The crystal structure of sinkankasite, a complex heteropolyhedral sheet mineral

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ABSTRACT

The crystal structure of sinkankasite, Z = 2, $Mn^{2+}(H_2O)_4[Al(PO_3OH)_2(OH)](H_2O)_2$, triclinic, a = 9.590(2), b = 9.818(2), c = 6.860(1) Å, $\alpha = 108.04(3)$, $\beta = 99.63(3)$, $\gamma = 98.87(3)^\circ$, V = 590.7(2) Å³, space group $P\overline{1}$, has been solved by direct methods and refined to an R index of 7.2% and an R_w index of 7.6% using 2170 unique observed reflections $||F_o|| \ge 5\sigma_F$) collected with MoK α X-ray radiation. The structure contains one distinct Mn^{2+} position, which is octahedrally coordinated by two O atoms and four H₂O groups. There are two distinct Al positions; each is octahedrally coordinated by four O atoms and two OH groups. The two distinct P positions are each tetrahedrally coordinated by three O atoms and one OH group. There are 15 unique H positions, which have been located and refined using soft constraints.

The structure consists of corner-sharing chains of $Al\phi_6$ octahedra, decorated on both sides by $P\phi_4$ tetrahedra. These heteropolyhedral chains are linked through the $Mn\phi_6$ octahedra to form continuous sheets of composition $MnAlP_2O_6(OH)_3 \cdot 4H_2O$. There are two interlayer anion positions, both of which are occupied by H_2O groups. There are eight unique H bonds between the interlayer cations and the sheets on either side, and no other bonding occurs between the sheets. A total of seven unique H bonds occur between anions of the same heteropolyhedral sheet.

INTRODUCTION

Sinkankasite was described as $H_2MnAl(PO_4)_2(OH)$. 6 H_2O by Peacor et al. (1984). It occurs in the Barker pegmatite, South Dakota, as an alteration product of triphylite, where it is associated with other secondary phosphates, including vivianite, hureaulite, carbonateapatite, strengite, barbosalite, and fluellite. Sinkankasite is multiply twinned by reflection on {100}. However, because sinkankasite is triclinic, there is only random overlap of reflections in reciprocal space (except for those occurring on the twin plane, i.e., {0kl} reflections). Hence single-crystal intensity data can be collected from one component of the twinned crystal if the cell of this component can be oriented. We have determined the crystal structure of sinkankasite in this fashion and report the results here.

EXPERIMENTAL METHODS

A small tabular crystal fragment of sinkankasite from the Barker pegmatite, near Keystone, Pennington County, South Dakota, was mounted on a Nicolet R3m automated four-circle diffractometer. Twenty-one reflections over the range $4^{\circ} \le 2\theta \le 18^{\circ}$ were centered using graphite-monochromated Mo $K\alpha$ X-radiation. The unitcell dimensions (Table 1) were derived from the setting angles of the 21 automatically aligned reflections by leastscan method with a 2.3° 2θ scan range and a variable scan rate ranging from 4 to 29.3° 20/min. A total of 3650 reflections was measured; the index ranges $0 \le h \le 13$, $-13 \le k \le 13$, and $-9 \le l \le 9$ were covered. Two standard reflections were measured every 50 reflections. There was a gradual decline in their intensities by several percent during data collection; we attribute this decline to crystal deterioration, and the data were scaled based on the standards. We recorded intensity data on a standard ruby crystal immediately after the sinkankasite experiment. There was no significant change in the standard reflections of the ruby crystal during data collection, supporting our assumption that the sinkankasite crystal deteriorated during exposure to the X-ray beam. An empirical absorption correction based on 36 ψ -scans for each of nine reflections over the range $8^\circ \le 2\theta \le 54^\circ$ was applied, reducing R (azimuthal) from 10.7 to 4.9%. This agreement index is higher than we normally encounter and is a result of the twinned nature of the crystal (all current absorption-correction routines assume that the data are collected from a single crystal). The data were corrected for absorption, Lorentz, polarization, and background effects; of the 3650 reflections measured, 2170 were classed as observed ($F \ge 5\sigma_F$).

squares refinement. Data were collected using the θ -2 θ

STRUCTURE SOLUTION AND REFINEMENT

Scattering curves for neutral atoms, together with anomalous-dispersion corrections, were taken from

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TABLE 1. Crystallographic data for sinkankasite

a (Å)	9.590(2)	Crystal size (mm)	0.04 × 0.24 × 0.12
b (Å)	9.818(2)	Radiation	ΜοΚα
c (Å)	6.860(1)	Total ref.	3650
α (°)	108.04(3)	$ F \geq 5\sigma_F$	2170
β (°)	99.63(3)	Final R	7.2%
γ (°)	98.87(3)	Final R.	7.6%
V (Å3)	590.7(2)		
F ₀₀₀	406	D _{calc}	2.243 g/cm ³
Space group	PĪ	μ	1.49 mm ⁻¹
Unit-cell conte	nts 2{Mn ²⁺ (H	20)4[AI(PO30H)2(0H)]	(H ₂ O) ₂ }
$R = \Sigma \left(F_o - \right)$			
$R_{w} = [\Sigma w (F_{c}$		F_0^2 ^{1/2} , w = 1/ σ_F^2	

Cromer and Mann (1968) and Cromer and Liberman (1970), respectively. The Siemens SHELXTL PLUS (PC version) system of programs was used throughout this study. R indices are of the form given in Table 1 and are expressed as percentages.

