollastonite (Si), corundum (Al), periclase (Mg), and magnetite (Fe). Data acquisition followed the procedure suggested by Foley (1989). A conversion from X-ray counts to oxide weight percentages (wt%) was carried out with the PAP method (Pouchou and Pichoir 1985). The relative analytical uncertainty is 1% for major elements and 4% for minor elements. Oxide weight percentages (wt%) averaged over 6 spots are reported in Table 1, also in comparison with those of the

 TABLE 1. Chemical composition (oxide wt%): comparison between

 1*M*- and 2*M*₁-BHG crystals

	2 <i>M</i> ₁ -BHG	1 <i>M</i> -BHG*
SiO ₂	35.7(3)	35.6(8)
Al ₂ O ₃	13.8(1)	13.9(3)
MgO	10.7(2)	10.7(4)
FeO	18.0(3)	18.5(8)
TiO ₂	6.3(2)	6.8(4)
MnO	0.06(4)	n.d.
NiO	0.02(2)	n.d.
K ₂ O	9.3(1)	9.4(1)
Na ₂ O	0.05(2)	0.02(2)
BaO	0.11(4)	0.02(2)
F	0.41(9)	0.0(1)
CI	0.25(2)	0.27(5)
Total	94.7	95.2
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Note: * from Schingaro et al. (2005); n.d. = not determined.

TABLE 2. Selected data about collection and refinements of the study crystal in C2/c, $C\overline{1}$, and C1 space groups

		2M₁-BHG	
Crystal size (mm ³)	0.48 x 0.23 x 0.03		
Space group	C2/c	C1	C1
a (Å)	5.352(1)		
b (Å)	9.268(1)		
c (Å)	20.168(1)		
β (°)	95.10(1)		
Cell volume (Å ³)	996.4(2)		
Z	4		
θ range for data collection	2 to 36°		
Reflections collected	23178		
Reflections unique	2389	4596	4596
$R_{\rm merging} (R_{\rm int})$ (%)	4.76	4.44	4.44
Reflections used $[l > 5\sigma(l)]$	1429	2287	2318
Range of h, k, l	$-8 \le h \le 8$		
	–15 ≤ <i>k</i> ≤ 15		
	–33 ≤ / ≤ 33		
No. of refined parameters	76	211	283
GooF*	1.033	1.076	1.096
R ₁ † [on F] (%)	7.96	4.38	4.16
$wR_2 \neq [on F^2]$ (%)	7.40	4.81	4.74
$\Delta \rho_{min} / \Delta \rho_{max} (e^{-} / Å^3)$	-2.63/2.02	-1.35/1.65	-0.83/1.59

* Goodness-of-fit = $\{\Sigma[w(F_o^2 - F_c^2)^2]/(N - P)\}^{1/2}$, where N and P are the number of reflections and parameters, respectively.

 $\dagger R_1 = \Sigma[|F_o| - |F_c|] / \Sigma |F_o|.$

 $\pm wR_2 = \{\Sigma[w(F_o^2 - F_c^2)^2]/\Sigma[w(F_o^2)^2]\}^{1/2}; w = \text{Chebyshev optimized weights.}$

1M polytype (Schingaro et al. 2005).

X-ray intensity data collections were performed by means of a Bruker AXS D8 APEXII automated diffractometer operating at 55 kV and 30 mA, with graphite monochromatized MoK α ($\lambda = 0.7107$ Å) radiation. The collection strategy was optimized by the Apex program suite (Bruker 2003a). A total of 5400 frames were collected with ω and φ scan steps, scan width 0.2°, exposure time 10s/frame, detector-to-sample distance 5 cm. For the extraction of the reflection intensities and Lorentz-polarization corrections, the SAINT package was used (Bruker 2003b). A semi-empirical absorption correction was accomplished using the SADABS software (Sheldrick 2003); the structure refinements were performed with the program CRYSTALS (Betteridge et al. 2003). Initial refinement was carried out in the space group C2/c starting from the atomic coordinates of Ruiz Peak biotite (Ohta et al. 1982). Reflections with $I > 5\sigma(I)$ were used and the refined parameters were: scale factor, atomic positions, cation occupancies, and anisotropic atomic displacement parameters. Ionized X-ray scattering curves were used for interlayer and octahedral sites, whereas ionized vs. neutral species were employed for Si and O to better evaluate electron densities at these sites (Hawthorne et al. 1995).

However, the analysis of reflections via XPREP showed the presence of 468 *h*0*l* reflections that violated the systematic absences required by the *c*-glide plane, out of which 242 had $I/\sigma > 3$. The mean intensity of these reflections was reported to be <I>=38.3, with $<I/\sigma>=4.8$. These data indicated a deviation from the *C*2/*c* symmetry. Therefore, further refinements were performed in different space groups as reported in the Results section below.

RESULTS

The studied crystal shows intra-crystalline homogeneity and its chemical composition (see Table 1) results very close to that of the 1M polytype from the same rock sample that has been previously analyzed by Schingaro et al. (2005).

In Table 2, the cell parameters of the studied crystal, a =5.352(1), b = 9.268(1), c = 20.168(1) Å, and $\beta = 95.10(1)^{\circ}$ are typical of the $2M_1$ polytype with ideal space group C2/c. The initial stage of the structure refinement was carried out in this space group. Therefore, the reflections violating the *c*-glide extinctions (XPREP analysis, see the Experimental section above) were neglected leaving 2389 unique reflections, out of which 1429 with $I > 5\sigma(I)$ were considered. Isotropic refinement with 61 refined parameters converged at $R_1 = 9.94\%$ and $wR_2 =$ 9.51% but evidenced anomalously low isotropic displacement parameters for the O31 and O32 oxygen sites. Fully anisotropic refinement with 112 refined parameters yielded $R_1 = 4.79\%$ and $wR_2 = 4.53\%$ but showed non positive definite displacement parameters for the O31, O32, and O4 oxygen atoms. The best refinement in this space group ($R_1 = 7.96$, $wR_2 = 7.40$, number of refined parameters = 76, see Table 2) was obtained by allowing the cations to vary anisotropically and the anions to vary isotropically. In this case, however, the isotropic displacement parameters of the octahedral O atoms (in particular O31 and O32 atoms) still refined at values lower than expected (see Table 3), because they are about one half than that of O4 oxygen.

