Castellaroite, Mn²⁺₃(AsO₄)₂·4.5H₂O, a new mineral from Italy related to metaswitzerite

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Abstract: Castellaroite (IMA2015-071), Mn²⁺₃(AsO₄)₂·4.5H₂O, is a new secondary arsenate mineral from the Monte Nero mine, Rocchetta Vara, La Spezia, Liguria, Italy and the Valletta mine, near Canosio, Cuneo, Piedmont, Italy. It crystallized from As- and Mn-rich hydrothermal fluids in an oxidizing environment. At Monte Nero, it is associated with coralloite, manganohörnesite, rhodochrosite, sarkinite, sterlinghillite, strashimirite and wallkilldellite. At Valletta, it is associated with braccoite, hematite, manganberzeliite, orthoclase and tiragalloite. Castellaroite occurs as thin blades, flattened on [001], striated and elongated parallel to [100] and exhibiting the forms {110}, {012} and {001}. The mineral is colourless and transparent with a vitreous to silky lustre and white streak. Crystals are flexible with a curved fracture, and one perfect cleavage on $\{001\}$. The Mohs' hardness is $2\frac{1}{2}$. The measured and calculated densities are 3.14(2) g·cm⁻³ and 3.164 g·cm⁻³, respectively. The mineral is easily soluble in dilute HCl at room temperature. Optically, castellaroite crystals are biaxial (-), with $\alpha = 1.644(1)$, $\beta = 1.662(1)$ and $\gamma = 1.667(1)$ (white light); 2V = 57(1)°; dispersion r < v, moderate; the optical orientation $Y = \mathbf{b}$; $Z \approx \mathbf{a}$. The Raman spectrum is dominated by features corresponding to the AsO_4 group and also confirms the presence of H_2O . Electron-microprobe analyses gave the empirical formula $Mn^{2+}_{3.02}(As_{1.94}P_{0.06})_{\Sigma2.00}O_{12.5}H_{8.96}$, based on 12.5 O *apfu*. Castellaroite is monoclinic, $P2_1/n$, with the unit-cell parameters: a = 8.7565(8), b = 13.4683(13), c = 18.652(2) Å, $\beta = 94.876(7)^{\circ}$, V = 2191.7(4) Å³ and Z = 8. The eight strongest lines in the X-ray powder diffraction pattern are [dobs/Å (I) (hkl)]: 10.90(100)(011), 9.27(67)(002), 6.97(42)(111), 3.323(47)(multiple), 3.043(87) $(\bar{1}34,204,\bar{2}32)$, 2.656(85)(multiple), 2.165(46)(multiple), and 1.5589(32)(multiple). The crystal structure was refined to $R_1 = 0.118$ for 2513 observed reflections $[F_0 > 4\sigma F]$. The structure contains kinked chains of edge-sharing MnO₆ octahedra parallel to [100]. The chains are linked to each other by corner-sharing, forming sheets parallel to $\{001\}$ and AsO₄ tetrahedra corner-link with octahedra in the sheet, forming a heteropolyhedral layer. Edge-sharing MnO₆-MnO₅ dimers share corners with octahedra and tetrahedra in adjacent layers, thereby linking them in the [001] direction. The heteropolyhedral layer is topologically identical to those in the structures of the phosphate minerals: angarfite, bakhchisaraitsevite, mejillonesite, metaswitzerite, rimkorolgite and switzerite. Overall, the structure is most similar to that of metaswitzerite.

Key-words: castellaroite; arsenate; crystal structure; Raman spectroscopy; Monte Nero mine; Valletta mine; Italy.

1. Introduction

Long inactive manganese deposits, also rich in arsenic, in the Alpine tectono-metamorphic terrains of north-western Italy have yielded a variety of interesting secondary hydrothermal minerals. Among these deposits, the Molinello and Gambatesa mines in Val Graveglia, Liguria, are well known (*cf.* Marchesini & Pagano, 2001), having yielded eight and six new mineral species, respectively. Two lesser-known mines, the Monte Nero mine in Liguria and the Valletta mine in Piedmont have also yielded new species: coralloite (Callegari *et al.*, 2012) from Monte Nero and braccoite (Cámara *et al.*, 2015), canosioite (Cámara *et al.*, 2016) and grandaite (Cámara *et al.*, 2014) from Valletta. Herein, we describe the new mineral castellaroite from both the Monte Nero mine and the Valletta mine.

The name castellaroite honours Fabrizio Castellaro (born 1970) of Mezzanego, Italy. Mr. Castellaro is an avid and sophisticated mineral collector specializing in the minerals of Liguria. He is the discoverer of castellaroite and has also discovered the new minerals balestraite, cerchiaraite-(Fe) and 1M O-D polytype of lavinskyite ("liguriaite"; IMA2014-035). Mr. Castellaro has agreed to the naming of the mineral in his honour.