The possible space groups of sinkankasite are P1 and $P\overline{1}$; reflection statistics favor the latter and the successful solution and refinement of the structure verifies that the space group of sinkankasite is $P\overline{1}$. The structure was solved by direct methods, followed by cycles of least-squares refinement and calculation of difference-Fourier maps. Refinement of the positional and isotropic displacement

parameters of the complete model (excluding H positions) gave an R index of 13%. Conversion to an anisotropic displacement model, together with refinement of all parameters, gave an R index of 7.5%. At this stage of the refinement, a difference-Fourier map was calculated, and the positions of 15 H atoms were determined. Subsequent cycles of refinement showed that the H positions were characterized by anomalously short O-H bond lengths, a common feature of H positions refined using X-ray data. The soft constraint that O-H distances should be ~ 0.96 Å was imposed by adding extra weighted observational equations to the least-squares matrix. Only the O-H distance is constrained, and each H position is free to seek its optimum position around the O atom. Refinement of all parameters, including a correction for isotropic extinction, gave a final R index of 7.2% and a $R_{\rm w}$ index of 7.6%. This relatively high final R index results from the twinned nature of the crystal and its influence on the absorption correction and from the deterioration of the crystal under the X-ray beam. The omission of 0kl reflections did not significantly improve the refinement. Final positional and displacement parameters are given in Table 2, observed and calculated structure factors in Table 3, selected interatomic distances and angles in Table 4, and a bond-valence analysis is given in Table 5.1

TABLE 2. Atomic coordinates and displacement factors for sinkankasite

	x	У	z	$U_{\rm eq}$ (Å ²)	U_{11}	U ₂₂	U ₃₃	U12	U ₁₃	U ₂₃
Mn	0.2887(2)	0.9520(2)	0.6821(2)	145(5)	240(9)	184(7)	39(5)	51(6)	54(5)	65(4)
P1	0.2971(3)	0.6566(2)	0.2606(3)	116(7)	186(12)	161(10)	41(8)	59(9)	41(7)	73(7)
P2	0.2914(3)	0.2566(2)	0.1038(3)	112(7)	187(12)	132(10)	38(8)	19(8)	40(7)	58(7)
AI1	1/2	1/2	1/2	82(12)	149(21)	123(18)	1(14)	30(13)	26(11)	56(11)
AI2	1/2	1/2	0	93(12)	190(21)	135(17)	1(14)	50(13)	49(12)	65(11)
01	0.2899(8)	0.8156(7)	0.3614(8)	178(21)	328(39)	192(27)	72(22)	113(27)	80(23)	79(20)
02	0.3909(7)	0.6426(6)	0.1003(7)	160(19)	312(36)	227(28)	23(20)	104(26)	97(21)	105(19)
03	0.3476(7)	0.5873(6)	0.4219(7)	131(18)	225(32)	202(27)	22(19)	63(24)	49(20)	98(18)
04	0.3857(7)	0.3101(6)	0.3259(8)	149(18)	261(34)	156(25)	41(20)	3(23)	29(21)	72(18)
05	0.3458(7)	0.3474(6)	-0.0230(8)	163(19)	293(36)	183(26)	48(21)	21(25)	39(22)	102(19)
06	0.2786(8)	0.0938(6)	0.9937(8)	170(20)	295(37)	172(27)	76(22)	53(26)	66(23)	77(19)
OH1	0.1363(8)	0.5740(8)	0.1438(10)	253(23)	211(37)	303(34)	240(30)	39(29)	-1(27)	125(26)
OH2	0.5789(6)	0.5165(6)	0.2761(7)	110(17)	190(30)	176(24)	1(18)	39(22)	36(19)	74(17)
OH3	0.1351(8)	0.2826(7)	0.1225(9)	205(21)	250(37)	224(29)	128(24)	-5(26)	52(23)	63(21)
OW1	0.2967(9)	0.7677(8)	0.7887(9)	270(26)	505(51)	274(34)	117(26)	155(34)	128(29)	121(24)
OW2	0.5186(9)	0.0040(9)	0.7596(12)	282(26)	255(41)	356(37)	354(35)	58(32)	71(31)	291(30)
OW3	0.2930(9)	0.1405(7)	0.5798(8)	229(23)	397(44)	232(31)	58(23)	23(30)	51(26)	71(21)
OW4	0.0573(6)	0.8887(9)	0.5972(12)	310(27)	291(44)	354(40)	308(35)	81(34)	121(32)	115(30)
OW5	0.0834(8)	0.9186(8)	0.1365(10)	242(23)	256(38)	307(34)	194(29)	53(29)	26(26)	145(25)
OW6	-0.0907(14)	0.5932(12)	0.4146(18)	613(46)	659(78)	542(60)	620(62)	-24(55)	222(56)	212(48)
H1	0.106(13)	0.470(2)	0.073(17)	300*		. ,	. ,	. ,	. ,	
H2	0.678(4)	0.547(12)	0.351(17)	300						
H3	0.063(9)	0.193(7)	0.055(17)	300						
H4	0.379(8)	0.737(14)	0.740(18)	300						
H5	0.290(15)	0.667(4)	0.710(16)	300						
H6	0.591(9)	0.071(10)	0.732(18)	300						
H7	0.580(11)	-0.041(12)	0.834(16)	300						
H8	0.397(1)	0.160(12)	0.613(18)	300						
Н9	0.288(13)	0.233(7)	0.676(14)	300						
H10	0.015(12)	0.785(3)	0.545(17)	300						
H11	0.000(11)	0.947(11)	0.675(16)	300						
H12	0.139(12)	0.858(11)	0.186(16)	300						
H13	0.081(1)	0.998(9)	0.259(10)	300						
H14	-0.148(12)	0.527(9)	0.465(17)	300						
H15	-0.067(14)	0.510(9)	0.319(14)	300						

