Chemical bonding and electronic structures of the Al$_2$SiO$_5$ polymorphs, andalusite, sillimanite, and kyanite: X-ray photoelectron- and electron energy loss spectroscopy studies

FUMIO S. OHUCHI,1 SUBRATA GHOSE,2 MARK H. ENGELHARD,3 AND DONALD R. BAER3

1Department of Materials Science and Engineering, Box 352120, University of Washington, Seattle, Washington 98195, U.S.A.
2Department of Earth and Space Science, Box 351310, University of Washington, Seattle, Washington 98195, U.S.A.
3Pacific Northwest National Laboratory, Environmental Molecular Science Laboratory, Richland, Washington 99352, U.S.A.

ABSTRACT

We have undertaken a detailed analysis of the X-ray photoelectron spectra obtained from the three polymorphs of Al$_2$SiO$_5$: andalusite, sillimanite, and kyanite. Comparison of the spectra was made based on the chemical bonding and structural differences in the Al- and Si-coordination within each polymorph. The spectra for Si(2p) for all three polymorphs are nearly identical, consistent with the fact that all the Si atoms are in 4-fold (tetrahedral) coordination, whereas the binding energies, peak shapes, and peak widths for Al(2p) vary depending on the type of polymorph. The upper-valence band for all three polymorphs is characterized by four main features derived from O(2p), Al(3s), Al(2p), Si(3s), and Si(3p), and the differences in their contributions are observed. The density of state of the Al$_2$SiO$_5$ polymorphs is relatively featureless compared to those observed from α-SiO$_2$ and α-Al$_2$O$_3$, suggesting that the orbital overlaps span a greater range in energy. The observed band gap energy for Al$_2$SiO$_5$ (sillimanite) was ~9.1eV, a value in between those for α-SiO$_2$ (~8.6eV) and α-Al$_2$O$_3$ (~9.6eV). The conduction band feature of Al$_2$SiO$_5$ was experimentally compared to those of α-SiO$_2$ and α-Al$_2$O$_3$, and shown that it is indeed intermediate between the α-SiO$_2$ and α-Al$_2$O$_3$ phases.

**Keywords:** Al$_2$SiO$_5$ polymorphs, X-ray photoelectron spectroscopy, low electron energy loss spectroscopy, valence- and conduction band structures, andalusite, sillimanite, kyanite

INTRODUCTION

The three polymorphs of Al$_2$SiO$_5$, sillimanite, andalusite, and kyanite, are geologically important minerals, whose crystal structures, and physical and thermodynamic properties have been extensively investigated (Kerrick 1990 and references therein). Crystal structures of sillimanite, andalusite, and kyanite were refined by Burnham (1963a), Burnham and Buerger (1961), and Burnham (1963b), respectively, following the structure determinations of kyanite by St. Naray Szabo et al. (1929), and sillimanite and andalusite by Taylor and Hey (1931). Further refinements of the sillimanite and andalusite structures were carried out using single-crystal neutron diffraction methods (Finger and Prince 1972). Winter and Ghose (1979) determined the thermal expansion and the structural changes of all three polymorphs as a function of temperature up to 1000 °C. The crystal structures of all three polymorphs have also been refined at high pressure [andalusite, Ralph et al. (1984); kyanite, Yang et al. (1997a) and Comodi et al. (1997); and sillimanite, Yang et al. (1997b)].

Sillimanite and andalusite are orthorhombic with space groups Pnma and Pbam, respectively, whereas kyanite is triclinic with P$ar{T}$. The crystal structures of the three polymorphs have one feature in common, namely, half of the Al atoms occur in sixfold (octahedral) coordination forming chains of edge-shared [AlO$_6$] octahedra parallel to the crystallographic c-axis. The differences in their stability relations depend critically on the differences in the chemical bonding of the remaining Al atoms in each polymorph: Al is in fourfold (tetrahedral) coordination in sillimanite (stable at low P and high T), fivefold coordination (trigonal bipyramidal) in andalusite (stable at low P and high T), and sixfold coordination (octahedral) in kyanite (stable at high P and high T). The Si atom is in fourfold (tetrahedral) coordination in all three polymorphs. The reader is referred to the paper by Winter and Ghose (1979) for the details of the crystal structures of all three polymorphs.

