The crystal structure of nissonite

LEE A. GROAT,* FRANK C. HAWTHORNE

Department of Geological Sciences, University of Manitoba, Winnipeg, Manitoba R3T 2N2, Canada

ABSTRACT

The crystal structure of nissonite, $[CuMg(PO_4)(OH)(H_2O)_2]_2(H_2O)$, monoclinic, a = 22.523(5), b = 5.015(2), c = 10.506(3)Å, $\beta = 99.62(2)^\circ$, V = 1169.9(6)Å³, space group C2/c, Z = 4, has been solved by direct methods and refined to an R index of 6.8% for 1271 observed reflections (>4 σ) measured on an automated four-circle single-crystal diffractometer using graphite-monochromated MoK α X-radiation. The structure consists of thick heteropolyhedral slabs parallel to (100), linked solely by a network of H bonds. Each slab consists of a sheet of corner-sharing $[Cu\phi_{10}]$ dimers sandwiched between two $[M(T\phi_4)\phi_3]$ sheets of corner-sharing (Mg ϕ_6) and (P ϕ_4) polyhedra; in each sheet, the polyhedra occupy the vertices of a 6³ net. One of the H₂O groups in the formula unit is not bonded to Mg, Cu, or P but is held in the structure solely by a network of H bonds. Graphically identical $[M(T\phi_4)\phi_3]$ sheets occur in the structure of newberyite and as components of the heteropolyhedral frameworks of the structures of metavariscite and variscite.

INTRODUCTION

Nissonite is a hydrous copper magnesium phosphate mineral described by Mrose et al. (1966). It occurs as thin crusts and occasional diamond-shaped crystals associated with such other Cu minerals as malachite, azurite, libethenite, turquoise, chrysocolla, and cuprite. The only other data in the literature are reported by Bayliss (1986), who gave improved X-ray powder data and cell dimensions; he also noted the paucity of material and derived the data by remeasuring the original Debye-Scherrer film of nissonite recorded by Mrose et al. (1966).

We have been investigating the graphical and topological aspects of the Cu²⁺ oxysalts for some time (Hawthorne, 1985a, 1985b, 1986; Hawthorne and Groat, 1985, 1986; Hawthorne and Eby, 1985; Hawthorne et al., 1989; Groat and Hawthorne, 1987; Eby and Hawthorne, 1989a, 1989b), and the structure of nissonite was of obvious interest.

EXPERIMENTAL

The material used in this work is from Panoche Valley, California, the original locality from which the mineral was described by Mrose et al. (1966) and the only locality from which it has been reported; the sample was obtained from a mineral dealer. As described by Bayliss (1986), the material available is extremely limited, and a considerable amount of time was spent examining unsuitable crystals. Finally, a crystal was found that gave only slightly diffuse spots, and this was used for the measurement of the metric and intensity data. four-circle diffractometer, and a random orientation photograph was used to locate 15 reflections that were automatically centered using graphite-monochromated MoK α X-radiation. The correct unit cell was chosen from an array of real space vectors corresponding to potential unit-cell axes, and least-squares refinement of the setting angles resulted in the cell parameters reported in Table 1. Intensity data were measured according to the method of Hawthorne and Groat (1985). A total of 1702 reflections was measured out to a maximum 2θ value of 60°. Two standard reflections were monitored every 48 reflections to check for stability and constancy of crystal alignment; no significant change was noted during data measurement. The intensity data were corrected for absorption, Lorentz, and polarization effects, and reduced to structure factors. A reflection was considered as observed if its magnitude exceeded that of three standard deviations based on counting statistics; of the 1702 unique reflections, 1271 were considered as observed.

The crystal was mounted on a Nicolet R3m automated

STRUCTURE SOLUTION AND REFINEMENT

Scattering curves for neutral atoms together with anomalous dispersion corrections were taken from the *International Tables for X-ray Crystallography* (1974); *R* indices are of the form given in Table 1 and are expressed as percentages. All calculations were done with the SHELXTL system of programs (Sheldrick, 1981).

