Subsolidus data for the join Ca_2SiO_4 — $CaMgSiO_4$ and the stability of merwinite.¹

By DELLA M. ROY, B.S., M.S., Ph.D.

College of Mineral Industries, The Pennsylvania State University, University Park, Pennsylvania, U.S.A.

[Communicated by Dr. M. H. Hey; taken as read 26 January 1956.]

Summary.—A study of phase equilibria among the orthosilicates on the Ca_2SiO_4 —CaMgSiO₄ join was undertaken with the use of hydrothermal techniques in the temperature range 600–900° C. at water pressures from 1000 to 15000 psi.² Monticellite, merwinite, and Ca_2SiO_4 crystallize as pure phases throughout this temperature range, and intermediate compositions contain mixtures of the end members. Powder X-ray diffraction data obtained on the mixtures indicate no change in lattice dimensions from the end members. Hence, very little or no solid solution is indicated at these temperatures. Ca_2SiO_4 was crystallized as the γ -form below 675° C. (at 2000 psi) and the β -form above this temperature (presumably having inverted from the α' -form on cooling to room temperature). Infra-red absorption spectra obtained on the compounds γ - Ca_2SiO_4 , β - Ca_2SiO_4 , merwinite, and monticellite show characteristic differences in the pattern within the 9–13 micron region. All the data thus appear to confirm merwinite as a unique compound stable in the temperature range studied.

O NE of the most elusive problems within the complex system CaO-MgO-SiO₂ has been the determination of the equilibria on the orthosilicate join Ca₂SiO₄-Mg₂SiO₄. On the basis of similarity in crystal structure, partial or complete solid solution might be expected between forsterite and monticellite, and similarly between monticellite and γ -Ca₂SiO₄, all of which possess the olivine structure. When O'Daniel (1942) assigned the olivine structure to γ -Ca₂SiO₄ he predicted that extensive solid solution should exist between γ -Ca₂SiO₄ and monticellite; and Bredig (1950) later suggested that this solid solution should be continuous. Recent work of Ricker and Osborn (1954) on the subsolidus region of the join CaMgSiO₄-Mg₂SiO₄ has shown that a solid solution series extends to a maximum of about 30 weight % from each compound at solidus temperatures, but the amount of solid solution decreases rapidly with decrease in temperature.

The presence of the intermediate compound merwinite (Ca₃MgSi₂O₈)

- ¹ Contribution no. 54-66, College of Mineral Industries.
- ² Pounds per square inch; this contraction is used throughout this paper.

must be explained, however, for it is a naturally occurring mineral (Larsen and Foshag, 1921), and has been shown by Phemister et al. (1942) and by Osborn (1943) to be a unique (incongruently melting) compound having a stability field in the system CaO-MgO-SiO₂. The presence of merwinite on the join Ca₂SiO₄-CaMgSiO₄ was explained by Bredig (1950) as a composition in the α' solid solution series, i.e. he considered merwinite to be merely an end point in a solid solution series extending from α' -Ca₂SiO₄¹ to the composition Ca₃MgSi₂O₈. Complete solid solution was presumed by Bredig to exist between the lower temperature polymorph (γ -Ca₂SiO₄) and monticellite. However, the possibility of a complete solid solution series between γ -Ca₂SiO₄ and monticellite, as well as the proposed α' -Ca₂SiO₄-merwinite solid solution series, was considered unlikely by Ricker and Osborn (1954) on the basis of theoretical considerations. A recent summary of the current knowledge on the chemistry and polymorphism of Ca₂SiO₄ is that of Nurse (1954), who believed that single-crystal X-ray data and powder X-ray data obtained at elevated temperatures favour the designation of monoclinic merwinite as a distinct compound, not a solid solution of orthorhombic a'-Ca₂SiO₄. Bredig (1954, see Nurse, 1954), in discussion of the above paper, suggested that merwinite (monoclinic, pseudoorthorhombic) may be considered to be a 'fifth polymorphous form of calcium orthosilicate, related to α' , and stabilized by large amounts of Mg,SiO₄', presumably formed on cooling from high temperatures.

The present study was undertaken with the application of hydrothermal techniques in order to aid crystallization and in the hope of attaining equilibrium at subsolidus temperatures. It was hoped to establish more positively the character of merwinite and its equilibrium relations with the other orthosilicates, as well as to obtain new information on the polymorphism of Ca_2SiO_4 .