Lattice dynamical studies of these minerals, including experimental measurements of phonon dispersion relations, density of states (DOS), and theoretical shell model calculations of their vibrational and thermodynamic properties, have been made by Rao et al. (1999). The electron density distributions in all three polymorphs have been determined experimentally from multipole refinements based on high resolution single crystal X-ray diffraction data collected at 100 K by a CCD detector (Dahaoui et al. 2001), and theoretically by ab initio quantum mechanical calculations using the LAPW method (Iglesias et al. 2001). The electric field gradient tensors at the 27Al sites in all three polymorphs calculated from the experimentally and theoretically derived electron density distributions match the values obtained from single-crystal 27Al NMR measurements (Raymond and Hafner 1970; Hafner et al. 1970; Hafner and Raymond 1967; Bryant et al. 1999).

The differences in the chemical bonding and electronic structures of the three Al$_2$SiO$_5$ polymorphs are fundamental to the understanding of their thermodynamic properties and
Al$_2$SiO$_5$ is available. Detailed analyses of photoelectron spectra of the chemical bonding for the three different polymorphs of $\alpha$ESCA microprobe (Physical Electronics Inc.) at Paci Laboratory Inc.) at the University of Washington, and a Quantum 2000 Scanning two different spectrometers: a modified SSL-300 spectrometer (Surface Science Laboratory Inc.) at the University of Washington, and a Quantum 2000 Scanning ESCA microprobe (Physical Electronics Inc.) at Pacific Northwest National Laboratories, Richland, Washington. In the first spectrometer, a focused monochromatic AlK$_\alpha$ X-rays with an X-ray spot of ~600 $\mu$m in diameter was used to measure core level data. In the second spectrometer, a monochromatic AlK$_\alpha$ X-ray beam (~100 $\mu$m diameter) was rastered over an area of ~1.4 mm by ~0.2 mm on the specimen to collect valence band data. During analysis, charge neutralization was achieved by using a low energy (3–4 eV) electron flood gun. In the Quantum 2000 spectrometer, a low energy Ar-ion beam was applied in addition to the electron flood gun. All spectra were acquired using a pass energy of 50 eV with the SSL-300, and 23.5 eV with the Quantum 2000 spectrometer. The binding energies were obtained from alumino-silicate glass have been reported (Yagi et al. 2001; Miura et al. 2000). In glass form, however, Al and Si polyhedra are all corner linked, therefore the present study from three different polymorphs can give further insight into the chemistry and electronic structure of bonding about edge- and face-linked polyhedrons. We have undertaken an X-ray photoelectron spectroscopic study of the three Al$_2$SiO$_5$ polymorphs to probe the differences in their chemical bonding, and compared the electronic structure of Al$_2$SiO$_5$ (sillimanite) with those of $\alpha$-Si$_2$O$_5$ (quartz) and $\alpha$-Al$_2$O$_3$ (corundum) using proper charge compensation techniques. Low electron energy loss spectroscopy (LEELS) was also used to study the conduction band structure of the sillimanite phase of Al$_2$SiO$_5$.

EXPERIMENTAL METHODS

Single-crystal samples of andalusite from Espiritu Santo, Brazil, sillimanite from Sri Lanka, and kyanite from Pizzo Forno, Switzerland were obtained from the Harvard Mineralogical Museum. Sample surfaces examined by XPS and LEELS were prepared by mechanically fracturing the samples in air, and mounted on a sample holder with indium solder. A typical duration for the fractured surface exposed in air (prior to evacuation in the sample load-lock chamber) was less than 30 min. The X-ray photoelectron spectroscopy experiments were performed using two different spectrometers: a modified SSL-300 spectrometer (Surface Science Laboratory Inc.) at the University of Washington, and a Quantum 2000 Scanning ESCA microprobe (Physical Electronics Inc.) at Pacific Northwest National Laboratories, Richland, Washington. In the first spectrometer, a focused monochromatic AlK$_\alpha$ X-rays with an X-ray spot of ~600 $\mu$m in diameter was used to measure core level data. In the second spectrometer, a monochromatic AlK$_\alpha$ X-ray beam (~100 $\mu$m diameter) was rastered over an area of ~1.4 mm by ~0.2 mm on the specimen to collect valence band data. During analysis, charge neutralization was achieved by using a low energy (3–4 eV) electron flood gun. In the Quantum 2000 spectrometer, a low energy Ar-ion beam was applied in addition to the electron flood gun. All spectra were acquired using a pass energy of 50 eV with the SSL-300, and 23.5 eV with the Quantum 2000 spectrometer. The binding energies were nominally calibrated using clean Cu and Au foils at Cu 2p$_{3/2}$ = 932.72 eV; Cu 3p$_{3/2}$ = 75.10, and Au 4f$_{7/2}$ = 83.98 eV. In the analysis, the absolute value of the binding energy was calibrated by setting the C(1s) peak for the adventitious hydrocarbons accumulated on the surface at 284.60 eV.