The structure refinement was then carried out in the $C\overline{1}$ space group. The refinement was restarted using initially the same geometrical constraints as occur in the C2/c space group, with occupancies being free to vary. All atoms were treated anisotropically. The independent reflections were 4596, out of which 2287 with $I > 5\sigma(I)$ were considered in the refinement, while the number of refined parameters was 129. The refinement converged at $R_1 = 5.6$, $wR_2 = 6.0\%$, with $\rho_{min} = -2.56$ and $\rho_{max} = 1.84 \text{ e}^{-}/\text{Å}^3$ with no anomalies in thermal parameters. At this stage, a cation ordering pattern was revealed from the analysis of the refined site-scattering powers. Specifically, for the M1 site, the position at z = 0.0 was almost fully occupied by Mg, whereas

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Site	Atom	x	У	Ζ	Occupancy	$U_{\rm iso/equiv}$	U ₁₁	U ₂₂	U ₃₃	U ₂₃	U ₁₃	U ₁₂
К	K+	0	0.0844(2)	1⁄4	1.001(1)	0.0358	0.0305(8)	0.0283(8)	0.049(1)	0.0000	0.0034(8)	0.0000
M1	Mg ²⁺	3/4	1⁄4	0	0.45(2)	0.0110	0.0085(5)	0.0092(6)	0.0157(7)	0.0000	0.0026(4)	0.0000
	Fe ²⁺	3/4	1⁄4	0	0.56(1)	0.0110	0.0085(5)	0.0092(6)	0.0157(7)	0.0000	0.0026(4)	0.0000
M2	Mg ²⁺	0.2435(2)	0.0813(1)	-0.00001(4)	0.54(2)	0.0123	0.0124(4)	0.0101(4)	0.0145(5)	0.0030(4)	0.0012(3)	0.0051(3)
	Fe ²⁺	0.2435(2)	0.0813(1)	-0.00001(4)	0.46(1)	0.0123	0.0124(4)	0.0101(4)	0.0145(5)	0.0030(4)	0.0012(3)	0.0051(3)
T1	Si, Si ⁴⁺	0.4624(2)	0.2505(1)	0.13768(7)	1.0(1)	0.0102	0.0064(4)	0.0084(5)	0.0158(6)	-0.0012(5)	0.0014(4)	0.0021(4)
T2	Si, Si4+	0.9639(2)	0.4177(1)	0.13749(6)	1.0(1)	0.0120	0.0108(5)	0.0114(5)	0.0141(5)	-0.0016(6)	0.0018(3)	0.0031(5)
011	O, O ²⁻	0.7386(7)	0.3184(4)	0.1653(2)	1.0(2)	0.0194(7)	-	-	-	-	-	-
O21	O, O ²⁻	0.2374(7)	0.3480(4)	0.1667(2)	1.0(2)	0.0199(7)	-	-	-	-	-	-
022	O, O ²⁻	0.4430(6)	0.0845(4)	0.1665(2)	1.0(2)	0.0192(6)	-	-	-	-	-	-
O31	O, O ²⁻	0.4307(5)	0.2620(3)	0.0543(2)	1.0(2)	0.0089(5)	-	-	-	-	-	-
O32	O, O ²⁻	0.9384(5)	0.4303(3)	0.0547(1)	1.0(2)	0.0078(5)	-	-	-	-	-	-
04	O, O ²⁻	0.9364(6)	0.0949(4)	0.0514(2)	1.0(2)	0.0130(5)	-	-	-	-	-	-

TABLE 3. Crystallographic coordinates, occupancies, equivalent isotropic (Å²), and anisotropic displacement parameters from the structure refinement of the study crystal in the C2/c space group

the position at z = 0.5 was fully occupied by Fe. Likewise, for the M2 site, the position at z = 0 displayed full occupancy of Mg scattering species, whereas the position at z = 0.5 was populated predominantly by Fe (~99%) and only ~1% by Mg. It was apparent from the analysis of site occupancies that those cation sites that were symmetry equivalent in the former C2/cspace group were no longer equivalent because of cation ordering. The following step was to relax the geometrical constraints of the C2/c space group. The number of parameters increased to 211, taking into account that the displacement parameters of atoms previously equivalent by symmetry were constrained to be equal. The refinement converged at $R_1 = 4.38$ and $wR_2 =$ 4.81% (see Table 2) but the equivalent isotropic displacement parameters for the O31, O31^[2], O32, and O32^[2] atoms were still lower than expected (see Table 4). However, attempts to release the constrains on anisotropic displacement parameter resulted in octahedral oxygen atoms to go negative. The final difference Fourier map gave $\rho_{min} = -1.35$ and $\rho_{max} = 1.65 \text{ e}^{-1}/\text{Å}^{-3}$, the most intense occurring close to the M1 atom. Relevant cation-anion bond lengths as determined by structure refinement in the $C\overline{1}$ space group are given in Table 5.

It was then hypothesized that because of the cation ordering and of the different steric and charge features of the substituting cations (Fe, Mg, and Ti above all) these atoms could shift from the ideal position (for instance M1 cation from the center of symmetry), so that refinement in space group C1 was attempted. As a first step, the geometrical constrains of space group C2/c were kept for all atoms but no convergence was obtained.

In the next step, a partial anisotropic refinement was performed, without geometrical constraints, refining isotropically the thermal parameters of the oxygen atoms. $R_1 = 4.61$ and wR_2 = 5.23 values were obtained using 2318 reflections and refining 223 parameters. No anomalies in isotropic equivalent displacement parameters were observed. A full anisotropic refinement was then carried out with 2318 reflections with $I > 5\sigma(I)$, and 283 refined parameters, in full matrix mode, using damping factors for atoms coordinates and occupancies. The N/p (number of reflection/number of parameter) lowered from 10 to 8, the refinement converged at $R_1 = 4.16\%$, $wR_2 = 4.74$ and difference Fourier synthesis yielded $\rho_{min} = -0.83$ and $\rho_{max} = 1.59 \text{ e}^{-1}/\text{Å}^3$ (see Table 2). Displacement of the M1 cation was observed from the ideal position at $(\frac{3}{4}, \frac{1}{4}, \frac{1}{2})$ and a slight deviation from centrosimmetry for most atoms (Table 6). Standard deviations of atomic positions were slightly higher than those refined in space group $C\overline{1}$ (compare Tables 4 and 6).

In conclusion, the values of agreement indexes, bond length

TABLE 4. Crystallographic coordinates, occupancies, equivalent isotropic ($Å^2$) and anisotropic displacement parameters from the structure refinement of the study crystal in the $C\overline{1}$ space group