The structure was determined using the weighted tangent formula method. The phase set with the maximum combined figure of merit resulted in an E-map, the most intense peaks of which could be interpreted in terms of the know ratios (and intensities) of the cations in nissonite. Most of the O atoms were located on the subsequent difference-Fourier map. Iterative full-matrix least-squares

^{*} Present address: Department of Geological Sciences, University of British Columbia, Vancouver, Lotus Land, British Columbia, Canada.

| TABLE I. IVISCEIIAREOUS GALA IOF RISSON | TABLE 1. | Miscellaneous | data | tor | nissonii |
|---|----------|---------------|------|-----|----------|
|---|----------|---------------|------|-----|----------|

| a | 22.523(5) Å | Crystal size (mm) | 0.10 × 0.16 × 0.16 |
|--------------------|--|------------------------------------|--------------------|
| b | 5.015(2) | Rad/mono | Mo/Gr |
| С | 10.506(3) | Total F _o | 1702 |
| β | 99.62(2)° | $ F_{o} > 4\sigma$ | 1271 |
| V | 1169.9(6) Å ³ | Final Robe | 6.8% |
| Space group | C2/c | Final R _{wobs} | 6.0% |
| Unit-cell cont | tents: 4{[CuMg | PO4)(OH)(H2O)2]2(H20 | O)} |
| $R = \Sigma (F_o$ | $ - F_{\rm c})/\Sigma \bar{F}_{\rm c} $ | | |
| $R_w = [\Sigma w($ | $(F_{o} - F_{o})^{2}/\Sigma$ | $wF_{o}^{2}]^{\frac{1}{2}}, w = 1$ | |

refinement and difference-Fourier maps located the remaining non-H atoms, eventually converging to an R index of 7.1% for an isotropic displacement model. A bond-valence calculation identified the O2-, OH, and H2O anions, and these were relabeled accordingly. Displacement factors were converted to anisotropic parameters, and the structure was refined to convergence. A difference-Fourier map calculated at this stage did not show any H-atom positions that are structurally reasonable. Consequently, a H-bond network was deduced on the basis of bond-valence and geometrical criteria, and H positions were derived; these were put into the structure model as fixed parameters. Subsequent full-matrix leastsquares refinement of all variables resulted in convergence at an R index of 6.0%; this value is higher than one would normally expect and reflects the rather inferior nature of the material. Final atomic parameters are given in Tables 2 and 3, and observed and calculated structure factors are deposited (Table 4)1. Selected interatomic distances and angles are given in Table 5, and an empirical bond-valence table, calculated with the curves of Brown (1981), is shown as Table 6.

¹ To receive a copy of Table 4, order Document Am-90-439 from the Business Office, Mineralogical Society of America, 1130 Seventeenth Street NW, Suite 330, Washington, DC 20036. Please remit \$5.00 in advance for the microfiche.

TABLE 2. Atomic parameters for nissonite

| Atom | x | У | z | $U_{\rm lso}^{*}$ | |
|--------|------------|-----------|------------|-------------------|--|
| Cu | 0.24534(5) | 0.1541(3) | 0.36677(9) | 0.96(3) | |
| Р | 0.3437(1) | 0.1768(6) | 0.2264(2) | 0.83(6) | |
| Mg | 0.1226(1) | 0.1547(8) | 0.4673(2) | 0.73(8) | |
| O(1) | 0.3631(3) | 0.103(1) | 0.3702(5) | 1.4(2) | |
| O(2) | 0.2770(3) | 0.296(1) | 0.2161(5) | 0.9(2) | |
| O(3) | 0.3846(3) | 0.388(1) | 0.1833(5) | 1.2(2) | |
| O(4) | 0.3405(3) | -0.072(2) | 0.1447(5) | 1.3(2) | |
| OH | 0.2136(3) | 0.055(1) | 0.5211(5) | 0.7(1) | |
| OW(1) | 0.0866(3) | -0.138(2) | 0.5790(5) | 1.5(2) | |
| OW(2) | 0.0359(3) | 0.288(2) | 0.4132(6) | 2.4(3) | |
| OW(3) | 1/2 | 0.201(3) | 1/4 | 2.8(4) | |
| H(OH) | 0.245 | 0 | 0.578 | 1.5 | |
| H(W1a) | 0.105 | 0 | 0.598 | 1.5 | |
| H(W1b) | 0.057 | 0.015 | 0.639 | 1.5 | |
| H(W2a) | 0.023 | 0.434 | 0.355 | 1.5 | |
| H(W2b) | 0 | 0.237 | 0.415 | 1.5 | |
| H(W3) | 0.460 | 0.268 | 0.227 | 1.5 | |

* $U_{\rm iso} = U_{\rm iso} \times 10^{\circ};$ H positions derived from bond-valence and stereo-chemical arguments.



Fig. 1. Details of the local polyhedral geometry in nissonite projected along [001].