The mineral and name (IMA2015–071) were approved by the IMA–CNMNC prior to publication. The description is based upon one holotype and three cotype specimens. The holotype and one cotype, both from the Monte Nero mine, are deposited in the collections of the Natural History Museum of Los Angeles County, 900 Exposition Boulevard, Los Angeles, CA 90007, USA, catalogue numbers 65603 (holotype) and 65604 (cotype). Two cotypes are deposited in the mineralogical collections of the Dipartimento di Scienze della Terra of the Università di Torino stored in the Museo Regionale di Scienze Naturali di Torino, Sezione di Mineralogia, Petrografia e Geologia, via Giovanni Giolitti 36, I-10123 Torino, Italy, catalogue numbers M/U 16950 (Monte Nero) and M/U 16951 (Valletta).

2. Occurrence

The new mineral was discovered by Fabrizio Castellaro on two specimens collected from the "classic" dump at the Monte Nero mine, Rocchetta Vara, La Spezia, Liguria, Italy (44°14'48"N, 9°45'27"E). The deposit is comprised of thin manganese stratiform ores that are located near the base of a chert sequence in the "Diaspri di Monte Alpe" formation, which overlies Jurassic ophiolites of the Bracco unit. Castellaroite is a secondary mineral that crystallized from As- and Mn-rich fluids, which circulated through a system of fractures during the final tectono-metamorphic stage of the deposit (Marescotti & Cabella, 1996). Other secondary minerals found with castellaroite include coralloite, manganohörnesite, rhodochrosite, sarkinite, sterlinghillite, strashimirite and wallkilldellite. The Monte Nero mine is also the type locality for coralloite (Callegari et al., 2012). More information on the deposit can be found in Callegari *et al.* (2012) and references therein.

Castellaroite has also been found on specimens recovered from the dumps of the Valletta mine, near Canosio, Piedmont, Italy (44°23'42"N, 7°5'42"E), where it was identified by SEM-EDX and Raman spectroscopy, and this is considered to be a cotype locality. The Valletta mine is a small Fe-Mn deposit that has not been worked in modern times and has never been studied geologically or petrologically. The geological setting is rather similar to that of the Monte Nero mine and has been summarized by Cámara *et al.* (2014) in their description of the new mineral grandaite, for which the Valletta mine is also the type locality. At the Valletta mine, castellaroite is a secondary mineral crystallized from hydrothermal fluids in an oxidizing environment. It is associated with braccoite, hematite, manganberzeliite, orthoclase and tiragalloite.

3. Physical and optical properties

At the Monte Nero mine, castellaroite occurs in radial aggregates up to about 5 mm in diameter consisting of very thin blades, and rarely as individual crystals up to about 2 mm in length (Fig. 1). At the Valletta mine, it forms aggregates of thin blades up to about 0.5 mm in length. Blades are flattened on [001], striated and elongated parallel to [100] and exhibit the forms {110}, {012} and {001} (Fig. 2). No twinning was observed. Crystals are colourless and transparent with a vitreous to silky lustre and white streak. The mineral is non-fluorescent. Crystals are flexible with a curved fracture, and one perfect cleavage on $\{001\}$. The Mohs' hardness is $2\frac{1}{2}$, based upon scratch tests. The density measured by floatation in Clerici solution is 3.14(2) g cm⁻³. The calculated density is 3.164 g \cdot cm⁻³, based on the empirical formula, and $3.174 \text{ g} \cdot \text{cm}^{-3}$, based on the ideal formula. The mineral is easily soluble in dilute HCl at room temperature. Optically, castellaroite crystals are biaxial (-), with $\alpha = 1.644(1), \beta = 1.662(1)$ and $\gamma = 1.667(1)$, measured in white light. The 2V measured directly on a spindle stage is 57(1)°; the calculated 2V is 55°. Dispersion is r < v, moderate and the optical orientation is $Y = \mathbf{b}$; $Z \approx \mathbf{a}$. The mineral is nonpleochroic. The Gladstone-Dale compatibility (Mandarino, 2007) is -0.019, in the range of superior compatibility, based upon the empirical formula.

4. Raman spectroscopy

The Raman spectra of castellaroite (Figs 3 and 4) from both the Monte Nero mine and the Valletta mine were obtained at the Dipartimento di Scienze della Terra (Università di Torino) using a micro/macro Jobin Yvon Mod. LabRam HRVIS, equipped with a motorized x-y stage and an Olympus microscope. The backscattered Raman signal was collected with $50 \times$ objective and the



Fig. 1. Castellaroite crystals from the Monte Nero mine. Field of view is 4 mm across.



Fig. 2. Crystal drawing of castellaroite, clinographic projection in nonstandard orientation, [100] vertical.

Raman spectra were obtained for non-oriented crystals. The 532 nm line of a solid-state Nd laser was used for excitation; laser power (250 mW) was controlled by means of a series of density filters. The minimum lateral and depth resolution was set to a few μ m. The system was calibrated using the 520.6 cm⁻¹ Raman band of silicon before each experimental session. The spectra were collected with multiple acquisitions (2–3) with single counting times ranging between 20 and 30 s. The

spectra were recorded using the LabSpec 5 program from 100 to 4000 cm⁻¹. Note that the spectrum obtained for the sample from Valletta is much weaker and, therefore, less well resolved; we only report it here from 100 to 1100 cm⁻¹. Band component analysis was undertaken using the Fityk software package (Wojdyr, 2010).