Note: displacement factors are multiplied by 104.

* Fixed during refinement.

TABLE 4.	sinkankasite	ic distances (A)	and angles () in
Mn-O1 Mn-OW1 Mn-OW4 Mn-O6° Mn-OW2° Mn-OW3°	2.195(5) 2.158(8) 2.137(8) 2.188(6) 2.120(8) 2.170(8)	P2-O4 P2-O5 P2-OH3 P2-O6 ^b (P2-O)	1.521(5) 1.521(7) 1.578(8) 1.516(6) 1.534
(Mn-O) P1-O1 P1-O2	2.170(0) 2.161 1.517(6) 1.521(7)	Al1-O3° Al1-O4° Al1-OH2° (Al1-O)	1.889(7) × 2 1.912(5) × 2 1.861(6) × 2 1.887
P1-03 P1-0H1 (P1-0)	1.520(7) 1.573(7) 1.533	AI2-O2 ^d AI2-O5 ^d AI2-OH2 ^d (AI-O)	1.904(7) × 2 1.884(6) × 2 1.865(5) × 2 1.884
	Mnø _s o	octahedron	
01-0W1 01-0W2 01-0W2 01-0W3 0W1-0W4 0W1-06 ^a 0W4-06 ^a 0W4-0W3 ^a 06 ^a -0W2 ^a 06 ^a -0W3 06 ^a -0W3 (0-0)	3.052(8) 2.129(11) 3.104(11) 3.035(9) * 3.004(12) 3.057	01-Mn-OW1 01-Mn-OW4 01-Mn-OW3 ^a 0W1-Mn-OW3 ^a 0W1-Mn-OW4 0W1-Mn-OW2 ^a 0W4-Mn-O6 ^a 0W4-Mn-O8 ^a 0W4-Mn-OW3 ^a 06 ^a -Mn-OW3 ^a 0W2 ^a -Mn-OW3 ^a 0W2 ^a -Mn-OW3 ^a 0W2 ^a -Mn-OW3 ^a	90.8(2) 88.2(3) 90.0(3) 89.3(2) 89.9(3) 91.7(2) 88.1(3) 89.8(3) 93.2(3) 92.2(3) 88.3(2) 88.3(2) 88.9(3) 90.0
01-02	2.497(9)	etrahedron 01-P1-02	110.6(4)
01-03 01-0H1 02-03 02-0H1 03-0H1 (0-0)	2.527(10) 2.453(8) 2.513(8) 2.523(10) 2.496(9) 2.501	01-P1-03 01-P1-0H1 02-P1-03 02-P1-0H1 03-P1-0H1 (0-P1-0)	112.6(3) 105.1(4) 111.5(4) 109.2(3) 107.6(4) 109.4
	P2ø₄ t	etrahedron	
04-05 04-0H3 04-06 ^b 05-0H3 05-06 ^b 0H3-06 ^b (0-0)	2.508(8) 2.498(9) 2.510(6) 2.481(10) 2.517(9) 2.508(10) 2.504	04-P2-05 04-P2-0H3 04-P2-06° 05-P2-0H3 05-P2-06° 0H3-P2-06° (0-P2-0)	111.1(3) 107.4(3) 111.5(4) 106.4(4) 111.9(3) 108.3(4) 109.4
		octahedron	
03°-04° 03°-0H2° 03°-04° 03°-0H2° 04°-0H2° 04°-0H2° (0-0)	2.693(9) × 2 2.681(9) × 2 2.682(8) × 2 2.621(8) × 2 2.621(8) × 2 2.672(9) × 2 2.665(6) × 2 2.669	03°-A11-04° 03°-A11-0H2° 03°-A11-04° 03°-A11-0H2° 04°-A11-0H2° 04°-A11-0H2° (0-A11-0)	90.2(2) × 2 91.3(3) × 2 89.8(2) × 2 88.7(3) × 2 90.1(2) × 2 89.9(2) × 2 90.0
001 051		octheedron	00.0(0)
02ª-05ª 02ª-0H2ª 02ª-05ª 02ª-0H2ª 05ª-0H2ª 05ª-0H2ª (0-0)	$\begin{array}{c} 2.697(8) \times 2 \\ 2.669(9) \times 2 \\ 2.659(10) \times 2 \\ 2.652(7) \times 2 \\ 2.679(7) \times 2 \\ 2.623(9) \times 2 \\ 2.665 \end{array}$	O2 ^d -Al2-O5 ^d O2 ^d -Al2-OH2 ^d O2 ^d -Al2-OH2 ^d O2 ^d -Al2-OH2 ^d O5 ^d -Al2-OH2 ^d O5 ^d -Al2-OH2 ^d (O-Al2-O)	$\begin{array}{l} 90.8(3) \times 2 \\ 90.2(3) \times 2 \\ 89.2(3) \times 2 \\ 89.8(2) \times 2 \\ 91.2(2) \times 2 \\ 88.8(2) \times 2 \\ 90.0 \end{array}$
			0.02/7
H1-OH1 OH1-OH3 H2-OH2 OH2-OW6° H3-OH3 OH3-OW5' H4-OW1 OW1-O4° H5-OW1	0.96(7)	H1OH3 OH1-H1-OH3 H2OW6° OH2-H2-OW6° H3OW5' OH3-H3-OW5' H4O4° OW1-H4-O4° H5O3	2.02(7) 139(9) 2.14(4) 161(6) 1.71(7) 153(9) 2.46(9) 170(10) 2.09(12)