Experimental procedure

Briefly, the experimental procedure was as follows: One or more samples contained in individual gold envelopes in a bomb were quickly brought up to the desired temperature. Water was pumped into the

¹ The complexity of the relations among polymorphic forms of Ca_2SiO_4 is, of course, involved in the equilibria; recent papers of Newman and Wells (1946), Van Valkenburg and McMurdie (1947), Trömel (1949), Trömel and Möller (1949), and Bredig (1950) give evidence of four polymorphic forms, from the highest to lowest temperature, respectively, hexagonal α , orthorhombic α' , monoclinic β , and the olivine γ . Bredig (1950) has further suggested that the β -modification is only metastable.

MERWINITE

bomb through pressure tubing to reach the desired water pressure. The bomb was then heated at this regulated pressure and temperature for a period varying from a few hours to several days. Most runs were made at low pressures, about 2000 psi, in an attempt to avoid entering an area of hydrate stability in the Ca_2SiO_4 -rich mixtures. After completion of the run the water pressure was released to zero and after $\frac{1}{2}$ to 1 minute the bomb was quenched in water. The samples were examined by X-ray diffraction and the petrographic microscope, and in a few cases infrared absorption spectra were obtained.

The hydrothermal equipment used was similar to that described by Roy, Roy, and Osborn (1950), and Roy and Osborn (1952). Stellite 'test-tube' bombs were used almost exclusively, and the water pressure was released just before quenching, as described above, in order to prevent hydrate formation during heating or cooling. A few 'dry' runs were made in a Pt-wound quenching furnace.

Starting materials used included mixed $CaCO_3$, $MgCO_3$, and silicic acid reacted at high temperatures, but most frequently a type of mixture which is much more intimate and almost 'amorphous to X-rays', prepared in the following manner: weighed amounts of $CaCO_3$ and MgO (or $MgCO_3$) were dissolved in dilute HNO₃ and added to an excess of absolute alcohol. Ethyl orthosilicate was measured out with a burette and dissolved in alcohol and the two solutions mixed, evaporated at 110° C., then heated to about 400° C, for a few hours, then to 700° to 900° C. overnight. Chemical analyses of the final products were not made, but compounds formed from stoichiometric ratios were found to be quite free from other phases when studied microscopically or by X-ray diffraction.¹

Results and discussion.

A summary of pertinent hydrothermal quenching data is given in table I, indicating phases present as detected by a combination of X-ray diffraction and optical data. The results are represented diagrammatically in fig. 1, which shows the phases present at the various temperatures in equilibrium with H_2O at a constant pressure of 2000 psi. The three orthosilicate compounds Ca_2SiO_4 , $Ca_3Mg(SiO_4)_2$, and $CaMgSiO_4$ crystallize as unique phases in the temperature range studied, 600–1000° C., and intermediate compositions show a mixture of monticellite and merwinite in the one case, and in the other a mixture of merwinite and Ca_2SiO_4 (the γ -form below 675° C. and β above).² Runs made at temperatures as low

¹ Mixtures of CaO or CaCO₃ and silica gel were found to be less satisfactory with respect to attainment of equilibrium.

² New information obtained on the polymorphism of Ca_2SiO_4 is the subject of a separate paper to be published soon. Briefly, the data obtained using 'amorphous' starting materials indicate that γ -Ca₂SiO₄ crystallizes below 675° C. (at 2000 psi) and that β forms above this temperature. The β -Ca₂SiO₄, however, is probably

TABLE I.

Summary of hydrothermal quenching data for the join Ca₂SiO₄---CaMgSiO₄.

The pressure was 2000 psi except for the runs at 475° C., for which it was 15 000 psi, and at 1472° and 1455° , for which it was atmospheric.

Abbreviations used: $X = Ca_2SiO_4$.hydrate, not completely identified. $C_2S = Ca_2SiO_4$; tr. = trace; Mer. = Merwinite; Mont. = Monticellite.