The Raman spectra confirmed the presence of $(AsO_4)^{3-1}$. The most characteristic features of the castellaroite spectra are observed in the region of $600-950 \text{ cm}^{-1}$ (values observed for Valletta material are reported in italics), where the spectra consist of four main overlapping bands observed at 800.5(800.4), 822.2(821.9), 846.8 (843.5) and 862.5(869.9) cm⁻¹, with two weaker shoulders at 910.6(910.2) and 933.8 cm^{-1} , the latter not well resolved in the Valletta spectrum (Fig. 3) because it is much weaker. The most intense bands of that group are the one at 862.5 cm⁻¹, which can be interpreted as the v_1 AsO₄ symmetric stretching vibration, and the one at 800.5 cm⁻¹, which can be interpreted as the v_3 AsO₄ antisymmetric stretching vibration. The band at 846.8 (843.5) is not well resolved for the Monte Nero sample and only weakly for the Valletta sample. The weak bands at 910.6(910.2) and 933.8 cm^{-1} could be interpreted as water vibrational modes (Martens et al., 2004). At wavenumbers 300–600 cm⁻¹, a prominent band centred at 425.6(426.5) cm⁻¹ is flanked on one side by two wellresolved peaks at 340(346.9) and 371.5(376.1) and on the other by a peak at 459.4(447.2) cm⁻¹. The bands at 425.6and 459.4 cm⁻¹ can be interpreted as the v₂ AsO₄ symmetric bending modes, and the ones at 340 and 371.5 cm^{-1} as the v₄ AsO₄ antisymmetric bending modes. Two weaker bands observed at 503.6 and 577.5 cm^{-1} are difficult to assign and may be related to the stretching of MnO_6 groups. In the lattice vibrations



Fig. 3. Raman spectra of castellaroite from the Monte Nero and Valletta mines in the $100-1100 \text{ cm}^{-1}$ range and, in the insert, the Raman spectrum of castellaroite from Monte Nero in the $1500-4000 \text{ cm}^{-1}$ range.



Fig. 4. Polyhedral layers parallel to {001} in castellaroite and metaswitzerite, viewed with [100] vertical. (online version in colour)

range (< 300 cm⁻¹), five weakly resolved bands can be recognized for the Monte Nero sample at 108.9, 149.9, 181.9, 215.1 and 256.9 cm⁻¹, with a shoulder at 143.2 cm⁻¹. The same bands are observed for the Valletta sample at *111.4, 149.6, 177.8, 215.9* and *252.1* cm⁻¹, with the shoulder shifted to lower wavenumbers and better resolved (*135.7* cm⁻¹) and a band at *177.8* that is very weak and split into two shoulders of adjacent bands at lower and higher wavenumbers (*165.0, 191.4*) (Fig. 3).

Regarding H_2O group bending and stretching vibrations, in the spectrum for the Monte Nero sample, a very weak band centred at 1662.6 cm⁻¹ is observed, which could be interpreted as the bending vibrations of the H_2O . The bands due to hydroxyl stretching modes of the H_2O groups are observed at 2925.1, 3115.9, 3240.9, 3491.4, 3757.7 and 3941.6 cm⁻¹, reflecting the complex nature of the hydrogen bonding on castellaroite (Fig. 3). Note that because of the paucity of material IR spectroscopy was not attempted.

Fig. 5. The crystal structures of castellaroite and metaswitzerite viewed parallel to $\{100\}$. The channel H₂O groups are shown as white spheres and are numbered for the castellaroite structure. The unit cells are shown by dashed lines. (online version in colour)

metaswitzerite

Table 1. Analytical data (in wt%) for castellaroite.

b

Constituent	Mean	Range	SD	Standard
$ \begin{array}{c} MnO\\ As_2O_5\\ P_2O_5\\ H_2O^*\\ Total \end{array} $	41.62 43.35 0.82 15.69 101.48	40.33–42.82 41.83–44.64 0.61–1.23	0.63 0.75 0.16	rhodonite syn. GaAs apatite

* Calculated on the basis of As + P = 2 *apfu*, charge balance and O = 12.5 apfu.

5. Chemical composition

Quantitative analyses (28 points on 7 crystals from Monte Nero) were performed at the University of Utah on a Cameca SX-50 electron microprobe with four wavelengthdispersive spectrometers (WDS). Analytical conditions were: 15 kV accelerating voltage, 20 nA beam current and a beam diameter of 10–15 μ m. Castellaroite exhibited minor damage under the electron beam. Counting times were 20 s on peak and 10 s on + and – background. No other elements than O, P, Mn and As were detected by energy-dispersive spectrometry. Other likely elements were sought by WDS,

Table 2. Powder X-ray data (d in Å) for castellaroite. Only calculated lines with I > 1 are shown.