TABLE 4. Selected interatomic distances (Å) and angles (°) in

TABLE	4	—Continue	d

.761(8) .95(14) .96(11) .760(11) .96(12) .837(11) .47(18) .96(2) .379(8)	OW1-H5-O3 H4-OW1-H5 H6O1° OW2-H6-O1° H7O6° OW2-H7-O6° H6-OW2-H7 H8O2°	126(9) 60(8) 1.81(11) 171(8) 1.89(12) 170(10) 100(9) 0.61(7)
9.96(11) 2.760(11) 9.96(12) 2.837(11) .47(18) 9.96(2)	H6O1° OW2-H6-O1° H7O69 OW2-H7-O69 H6-OW2-H7	1.81(11) 171(8) 1.89(12) 170(10) 100(9)
2.760(11) 0.96(12) 2.837(11) .47(18) 0.96(2)	OW2-H6-O1° H7O6º OW2-H7-O6º H6-OW2-H7	171(8) 1.89(12) 170(10) 100(9)
0.96(12) 2.837(11) .47(18) 0.96(2)	H7····O69 OW2-H7-O69 H6-OW2-H7	1.89(12) 170(10) 100(9)
2.837(11) .47(18) 0.96(2)	OW2-H7-O69 H6-OW2-H7	170(10) 100(9)
.47(18) .96(2)	H6-OW2-H7	100(9)
.96(2)		
	H8O2°	0 61(7)
3.379(8)		2.61(7)
	OW3-H8-O2°	137(9)
.46(8)	H905 ⁿ	1.95(8)
.757(7)	OW3-H9-O5"	140(9)
.41(14)	H8-OW3-H9	95(9)
.96(4)	H10OW6	1.86(4)
2.807(13)	OW4-H10-OW6	170(3)
).96(11)	H11···OW5	1.91(11)
2.866(11)	OW4-H11-OW5	170(10)
.61(11)	H10-OW4-H11	114(9)
.96(12)	H12···01	1.90(12)
2.808(11)	OW5-H12-O1	156(9)
).96(8)	H13 OW4	2.04(12)
2.866(11)	OW5-H13-OW4	142(10)
.54(16)	H12-OW5-H13	106(9)
).97(12)	H14…O3	2.43(12)
3.359(15)	OW6-H14-O3	161(9)
).96(9)	H15···OH1	2.55(13)
3.080(16)	OW6-H15-OH1	115(6)
.35(18)	H14-OW6-H15	89(9)
	.41(14) .96(4) .807(13) .96(11) .866(11) .61(11) .96(12) .808(11) .96(8) .866(11) .54(16) .97(12) .359(15) .96(9) .080(16) .35(18) + 1, z; b = x, y,	.41(14) H8-OW3-H9 .96(4) H10···OW6 .807(13) OW4-H10-OW6 .96(11) H11···OW5 .866(11) OW4-H11-OW5 .61(11) H10···OW4-H11 .96(12) H12···O1 .808(11) OW5-H12-O1 .96(8) H13···OW4 .866(11) OW5-H13-OW4 .866(11) OW5-H13-OW4 .54(16) H12-OW5-H13 .97(12) H14···O3' .359(15) OW6-H14-O3' .96(9) H15···OH1 .080(16) OW6-H15-OH1

DESCRIPTION OF THE STRUCTURE

$MnO_2(H_2O)_4$ and $AlO_4(OH)_2$ octahedra

The sinkankasite structure contains one symmetrically unique Mn position that is octahedrally coordinated by two O atoms and four H₂O groups. The H₂O groups are at the equatorial positions of the octahedron, which has a (Mn-O) distance of 2.161 Å and bond lengths ranging from 2.120 to 2.195 Å. Peacor et al. (1984) reported chemical analyses for sinkankasite from the Barker pegmatite. The material they studied contains significant Fe2+; the ratio of Mn^{2+} to Fe^{2+} at this site is about 1.0 to 0.6. The Mn²⁺ and Fe²⁺ species have very similar X-ray scattering factors, and it is not possible to refine their ratios directly using X-ray data. However, the sums of the empirical radii of Shannon (1976) for 161Mn2+, 161Fe2+, and ^[3]O²⁻ are 2.19 and 2.14 Å for the Mn²⁺-O and Fe²⁺-O bonds, respectively. The (Mn-O) bond length of 2.161 Å reported here is consistent with the occurrence of both Mn²⁺ and Fe²⁺ at this site.