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Site	Atom	х	У	Ζ	Occupancy	$U_{\rm iso/equiv}$	U ₁₁	U ₂₂	U ₃₃	U ₂₃	U ₁₃	U ₁₂
К	K+	-0.0005(2)	0.0843(1)	0.24990(8)	1.0189(9)	0.0370	0.0309(5)	0.0305(5)	0.0495(7)	-0.0001(4)	0.0033(4)	-0.0001(4)
M1	Mg ²⁺	3/4	1/4	0	0.8977(8)	0.0117	0.0105(6)	0.0116(6)	0.013(1)	-0.0012(6)	0.0024(6)	-0.0013(5)
	Fe ²⁺	3⁄4	1/4	0	0.1023(8)	0.0117	0.0105(6)	0.0116(6)	0.013(1)	-0.0012(6)	0.0024(6)	-0.0013(5)
M1 ^[2]	Fe ²⁺	1⁄4	1/4	1/2	1.0005(7)	0.0103	0.0075(3)	0.0068(3)	0.0167(5)	0.0011(3)	0.0026(3)	0.0020(2)
M2	Mg ²⁺	0.2390(2)	0.0798(1)	-0.00001(6)	1.0002(7)	0.0132	0.0150(6)	0.0130(5)	0.012(1)	0.0006(5)	0.0002(6)	-0.0008(2)
M2 ^[2]	Mg ²⁺	-0.24546(9)	0.08188(5)	0.49997(3)	0.0868(8)	0.0115	0.0102(2)	0.0079(2)	0.0166(4)	-0.0006(2)	0.0016(2)	-0.0008(2)
	Fe ²⁺	-0.24546(9)	0.08188(5)	0.49997(3)	0.9131(8)	0.0115	0.0102(2)	0.0079(2)	0.0166(4)	-0.0006(2)	0.0016(2)	-0.0008(2)
T1	Si⁴+, Si	0.4630(2)	0.2499(1)	0.13771(6)	0.995(1)	0.0118	0.0106(4)	0.0125(4)	0.0123(5)	-0.0001(4)	0.0009(4)	-0.0003(3)
T1 ^[2]	Si4+, Si	-0.4617(1)	0.25022(9)	0.36246(6)	0.989(1)	0.0085	0.0034(3)	0.0038(3)	0.0183(6)	-0.0004(3)	0.0016(3)	-0.0008(3)
T2	Si⁴+, Si	0.9651(2)	0.4172(1)	0.13766(6)	0.995(1)	0.0150	0.0169(5)	0.0178(5)	0.0102(6)	0.0000(4)	0.0009(4)	0.0006(4)
T2 ^[2]	Si⁴+, Si	-0.9626(2)	0.41704(9)	0.36252(6)	0.985(1)	0.0092	0.0052(4)	0.0045(4)	0.0183(7)	-0.0004(4)	0.0023(4)	-0.0002(3)
011	0²-, 0	0.7372(4)	0.3192(3)	0.1654(2)	0.3255(7)	0.0205	0.017(1)	0.026(1)	0.018(1)	0.002(1)	0.000(1)	-0.004(1)
O11 ^[2]	0²-, 0	-0.7382(4)	0.3188(3)	0.3345(2)	0.4462(7)	0.0197	0.015(1)	0.024(1)	0.020(2)	-0.002(1)	0.001(1)	0.009(1)
021	0²-, 0	0.2378(4)	0.3479(3)	0.1666(2)	0.4798(7)	0.0208	0.016(1)	0.027(1)	0.019(2)	-0.001(1)	0.001(1)	0.004(1)
O21 ^[2]	0²-, 0	-0.2379(4)	0.3489(3)	0.3336(2)	0.3359(7)	0.0198	0.016(1)	0.020(1)	0.023(2)	0.002(1)	0.001(1)	-0.009(1)
022	0²-, 0	0.4438(5)	0.0839(3)	0.1662(2)	0.3506(7)	0.0209	0.028(2)	0.016(1)	0.020(2)	-0.0005(1)	0.006(1)	0.000(1)
O22 ^[2]	0²-, 0	-0.4428(5)	0.0843(3)	0.3334(2)	0.2173(7)	0.0195	0.029(1)	0.010(1)	0.021(2)	0.000(1)	0.005(1)	0.000(1)
031	0²-, 0	0.4322(4)	0.2504(3)	0.0552(2)	0.6198(7)	0.0189	0.021(1)	0.020(1)	0.016(2)	0.000(1)	0.003(1)	-0.002(1)
O31 ^[2]	0²-, 0	-0.4296(4)	0.2499(2)	0.4459(2)	0.4898(7)	0.0090	0.0061(9)	0.0069(9)	0.014(1)	0.0013(8)	0.0022(9)	0.0013(7)
032	0²-, 0	0.9413(4)	0.4191(3)	0.0555(2)	0.4229(7)	0.0190	0.019(1)	0.022(1)	0.016(2)	-0.001(1)	0.001(1)	-0.001(1)
O32 ^[2]	0²-, 0	-0.9358(4)	0.4187(2)	0.4458(2)	0.5837(7)	0.0085	0.0042(8)	0.0054(8)	0.016(2)	0.0001(8)	0.0031(9)	0.0016(7)
041	0²-, 0	0.9388(4)	0.0815(3)	0.0516(2)	0.5853(7)	0.0219	0.021(1)	0.019(1)	0.025(2)	-0.001(1)	0.004(1)	0.002(1)
O41 ^[2]	0 ²⁻ , 0	-0.9346(4)	0.0821(3)	0.4487(2)	0.8919(7)	0.0173	0.016(1)	0.015(1)	0.021(2)	-0.001(1)	0.000(1)	-0.0002(9)
Note S	unerscrin	t ^[2] indicates at	tomic position	obtained by th	e twofold sym	metry oper	ation					

TABLE 5. Selecte	ed bond distances (Å) as determined by strue	cture refine-	TABLE 5.—CO	NTINUED		
ment i	in the C1 space gro	up		M2-O32	2.126(3)	M2 ^[2] -O32 ^[2]	2.108(2)
	2 <i>M</i>	-BHG		M2-O41	1.988(3)	M2 ^[2] -O41 ^[2] M2 ^[2] -O41 ^[2]	2.033(3) 2.037(3)
	T1 tetr	ahedron		M2-O41	2.011(3)		
T1-011	1.653(3)	T1 ^[2] -O11 ^[2]	1.662(3)	<m2-0></m2-0>	2.078	<m2<sup>[2]-O></m2<sup>	2.073
T1-021	1.655(3)	T1 ^[2] -O21 ^[2]	1.653(3)				
T1-O22	1.649(3)	T1 ^[2] -O22 ^[2]	1.652(3)		Inter	layer	
T1-O31	1.659(4)	T1 ^[2] -O31 ^[2]	1.675(3)		K-O11	3.032(3)	
<t1-0></t1-0>	1.654	<t1<sup>[2]-O></t1<sup>	1.661		K-011	3.308(3)	
	T2 tetr	ahedron			K-O11 ^[2]	3.030(3)	
T2-011	1.657(3)	T2 ^[2] -O11 ^[2]	1.647(3)		K-O11 ^[2]	3.312(3)	
T2-021	1.653(3)	T2 ^[2] -O21 ^[2]	1.660(3)		K-O21	3.028(3)	
T2-022	1.656(3)	T2 ^[2] -O22 ^[2]	1.664(3)		K-O21	3.287(3)	
T2-032	1.650(4)	T2 ^[2] -O32 ^[2]	1.673(3)		K-O21 ^[2]	3.027(3)	
<t2-0></t2-0>	1.654	<t2<sup>[2]-O></t2<sup>	1.661		K-O21 ^[2]	3.296(3)	
	M1 oct	ahedron			K-022	3.036(3)	
M1-O31 (x2)	2 114(3)	M1 ^[2] -O31 ^[2] (×2)	2 113(2)		K-022	3.287(3)	
M1-032 (x2)	2.111(3)	M1 ^[2] -O32 ^[2] (×2)	2.115(2)		K-O22 ^[2]	3.024(3)	
M1-O41 (x2)	2.087(3)	M1 ^[2] -O41 ^[2] (×2)	2.070(3)		K-O22 ^[2]	3.294(3)	
<m1-0></m1-0>	2.007(0)	<m1<sup>[2]_O></m1<sup>	2.07 0(3)		<k-o>_{inner}</k-o>	3.030	
	Z.112 M2 oct	abedron	2.090		<k-o>_{outer}</k-o>	3.297	
M2-031	2 001(3)	M2 ^[2] -O31 ^[2]	2 080(3)		<k-o></k-o>	3.164	
M2_031	2.001(0)	M2 ^[2] _O31 ^[2]	2.000(3)	Note: Superscr	int [2] indicates atomic r	position obtained by the	- twofold sym-
M2-032	2.104(3)	M2 ^[2] _O32 ^[2]	2.079(3)	metry operatio	n.		