I _{obs}	$d_{\rm obs}$	$d_{\rm calc}$	I_{calc}	hkl	I _{obs}	$d_{\rm obs}$	$d_{\rm calc}$	I_{calc}	hkl
100	10.90	10.9056	100	011			2.1812	4	400
67	9.27	9.2922	84	002	16	0.165	2.1726	6	315
		7.6485	4	012	40	2.105	2.1648	10	$\bar{4} 0 2$
42	6.97	6.9849	28	$\bar{1} \ 1 \ 1$			2.1539	6	161
26	6.63	6.6529	22	111	7	2 000	(2.0925	2	325
		5.6280	3	013	/	2.088	2.0848	2	047
		5.5618	5	112			2.0632	2	038
11	5.20	5.1963	7	$\bar{1} \ 2 \ 1$			(2.0424	2	$\bar{4} 0 4$
		4,7313	5	$\bar{1}$ 2 2	10	2.0355	2.0378	2	163
9	4,609	4.5714	10	113			2.0261	3	$\frac{1}{3}$ 1 7
10	4.373	4.3624	7	200	_		(1.9660	2	237
10	110 / 0	4.1487	3	$\frac{1}{1}$ $\frac{2}{2}$ $\frac{3}{3}$	7	1.9505	1.9357	2	431
12	4 031	4 0424	8	032	6	1 9218	1 9129	2	404
12	1.051	3 9920	2	130	9	1 8951	1 8906	3	432
		3 8721	2	131		1.0901	1 8646	2	317
11	3 830	3 8258	6	202			1.8593	2	$\frac{3}{4}0.6$
11	5.050	(3 7207	8	$\frac{2}{1}$ 3 2	18	1.8590	1.8585	6	0.0.10
17	3.692	3 6352	11	033			1.8576	3	354
		3 6172	3	132			1.0570	2	265
24	3 50/	3 5963	22	$\frac{1}{1} \frac{3}{2} \frac{2}{4}$			(1.7603	4	049
0	3 453	3 1158	18	124	6	1.7504	1 7464	2	158
9	5.455	(2 2244	10	$\frac{1}{2}$ $\frac{2}{1}$ $\frac{4}{1}$			(1.7404	2	1 3 0
47	3.323	2 2121	22	204	9	1.7167	1.7223	2	443 220
		2 2295	5	041			(1.0900	2	329
		2 2220	2	115	12	1 6920	1.0005	2 5	262
		5.2229	2	220	15	1.0850	1.0651	2	203
20	3.115	2 1261	12	$\frac{2}{1}250$			1.0704	2	522
		2 0979	12	$\frac{1}{1} \frac{2}{2} \frac{3}{4}$			1.0364	2	$\frac{3}{2}$
		5.0878	50	1 5 4	18	1.6295	1.0300	2	505 E 24
87	3.043	2 0212	50 15	204			1.0312	2	024
		(3.0213	15	232			1.0240	4	500
17	2.961	2.9773	0	125			1.590/	2	523
		(2.9383	5	043			1.5928	3	249
10	2 00 4	2.9119	9	$\frac{2}{5}$ $\frac{3}{2}$	20	1.5819	1.5844	4	3 2 10
18	2.894	2.8872	3	$\frac{2}{2}$ $\frac{1}{2}$ $\frac{5}{2}$			(1.5828	3	084
		2.8638	8	$\frac{2}{1}$ $\frac{3}{2}$ $\frac{3}{5}$			1.5/50	2	329
		2.7407	6	126			1.5/1/	4	365
		2.7264	4	044 $\bar{2}21$			1.5/06	2	280
		2.6/23	8	321	22	1 5500	1.5691	3	$\frac{281}{447}$
0.5	0.656	2.6699	12	320	32	1.5589	1.5600	2	44/
85	2.656	2.6638	5	215			1.5588	2	438
		2.6579	18	$\frac{241}{2}$			1.5565	4	282
1.5	2 (14	2.6334	18	206			1.54/2	3	524
17	2.614	2.6197	7	126			1.5371	4	526
8	2.566	2.5872	6	052	15	1.5290	(1.5336	3	085
		(2.5495	8	036			1.5133	4	169
15	2.492	2.4954	13	045	6	1.4760	1.4816	3	169
		2.4646	3	$\frac{1}{5}$ 5 2			1.4617	2	2411
3	2.418	2.4033	4	332	19	1.4550	(1.4557	2	$\frac{6}{2}$ 0 2
4	2.327	2.3303	4	054			1.4395	2	193
•	2.321	2.3231	2	008	10	1.4296	1.4229	2	2_3 12
		2.2852	2	037	12	1.3783	1.3807	4	561
7	2.216	2.2285	2	0 6 1					
/	2.216	2.2214	3	245					

but none were above the detection limits. Raw X-rav intensities were corrected for matrix effects with a ϕ (pz) algorithm (Pouchou & Pichoir, 1991). Because insufficient material was available for a direct determination of H₂O, the amount of water was calculated on the basis of As + P = 2 atoms per formula unit (*apfu*), charge balance and O = 12.5 apfu, as determined by the crystal-structure analysis which also confirmed the divalent state of Mn (see below). Analytical data are given in Table 1. $Mn^{2\bar{+}}_{3.02}(As_{1.94}P_{0.06})$ The empirical formula is $\Sigma_{2.00}O_{12.5}H_{8.96}$. The simplified structural formula is Mn² $(AsO_4)_2 \cdot 4.5H_2O$, which requires MnO 40.63, As_2O_5 43.89, H₂O 15.48, total 100 wt%.