Mol. %	Temp.	Time	
Ca ₂ SiO ₄ .	° C.	hrs.	Phases present.
100	640	24	γ -C ₂ S+tr. X
100	662	48	$\gamma - C_2 S + \text{some } X$
100	670	96	$\gamma - C_2 S + \beta - C_2 S + \text{some } X$
100	685	48	$\beta - C_2 S + X$
100	705	24	Mostly β -C ₂ S
100	870	24	β -C ₂ S
100	1472	20	γ –C ₂ S on quench
75	655	96	Mer. + γ -C ₂ S + X
75	665	72	Mer. $+\gamma$ -C ₂ S + tr. X
75	685	48	$Mer. + \gamma - C_2S + \beta - C_2S$
75	743	24	Mer. $+\beta$ -C ₂ S
75	885	96	Mer. $+\beta$ -C ₂ S
50	475	168	Mont. + hydrate
50	625	96	Mer. + tr. Mont.?
50	645	24	Mer.
50	650	96	Mer.
50	744	120	Mer.
50	870	20	Mer.
50	1472	20	Mer.
25	655	96	Mont. + Mer.
25	750	24	Mont. + Mer.
25	850	72	Mont. + Mer.
0	475	168	Mont.
0	610	72	Mont.
0	705	24	Mont.
0	860	24	Mont.
0	1455	2	Mont.

as 475° C. at 15 000 psi indicated that monticellite is stable under these conditions. Merwinite was not formed at 475° C. and 15 000 psi from an 'amorphous' mixture of the merwinite composition, which gave a mixture of monticellite and a dicalcium silicate hydrate. This is not positive evidence, however, that merwinite is no longer stable under these conditions.

formed only metastably, through a fairly rapid inversion on cooling, from α' -Ca₂SiO₄. A further complication is the formation of a small amount of a limesilica hydrate below about 700° C., both in pure Ca₂SiO₄ mixtures and the 75 mol. % Ca₂SiO₄ mixture.

MERWINITE

Within the experimental error, the transition temperature from γ to α' Ca₂SiO₄ is unchanged in the 75 mol. % Ca₂SiO₄ mixtures from that for pure Ca₂SiO₄, i.e. $675\pm10^{\circ}$ C. at 2000 psi. The *d*-spacings of phases present in the intermediate mixture show no change (within the experimental error of $\pm 0.02^{\circ}$ in 2θ) from those of the pure components Ca₂SiO₄ and merwinite. There is similarly no change in the spacings of merwinite and monticellite found in the 25 mol. % Ca₂SiO₄ mixture from those of the pure end-members.

Measurements from typical X-ray diffraction patterns of the end member compounds and intermediate compositions are given in table II.

Most of these data were obtained with less than maximum precision in recording, but they demonstrate clearly that the intermediate compositions contain mixtures of the end-members. not an intermediate solid solution. More precise measurements obtained with slower speed of recording of a few selected reflections, which are diagnostic,¹ and in which there is no shift due to overlapping of peaks, demonstrate that very little, if any, solid solution exists among the various compounds and polymorphs.

Although the crystals formed, especially at low temperatures, are too small for accurate determination of optical properties, the differences in index of refrac-



FIG. 1. Subsolidus phase equilibrium diagram for the join Ca_2SiO_4 —CaMgSiO₄ at a water pressure of 2000 psi. Symbols used: $\bigcirc =$ merwinite, $\blacklozenge =$ monticellite, $\bigtriangledown = \gamma$ -Ca₂SiO₄, $\blacktriangle = \alpha'$ -Ca₂SiO₄ (inverted to β at room temperature), $\blacklozenge =$ hydrate.

tion are distinctive. The low index (c. 1.644) crystals of monticellite stand out clearly against the higher index (c. 1.700) merwinite crystals in the mixture with 25 mol. % Ca₂SiO₄, and the low index (c. 1.650) γ -Ca₂SiO₄

¹ Peaks at 2θ values (Cu-K α) of 50·10° and 30·405° for monticellite, at 34·35°, 32·60°, 32·02°, and 32·16° for β -Ca₂SiO₄, at 32·86° and 29·715° for γ -Ca₂SiO₄, and at 48·445°, 33·78°, and 33·43° for merwinite were unchanged within the experimental error of \pm 0·03° in 2θ in intermediate compositions showing mixtures of merwinite plus one of the other phases.