For LEELS, the spectra were excited by unpolarized electrons with energies ~250 eV incident normal to the surface, and were measured with a single pass cylindrical mirror analyzer (CAM). The current density of the incident electron beam was minimized to prevent electron beam damage of the sample surfaces. A pulse counting mode was used to collect the data. To obtain a constant system resolution, the full width at half maximum (FWHM) of the incident beam was set at a value of about 0.7 eV. For LEELS measurements, only a sillimanite sample was examined, and the results are compared to those from $\alpha$-Al$_2$O$_3$ (corundum) and $\alpha$-Si$_2$O$_5$ (quartz) phases. The binding energies of O(1s) found in $\alpha$-Al$_2$O$_3$ and $\alpha$-Si$_2$O$_5$ are 531.2 eV and 532.5 eV, respectively, which are similar to those reported previously (531.2 eV for $\alpha$-Al$_2$O$_3$ and 532.8 eV for $\alpha$-Si$_2$O$_5$) (Barr 1994).

RESULTS AND DISCUSSION

Si(2p) core-level spectra

The Si atoms in all three polymorphs are tetrahedrally bonded to O atoms, and form Al-O-Si linkages. The Si(2p) core level spectra were carefully measured from sillimanite, andalusite, and kyanite, and compared with that measured from $\alpha$-Si$_2$O$_5$ (Fig. 1). Analyses indicate that all the spectra reasonably fall into a single Gaussian peak with similar binding energy (E_b) at around 102.28 eV (see Table 1), suggesting that the electronic environment around Si atoms in all three polymorphs of Al$_2$SiO$_5$ and $\alpha$-quartz ($\alpha$-Si$_2$O$_5$) is very similar, and that the second nearest-neighbor Al atoms do not significantly influence the electronic structure of the tetrahedrally-bonded Si atoms.

Al(2p) core-level spectra

Figure 2 shows the Al(2p) spectra obtained from the three polymorphs. In kyanite, there are four symmetry-independent Al atoms present, all in octahedral coordination. The Al-O bond lengths vary within narrow limits (1.816 to 1.997 Å) (Winter and Ghose 1979); therefore, the Al(2p) peak is expected to be reasonably symmetric with a relatively narrow width. Curve fitting to a single Gaussian shape results in a peak maximum at 74.48 ± 0.05 eV with FWHM of 1.68 eV (Fig. 2). Since the Al atomic arrangement in $\alpha$-Al$_2$O$_3$ is also characterized by octahedral coordination with the Al-O bond lengths similar to those found in kyanite, Al(2p) spectra from kyanite and $\alpha$-Al$_2$O$_3$ may be compared (Fig. 3). As shown in Fig. 3, the two Al(2p) peaks are nearly identical in terms of the relative binding energy. Note, however, that the Al(2p) peak from kyanite is slightly skewed in its low E_b tail, whereas a nearly perfect Gaussian peak was observed for $\alpha$-Al$_2$O$_3$. This small difference may be due to the influence of the second nearest-neighbor tetrahedral Si atoms in kyanite, and/or greater variation of the Al-O bond distances in kyanite (1.816 to 1.997 Å; Winter and Ghose 1979) vs. $\alpha$-Al$_2$O$_3$ (1.857 Å and 1.969 Å for short and long Al-O bond lengths, respectively; Chiang and Xu 1994).

The Al(2p) peak for sillimanite appears at a similar position in binding energy; however, the peak is considerably broader, and
the distribution extends toward the lower \( E_b \) side, suggesting two kinds of Al atoms: one similar to that in kyanite (more ionically bonded) and another less ionically bonded (or more covalently bonded). Based on the crystal structure of sillimanite (Winter and Ghose 1979), the Al atoms are in octahedral and tetrahedral coordination with equal fractions. We therefore fitted the curve to two peaks with their ratio constrained to 1:1 (Fig. 2c), resulting in \( E_b \) (FWHM) values 74.65 (1.70) and 74.06 (1.71). The analysis indicates that ~50% of the octahedrally coordinated Al atoms in sillimanite are slightly more ionic than those found in kyanite based on their binding energy values (75.65 eV for sillimanite vs. 74.48 eV for kyanite), whereas the other ~50% in tetrahedral coordination are less ionic (\( E_b \sim 74.06 \)).