STRUCTURE DESCRIPTION

There is one unique P position in nissonite, tetrahedrally coordinated by four O atoms; the resulting oxyanion shows a typical range of bond lengths and angles (cf. Baur, 1974). The variation in P-O bond lengths cor-



Fig. 2. Polyhedral representation of the structure of nissonite projected down [010]. (Mg ϕ_6) and (P ϕ_4) polyhedra are shown ruled, and (Cu ϕ_6) polyhedra are shown by dashed shading; the small circles in the upper part of the figure represent Cu atoms, and Cu-O bonds are indicated by the broken lines; the large circles represent the H₂O groups that are not directly bonded to Mg, Cu, or P, and the dotted lines represent H bonds with the donor-acceptor character indicated by the arrows. One unit cell is outlined.

relates well with the local bond-valence requirements of the corresponding O anions. There is one Mg position, coordinated by six anions in a (distorted) octahedral arrangement; of these, three are (phosphate) O atoms, one is an OH, and the remaining two are H₂O groups, the latter being in a *cis* relationship. There is one unique Cu²⁺ position, surrounded by six anions in a very distorted octahedral arrangement, the Cu- ϕ distances varying from 1.94–2.66 Å. This type of geometry is typical for Cu²⁺, being a result of the Jahn-Teller effect. Details of the local geometry are shown in Figure 1; note that the three unique polyhedra are connected by shared edges, all of which are shortened relative to the unshared edges of the same polyhedra.

TABLE 3. Anisotropic displacement factor coefficients for nissonite

| Atom | U11* | U22 | U ₃₃ | U ₂₃ | U ₁₃ | U12 |
|--------|-----------------------|---------|-----------------|-----------------|-----------------|---------|
| Cu | 86(5) | 131(6) | 55(4) | -16(5) | -33(3) | -23(6) |
| P | 61(10) | 115(13) | 56(8) | 2(10) | -43(7) | 13(12) |
| Mg | 49(13) | 108(16) | 44(11) | -1(13) | -40(10) | -10(16) |
| O(1) | 177(34) | 167(43) | 59(24) | 13(25) | -63(23) | 59(32) |
| O(2) | 58(28) | 123(40) | 88(25) | 5(25) | -23(21) | 34(29) |
| O(3) | 83(30) | 129(41) | 122(26) | 35(26) | - 19(23) | -5(30) |
| O(4) | 85(32) | 169(39) | 123(28) | -36(26) | -38(24) | 8(31) |
| OH | 71(12) | . , | . , | . , | . , | . , |
| OW(1) | 127(31) | 176(39) | 131(27) | -8(31) | -47(23) | -47(36) |
| OW(2) | 59(32) | 355(56) | 263(36) | 0 | -55(27) | 0 |
| OW(3) | 90(47) | 292(76) | 388(57) | 0 | -134(42) | 0 |
| * U, = | $U_{\mu} \times 10^3$ | | | | | |