6. X-ray crystallography and crystal-structure determination

Powder X-ray diffraction data for castellaroite were obtained on a Rigaku R-Axis Rapid II curved–imagingplate microdiffractometer utilising monochromatised MoKa radiation. A Gandolfi-like motion on the φ and ω axes was used to randomize the sample. Observed *d* spacings and intensities were derived by profile fitting using JADE 2010 software (Materials Data Inc.). Data are given in Table 2. The unit-cell parameters refined from the powder data using JADE 2010 with whole-pattern fitting are: a = 8.733(4), b = 13.438(4), c = 18.643(4) Å, $\beta = 95.006(11)^{\circ}$ and V = 2179.5(13) Å³.

Single-crystal structure data were obtained on the same instrument noted above. The Rigaku CrystalClear software package was used for processing of structure data, including the application of an empirical multi-scan absorption correction using ABSCOR (Higashi, 2001). The structure was solved by direct methods using SIR2011 (Burla et al., 2012). SHELXL-97 (Sheldrick, 2008) was used for the refinement of the structure, with scattering curves for neutral atoms (Wilson, 1992). The As and Mn sites were refined with fixed full occupancy. Two O sites in the channel of the structure refined to approximately 1/2 occupancy and were assigned exactly 1/2 occupancy in the final refinement. The poor diffraction quality of crystals caused by a strong tendency toward crystal warping resulted in the high R factors and large electron-density residuals. In spite of the poor diffraction quality, it was possible to refine anisotropic displacement parameters for all atoms. The very reasonable bond lengths, polyhedral geometries and bond-valence sums corroborate the correctness of the structure. Data collection and refinement details are given in Table 3, atom coordinates and displacement parameters in Table 4, selected bond distances in Table 5 and a bond valence analysis in Table 6. The Crystallography Information File (CIF), including reflection data, is available online as supplementary material linked to this article on the GSW website of the journal, http://eurjmin.geoscienceworld. org.

Table 3. Data collection and structure refinement details for castellaroite.

Diffractometer	Rigaku R-Axis Rapid II
X-ray radiation/power	MoKa ($\lambda = 0.71075$ Å)/50 kV, 40 mA
Temperature	293(2) K
Structural formula	$Mn_3(AsO_4)_2 \cdot 4.5H_2O$
Space group	$P2_1/n$
Unit-cell dimensions	a = 8.7565(8) Å
	b = 13.4683(13) Å
	c = 18.652(2) Å
	$\beta = 94.876(7)^{\circ}$
V	2191.7(4) Å ³
Ζ	8
Density (for above formula)	$3.174 \text{ g} \cdot \text{cm}^{-3}$
Absorption coefficient	9.458 mm^{-1}
F(000)	2000
Crystal size	$240 \times 45 \times 5 \ \mu m$
θ range	3.03–25.25°
Index ranges	$-10 \le h \le 10, -15 \le k \le 15, -22 \le l \le 22$
Refls. collected/unique	$17418/3407; R_{int} = 0.093$
Reflections with $F > 4\sigma(F)$	2513
Completeness to $\theta = 19.96^{\circ}$	85.9 %
Max. and min. transmission	0.954 and 0.210
Refinement method	Full-matrix least-squares on F^2
Parameters/restraints	325/0
GoF	1.346
Final R indices $[F_o > 4\sigma(F)]$	$R_1 = 0.1182, wR_2 = 0.3270$
R indices (all data)	$R_1 = 0.1365, wR_2 = 0.3503$
Largest diff. peak/hole	$+4.87/-3.68 \ e/A^3$

 $*R_{\text{int}} = \Sigma [F_o^2 - F_o^2(\text{mean})] / \Sigma [F_o^2]. \text{ GoF} = S = \{ \Sigma [w(F_o^2 - F_c^2)^2] / (n-p) \}^{1/2}. R_1 = \Sigma [|F_o| - |F_c|] / \Sigma [F_o|. wR_2 = \{ \Sigma [w(F_o^2 - F_c^2)^2] / \Sigma [w(F_o^2)^2] \}^{1/2}; w = 1 / [\sigma^2 (F_o^2) + (aP)^2 + bP] \text{ where } a \text{ is } 0.2, b \text{ is } 0 \text{ and } P \text{ is } [2F_c^2 + \text{Max}(F_o^2, 0)] / 3.$