The structure has two symmetrically distinct Al positions, and both occur on a center of symmetry. Each Al is octahedrally coordinated by four O atoms and two OH groups, and the latter are in a trans arrangement. The \langle Al1-O \rangle and \langle Al2-O \rangle distances are 1.887 and 1.884 Å, respectively, fairly close to the sum of the ionic radii for ^[6]Al, ^[3]O, and ^[2]OH (0.535 + 1.353 = 1.888 Å: Shannon,

¹ A copy of Table 3 may be ordered as Document AM-95-588 from the Business Office, Mineralogical Society of America, 1130 Seventeenth Street NW, Suite 330, Washington, DC 20036, U.S.A. Please remit \$5.00 in advance for the microfiche.

 TABLE 5.
 Bond-valence analysis (vu) for sinkankasite

	01	02	O3	04	O5	06	OH1	OH2	OH3	OW1	OW2	OW3	OW4	OW5	OW6	Σ
Mn	0.335					0.341				0.370	0.410	0.358	0.391			2.205
AI1			0.526ײ↓	0.494ײ↓				0.567×2↓								3.174
Al2		0.505×2↓			0.533ײ↓			0.561×2 į								3.198
P1	1.310	1.296	1.300				1.126									5.032
P2				1.296	1.296	1.314			1.111							5.017
H1							0.85		0.15							1.0
H2 H3								0.90							0.10	1.0
H4				0.05					0.80					0.20		1.0
H5			0.12	0.05						0.95						1.0
H6	0.20		0.12							0.88	0.00					1.0
H7	0.20					0.20					0.80					1.0
H8		0.05				0.20					0.80	0.95				1.0
H9		0.00			0.15							0.95				1.0 1.0
H10					0.10							0.05	0.80		0.20	1.0
H11													0.80	0.20	0.20	1.0
H12	0.20												0.00	0.80		1.0
H13													0.15	0.85		1.0
H14			0.05											0.00	0.95	1.0
H15							0.05								0.95	1.0
Σ	2.045	1.851	1.996	1.840	1.979	1.855	2.026	2.028	2.061	2.200	2.010	2.158	2.141	2.050	2.200	
	_	ce paramete				_	2.020	2.020	2.001	2.200	2.010	2.100	2.141	2.000	2.200	

1976). Thus, both the site scattering and the observed $\langle Al-O \rangle$ distances are consistent with complete occupancy of Al1 and Al2 by Al, as indicated by the electron-micro-probe analysis of Peacor et al. (1984).

PO₃OH tetrahedra

There are two symmetrically distinct P positions in sinkankasite. Each P is tetrahedrally coordinated by three O atoms and an OH group, thus forming acid phosphate groups. The $\langle P1-O \rangle$ and $\langle P2-O \rangle$ distances are 1.533 and 1.534 Å, respectively. The tetrahedra show significant bond-length distortion, with bond lengths ranging from 1.517 to 1.573 Å and 1.516 to 1.578 Å in the P1 ϕ_4 ($\phi =$ unspecified ligand) and P2 ϕ_4 tetrahedra, respectively; in each case, the longest bond is to the OH group, as is usually observed for acid phosphate groups because of local bond-valence requirements around the acid ligand (Table 5).

Structure topology

The Al ϕ_6 octahedra share one set of trans vertices to form an Al ϕ_5 chain extending along the c direction. Equatorial vertices of each chain are linked by P ϕ_4 tetrahedra (Fig. 1a) to form a heteropolyhedral chain of stoichiometry Al(PO₃OH)₂(OH). These chains are cross-linked into sheets of the form Mn²⁺(H₂O)₄[Al(PO₃OH)₂(OH)] by linkage of the trans vertices of the Mn²⁺ ϕ_6 octahedra with vertices of P ϕ_4 tetrahedra of adjacent chains (Fig. 2). These sheets lie in the (100) plane and are cross-linked into a three-dimensional structure by a network of H bonds (Fig. 3); this arrangement accounts for the perfect {100} cleavage of sinkankasite (Peacor et al., 1984).

H bonding

The H atom positions obtained by constrained leastsquares refinement of the positional parameters are reasonable, both in terms of donor and acceptor bond lengths and bond angles (Table 4) and the bond-valence requirements of the donor and acceptor anions (Table 5). There are 15 symmetrically unique H positions in sinkankasite, and H bonding plays two distinct roles in the structure: (1) H bonds bridge between the heteropolyhedral sheets and the interlayer anions, providing the only linkage be-

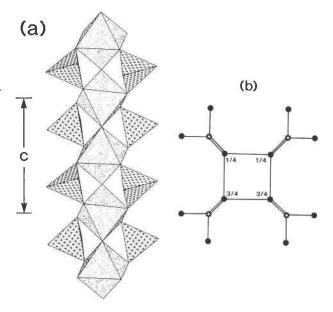


Fig. 1. (a) The Al(PO₃OH)₂(OH) chain in sinkankasite; Al ϕ_6 octahedra are shown by irregular dot shading; P ϕ_4 tetrahedra are shaded with crosses; (b) the type VI chain of Moore (1970), showing the octahedral chain viewed down its length, with the relative heights of the tetrahedra shown as fractions of the chain-axis length of ~7 Å.

