## SHORT COMMUNICATIONS

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## On the occurrence of Chlorapatite at the Emerald Mine, Rajagarh village, Ajmer

THIS paper describes the mineralogy, chemistry, X-ray, and trace element study of a chlorapatite (asparagus stone) collected from the rocks exposed at the emerald mine, near the village Rajagarh, Ajmer. The mica schists and gneisses of Aravalli age (Archaean), constitute the country rocks in the area. The schists are associated with altered peridotites and intrusive granite and pegmatite. The asparagus stone occurs at the contact of schists and pegmatite. A few crystals of chlorapatite, which were first mistaken for emerald, were collected from the contact zone during the course of a geological excursion.

The crystals measured  $1\frac{1}{4}'' \times \frac{3}{4}''$ , yellowish-green, semi-transparent with uneven to sub-conchoidal fracture and vitreous lustre; the specific gravity 3.17.

A crystal was broken, crushed to -80 mesh, and separated to remove any associated gangue mineral, using Clerici's solution of varying dilutions. The mineral separated was almost pure (99 %) excepting the minute inclusions of sericite, which could not be separated. In order to estimate the amount of sericite, grain mounts of the separated mineral were studied using refractive index liquid of 1.645, with the help of a point counter. The sericite inclusions were estimated at 0.93 %.

Optical study: The mineral is colourless in thin section, with acicular and fluid inclusions(?); optically negative, it displays a pseudo-biaxial figure (2E 12°), probably due to strain. The refractive indices were determined by immersion method:  $\epsilon = 1.642$ ,  $\omega = 1.649$ , slightly higher than the values for common apatite, and comparing well with the values cited for chlorapatites by Deer *et al.* (1963). The birefringence is low, 0.008.

*X-ray study*: The material was scanned in a diffractometer using Cu-K radiation from  $0^{\circ}$  to  $60^{\circ}$  at  $1^{\circ}$ /minute; the *d*-values and intensities compare well with the Powder Diffraction File values for chlorapatite (Table I).

Chemistry: CaO and  $P_2O_5$  were determined by gravimetry; fluorine and chlorine by volumetry; alkalies by flame photometry and the rest by atomic-absorption spectrophotometry. Water was determined, following a suggestion by Dr. M. H. Hey, by igniting the mineral with about 4 times its weight of freshly ignited CaO, at about 900 °C, and determining the loss in weight. The analysis compares fairly well with the published analysis of chlorapatite, with slight variations: it contains less Mn, F, and Cl, and more CaO. The analysis and atomic ratios are given in Table II.

### SHORT COMMUNICATIONS

Trace element study: The several trace elements determined are presented in Table II. The elements V, Yt, Zr, and Ce were determined by emission spectroscopy and the results are semiquantitative. The trace element study was undertaken to decide

hkil  $d_{\text{cale}}$  $d_{\rm meas}$ I hkil  $d_{\text{cale}}$ I  $d_{\text{meas}}$ 58 2242 1.956 1121 3.913 3.917 7 1.950 18 0002 3.420 3142 1.905 1.900 23 3.372 1.837 56 1012 3.100 10 3251 1.853 3.125 2131 2.850 2.850 85 4150 1.816 1.800 18 3030 2.774 2.768 100 4042 1.770 1.768 8 1.686 2022 2.620 2.643 25 0004 1.686 15 8 1.600 3031 2.565 2.540 4152 1.599 10 7 2.110

TABLE I. X-ray powder data for a chlorapatite (chlorian hydroxyapatite) fromAjmer, India. Indexed on a cell with a 9.598, c 6.739 Å, c/a 0.702

TABLE II. Chemical analysis and atomic ratios for a chlorapatite from Ajmer, India

		Trace elements in p.p.m.		Atomic ratios to 26 (O, OH, F, Cl)	
$\begin{array}{c} P_2O_5\\ CaO\\ MnO\\ MgO\\ FeO\\ Na_2O\\ F\\ Cl\\ H_2O^+\\ H_2O^-\\ Others^*\\ Sum\\ Less\\ O\equiv F, Cl\\ Net sum\\ \end{array}$	41.50 55.69 0.09 0.11 0.10 0.45 1.31 0.73 0.02 0.76 100.86 0.48	Si Al Mn Mg Sr Cu Ni V Zr Yt Ce	$2200 \\ 1300 \\ 700 \\ 665 \\ 284 \\ 5 \\ 5 \\ 10 \\ (> 500 \\ < 500 \\ < 500 \\ < 250 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 \\ < 500 $	$ \begin{array}{cccc} P & 6 \cdot 0 I \\ Ca & I0 \cdot 15 \\ Mn & 0 \cdot 01 \\ Mg & 0 \cdot 03 \\ Fe & 0 \cdot 01 \\ Na & 0 \cdot 03 \\ F & 0 \cdot 24 \\ Cl & 0 \cdot 38 \\ OH & 0 \cdot 85 \\ O & 24 \cdot 53 \end{array} I \cdot 47 $	

\* SiO<sub>2</sub>, Al<sub>2</sub>O<sub>3</sub>, and K<sub>2</sub>O present as sericite.

whether the mineral is igneous or metamorphic in parentage (Cruft, 1966). The trace element distribution compares very well with the trace elements of apatites from pegmatites and granitic pegmatites cited by Cruft (1966, p. 380) (samples 9, 10, 22 and 23). According to Cruft 'many apatite samples from granitic pegmatites contain high Mn, a high ratio of Y to Ce+La and a low total rare earths in contrast with samples from plutonic and igneous intrusives and metamorphic environments'. The high amount of manganese and the predominance of yttrium over cerium and the low

#### SHORT COMMUNICATIONS

content of strontium indicate that the present mineral belongs to the granitic pegmatite. Further, when the cerium and silicon (in ppm) are plotted on the diagram given by Cruft (1966, p. 393), the point falls in the field given for pegmatitic apatites.

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Dept. of Geology University of Saugar Sagar, Madhya Pradesh, India S. K. BABU

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# Helvine in the gondite of Kajlidongri, Madhya Pradesh, India

THE gondite wall-rock around the epigenetic veins of Kajlidongri, Jhabua District, Madhya Pradesh, India (22° 57' N., 74° 28' E.), contains dark brown well-developed dodecahedral grains filling fractures at the crest of antiforms, and ranging in diameter from I to 2.5 mm; occasionally they exhibit traces of poorly developed {III} cleavage. Chemical tests showed that they contain iron, manganese, zinc, and beryllium. Because of the very small size of the included quartz grains, it is extremely difficult to obtain a pure sample of the mineral for analysis; a partial chemical analysis of two samples estimated to contain about 10 % impurity by microscopic examination gave: MnO 29.56, 30.50, FeO 14.05, 14.67, ZnO 4.15, 4.49 %, with refractive index 1.74,  $1.76\pm0.001$  and cell-size a 8.16 and 8.18 Å. The only other sulphur-bearing mineral in and around the epigenetic veins in the area is baryte, forming fracture-filled veins in a gondite composed of porphyroclasts of quartz and recrystallized grains of spessartine and alurgite. The gondite in this area includes many textural types of different permeabilities, and the baryte is restricted to gondite wall-rocks of considerable porosity, while helvine is restricted to gondite with recrystallized quartz where the intergranular permeability is almost absent. Presumably a higher  $f_{0_n}$  is responsible

974