Table 4. A	tom fractional coo	rdinates and displa	acement parameter.	s (Å ²) for castel	llaroite.					
	x	у	ы	$U_{\rm eq}$	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
Mn1	0.9614(3)	0.58628(18)	0.26437(14)	0.0238(7)	0.0210(15)	0.0110(14)	0.0400(17)	0.0017(9)	0.0063(10)	-0.0004(10)
Mn2	0.5628(3)	0.58800(18)	0.26092(14)	0.0233(7)	0.0200(14)	0.0149(15)	0.0349(16)	0.0021(9)	0.0018(10)	-0.0001(10)
Mn3	0.4562(3)	0.31797(18)	0.25484(14)	0.0235(7)	0.0208(14)	0.0109(15)	0.0395(17)	-0.0016(9)	0.0061(10)	-0.0001(10)
Mn4	0.0640(3)	0.31829(17)	0.25889(14)	0.0232(7)	0.0177(14)	0.0118(15)	0.0401(17)	-0.0010(9)	0.0027(10)	-0.0029(10)
Mn5	0.1037(3)	0.31049(19)	0.45711(15)	0.0275(7)	0.0219(15)	0.0221(16)	0.0384(17)	-0.0039(10)	0.0021(11)	-0.0003(10)
Mn6	0.3812(3)	0.42197(18)	0.54981(14)	0.0267(7)	0.0263(15)	0.0162(15)	0.0373(17)	0.0000(10)	0.0009(10)	-0.0040(10)
As1	0.21688(18)	0.58951(13)	0.64641(9)	0.0228(6)	0.0208(11)	0.0128(11)	0.0353(12)	0.0004(6)	0.0048(7)	-0.0006(6)
As2	0.23328(19)	0.23738(13)	0.64198(9)	0.0244(6)	0.0211(11)	0.0162(11)	0.0363(12)	0.0015(6)	0.0046(7)	-0.0005(7)
As3	0.28564(18)	0.48655(12)	0.36085(9)	0.0215(6)	0.0183(10)	0.0115(11)	0.0350(12)	-0.0011(6)	0.0033(7)	-0.0014(6)
A_{S4}	0.27166(18)	0.14694(14)	0.35133(9)	0.0245(6)	0.0208(11)	0.0158(11)	0.0372(12)	-0.0007(6)	0.0052(7)	0.0007(7)
01	0.6658(13)	0.3246(8)	0.3185(7)	0.031(3)	0.026(6)	0.016(6)	0.050(8)	-0.002(5)	-0.008(5)	-0.006(5)
02	0.2718(12)	0.5586(8)	0.5659(6)	0.032(3)	0.020(6)	0.021(7)	0.056(9)	-0.001(5)	0.014(5)	0.012(5)
03	0.7709(12)	0.5142(8)	0.3035(6)	0.026(3)	0.019(6)	0.008(6)	0.051(8)	0.007(4)	0.003(5)	0.005(5)
04	0.9644(11)	0.3663(8)	0.3585(6)	0.023(2)	0.020(6)	0.016(6)	0.033(7)	-0.007(4)	0.000(4)	0.002(4)
05	0.2281(12)	0.1159(9)	0.6288(6)	0.029(3)	0.024(6)	0.028(7)	0.036(7)	-0.006(5)	0.006(5)	0.006(5)
90	0.5765(12)	0.2280(8)	0.1810(6)	0.026(3)	0.020(6)	0.014(6)	0.045(8)	-0.006(4)	0.012(5)	-0.002(5)
07	0.2409(13)	0.3026(8)	0.5653(6)	0.029(3)	0.033(7)	0.026(7)	0.027(7)	-0.009(5)	0.001(5)	0.001(5)
08	0.8912(11)	0.2335(8)	0.1962(6)	0.025(3)	0.009(5)	0.018(6)	0.048(8)	-0.009(5)	0.007(4)	-0.004(4)
60	0.1599(13)	0.5728(8)	0.3372(6)	0.028(3)	0.030(7)	0.011(6)	0.043(8)	-0.004(5)	-0.004(5)	-0.003(5)
010	0.2727(13)	0.3943(8)	0.3008(6)	0.030(3)	0.029(7)	0.029(7)	0.033(7)	0.001(5)	0.004(5)	-0.005(5)
011	0.2603(12)	0.4389(9)	0.4422(6)	0.027(3)	0.025(6)	0.039(8)	0.020(6)	0.000(5)	0.010(4)	-0.006(5)
012	0.4622(12)	0.5394(8)	0.3616(6)	0.027(3)	0.031(6)	0.017(6)	0.031(7)	0.001(4)	0.000(5)	-0.005(5)
013	0.2772(12)	0.2173(8)	0.4243(6)	0.031(3)	0.022(6)	0.027(7)	0.045(8)	-0.005(5)	0.008(5)	-0.005(5)
014	0.2710(13)	0.0260(8)	0.3696(6)	0.027(3)	0.033(7)	0.017(6)	0.031(7)	0.003(4)	0.003(5)	0.001(5)
015	0.4225(11)	0.1762(8)	0.3056(6)	0.024(2)	0.014(5)	0.019(7)	0.040(7)	-0.004(4)	0.009(4)	-0.005(5)
016	0.1098(12)	0.1733(8)	0.2983(6)	0.028(3)	0.020(6)	0.029(7)	0.034(7)	-0.003(5)	-0.010(4)	-0.002(5)
OW17	0.4943(12)	0.4560(8)	0.1905(6)	0.027(3)	0.023(6)	0.015(6)	0.042(7)	-0.002(4)	0.000(5)	0.004(5)
OW18	0.2432(12)	0.2744(8)	0.1769(6)	0.026(3)	0.025(6)	0.013(6)	0.040(7)	-0.010(4)	0.000(5)	0.011(5)
0W19	0.9946(12)	0.4539(8)	0.1954(6)	0.027(3)	0.024(6)	0.003(6)	0.053(8)	0.005(4)	0.006(5)	-0.005(5)
OW20	0.7623(12)	0.1379(8)	0.3212(7)	0.032(3)	0.026(7)	0.014(7)	0.058(9)	-0.006(5)	0.013(5)	0.006(5)
OW21	0.9554(16)	0.3867(13)	0.5258(8)	0.066(5)	0.056(10)	0.100(14)	0.042(9)	-0.014(8)	0.001(7)	0.038(9)
OW22	0.9491(16)	0.1751(10)	0.4646(9)	0.058(4)	0.049(9)	0.037(9)	0.089(13)	0.006(7)	0.013(7)	-0.001(7)
OW23	0.5233(13)	0.3192(9)	0.4894(7)	0.037(3)	0.025(7)	0.024(8)	0.065(10)	-0.003(5)	0.010(6)	0.000(5)
0W24	0.1836(18)	0.3389(11)	0.0422(8)	0.064(4)	0.086(12)	0.046(10)	0.061(11)	0.012(7)	0.000(8)	-0.007(8)
0W25*	0.127(3)	0.0263(19)	0.5063(12)	0.043(6)	0.067(18)	0.038(16)	0.023(14)	0.006(10)	0.005(11)	-0.003(13)
OW26*	0.379(4)	0.028(2)	0.5217(16)	0.066(9)	0.09(2)	0.05(2)	0.06(2)	-0.008(14)	0.019(17)	-0.020(17)
* OW25 an	d OW26 are half occ	upied.								