| Cu-O(1) Cu-O(2) Cu-OH Cu-O(2)a Cu-O(2)a Cu-O(4)b Cu-OHc (Cu-O) P-O(1) | 2.660(7) 1.975(6) 1.943(6) 2.021(7) 2.360(7) 2.003(6) 2.160 1.545(6) | Mg-OH Mg-OW(1) Mg-OW(2) Mg-O(1)c Mg-O(3)a Mg-O(4)b ⟨Mg-O⟩ | 2.095(7) 2.122(8) 2.051(7) 2.075(7) 2.075(7) 2.077(8) 2.079 |
|--|--|---|--|
| P-O(2) P-O(3) P-O(4) | 1.604(7) 1.523(7) 1.507(7) | | |
| (P-O) | 1.545 | | |
| O(1)-O(2)* O(1)-OH O(1)-OHC O(2)-O(2)a O(2)-O(2)a O(2)-O(4)b O(2)-OHc OH-O(2)a OH-O(4)b* OH-OHc* O(2)a-O(4)b O(4)b-OHc | 2.504(8) 3.963(9) 3.493(9) 2.81(1) 2.927(7) 3.299(9) 2.834(8) 2.847(8) 2.707(9) 2.64(1) 3.608(9) 2.937(9) | O(1)-Cu-O(2)* O(1)-Cu-OH O(1)-Cu-O(2)a O(1)-Cu-OHc O(2)-Cu-O(2)a O(2)-Cu-O(2)a O(2)-Cu-OHc OH-Cu-O(2)a OH-Cu-O(4)b* OH-Cu-OHc* O(2)a-Cu-O(4)b O(4)b-Cu-OHc | 63.4(2) 118.0(2) 95.6(2) 72.6(2) 94.2(2) 98.7(2) 90.8(3) 91.8(3) 77.3(2) 83.9(3) 110.6(2) 84.2(2) |
| (O-O Cu) | 3.05 | (O-Cu-O) | 90.1 |
| OH-OW(1) OH-O(1)c OH-O(3)a OH-O(4)b* OW(1)-OW(2) OW(1)-O(1)c OW(1)-O(3)a OW(2)-O(1)c OW(2)-O(3)a OW(2)-O(4)b O(1)c-O(4)b O(1)c-O(4)b O(3)a-O(4)b | 3.173(9) 2.81(1) 2.935(8) 2.707(9) 2.87(1) 2.92(1) 2.938(8) 2.986(8) 2.986(8) 2.986(1) 3.03(1) 3.015(9) 2.890(9) | OH-Mg-OW(1) OH-Mg-O(1)c OH-Mg-O(3)a OH-Mg-O(4)b* OW(1)-Mg-OW(2) OW(1)-Mg-O(1)c OW(1)-Mg-O(1)c OW(2)-Mg-O(1)c OW(2)-Mg-O(1)c OW(2)-Mg-O(4)b O(1)c-Mg-O(4)b O(3)a-Mg-O(4)b | 97.6(3) 84.7(3) 90.0(3) 81.1(3) 86.8(3) 88.3(3) 89.3(3) 92.7(3) 92.8(3) 94.6(3) 93.3(3) 88.9(3) |
| (O-O Mg) | 2.94 | (O-Mg-O) | 90.0 |
| O(1)-O(2)* O(1)-O(3) O(1)-O(4) O(2)-O(3) O(2)-O(4) O(3)-O(4) | 2.504(8) 2.538(9) 2.497(8) 2.546(9) 2.52(1) 2.52(1) | O(1)-P-O(2)* O(1)-P-O(3) O(1)-P-O(4) O(2)-P-O(3) O(2)-P-O(4) O(3)-P-O(4) | 105.3(4) 111.7(4) 109.8(4) 109.0(4) 108.4(4) 112.3(4) |
| (O-O P) | 2.52 | (O-P-O) | 109.4 |
| OH-H(OH) OH-O(2)d OH-O(4)d H(OH)-O(2)d H(OH)-O(4)d | 0.89 2.89 2.94 2.12 2.18 | OW(2)-H(W2a) OW(2)-H(W2b) OW(2)-OW(3)f OW(2)-OW(1)e H(W2a)-OW(3)f H(W2b)-OW(1)e | 0.97 0.85 2.73 2.87 1.76 2.02 |
| OW(1)-H(W1a) OW(1)-H(W1b) OW(1)-O(1)c OW(1)-O(3)c H(W1a)-O(1)c H(W1b)-O(3)c | 0.81 1.25 2.93 2.77 2.12 2.15 | OW(3)-H(W3) OW(3)-O(3)g H(W3)-O(3)g | 0.95 2.74 1.79 |

Note: a: 1/2 - x, y - 1/2, 1/2 - z; b: 1/2 - x, 1/2 + y, 1/2 - z; c: 1/2 - x, 1/2 - y, 1 - z; d: x, -y, 1/2 + z; e: -x, -y, 1 - z; f: -1/2 + x, 1/2 + y, 1/2 - z; g: 1 - x, y, 1/2 - z. * Shared edge.

Figure 2 shows the structure projected down [010]. In this orientation, the structure exhibits a very layered character. The $(Cu\phi_6) (\phi = unspecified ligand), (Mg\phi_6)$ and $(P\phi_4)$ polyhedra form thick slabs parallel to (100). Intercalated between these slabs are planar layers of (H₂O)

TABLE 5. Interatomic distances (Å) and angles (°) in nissonite







Fig. 4. Geometrically idealized versions of the polyhedral sheets shown in Figure 3. (a) The $[M(T\phi_4)\phi_3]$ sheet and (b) the $[Cu\phi_n]$ sheet.

Fig. 3. The sheet elements of the nissonite structure viewed down [100]; the legend is as in Figure 2. (a) The [Mg(PO₄)(OH)- $(H_2O)_2$] sheet, consisting of alternating (Mg ϕ_6) and (P ϕ_4) polyhedra arranged at the vertices of a 6³ net and (b) the [Cu ϕ_n] sheet, consisting of edge-sharing dimers of (Cu ϕ_6) octahedra that share corners to form an array with the polyhedra at the vertices of a 6³ net.

groups that provide the necessary interslab linkage by a H-bond network; this overall structure accounts nicely for the $\{100\}$ cleavage shown by nissonite. As is apparent

from Figure 2, each slab consists of three polyhedral sheets, a central sheet of $(Cu\phi_6)$ octahedra sandwiched between heteropolyhedral sheets of $(Mg\phi_6)$ octahedra and $(P\phi_4)$ tetrahedra.