In the andalusite structure, the octahedral Al-coordination is severely distorted: four closely bonded O atoms within a square planar configuration with two each at 1.827 and 1.891 Å, and two farther apical O atoms at 2.086 Å (298 K) (Winter and Ghose 1979). The other Al atom in fivefold coordination is closely bonded to two O atoms at 1.814 Å and one at 1.816 Å forming a trigonal pyramid, while two further O atoms at 1.839 and 1.899 Å complete the trigonal bipyramid (298 K). In view of the structural complexity and hence, the complex nature of the chemical bonding, it does not seem justified to resolve the Al(2p) peak into contributions from the two types of Al atoms in distorted octahedral and the fivefold trigonal bipyramidal coordination. Therefore, we have fitted the Al(2p) spectra to a single peak. A reasonably good fit was found with a peak maximum at 74.06 eV with FWHM of 1.71 eV, as shown in Figure 2. These values are very similar to that for the tetrahedral Al atom in sillimanite (\( E_b \) 74.06) (Table 1). Furthermore, the overall Al(2p) peak, shifted toward the lower binding energy by ~0.4 eV compared to kyanite (\( E_b \) 74.48), suggests that the bonding for the Al atoms in andalusite is less ionic, compared to the octahedral Al atoms in either kyanite or sillimanite.

### Table 1. Binding energy (\( E_b \) in eV), FWHM (eV) and relative fraction (%) for Si(2p), Al(2p), and O(1s) orbitals observed for the three polymorphs of Al\(_2\)SiO\(_5\)

<table>
<thead>
<tr>
<th>Material</th>
<th>Si(2p) ( E_b ) (eV)</th>
<th>FWHM (eV)</th>
<th>Relative fraction %</th>
<th>Al(2p) ( E_b ) (eV)</th>
<th>FWHM (eV)</th>
<th>Relative fraction %</th>
<th>O(1s) ( E_b ) (eV)</th>
<th>FWHM (eV)</th>
<th>Relative fraction %</th>
</tr>
</thead>
<tbody>
<tr>
<td>Kyanite</td>
<td>102.27</td>
<td>1.81</td>
<td>100</td>
<td>74.48</td>
<td>1.68</td>
<td>100</td>
<td>531.39</td>
<td>1.97</td>
<td>80</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>530.59</td>
<td>1.98</td>
<td>20</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Andalusite</td>
<td>102.31</td>
<td>1.80</td>
<td>100</td>
<td>74.07</td>
<td>1.70</td>
<td>100</td>
<td>531.69</td>
<td>1.98</td>
<td>20</td>
</tr>
<tr>
<td>Sillimanite</td>
<td>102.28</td>
<td>1.91</td>
<td>100</td>
<td>74.65</td>
<td>1.70</td>
<td>50</td>
<td>532.02</td>
<td>1.94</td>
<td>40</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>530.85</td>
<td>1.92</td>
<td>60</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>( \alpha )-SiO(_2)</td>
<td>102.28</td>
<td>1.92</td>
<td>100</td>
<td>74.48</td>
<td>1.78</td>
<td>100</td>
<td>531.21</td>
<td>1.98</td>
<td>100</td>
</tr>
<tr>
<td>( \alpha )-Al(_2)O(_3)</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

**Figure 2.** Overlay of Al(2p) spectra observed from three polymorphs, and the results from curve fitting. See text for details.

**Figure 3.** Comparison of Al(2p) for kyanite and \( \alpha \)-Al\(_2\)O\(_3\).
oxygen atom may not vary significantly, we do not expect to be able to resolve all these binding energy differences; however, the overall shapes of the O(1s) peaks observed in the three polymorphs are different.

In kyanite, of the 20 O atoms in the unit cell, 16 are involved in Si-O-Al linkage, and 4 in Al-O-Al type linkages. The O atoms in the latter type are expected to be slightly more ionic than the former. We have therefore fitted the spectra with two Gaussian peaks with the ratio of 4:1, and obtained O(1s) peaks corresponding to Si-O-Al linkage (80%) and Al-O-Al linkage (20%) at 531.39 eV and 530.59 eV, respectively (see Table 1). Although the full-width at half-maximum (FWHM) of O(1s) peak ranges from 1.92 to 1.98 eV based on the Gaussian curve fitting, the separation of the peak maxima for the two-types of O(1s) is large enough to distinguish one peak from another.