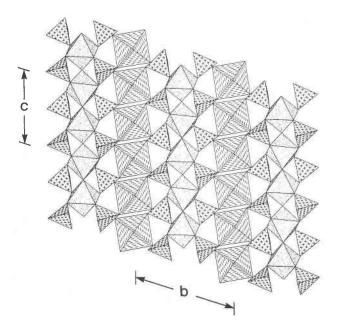


Fig. 2. The $Mn^{2+}(H_2O)_4[Al(PO_3OH)_2(OH)]$ sheet in the structure of sinkankasite; $Mn\phi_6$ octahedra are shaded with dashed lines; other polyhedral shadings are as in Fig. 1.

tween sheets, and (2) H bonding occurs within the heteropolyhedral sheets.

There are eight H positions that provide linkage between the heteropolyhedral sheets and the interlayer OW5 and OW6 positions (Fig. 4). These H positions participate in the following bonds: $OH2-H2\cdots OW6^{e}$, $OH3-H3\cdots$ $OW5^{f}$, $OW4-H10\cdots OW6$, $OW4-H11\cdots OW5^{i}$, OW5- $H12\cdots O1$, $OW5-H13\cdots OW4^{i}$, $OW6-H14\cdots O3^{j}$, and $OW6-H15\cdots OH1$. Acceptor distances range from 1.71 Å (strong) to 2.55 Å (weak).

Each of the OW5 and OW6 anions bond to four H atoms. The OW5 anion donates the OW5-H12...O1 and OW5-H13...OW4ⁱ bonds, one of which bonds to the heteropolyhedral sheet on either side, with acceptor distances of 1.90 and 2.04 Å, respectively. The OW5 position also accepts the OH3-H3...OW5^f and OW4-H11...OW5ⁱ bonds, with acceptor distance of 1.71 and 1.91 Å, respectively, and each of these bonds is to the same heteropolyhedral sheet (Fig. 4).

The OW6 anion donates the OW6-H14 \cdots O3^j and OW6-H15 \cdots OH1 bonds, one of which bridges to the heteropolyhedral sheet on the other side, with acceptor distances of 2.43 and 2.55 Å, respectively, which indicate weak interactions only. The OW6 anion accepts a H bond from each adjacent sheet; the fairly strong OH2-H2 \cdots OW6^c and OW4-H10 \cdots OW6 bonds have acceptor distances of 2.14 and 1.86 Å, respectively.

Seven H positions participate in H bonds between two anions of the same heteropolyhedral sheet (Fig. 5). These bonds are as follows: $OH1-H1\cdots OH3$, $OW1-H4\cdots O4^{\circ}$, $OW1-H5\cdots O3$, $OW2-H6\cdots O1^{\circ}$, $OW2-H7\cdots O6^{\circ}$, OW3- $H8\cdots O2^{\circ}$, and $OW3-H9\cdots O5^{h}$, with acceptor distances

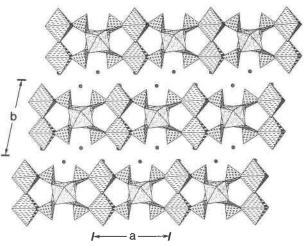


Fig. 3. The structure of sinkankasite projected down [001]; shadings as in Fig. 2; interstitial H_2O groups shown by shaded circles.

ranging from 1.81 to 2.61 Å. These H bonds increase the sheet rigidity and satisfy the bond-valence requirements of the donor and acceptor anions.

General structural relations

The principal structure motif in sinkankasite is the Al(PO₃OH)₂OH chain that we can generalize to the form $M(T\phi_4)_2\phi$, where M is an octahedrally coordinated cation and T is a tetrahedrally coordinated cation. In a paper of fundamental importance, Moore (1970) considered the structural hierarchy of octahedral-tetrahedral chains based on an $M\phi_5$ backbone. He developed a hierarchy of all possible arrangements with a chain of repeat distance of ~7.2 Å and showed that there is a fairly large family of minerals based on these chains.

Sinkankasite has the type VI chain (Fig. 1b) of Moore (1970), designated as the tancoite chain by Hawthorne (1985). This chain is found in a number of structures (Table 6), all of which show the typical ~7.2-Å repeat (6.9-7.3 Å), except for tancoite, which has a doubled repeat distance of 14.07 Å. Sinkankasite shows most affinity with the segelerite-group structures. The latter are treated in some detail by Moore and Araki (1977), who show the relationship of segelerite to jahnsite in terms of cross linkage of dense X²⁺M³⁺(PO₄)₂OH slabs by $M^{2+}O_2(H_2O)_4$ octahedra. The solution of the sinkankasite structure raises another aspect of geometrical isomerism (Hawthorne, 1983a) with regard to these structures. Sinkankasite consists of slabs constructed from $M(T\phi_4)_2\phi$ chains that are cross-linked by MO2(H2O)4 octahedra, and there is ample opportunity for geometrical isomerism in this general type of arrangement. The possibilities are investigated in Figure 6 for repeat distances of one (~9 Å) or two (~18 Å) chains in the direction of linkage. As is apparent from Figure 1, the $M\phi_5$ chain is modulated such