The constitution of the sheets is illustrated in Figure 3. The sandwiching Mg-P nets are symmetrically equivalent; one is illustrated in Figure 3a. It consists of a cornersharing array of $(Mg\phi_6)$ and $(P\phi_4)$ polyhedra with unit stoichiometry $[Mg(PO_4)(OH)(H_2O)_2]$. In this orientation,

| TABLE 6. Empirical bond-valence* ta | ble for nissonite |
|-------------------------------------|-------------------|
|-------------------------------------|-------------------|

| | Cu | Mg | Р | H(1) | H(2) | H(3) | H(4) | H(5) | H(6) | Σ |
|-------------------------|----------------|-------------------------|----------------|--------------|------|------|----------------|--------------|----------------|-------------------------------|
| O(1) O(2) | 0.072 | 0.345 | 1.232 1.046 | 0.15 | 0.13 | | | | | [1.779] [2.004] |
| O(3) O(4) OH | 0.149 0.476 | 0.359 0.348 0.328 | 1.321 1.379 | 0.12 0.73 | | 0.19 | | | 0.19 | [2.060] [1.996] [1.933] |
| OW(1) OW(2) OW(3) | 0.399 | 0.323 0.355 | | | 0.87 | 0.81 | 0.80 0.20≚² | 0.16 0.84 | 0.81 ≚² | [2.163] [1.995] [2.020] |
| Σ | 1.904 | 2.058 | 4.978 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | 1.0 | |
| * Calcu | lated from the | e curves of Bro | own (1981). | | | | | | | |



Fig. 5. The structural units in (a) newberyite (Sutor, 1967); (b) metavariscite (Kniep and Mootz, 1973); and (c) variscite (Kniep et al., 1977).

the sheet links outward (along a) to the $[Cu\phi_n]$ sheet. It is notable that the connectivity of the polyhedra is that of a 6³ net. A geometrically idealized version of this sheet is shown in Figure 4a.

The central $[Cu\phi_n]$ sheet consists of edge-sharing [Cu₂O₈(OH)₂] dimers that link by corner-sharing [through O(2) atoms] to form the sheet shown in Figure 3b. This sheet may be geometrically idealized to that shown in Figure 4b, where the polyhedral connectivity is based on that of a 6³ net. The relationship between the actual sheet and its idealized analogue is immediately apparent by comparing Figures 3b and 4b; the observed sheet has a commensurate modulation along c. The origin of this modulation stems from the requirement that this sheet link to the sandwiching $[Mg(P\phi_4)\phi_3]$ sheets. In the idealized $[Cu\phi_n]$ sheet, the anions that link to the neighboring $[Mg(P\phi_4)\phi_3]$ sheet form a (geometrically distorted) 6³ net. The modulation brings this array of anions much closer to a geometrically ideal 63 net, matching it with the corresponding array in the adjacent $[Mg(P\phi_4)\phi_3]$ net.

COMPARISON WITH RELATED STRUCTURES

As discussed above, the strongly-bonded structural unit in nissonite consists of a sheet of $(Cu\phi_6)$ octahedra, sandwiched between heteropolyhedral $[Mg(PO_4)(OH)(H_2O)_2]$ sheets; the latter may be written in the general form $[M(T\phi_4)\phi_3]$, where M = six-coordinated cations, T = fourcoordinated cations, ϕ = unspecified anions. Graphically identical $[M(T\phi_4)\phi_3]$ sheets are found in several other minerals. A similar sheet constitutes the structural unit in newberyite, $[Mg(PO_3OH)(H_2O)_3]$; this is illustrated in Figure 5a, and a comparison with the sheet of Figure 3a shows a subtle difference between the two sheets. In nissonite, each tetrahedron has one face in the plane of the sheet, and all tetrahedra point in the same direction, toward the central $[Cu\phi_n]$ sheet. In newberyite, each tetrahedron has one face in the plane of the sheet, but alternate tetrahedra point different ways (along b). Consequently, the sheets in nissonite and newberyite are graphically identical but are geometrical isomers (Hawthorne, 1983, 1985c).