TABLE 5. Selected bond distances (Å) as determined by structure refine

TABLE 6. Crystallographic coordinates, occupancies, equivalent isotropic (Å²) and anisotropic displacement parameters from the structure refinement of the study crystal in the C1 space group

Site	Atom	Х	у	Ζ	Occupancy	$U_{\rm iso/equiv}$	<i>U</i> ₁₁	U ₂₂	U ₃₃	U ₂₃	U ₁₃	<i>U</i> ₁₂
K	K+	-0.0150(5)	0.0842(3)	0.2500(3)	1.026(3)	0.0366	0.0311(5)	0.0306(5)	0.0480(6)	0.0003(4)	0.0029(4)	-0.0006(4)
K ^[1]	K+	-0.0139(5)	-0.0848(3)	0.7499(2)	1.028(3)	0.0366	0.0311(5)	0.0306(5)	0.0480(6)	0.0003(4)	0.0029(4)	-0.0006(4)
M1	Mg ²⁺	3/4	1/4	0	0.768(2)	0.0107	0.0070(3)	0.0081(3)	0.0175(3)	0.0000(3)	0.0032(3)	0.0017(2)
	Fe ²⁺	3⁄4	1⁄4	0	0.232(2)	0.0107	0.0070(3)	0.0081(3)	0.0175(3)	0.0000(3)	0.0032(3)	0.0017(2)
M1 ^[2]	Fe ²⁺	0.2369(4)	0.2411(2)	0.4996(2)	0.9990(9)	0.0107	0.0070(3)	0.0081(3)	0.0175(3)	0.0000(3)	0.0032(3)	0.0017(2)
M2	Mg ²⁺	0.2074(4)	0.0798(3)	0.0004(2)	0.833(2)	0.0106	0.0102(7)	0.0080(5)	0.0134(6)	0.0020(4)	0.0003(6)	0.0060(5)
_	Fe ²⁺	0.2074(4)	0.0798(3)	0.0004(2)	0.167(2)	0.0106	0.0102(7)	0.0080(5)	0.0134(6)	0.0020(4)	0.0003(6)	0.0060(5)
M2 ^[1]	Mg ²⁺	-0.2756(4)	-0.0808(3)	0.0006(2)	0.930(2)	0.0106	0.0102(7)	0.0080(5)	0.0134(6)	0.0020(4)	0.0003(6)	0.0060(5)
	Fe ²⁺	-0.2756(4)	-0.0808(3)	0.0006(2)	0.070(2)	0.0106	0.0102(7)	0.0080(5)	0.0134(6)	0.0020(4)	0.0003(6)	0.0060(5)
M2 ^[2]	Mg ²⁺	-0.2592(3)	0.0725(2)	0.5001(2)	0.172(2)	0.0105	0.0080(2)	0.0077(2)	0.0161(3)	-0.0008(2)	0.0017(2)	-0.0020(2)
	Fe ²⁺	-0.2592(3)	0.0725(2)	0.5001(2)	0.828(2)	0.0105	0.0080(2)	0.0077(2)	0.0161(3)	-0.0008(2)	0.0017(2)	-0.0020(2)
M2 ^[c]	Mg ²⁺	0.2315(3)	-0.0907(2)	0.5000(2)	0.119(2)	0.0105	0.0080(2)	0.0077(2)	0.0161(3)	-0.0008(2)	0.0017(2)	-0.0020(2)
	Fe ²⁺	0.2315(3)	-0.0907(2)	0.5000(2)	0.881(2)	0.0105	0.0080(2)	0.0077(2)	0.0161(3)	-0.0008(2)	0.0017(2)	-0.0020(2)
T1	Si⁴+, Si	0.4478(4)	0.2522(3)	0.1386(2)	0.996(2)	0.0100	0.0110(4)	0.0061(4)	0.0129(5)	-0.0003(4)	0.0003(4)	-0.0003(3)
T1 ^[ī]	Si⁴+, Si	-0.4792(4)	-0.2478(3)	-0.1366(2)	1.000(2)	0.0100	0.0110(4)	0.0061(4)	0.0129(5)	-0.0003(4)	0.0003(4)	-0.0003(3)
T1 ^[2]	Si ⁴⁺ , Si	-0.4787(4)	0.2451(3)	0.3629(2)	0.995(2)	0.0090	0.0044(3)	0.0060(4)	0.0165(5)	-0.0005(4)	0.0013(3)	-0.0013(3)
T1 ^[c]	Si ⁴⁺ , Si	0.4446(4)	-0.2551(3)	0.6381(2)	0.999(2)	0.0090	0.0044(3)	0.0060(4)	0.0165(5)	-0.0005(4)	0.0013(3)	-0.0013(3)
T2	Si⁴+, Si	0.9515(4)	0.4194(3)	0.1376(2)	1.000(2)	0.0120	0.0147(5)	0.0087(5)	0.0126(5)	0.0005(5)	0.0011(4)	0.0006(4)
T2 ^[1]	Si ⁴⁺ , Si	-0.9792(4)	-0.4150(3)	-0.1374(2)	1.000(2)	0.0120	0.0147(5)	0.0087(5)	0.0126(5)	0.0005(5)	0.0011(4)	0.0006(4)
T2 ^[2]	Si4+, Si	-0.9784(4)	0.4108(3)	0.3633(2)	1.000(2)	0.0096	0.0060(4)	0.0079(4)	0.0152(5)	-0.0004(4)	0.0027(4)	-0.0009(3)
T2 ^[c]	Si ⁴⁺ , Si	0.9453(4)	-0.4234(3)	0.6383(2)	0.993(2)	0.0096	0.0060(4)	0.0079(4)	0.0152(5)	-0.0004(4)	0.0027(4)	-0.0009(3)
011	O, O ²⁻	0.7256(6)	0.3220(5)	0.1668(4)	1.000(2)	0.0196	0.018(1)	0.022(1)	0.019(1)	0.002(1)	0.002(1)	-0.006(1)
011 ^[ī]	O, O ²⁻	-0.7492(6)	-0.3166(5)	-0.1644(4)	1.000(2)	0.0196	0.018(1)	0.022(1)	0.019(1)	0.002(1)	0.002(1)	-0.006(1)
O11 ^[2]	O, O ²⁻	-0.7549(6)	0.3135(5)	0.3350(4)	1.000(2)	0.0192	0.017(1)	0.022(1)	0.019(1)	-0.003(1)	0.001(1)	0.007(1)
O11 ^[c]	O, O ²⁻	0.7214(6)	-0.3247(5)	0.6660(4)	1.000(2)	0.0192	0.017(1)	0.022(1)	0.019(1)	-0.003(1)	0.001(1)	0.007(1)
O21	O, O ²⁻	0.2234(6)	0.3545(5)	0.1672(4)	1.000(2)	0.0179	0.017(1)	0.018(1)	0.019(1)	-0.001(1)	-0.000(1)	0.005(1)
O21 ^[ī]	O, O ²⁻	-0.2513(6)	-0.3419(5)	-0.1659(4)	1.000(2)	0.0179	0.017(1)	0.018(1)	0.019(1)	-0.001(1)	-0.000(1)	0.005(1)
O21 ^[2]	O, O ²⁻	-0.2568(6)	0.3417(5)	0.3348(4)	1.000(2)	0.0183	0.016(1)	0.020(1)	0.020(1)	0.002(1)	0.000(1)	-0.010(1)
O21 ^[c]	O, O ²⁻	0.2188(6)	-0.3555(5)	0.6673(4)	1.000(2)	0.0183	0.016(1)	0.020(1)	0.020(1)	0.002(1)	0.000(1)	-0.010(1)
022	O, O ²⁻	0.4274(6)	0.0896(4)	0.1652(4)	1.000(2)	0.0177	0.026(1)	0.008(1)	0.019(1)	0.002(1)	0.004(1)	0.002(1)
O22 ^[1]	O, O ²⁻	-0.4612(6)	-0.0788(4)	-0.1674(4)	1.000(2)	0.0177	0.026(1)	0.008(1)	0.019(1)	0.002(1)	0.004(1)	0.002(1)
O22 ^[2]	O, O ²⁻	-0.4616(6)	0.0797(4)	0.3346(4)	1.000(2)	0.0192	0.029(2)	0.010(1)	0.020(1)	0.000(1)	0.008(1)	-0.002(1)
O22 ^[c]	O, O ²⁻	0.4245(6)	-0.0894(5)	0.6674(4)	1.000(2)	0.0192	0.029(2)	0.010(1)	0.020(1)	0.000(1)	0.008(1)	-0.002(1)
O31	O, O ²⁻	0.4101(6)	0.2509(5)	0.0567(4)	1.000(2)	0.0137	0.015(1)	0.012(1)	0.014(1)	-0.001(1)	0.002(1)	0.0040(9)
O31 ^[ī]	O, O ²⁻	-0.4558(6)	-0.2479(5)	-0.0543(4)	1.000(2)	0.0137	0.015(1)	0.012(1)	0.014(1)	-0.001(1)	0.002(1)	0.0040(9)
O31 ^[2]	O, O ²⁻	-0.4474(5)	0.2386(4)	0.4456(4)	1.000(2)	0.0103	0.0077(9)	0.009(1)	0.015(1)	-0.001(1)	0.004(1)	-0.0009(8)
O31 ^[c]	O, O ²⁻	0.4086(5)	-0.2607(4)	0.5535(4)	1.000(2)	0.0103	0.0077(9)	0.009(1)	0.015(1)	-0.001(1)	0.004(1)	-0.0009(8)
032	O, O ²⁻	0.9185(6)	0.4185(5)	0.0564(4)	1.000(2)	0.0133	0.012(1)	0.012(1)	0.015(1)	-0.000(1)	0.000(1)	0.0047(9)
O32 ^[1]	O, O ²⁻	-0.9631(6)	-0.4196(5)	-0.0549(4)	1.000(2)	0.0133	0.012(1)	0.012(1)	0.015(1)	-0.000(1)	0.000(1)	0.0047(9)
O32 ^[2]	O, O ²⁻	-0.9573(5)	0.4097(5)	0.4444(4)	1.000(2)	0.0103	0.0058(9)	0.012(1)	0.013(2)	-0.002(1)	0.002(1)	-0.0026(8)
O32 ^[c]	O, O ²⁻	0.9148(5)	-0.4280(5)	0.5529(4)	1.000(2)	0.0103	0.0058(9)	0.012(1)	0.013(2)	-0.002(1)	0.002(1)	-0.0026(8)
O41	O, O ²⁻	0.9100(6)	0.0797(5)	0.0516(4)	1.000(2)	0.0148	0.011(1)	0.014(1)	0.020(2)	0.001(1)	-0.002(1)	0.006(1)
041 ^[ī]	O, O ²⁻	-0.9681(6)	-0.0821(5)	-0.0526(4)	1.000(2)	0.0148	0.011(1)	0.014(1)	0.020(2)	0.001(1)	-0.002(1)	0.006(1)
O41 ^[2]	O, O ²⁻	-0.9545(6)	0.0717(5)	0.4488(4)	1.000(2)	0.0144	0.012(1)	0.013(1)	0.018(2)	-0.001(1)	0.001(1)	-0.0031(9)
O41 ^[c]	O, O ²⁻	0.9131(6)	-0.0903(5)	0.5502(4)	1.000(2)	0.0144	0.012(1)	0.013(1)	0.018(2)	-0.001(1)	0.001(1)	-0.0031(9)
Note: Si	inerscript	^[2] indicates at	omic position	obtained by	the twofold svr	nmetry one	ration: super	script ^[] indic	ates atomic	position obta	ined by inve	rsion center