Table 5. Selected bond distances (Å) for castellaroite.

2.105(11)				
2.105(11)	Mn4–O16	2.113(11)	As1–O1	1.646(11)
2.113(11)	Mn4–O8	2.158(10)	As1–O2	1.668(11)
2.121(11)	Mn4010	2.181(11)	As1–O3	1.679(10)
2.202(11)	Mn4–O4	2.215(10)	As1–O4	1.690(10)
2.232(11)	Mn4–OW19	2.233(11)	<as1–o></as1–o>	1.671
2.517(12)	Mn4–OW18	2.357(11)		
2.215	<mn4–o></mn4–o>	2.210	As2–O5	1.655(12)
			As2–O6	1.674(10)
2.131(10)	Mn5-O13	2.101(11)	As2–O7	1.684(12)
2.142(11)	Mn5–OW21	2.159(14)	As2–O8	1.689(11)
2.166(10)	Mn5-O11	2.239(11)	<as2–o></as2–o>	1.676
2.238(11)	Mn5–O4	2.249(11)		
2.261(11)	Mn5–O7	2.262(12)	As3–O9	1.635(11)
2.354(11)	Mn5–OW22	2.282(14)	As3–O10	1.670(11)
2.215	<mn5–o></mn5–o>	2.215	As3011	1.679(11)
			As3012	1.701(10)
2.101(11)	Mn6–O7	2.059(12)	<as3–o></as3–o>	1.671
2.147(12)	Mn6–O2	2.107(11)		
2.163(11)	Mn6-O12	2.121(11)	As4013	1.656(11)
2.173(10)	Mn6-O11	2.200(11)	As4–O14	1.664(11)
2.253(11)	Mn6–OW23	2.230(13)	As4015	1.679(10)
2.340(11)	<mn6–o></mn6–o>	2.143	As4–O16	1.695(10)
2.196			<as4–o></as4–o>	1.674
2.619(16)	OW20O1	2.651(15)	OW23–O2	2.704(17)
2.664(15)	OW20015	3.009(15)	OW23013	2.749(16)
2.667(18)	OW21–O2	2.617(18)	OW24–OW23	2.69(2)
2.864(14)	OW21-O9	2.878(18)	OW24–O14	3.018(19)
2.655(15)	OW22–OW24	2.85(2)		
2.716(15)	OW22–OW20	3.06(2)		
	$\begin{array}{c} 2.113(11)\\ 2.121(11)\\ 2.202(11)\\ 2.232(11)\\ 2.517(12)\\ 2.215\\ \hline\\ 2.131(10)\\ 2.142(11)\\ 2.166(10)\\ 2.238(11)\\ 2.261(11)\\ 2.261(11)\\ 2.354(11)\\ 2.215\\ \hline\\ 2.101(11)\\ 2.147(12)\\ 2.163(11)\\ 2.173(10)\\ 2.253(11)\\ 2.340(11)\\ 2.196\\ \hline\\ 2.619(16)\\ 2.664(15)\\ 2.667(18)\\ 2.864(14)\\ 2.655(15)\\ 2.716(15)\\ \hline\end{array}$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$

Table 6. Bond-valence analysis for castellaroite. Values are expressed in valence units.