The bonding differences in the various O atoms in sillimanite and andalusite can be analyzed in terms of the total charge each receives from the surrounding cations. In sillimanite, 12 O atoms are either charge balanced or slightly over bonded, and 8 are slightly under bonded. In andalusite, 4 are under bonded and 16 are either charge balanced or slightly over bonded. Following these considerations, the curve fittings for sillimanite and andalusite were made in 4:1, 2:3, and 1:4 ratios, respectively. Curve fitted results are shown for each polymorph in Figure 4 and Table 1. Our analysis indicates that 60% of the total O atoms in sillimanite (Eb ~530.85) are more ionic compared to the other 40% (Eb ~532.02). This is consistent with the result obtained from Al(2p), in which one half of the total Al atoms is similar to those in kyanite, which is more ionic.

In andalusite, a major peak (>80% of the total) is centered around 530.84 eV, whereas <20% of the oxygen peak appeared with a peak maximum at 531.69 eV, suggesting that the majority of the O atoms are more ionic than those tetrahedrally bonded to Al in sillimanite (532.02). However, this result appears to be inconsistent with the Al(2p) value (74.07), which is comparable to that of tetrahedral Al in sillimanite (74.06). We currently do not have an explanation for this anomaly, and can only suggest that the explanation may lie in the structural complexity and hence, the complex nature of the chemical bonding of andalusite compared to that of kyanite and sillimanite.

Valence-level spectra

Valence-level X-ray photoemission spectra are useful to characterize the nature of the chemical bonding, and may be directly compared with the electron density of state (DOS) calculations. In materials like Al₂SiO₅, an upper valence band (upper-VB) is characterized by a superposition of Al-Si-O hybridized orbitals and non-bonding O(2p) orbitals, whereas a lower valence band (lower-VB) is mainly due to O(2s) orbitals.

The upper-VB spectra obtained from the three different polymorphs of Al₂SiO₅ are shown in Figure 5. The valence band width is very similar for all three phases, around ~11.0 eV, and the top of the valence band positions are all ~3.9eV below the Fermi energy. Within the VB spectra, however, noticeable differences are observed from each polymorph, which are characterized by four main features labeled a, b, c, and d. The locations of these features in terms of Eb are similar, but the relative contributions vary for different polymorphs, obviously reflecting the types of bonding and coordination for Al, Si, and O atoms in the three polymorphs.

So far, the only DOS calculations available in the literature are those of Iglesias et al. (2001). These authors claimed that the cross-features of the DOS were similar for all atoms in all the three polymorphs. The total upper-VB DOS consists of mainly of O(2p), Al(3s), Al(2p), Si(3s), and Si(3p), in which the Si states appear at the lower energies, and the s-p splitting is more pronounced than for the corresponding Al states. All five orbitals are strongly mixed over the entire upper-VB regions. Note that Iglesias et al. (2001) assumed the top of the valence band as the Fermi-energy for the calculated partial density of states (PDOS), whereas the experimentally observed valence band starts around ~4 eV in binding energy. In addition, the calculated binding energies are somewhat smaller than those observed experimentally, although the overall nature of the calculated and experimentally observed valence band structures is very similar.

Based on these calculations, the feature labeled a (around 5-6 eV in Fig. 5) consists of mainly O(2p) contribution with little overlap from the Si and Al orbitals. Feature b around 9eV is the Si(2p and 3p) and Al(2p and 3s) overlapped with O(2p), whereas the s-orbitals from both Al and Si are not strongly involved. Feature c at around 11 eV is similar to b, but has more contribution from the Al(3s) orbitals; therefore, it varies for the three different polymorphs. In our measurements, this was seen...
we were unable to clearly identify a shift based on the leading of increased ionicity in some materials (Barr 1991). However, in the valence band toward the Fermi energy is a direct measure of this. Nevertheless, this difference in terms of binding energies, but appear ~1 eV higher than that of Al2SiO5. For both α-SiO2 and α-Al2O3, the upper-VB consists of mainly two features; a peak located at lower E_b side (labeled as e or g), and a peak at higher E_b side (labeled as f or h). Based on the DOS calculations, the features, e and g, are predominantly O(p)-non-bonding states with little contribution from O(s). These non-bonding states are spatially localized on the atomic sites and oriented transversally with respect to the Si-O and Al-O bonding direction for the quartz and corundum structures, respectively. The features at higher E_b side, f and h, mainly consist of strongly hybridized orbitals. The calculations indicate that these two features are well separated in both quartz and corundum, reflecting two distinct features found in calculated DOS (Iglesias et al. 2001). Observed DOS calculations for Al2SiO5 are relatively featureless compared to those observed for α-SiO2 and α-Al2O3, suggesting that the orbital overlaps are much more spread out over the energy range, as previously described. It should also be pointed out that the valence band edges from the three materials occur nearly at the same binding energy (3.80 ± 0.10 eV).