^{1]} indicates atomic position obtained by inversion center; twofold symmetry operation; superscript ¹ superscript ^[c] indicates atomic position obtained by the *c*-glide.

distances and mean atomic number were generally similar for the structure refinements in both C1 and $\overline{C1}$ space groups (see Tables 2 and 8, and compare Tables 5 and 7). However, the structural model in C1 space group appears more consistent with crystal-chemical considerations about cation ordering in the z = 0.0 and z = 0.5 layer (see the Discussion and Conclusion section below).

DISCUSSION AND CONCLUSION

The EPMA (see Table 1) and SCXRD data were combined with the Mössbauer results [$^{VI}Fe^{2+} = 97.0(7)$; $^{VI}Fe^{3+} = 3.0(7)$] reported in the 1*M*-BHG phlogopite study (see Schingaro et al. 2005) and OH⁻ content geometrically estimated basing on the *c* vs. OH⁻ relationship for 2*M*₁ micas (Scordari et al. 2012) to calculate, on the basis of (O_{12-x-y-z}OH_xCl_yF_z), the following structural formula:

 $\begin{array}{l}(K_{0.95}Na_{0.01})(Al_{0.15}Mg_{1.27}Fe_{1.16}^{2+}Fe_{0.04}^{3+}Ti_{0.38}^{4+})(Si_{2.85}Al_{1.15})\\O_{10.76}F_{0.11}Cl_{0.03}OH_{1.10}.\end{array}$

The Ti was considered all belonging to the tetravalent species

 TABLE 7.
 Selected bond distances (Å) as determined by structure refinement in the C1 space group