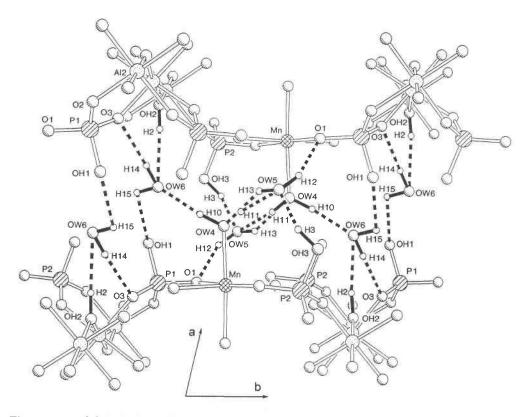


Fig. 4. The structure of sinkankasite projected down [001], showing intersheet H bonds. Al atoms are stippled, Mn atoms are crosshatched, P atoms are parallel ruled, O atoms are shaded in the lower left corner, and H atoms are given as small open circles. Donor H and acceptor H bonds are drawn as heavy solid and broken lines, respectively.

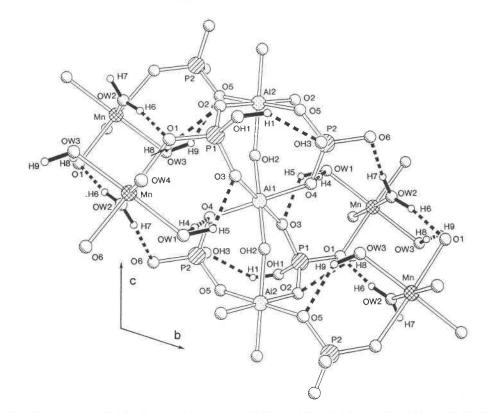


Fig. 5. The structure of sinkankasite projected down [100], showing intrasheet H bonds; legend as in Fig. 4.

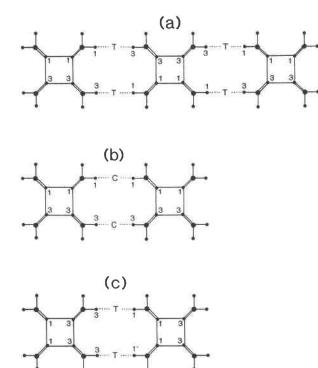
Mineral	Formula	a (Å)	b (Å)	c (Å)	β (°)	Sp. Gr.	Refer- ence
Tancoite	Na2Li[AI(PO4)2(OH)]H	6.948(2)	14.089(4)	14.065(3)		Pbcb	1
Sideronatrite*	Na ₂ [Fe ³⁺ (SO ₄) ₂ (OH)]·3H ₂ O	7.29(1)	20.56(2)	7.17(2)		Pnn2	2
Jahnsite	CaMnMg ₂ [Fe ³⁺ (PO ₄) ₂ (OH)] ₂ ·8H ₂ O	14.94(2)	7.14(1)	9.93(1)	110.16(8)	P2/a	3
Whiteite	CaFe ²⁺ Mg ₂ [Al(PO ₄) ₂ (OH)] ₂ -8H ₂ O	14.90(4)	6.98(2)	10.13(2)	113.11(9)	P2/a	4
Lun'okite	Mn ₂ (Mg,Fe ²⁺ ,Mn) ₂ [Al(PO ₄) ₂ (OH)] ₂ ·8H ₂ O	14.95	18.71	6.96		Pbca	_
Overite	Ca2Mg2[AI(PO4)2(OH)]2 8H2O	14.72(1)	18.75(2)	7.107(4)		Pbca	5
Segelerite	Ca2Mg2[Fe3+(PO4)2(OH)]2.8H2O	14.826(5)	18.751(4)	7.307(1)		Pbca	5
Wilhemvierlingite	Ca2Mn2[Fe3+(PO4)2(OH)]2 ·8H2O	14.80(5)	18.50(5)	7.31(2)		Pbca	
Sinkankasite	Mn ²⁺ (H ₂ O),[A!(PO ₃ OH) ₂ (OH)](H ₂ O) ₂	9.590(2)	9.818(2)	6.860(1)	99.63(3)	PT	6
Guildite	Cu ²⁺ [Fe ³⁺ (SO ₄) ₂ (OH)] 4H ₂ O	9.786(2)	7.134(1)	7.263(1)	105.38(1)	$P2_1/m$	7
Yftisite	Y ₄ [Ti(SiO ₄) ₂ O](F,OH) ₆	14.949(4)	10.626(2)	7.043(2)		Cmcm	8

TABLE 6. Minerals based on the type VI $M(T\phi_4)_2\phi$ chain of Fig. 1

Note: references are as follows: 1 = Hawthorne (1983b); 2 = Scordari (1981); 3 = Moore and Araki (1974); 4 = Moore and Ito (1978); 5 = Moore and Araki (1977); 6 = this study; 7 = Wan et al. (1978); 8 = Balko and Bakakin (1975). * Structure not definitively established.