The lower-VB spectra observed from the three materials are also shown in Figure 6. The DOS calculations reveal that the lower valence band is mainly O(2s) with small overlap from 2p-states associated with Si and/or Al. Thus, the differences in bonding and coordination are reflected in the peak locations and their shapes. The lower-VB spectra observed for sillimanite are similar to those of α-SiO2 in terms of binding energies, but appear ~1 eV higher than that of α-Al2O3. Compared with O(1s) peaks, the lower VB peaks are considerably broader (~4 eV of FWHM), and therefore, detailed structures cannot be resolved. Nevertheless, this difference implies that the bonding character for sillimanite is more covalent than that of α-Al2O3, but similar to that of α-SiO2.

Low energy electron loss spectroscopy (LEELS) is a technique for studying unoccupied states since the core electrons are excited into the unoccupied states in the conduction band (Koma and Ludeke 1975; Araki et al. 1976). Because the core states have well-defined energies and angular momenta, LEELS can probe the variation in the angular-momentum-resolved density of the conduction band state as joint density of states (JDOS). Shown in Figure 7 is LEELS spectra obtained from α-Al2O3, sillimanite, and α-SiO2; the spectra for α-Al2O3 and α-SiO2 are excited from the Al(2p) and Si(2p) core levels, respectively, and for sillimanite from both Al(2p) and Si(2p) core levels. The LEELS spectra are then aligned by setting the threshold above 0 eV (Fermi energy) to an amount equal to the loss energy minus the corresponding
core binding energy previously determined by XPS. The loss peak distribution above 0 eV represents JDOS of the conduction band. The onset energy of the JDOS corresponds to the energy difference between the bottom of the conduction band and the Fermi level of the material. The valence band edge occurs at 3.80 ± 0.10 eV for all three materials. Therefore, the band gap energies for α-Al2O3, sillimanite, and α-SiO2 are determined to be ~9.6 eV (3.8 + 5.8), ~9.1 eV (3.8 + 5.3), and ~8.6 eV (3.8 + 4.8), respectively. Published values for the band gap ranges from 5.67 to 11.5 eV for SiO2 (Young-nian et al. and references therein) and from ~7 to ~10 eV for Al2O3 (French et al. 1994; Gillet and Ealet 1992). Explanations for the range include the formation of excitons, presence of defects and temperature dependent Urbach tails. The band gap energy for Al2SiO5, sillimanite, is intermediate to those of α-Al2O3 and α-SiO2, which is reasonable based on the mixture of Si-O and Al-O bond configurations in these materials.

The JDOS spectra mainly consist of three primary regions labeled i (at ~4–8 eV), j (~10–15 eV), and k (~20–26 eV) above the Fermi energy. Differences in the relative height for each region are related to the transition probability between the core and unoccupied conduction levels. Although transitions that give rise to the core-loss edge are governed by the atomic-dipole selection rules for electronic transitions, the spin-orbit split for Al(2p) and Si(2p) cannot be resolved in the present case. The resultant JDOS deduced from the LEELS, therefore, consists of all s-, p-, and d-like structures of the unoccupied states of the material. The region i for the three materials is similar, but the peak from α-SiO2 is shifted toward higher energy. In the regions j and k for α-SiO2, is less pronounced compared to α-Al2O3. Sillimanite follows both features shown by α-SiO2 and α-Al2O3, implying that the conduction band for sillimanite is indeed intermediate to those of α-SiO2 and α-Al2O3 phases. Given the broad nature of the conduction band for sillimanite, it is expected that the conduction bands in andalusite and kyanite should be similar to that found in sillimanite.

**ACKNOWLEDGMENTS**

We are indebted to Carl Francis, Harvard Mineralogical Museum, Cambridge, Massachusetts for the single-crystal samples of andalusite (Embilipitiya, Brazil HMM no. 129874), sillimanite (Sri Lanka), and kyanite (Pizzo Forno, Switzerland). Parts of the work were conducted in the Environmental Molecular Sciences Laboratory a DOE User Facility operated by Battelle for the DOE Office of Biological and Environmental Research.

**REFERENCES CITED**