	1 3	2A	A-BHG			
		T1 Tot	trahadran			
T1-011 1674(4)	T1 ^[]] -O11 ^[]	1 633(4)	T1 ^[2] -O11 ^[2]	1 660(4)	T1 ^[c] -O11 ^[c]	1 667(4)
$T1_071_{1.074(4)}$	T10-0210	1.652(5)	T1 ^[2] -O21 ^[2]	1.600(4)	T1 ^[c] -O21 ^[c]	1.673(5)
$T1_027_1.072(5)$	T1 ^[1] =022 ^[1]	1.691(5)	T1 ^[2] -O22 ^[2]	1.620(5)	T1 ^[c] -O22 ^[c]	1.652(5)
T1-O31 1 646(9)	T1 ^[1] -O31 ^[1]	1.654(9)	T1 ^[2] -O31 ^[2]	1.642(3)	T1 ^[c] -031 ^[c]	1 701(8)
<t1-0> 1.645</t1-0>	< T1 ^[1] -O>	1 658	<t1<sup>[2]-O></t1<sup>	1 648	<t1<sup>[c]-O></t1<sup>	1 673
1.015		T2 Tet	trahedron	1.010		1.075
T2-O11 1.658(5)	T2 ^[ī] -O11 ^[ī]	1.661(5)	T2 ^[2] -O11 ^[2]	1.640(5)	T2 ^[c] -O11 ^[c]	1.644(5)
T2-O21 1.638(4)	T2 ^[1] -O21 ^[1]	1.662(4)	T2 ^[2] -O21 ^[2]	1.675(4)	T2 ^[c] -O21 ^[c]	1.652(4)
T2-022 1.681(6)	T2 ^[ī] -O22 ^[ī]	1.639(6)	T2 ^[2] -O22 ^[2]	1.674(5)	T2 ^[c] -022 ^[c]	1.654(5)
T2-O32 1.631(9)	T2 ^[ī] -O32 ^[ī]	1.660(9)	T2 ^[2] -O32 ^[2]	1.629(8)	T2 ^[c] -O32 ^[c]	1.716(8)
<t2-o> 1.652</t2-o>	<t2<sup>[ī]-O></t2<sup>	1.656	<t2<sup>[2]-O></t2<sup>	1.655	<t2<sup>[c]-O></t2<sup>	1.667
		M1 Oc	tahedron			
	M1-031	2.233(5)	M1 ^[2] -O31 ^[2]	2.091(6)		
	M1-O31 ^[1]	1.996(5)	M1 ^[2] -O31 ^[c]	2.147(6)		
	M1-032	2.089(5)	M1 ^[2] -O32 ^[2]	2.134(5)		
	M1-O32 ^[1]	2.185(5)	M1 ^[2] -O32 ^[c]	2.083(6)		
	M1-041	2.037(5)	M1 ^[2] -O41 ^[2]	2.092(6)		
	M1-O41 ^[1]	2.165(5)	M1 ^[2] -O41 ^[c]	2.050(6)		
	<m1-0></m1-0>	2.118	<m1<sup>[2]_O></m1<sup>	2.100		
		M2 Oc	tahedron			
M2-O31 2.182(6)	M2 ^[ī] -O31	2.121(6)	M2 ^[2] -O31 ^[2]	2.097(5)	M2 ^[c] -O31 ^[2]	2.107(5)
M2-O31 ^[1] 2.088(6)	M2 ^[1] -O31 ^[1]	2.089(6)	M2 ^[2] -O31 ^[c]	2.047(5)	M2 ^[c] _O31 ^[c]	2.089(5)
M2-O32 2.136(5)	M2 ^[ī] -O32	2.068(6)	M2 ^[2] -O32 ^[2]	2.110(5)	M2 ^[c] _O32 ^[2]	2.089(5)
M2-O32 ^[1] 2.169(6)	M2 ^[ī] -O32 ^[ī]	2.072(6)	M2 ^[2] -O32 ^[c]	2.124(5)	M2 ^[c] -O32 ^[c]	2.046(5)
M2-O41 1.972(6)	M2 ^[1] -O41	2.019(6)	M2 ^[2] -O41 ^[2]	2.007(6)	M2 ^[c] _O41 ^[2]	2.035(5)
M2-O41 ^[1] 2.024(6)	M2 ^[1] -O41 ^[1]	2.043(6)	M2 ^[2] -O41 ^[c]	1.996(5)	M2 ^[c] -O41 ^[c]	2.058(5)
<m2-o> 2.095</m2-o>	<m2<sup>[1]-O></m2<sup>	2.069	<m2<sup>[2]-O></m2<sup>	2.064	<m2<sup>[2]-O></m2<sup>	2.071
		Int	erlayer			
	K-011	3.031(6)	K ^[1] -O11 ^[1]	3.029(6)		
	K-O11	3.280(6)	K ^[1] -O11 ^[1]	3.337(6)		
	K-011 ^[2]	2.996(6)	K ^[1] -O11 ^[c]	3.064(6)		
	K-011 ^[2]	3.358(6)	K ^[1] -O11 ^[c]	3.262(6)		
	K-O21	2.977(6)	K ^[1] -O21 ^[1]	3.082(6)		
	K-021	3.327(6)	K ^[1] -O21 ^[1]	3.249(6)		
	K-O21 ^[2]	3.077(6)	K ^[1] -O21 ^[c]	2.986(6)		
	K-O21 ^[2]	3.271(7)	K ^[1] -O21 ^[c]	3.315(6)		
	K-022	3.041(6)	K ^[1] -O22 ^[1]	3.037(6)		
	K-022	3.305(5)	K ^[1] -O22 ^[1]	3.265(5)		
	K-O22 ^[2]	3.058(6)	K ^[1] -O22 ^[c]	2.995(6)		
	K-O22 ^[2]	3.285(5)	K ^[1] -O22 ^[c]	3.304(5)		
	<k-o>_{inner}</k-o>	3.030	<k<sup>[1]-O>_{inner}</k<sup>	3.032		
	<k-o>_{outer}</k-o>	3.304	<k<sup>[1]-O>_{outer}</k<sup>	3.288		
	<k-o></k-o>	3.167	<k<sup>[1]_O></k<sup>	3.160		

Note: Superscript ^[2] indicates atomic position obtained by the twofold symmetry operation; superscript ^[7] indicates atomic position obtained by inversion center; superscript ^[c] indicates atomic position obtained by the *c*-glide.

according to recent developments of data analysis based on a more rigorous treatment of the background for X-ray photoelectron spectra of BHG-1*M* crystals (Castle and Salvi 2001; Mesto et al. 2008).

The structural formula, above, is also in agreement with the X-ray data (see Table 8) and is balanced by the following substitution mechanisms: Al, Fe³⁺-Tschermak [^{VI}M²⁺ + ^{IV}Si⁴⁺ \leftrightarrow ^{VI}M³⁺ + ^{IV}Al³⁺]; Ti-oxy [^{VI}M²⁺ + 2 (OH)⁻ \leftrightarrow ^{VI}Ti⁴⁺ + 2 (O²⁻) + H₂↑]; and ^{XII}K⁺ + ^{IV}Al³⁺ \leftrightarrow ^{IV}Si⁴⁺ + ^{XII}□.

In comparison to all the other literature $2M_1$ micas (Scordari et al. 2012 and references therein), in the classification plot of Figure 1 the study $2M_1$ -BHG mica follows the general trend, showing a slight deviation toward the annite end-member.

The sketch shown in Figure 2 illustrates the structural features of the study crystal. The analysis of the details of the TOT layer at z = 0 as compared with the TOT layer at z = 0.5 evidences that:

(1) An octahedral ordering is clearly apparent from a chemical viewpoint, i.e., if refined site-scattering powers are taken into account. This can be seen from the comparison of the occupancies for the sites M1, M2 (Mg-rich) and M1^[2], M2^[2] (Fe-rich) in $C\overline{1}$

space group (see Table 4) and for the sites M1, M2, M2^[1] (Mg-rich) and M1^[2], M2^[2], M2^[c] (Fe-rich) in C1 space group (see Table 6). As a further detail, scattering power at M2 site in $C\overline{1}$ refines to 12 e⁻, whereas in C1 M2 refines to 14.34 and $M2^{[\bar{1}]}$ to 12.98 e⁻ with a difference of 1.36 e⁻, which indicates significantly different chemical content in the latter two sites. From a geometrical viewpoint (i.e., average bond distances and octahedral volumes), there seems to be no evidence of octahedral ordering. For instance, $\langle M-O \rangle_{z=0}$ is slightly larger than $\langle M-O \rangle_{z=0.5}$ (see Tables 5 and 7) despite the fact that the former site is predominantly occupied by Mg (ionic radius 0.72 Å from Shannon 1976) and the latter by Fe^{2+} (ionic radius 0.78 Å from Shannon 1976). This may be due to the covalent character of the Fe-O as to the Mg-O bonds.