Graphically identical sheets are also found in the structures of metavariscite, [Al(PO₄)(H₂O)₂], and variscite, $[Al(PO_4)(H_2O)_2]$ (Figs. 5a and 5b); in both of these structures, adjacent [M(T ϕ_4) ϕ_3] sheets condense to form heteropolyhedral frameworks. Both sheets have the same tetrahedral conformation as the heteropolyhedral sheet in nissonite, with alternate tetrahedra pointing in opposite directions orthogonal to the plane of the sheet. However, there are still geometrical differences between the sheets. In metavariscite and variscite, three tetrahedra link to each octahedron such that the common vertices are all cis to each other (Figs. 5b and 5c). In nissonite, the tetrahedra link to each octahedron such that the linking vertices are both cis and trans. Consequently, the sheets in nissonite and in metavariscite and variscite are also geometrical isomers.

ACKNOWLEDGMENTS

We thank Dan Appleman and Paul Moore for their helpful comments. This work was supported by the Natural Sciences and Engineering Research Council of Canada in the form of a Graduate Fellowship to L.A.G., Operating, Major Equipment and Infrastructure Grants, and a University Research Fellowship to F.C.H.

References cited

- Baur, W.H. (1974) The geometry of polyhedral distortions. Predictive relationships for the phosphate group. Acta Crystallographica, B30, 1195-1215.
- Bayliss, Peter (1986) X-ray powder data for nissonite and waylandite. Powder Diffraction, 1, 331-333.
- Brown, I.D. (1981) The bond-valence method: An empirical approach to chemical structure and bonding. In M. O'Keeffe and A. Navrotsky, Eds., Structure and bonding in crystals, vol. II, p. 1-30. Academic Press, New York.
- Eby, R.K., and Hawthorne, F.C. (1989a) Euchroite, a heteropolyhedral framework structure. Acta Crystallographica, C45, 1479–1482.

- Groat, L.A., and Hawthorne, F.C. (1987) Refinement of the crystal structure of papagoite. Tschermaks Mineralogische und Petrographische Mitteilungen, 87–96.
- Hawthorne, F.C. (1983) Graphical enumeration of polyhedral clusters. Acta Crystallographica, A39, 724-736.
- ——(1985a) The crystal structure of botallackite. Mineralogical Magazine, 49, 85–91.
- (1985b) The crystal structure of stringhamite. Tschermaks Mineralogische und Petrographische Mitteilungen, 34, 15–34.
- (1985c) Towards a structural classification of minerals: The ${}^{vi}M^{iv}T_2\Phi n$ minerals. American Mineralogist, 70, 455–473.
- (1986) Lammerite, a modulated close-packed structure. American Mineralogist, 71, 206–209.
- Hawthorne, F.C., and Eby, R.K. (1985) Refinement of the structure of lindgrenite. Neues Jahrbuch der Mineralogie Monatschefte, 234--240.
- Hawthorne, F.C., and Groat, L.A. (1985) The crystal structure of wroewolfeite, a mineral with [Cu₄(SO₄)(OH)₆(H₂O)] sheets. American Mineralogist. 70, 1050–1055.
- (1986) The crystal structure and chemical composition of cumengeite. Mineralogical Magazine, 50, 157–162.

Hawthorne, F.C., Groat, L.A., and Eby, R.K. (1989) Antlerite, a modu-

lated heteropolyhedral framework structure. Canadian Mineralogist, 27, 205-209.

- International tables for X-ray crystallography, vol. 3 (1974). The Kynoch Press, Birmingham, England.
- Kniep, R., and Mootz, D. (1973) Metavariscite A redetermination of its crystal structure. Acta Crystallographica, B29, 2292–2294.
- Kniep, R., Mootz, D., and Vegas, A. (1977) Variscite. Acta Crystallographica, B33, 263-265.
- Mrose, M.E., Meyrowitz, Robert, Alfors, J.T., and Chesterman, C.W. (1966) Nissonite, CuMg(PO₄)(OH)·2.5H₂O, a new hydrous copper magnesium phosphate mineral from the Panoche Valley, California (abs.). Geological Society of America Programs, Annual Meeting 1966, 145–146.
- Sheldrick, G.M. (1981) Nicolet SHELXTL operations manual. Nicolet XRD Corporation, Madison, Wisconsin.
- Sutor, D.J. (1967) The crystal and molecular structure of newberyite. Acta Crystallographica, 23, 418–422.

Manuscript received February 21, 1989 Manuscript accepted July 9, 1990