	Mn1	Mn2	Mn3	Mn4	Mn5	Mn6	As1	As2	As3	As4	hydrogen bonds	Σ_{a}
01			0.43				1.39				+0.25	2.07
02						0.42	1.31				+0.28, +0.22	2.23
O3	0.42	0.36					1.27					2.05
04				0.32	0.29		1.23					1.84
05								1.35			+0.27, +0.22	1.84
06	0.33		0.36					1.29				1.98
07					0.28	0.48		1.25				2.01
08		0.39		0.37				1.23				1.99
09	0.41								1.43		+0.16, +0.16	2.16
O10			0.38	0.35					1.30			2.03
011					0.30	0.33			1.27			1.90
012		0.30				0.41			1.20			1.91
013					0.43					1.35	+0.20	1.98
014										1.32	+0.25, +0.25, +0.13	1.95
015	0.43		0.36							1.27	+0.13	2.19
016		0.40		0.42						1.21		2.03
OW17		0.28	0.29								-0.25,-0.27	0.05
OW18			0.23	0.22							-0.16,-0.24	0.05
OW19	0.30			0.30							-0.22,-0.25	0.13
OW20	0.14	0.22									+0.12, -0.13, -0.25	0.10
OW21					0.37						-0.16,-0.28	-0.07
OW22					0.26						-0.12,-0.17	-0.03
OW23						0.30					+0.23, -0.20, -0.22	0.11
OW24											+0.17, +0.24, -0.13, -0.23	0.05
Σ_{c}	2.16	1.97	2.02	1.96	1.94	1.90	5.13	4.94	5.32	5.01		

Bond valence parameters from Brown & Altermatt (1985). Hydrogen-bond strengths based on O–O bond distances from Ferraris & Ivaldi (1988). Halfoccupied OW25 and OW26 sites are not included. Donated and accepted hydrogen bonds are indicated by – and +, respectively.

7. Discussion of the structure

In the structure of castellaroite, there are four different AsO_4 tetrahedra, five different MnO_6 octahedra and one MnO_5 square pyramid. The octahedra centred by Mn1, Mn2, Mn3 and Mn4 share edges forming kinked chains parallel to [100]. The chains are linked to each other by corner-sharing, forming sheets parallel to {001}. All four AsO₄ tetrahedra corner-link with octahedra in the sheet, forming a heteropolyhedral layer (Fig. 4). The Mn5O₆–Mn6O₅ edge-sharing dimers are located between the heteropolyhedral layers and share corners with octahedra and tetrahedra in adjacent layers, thereby linking the layers in the [001] direction.

The heteropolyhedral layer is topologically identical to those in the structures of several phosphate minerals: angarfite, NaFe³⁺₅(PO₄)₄(OH)₄·4H₂O (Kampf *et al.*, 2012). bakhchisaraitsevite. $Na_2Mg_5(PO_4)_4 \cdot 7H_2O$ (Liferovich et al., 2000; Yakubovich et al., 2000), mejillonesite. NaMg₂(PO₃OH)(PO₄)(OH)·(H₃O + H₂O) (Atencio et al., 2012), metaswitzerite, Mn₃(PO₄)₂·4H₂O (formerly switzerite; Fanfani & Zanazzi, 1979), rimkorolgite, BaMg₅(PO₄)₄·8H₂O (Krivovichev et al., 2002) and switzerite, Mn₃(PO₄)₂·7H₂O (Zanazzi et al., 1986). Castellaroite is the first arsenate with this type of heteropolyhedral layer. The interlayer region in the structure of castellaroite is different from that in any of these structures, but it is most similar to that in metaswitzerite, which except for the configuration of Mn5O₆ and Mn6O₅ polyhedra, could be regarded as the PO₄ analogue of castellaroite.

In the metaswitzerite structure, pairs of $Mn6O_5$ pyramids form edge-sharing dimers and $Mn5O_6$ octahedra do not share edges with any other polyhedra. Although the linkages in both structures result in heteropolyhedral frameworks, the castellaroite framework contains continuous channels along [100], while the metaswitzerite framework contains discontinuous cavities. Isolated H₂O groups are contained in the channels in castellaroite (O24, O25 and O26) and the cavities in metaswitzerite are compared in Fig. 5.

The layers in the structure of switzerite, including outlying MnO_6 octahedra, are linked to adjacent layers only by hydrogen bonds through isolated H₂O groups in the interlayer region. Switzerite transforms readily to metaswitzerite by losing ³/₄ of its isolated H₂O groups resulting in the condensation of the layers into the framework noted above (Zanazzi *et al.*, 1986). An arsenate equivalent of switzerite, which could correspond to a hydrated precursor to castellaroite, is not known. Manganohörnesite, $Mn_3(AsO_4)_2 \cdot 8H_2O$, a member of the vivianite group, has a structure very different from that of switzerite and cannot transition to castellaroite through dehydration via simple structural condensation.

It is worth noting that sterlinghillite, Mn_3 (AsO₄)₂·3H₂O (Dunn, 1981; Matsubara *et al.*, 2000), which has been found with castellaroite, has 1½ less

 H_2O groups *pfu*. Its structure has never been determined; however, its powder XRD pattern is quite distinct from that of castellaroite. The sterlinghillite from the Monte Nero mine that we examined occurs as subparallel intergrowths of tiny plates. The plates are distinctly brittle and much harder than castellaroite. Unfortunately, efforts to separate a plate suitable for single-crystal study were unsuccessful.

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