The octahedral cation distributions for both $C\overline{1}$ and C1 space groups give good agreement (i.e., differences are within ~0.5 e⁻/site) between mean atomic numbers as assessed by chemical determinations and structure refinement (see Table 9). However, note that the octahedral cation partitioning for the $C\overline{1}$ space group provides 0.16 and 0.60 Ti atoms per formula unit (apfu) at z = 0 and z = 0.5 layer, respectively. Conversely, the proposed cation distribution for the C1 space group gives 0.51 and 0.25 Ti apfu at z = 0 and z = 0.5 layer, respectively. The latter result is compatible with a greater extent of the Ti-oxy component for the z = 0 layer as suggested by the analysis of diagnostic geometric parameters (see below). For such a reason, we consider the C1as the most suitable space group for the description of the study crystal.

The oxy-type mechanism has been previously proposed by Schingaro et al. (2005) as the most likely for the entry of Ti in the structure of the 1*M*-BHG micas. As stated above, this appears also to

TABLE 8. Mean atomic numbers (electrons, e⁻) of cation sites, octahedral and tetrahedral mean distances (Å), as determined by structure refinements (X-ref) in the CT and C1 space group and chemical analyses (EPMA)

	CĪ	C1
<t-o>_{X-ref}</t-o>	1.658	1.657
<t-o>_{EPMA}</t-o>	1.657	1.657
<m-o>_{X-ref}</m-o>	2.085	2.086
<m-o>_{EPMA}</m-o>	2.079	2.079
T e ⁻ _{X-ref}	13.87	13.97
T e ⁻ _{EPMA}	13.71	13.71
e ⁻ (M1) X-ref	39.44	41.22
e ⁻ _{(M2) X-ref}	36.78	37.62
e-(M1+2M2) X-ref	56.50	58.23
e-(M1+2M2) EPMA	56.75	56.75
K e ⁻ _{X-ref}	19.36	19.51
K e ⁻ _{EPMA}	18.16	18.16

Note: Average error for X-ref mean atomic numbers is $\pm 0.5 \text{ e}^{-}$ /site; EPMA mean atomic numbers are reproducible within $\pm 0.01 \text{ e}^{-}$ for tetrahedral site, $\pm 0.4 \text{ e}^{-}$ for octahedral site and $\pm 0.3 \text{ e}^{-}$ for interlayer.

TABLE 9. Cation distribution for $C\overline{1}$ and C1 space groups

		e ⁻ _{X-ref}	e ⁻ EPMA
	Octahedral sheet		
C1space group			
Layer at $z = 0$	(Mg _{2.54} Al ³⁺ _{0.30} Ti _{0.16})	37.44	37.90
Layer at $z = 0.5$	(Fe ²⁺ _{2.32} Fe ³⁺ _{0.08} Ti _{0.60})	75.58	75.60
C1 space group			
Layer at $z = 0$	(Mg _{2.49} Ti _{0.51})	42.57	41.10
Layer at $z = 0.5$	$(AI_{0.30}^{3+}Mg_{0.05}Fe_{2.32}^{2+}Fe_{0.08}^{3+}Ti_{0.25})$	73.90	72.40
	Tetrahedral sheet		
C1 space group			
Layer at $z = 0$	(Si _{2.98} Al _{1.02})	55.94	54.98
Layer at $z = 0.5$	(Si _{2.72} Al _{1.28})	55.82	54.72
Note: for errors associ	ated to e_{X-ref}^{-} and e_{EPMA}^{-} see footnote	e of Table 8.	

be the case of the $2M_1$ -BHG studied crystal. Three out of the most sensitive octahedral geometrical parameters reflective of the Ti-oxy mechanism are the $\langle M2-O4 \rangle$ distance, the shift_{M2} parameter and the BLD_{M2} , bond-length distortion parameter of the M2 site (see Tables 7 and 10). Indeed, Figure 3 shows that the study crystal together with all the other literature $2M_1$ micas define rough a negative trend, as expected for micas interested by oxy mechanisms (Scordari et al. 2012; Lacalamita et al. 2011). Note that the average M2-O4 distance assumes different values for the z = 0 (2.015 Å) and z = 0.5 layers (2.024 Å), the lower the value the greater the amount of the oxy component. In Figure 4, the shift_{M2} parameter becomes larger when the Ti-oxy component increases. For the study sample, the greatest $shift_{M2}$ value (on average 0.083 Å) for the z = 0 layer testifies for a stronger oxy component affecting this layer in comparison to that (on average 0.061 Å) at z = 0.5. In addition, the average BLD_{M2} parameter calculated for the z = 0 layer (2.215) is higher than that (1.718) obtained for the z = 0.5 layer (see Table 10). Also in this case the interpretation is straightforward if it is recalled that direct proportionality exists between the value of the BLD_{M2} parameter and the oxy component (Cruciani and Zanazzi 1994).

(2) The T sheets at z = 0 and z = 0.5 are quite homogeneous in composition. Indeed, only slight differences occur between the occupancies of the sites T1, T1^[T], T2, and T2^[T] on one side and T1^[2], T1^[c], T2^[2], and T2^[c] on the other side (see Table 6), the average distances <T-O>_{z=0} (1.653 Å) and <T-O>_{z=0.5} (1.661 Å) and the in plane ditrigonal angles, α (see Table 10). Accordingly, the tetrahedral cation distribution at z = 0 and z = 0.5 are



FIGURE 1. Chemical classification diagram modified after Scordari et al. (2012) including the study crystal and all the literature trioctahedral $2M_1$ micas belonging to the phlogopite-annite solid solution. *feal* = (Fe_{tot} + Mn + Ti - ^{VI}Al); *mgli* = (Mg-Li). Symbols: solid symbols = study crystal (circle: BHG); open symbols = igneous phlogopites from literature (square: Takeda and Ross 1975; circle: Ohta et al. 1982; pointing upward triangle: Bigi et al. 1993; pointing downward triangle: Bigi and Brigatti 1994; diamond: Brigatti et al. 2005); open symbols with horizontal line = other igneous phlogopites from literature (square: Laurora et al. 2007; circle: Pini et al. 2008; pointing upward triangle: Scordari et al. 2012); open symbols with vertical line = metamorphic phlogopites from literature (square: Bohlen et al. 1980; circle: Brigatti et al. 2000; pointing upward triangle: Brigatti et al. 2008).



FIGURE 2. Representation of the $2M_1$ -BHG structure: the octahedral ordering at z = 0 and z = 0.5 layers is highlighted, according to the cation partition shown in Table 9. (Color online.)

fairly similar (see Table 9). To be more precise, on average, the tetrahedral sheet at z = 0.5 shows a slightly greater Al content than that at z = 0.