that tetrahedra can link to vertices of adjacent octahedra. This modulation can occur orthogonal to or in the plane of the sheet formed by the $M(T\phi_4)_2\phi$ chains and their linking $MO_2(H_2O)_4$ octahedra. It is this difference in modulation orientation that combines with the cis and trans possibilities of the linking $MO_2(H_2O)_4$ octahedra to give the possible arrangements of Figure 6. We may divide them into three groups: (1) those arrangements with

the modulation orthogonal to the sheet and only one type of octahedral cross-linkage (Fig. 6a and 6b), where Figure 6a corresponds to the arrangement in sinkankasite, (2) those arrangements with the modulation in the plane of the sheet (Fig. 6c and 6d), corresponding to the arrangements in the hypothetical structure in Figure 4c of Moore and Araki (1977) and segelerite, respectively, and (3) those arrangements with both modulations present (Fig. 6e and



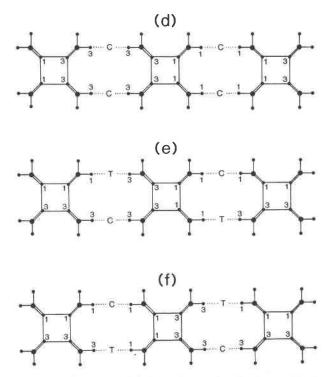


Fig. 6. Possible topologies involving $M(T\phi_4)_2\phi$ chains (viewed end-on and represented as squares) cross-linked by $M\phi_6$ octahedra with the linkages in cis (M_c) or trans (M_1) arrangements, with the modulation direction of the $M(T\phi_4)_2\phi$ chains taken into account; the heights of the tetrahedral linkages are given by symbols 1 (=¼) and 3 (=¾), which are the approximate z coordinates of the T cations for a chain-repeat length of ~7 Å and M cations

at z = 0 and $\frac{1}{2}$; (a) modulation orthogonal to the sheet and only trans linkages (sinkankasite); (b) modulation orthogonal to the sheet and only cis linkages (not yet known); (c) modulations in the plane of the sheet and only trans linkages (hypothetical structure of Moore and Araki, 1977); (d) modulation in the plane of the sheet and only cis linkages (segelerite, overite); (e), (f) mixed modulations and mixed linkages (not yet known).

6f) (these have not yet been found in crystals). Thus we have a coherent hierarchy of arrangements, two of which have so far been found in structural arrangements of minerals.

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References cited

- Balko, V.P., and Bakakin, V.V. (1975) The crystal structure of the natural yttrium and rare-earth fluortitano-silicate $(Y,Tr)_4(F,OH)_6TiO(SiO_4)_2$ (Yftisite). Zhurnal Strukturnoi Khimii, 16, 837–842.
- Brown, I.D., and Altermatt, D. (1985) Bond-valence parameters obtained from a systematic analysis of the inorganic crystal structure database. Acta Crystallographica, B41, 244–247.
- Cromer, D.T., and Liberman, D. (1970) Relativistic calculation of anomalous scattering factors for X rays. Journal of Chemical Physics, 53, 1891-1898.
- Cromer, D.T., and Mann, J.B. (1968) X-ray scattering factors computed from numerical Hartree-Fock wave functions. Acta Crystallographica, A24, 321-324.
- Hawthorne, F.C. (1983a) Graphical enumeration of polyhedral clusters. Acta Crystallographica, A39, 724–736.
- (1983b) The crystal structure of tancoite. Tschermaks mineralogisch-petrographische Mitteilungen, 31, 121–135.

(1985) Towards a structural classification of minerals: the ${}^{v_1}M^{iv}T_2\phi_n$ minerals. American Mineralogist, 70, 455–473.

Moore, P.B. (1970) Structural hierarchies among minerals containing octahedrally coordinating oxygen: I. Stereoisomerism among corner-sharing octahedral and tetrahedral chains. Neues Jahrbuch für Mineralogie Monatshefte, 163–173.

Moore, P.B., and Araki, T. (1974) Jahnsite, CaMn²⁺Mg₂(H₂O)_s-Fe₂³⁺(OH)₂[PO₄]₄: A novel stereoisomerism of ligands about octahedral corner-chains. American Mineralogist, 59, 964–973.

- —— (1977) Overite, segelerite, and jahnsite: A study in combinational polymorphism. American Mineralogist, 62, 692–702.
- Moore, P.B., and Ito, J. (1978) Whiteite, a new species, and a proposed nomenclature for the jahnsite-whiteite complex series. Mineralogical Magazine, 42, 309-316.
- Peacor, D.R., Dunn, P.J., Roberts, W.L., Campbell, T.J., and Simmons, W.B. (1984) Sinkankasite, a new phosphate from the Barker pegmatite, South Dakota, American Mineralogist, 69, 380–382.
- Scordari, F. (1981) Sideronatrite: A mineral with a {Fe₂(SO₄)₄(OH)₂} guildite type chain. Tschermaks mineralogisch-petrographische Mitteilungen, 28, 315–319.
- Shannon, R.D. (1976) Revised effective ionic radii and systematic studies of interatomic distances in halides and chalcogenides. Acta Crystallographica, A32, 751–767.
- Wan, C., Ghose, S., and Rossman, G.R. (1978) Guildite, a layer structure with a ferric hydroxy-sulphate chain and its optical absorption spectra. American Mineralogist, 63, 478–483.

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