As mentioned above, the study $2M_1$ -BHG phlogopite has

				2M1	1-BHG				
		Layer a	nt z = 0			Layer a	at z = 0.5		
	T1 ^[1] lower	T2 ^[1]	T1 _{upper}	T2 upper	T1 ^[2] lower	T2 ^[2] lower	T1 ^[c]	T2 ^[c]	
BLD _T	1.024	0.469	1.419	1.071	0.800	1.209	0.819	1.487	
<i>V</i> ⊤ (ų)	2.331	2.328	2.302	2.313	2.296	2.325	2.402	2.373	
TQET	1.001	1.001	1.001	1.001	1.001	1.000	1.001	1.001	
TAV _⊤ (°)	3.39	1.60	1.77	2.08	1.69	1.20	2.63	1.55	
τ (°)	110.8	110.3	109.8	110.4	110.1	110.1	110.3	110.4	
∆z _{lower} (Å)	0.060				0.008				
∆z _{upper} (Å)	0.040				0.028				
α _{lower} (°)	5.13				6.07				
α _{upper} (°)	6.27				5.71				
D.M. _{lower} (Å)	0.424				0.501				
D.M. _{upper} (Å)	0.442				0.507				
t _{tet} (Å) _{lower}	2.236				2.214				
t _{tet} (Å) _{upper}	2.207				2.284				
	M1				M1 ^[2]				
ψ _{M1} (°)	58.18				58.82				
BLD _{M1}	3.629				1.312				
V _{M1} (Å ³)	12.410				12.063				
OQE _{M1}	1.015				1.015				
OAV _{M1} (°)	44.78				48.52				
ELD _{M1}	4.974				5.683				
e _u M1/e _s M1	1.105		~		1.122				
	M2	M	2 ^[1]		M2 ^[2]	N	12 ^[c]		
ψ _{M2} (°)	57.80	57	.34		58.21	57	7.37		
BLD _{M2}	3.194	1.2	236		2.271	1.	165		
ELD _{M2}	4.568	4.0)62		5.053	5.	204		
V _{M2} (A ³)	12.070	11.	674		11.504	11	.621		
OQE _{M2}	1.012	1.0	800		1.012	1.	012		
OAV _{M2} (°)	35.99	24	.49		38.87	40).41		
e _u M2/e _s M2	1.096	1.0)85		1.106	1.	110		
Shift _{™2} (Å)	0.122	0.0)43		0.080	0.	041		
t _{oct} (Å)	2.233				2.174				
t _{int} (Å)	3.383				3.364				
Δ _{κ-0} (Å)	0.274				0.256				
t _{K∙O4 Lower} (Å)	3.980				3.996				
t _{K-O4 Upper} (Å)	3.970				3.953				

 TABLE 10.
 Selected distortional parameters from the structure refinements of the study crystal in the C1 space group

Notes: Superscript ^[2] indicates atomic position obtained by the twofold symmetry operation; superscript ^[1] incates atomic position obtained by inversion center; superscript ^[c] indicates atomic position obtained by c-glide; subscript "lower" indicates the lower tetrahedral layer of a TOT unit; subscript "upper" indicates the upper tetrahedral laver of a TOT unit. t...: tetrahedral sheet thickness calculated from z coordinates of basal and apical O atoms; TQE: tetrahedral quadratic elongation (Robinson et al. 1971); TAV: tetrahedral angle variance (Robinson et al. 1971); T: tetrahedral flattening angle; C: tetrahedral rotation angle (Hazen and Burnham 1973); Δz : departure from complanarity of the basal O atoms (Güven 1971); D.M.: dimensional misfit between tetrahedral and octahedral sheets (Torava 1981); u: octahedral flattening angles (Donnav et al. 1964a, 1964b); BLD: bond-length distortions (Renner and Lehmann 1986); ELD: edge-length distortion (Renner and Lehmann 1986); Shift_{M2}: off-center shift of the M2 cation defined as the distance between the M2 cation and the polyhedron centroid; OQE: octahedral quadratic elongation (Robinson et al. 1971); OAV: octahedral angle variance (Robinson et al. 1971); eu, es: mean lengths of unshared and shared edges, respectively (Toraya 1981); toct: octahedral sheet thickness (Toraya 1981); t_{int} calculated from the z coordinates of basal O atoms; $\Delta_{K-O} = \langle K-O \rangle_{outer} - \langle K-O \rangle_{inner}$; t_{K-O4} : projection of K-O4 distance along c*. Errors on distortion parameters, estimated by varying the refined positional parameters within one standard deviation, are in the following ranges: <0.5% for volumes, thicknesses, projected bond lengths, shifts; 0.1-13% for angles, bond/edge lengths distortions, sheet corrugations, D.M., Δ_{K-O}

been found in coexistence in the same rock sample with a 1*M* polytype (Schingaro et al. 2005). By comparing the 1*M*- and 2*M*₁-BHG structures in the 2*M*₁ setting (space group *C*2/*c*), it is easily calculated that the only significant differences between atomic positions are those relevant to the *y* coordinate of the O31, O32, and O4 oxygen atoms (the absolute value of $\Delta y_{O31,O32}$ = 0.012, Δy_{O4} = 0.013). Analogous values calculated by Takeda and Ross (1975) for their 1*M* and 2*M*₁ coexisting biotite are $\Delta y_{O31,O32,O4}$ = 0.009. As also previously reported by these authors,



FIGURE 3. <M2-O4> distance vs. Ti content plot. Symbols as in Figure 1. Black solid symbol = $2M_1$ -BHG sample at z = 0; gray solid symbol = $2M_1$ -BHG sample at z = 0.5.



FIGURE 4. Shift_{M2} parameter vs. Ti content plot modified after Scordari et al. (2012). Symbols as in Figure 1. Black solid symbol = $2M_1$ -BHG sample at z = 0; gray solid symbol = $2M_1$ -BHG sample at z = 0.5.

the symmetry lowering shown by the study $2M_1$ phlogopite may be a consequence of the $2M_1$ layer stacking, which affects the position of the octahedral oxygen atoms along the $\pm b$ direction. However, the symmetry lowering of the study crystal may be also due to the octahedral cation ordering. In a very recent paper on the coexistence of 1M- and $2M_1$ -polytypes, it was hypothesized that the probability of occurrence of $2M_1$ polytype depends on the degree of cation ordering at octahedral sites but may be also related to the extent of oxy-type substitutions that destabilize the octahedral sheet (Lacalamita et al. 2012). All the $2M_1$ -phlogopites reported to date in literature are completely (Ohta et al. 1982; Laurora et al. 2007; BU1 8, BU1 14, and BU1 16 crystals from Lacalamita et al. 2012) or partially (Takeda and Ross 1975; Bigi and Brigatti 1994; Brigatti et al. 2008; BU1 15 and BU1 17 crystals from Lacalamita et al. 2012) disordered at octahedral sites. Also the 1M-BHG polytype coexisting with the study sample show disordered cation distribution at octahedral sites, apart from the partial ordering (Ti at M2 site) directly caused by the Ti-oxy substitution. The peculiarity of the studied $2M_1$ - polytype stems from the remarkable ordering of the octahedral layers, that results in the *c* cell parameter becoming twice as that of the 1*M*-BHG in Schingaro et al. (2005), as well as in notably different chemical composition of the z = 0 and z = 0.5 layers. The results point to the description of the present samples as a "configurational polytype" (Ferraris et al. 2004). Really, with respect to the strict definition of polytypes, characterized by constant or nearly constant composition (up to 0.25 atoms per formula unit) of the stacking layers (Bailey 1984; Guinier et al. 1984), configurational polytypes have been described as having the same layer configuration and layer-stacking principle, but radically different chemical compositions (Makovicky 1997; Ferraris et al. 2004).

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