actometer		ticellite.	$I I_n$	1.0	i ei O	0.8 0	0.4	0.4	0.05	1.0	0.6	0.15	0.3	0.05	0.1	0.02	0.3	0.02	0.02	0.1	0.02	0.2	0.05	0.05	1	ł		1	1]	ł	ļ]			
angle diffr		Mon	d (Å.)	5.56	4.19	3.640	3.192	2.939	2.771	2.676	2.598	2.546	2.404	2.359	2.213	1.920	1.819	1-777	1.749	1.725	1.689	1.600	1.508	1.390	1	!	!	1	1]		1	1	l]	
leo High-	inite+	icellite.	I/I_a	0.05	0.02	0.1	0.05	0.4	0.05	0.1	0.05	0.05	0.05	0.25	0.1	0.2	0·1	1.0	6.0	0-4	0·3	0.1	0.05	0.05	0.1	0.05	0-1	0.05	0.05	0.3	0.2 2	0.05	0.05	0.05	0.05	0.05
l on Nore]	Merw	Mont	d (Å.)	5.55	5.12	4.20	3.898	3.645	3.326	3.199	3.169	3.052	2.997	2.940	2.873	2.755	2.714	2.686	2.679	2.656	2-592	2.549	2.517	2.465	2.405	2.317	2.216	2.173	2.029	1.912	1.882	1.781	1.752	1.724	1.602	1.595
measured		inite.	I/I_n	0.05	0.05	0.1	0.1	0.1	0.2	1·0	6.0	0.3	0.02	0.05	0.05	0.1	0.1	0.1	0.1	0.05	0.1	0·3	0.2	0.02	0.02	0.05	0.05	0.1	1	ļ	ļ	l	ļ]	ļ	
Spacings on.		Merw	d $(Å.)$	3.883	3.633	3.340	3.173	2.868	2.749	2.682	2.675	2.652	2.618	2.589	2.398	2.318	2.306	2.214	2.171	2.060	2.030	1.909	1.877	1.722	1.705	1.579	1.549	1.535	1	ļ		l]	l	ļ	
SiO4 join. Cu radiati	$+_{a}$ 0i	nife.	I/I_0	0.02	0.3	1.0	1.0	0.8	0.5	0.5	1.0	6.0	0.2	0.3	0.1	0.2	0.2	0.3	0.1	0.05	0.1	0.25	0.1	0.1	0.1	1.0	0.05	ļ	ļ	I	ļ]]	ļ	1	
⁴ , CaMg	β -Ca ₂ S	Merwi	d (Å.)	3.818	2.876	2.790	2.782	2.744	2.730	2.722	2.681	2.671	2.655	2.609	2.547	2.283	2.214	2.189	2.166	2.027	1.983	1.910	1.878	1.630	1.610	1.535	1-414	1	-	1	1	1	1		1	
the Ca ₂ SiC usin	04+	nite.	I/I_0	0.05	1.0	0-3	0.2	0.3	0.35	0.1	0.1	0.4	0.3	1.0	0.1	0.05	0·1	0.1	0.05	0.4	0.2	0.1	0.1	0.1	0.15	1	ļ	1	ł	ļ	1	ł	1	ļ	!	
hases on t	γ -Ca ₂ Si	Merwir	d (Å.)	5.62	4.33	3.824	3.345	3.043	3.010	2.990	2.780	2.750	2.726	2.685	2.660	2.522	2.321	2.212	2.032	1.911	1.879	1.809	1.757	1.689	1.535		ļ	ļ	1]	ļ		1		l	
data for J		i0 ₄ .	I/I_0	0.2	0.4	0.3	0.5	1.0	0.1	0.2	0.1	6.0	1.0	0.1	0.1	0.1	0.2	0·I	0.1	0.05	0·1	0.15	0.8	0·3	0.05	0.2	0.2]		ł	1		Į	-]	
iffraction		γ -Ca ₂ Si	d (Å.)	5.62	4.33	4.058	3.821	3.014	2.898	2.840	2.780	2.750	2.728	2.686	2.589	2.536	2.519	2.475	2.327	2.249	2.189	2.026	1.910	1.806	1.755	1.689	1.637	1	1	ļ	ļ	ļ		1		ļ
. X-ray d		i0 <u>.</u> .	I/I_0	0.05	0.05	0.1	0.05	0.1	1.0	6.0	6.0	0.4	0.4	0.1	0.05	0.2	0.2	0-4	0.1	0.05	0·1	0.2	0.05	0.02	0.05	0.05	0.1	0.02	0.02	1	l]	Į	!	ļ	
TABLE II		B-Ca ₂ S	d (Å.)	3.807	3.381	3.243	3.048	2.876	2.791	2.783	2.746	2.726	2.607	2.545	2.449	2.406	2.283	2.189	2.164	2.047	2.022	1.985	1.895	1.804	1.708	1.631	1.608	1.526	1.486					l	1]

192

D. M. ROY ON

MERWINITE

crystals are distinct from those of merwinite in the 75 mol. % Ca_2SiO_4 mixture treated below 675° C. The higher index β - Ca_2SiO_4 crystals, formed in the latter mixture just above 675° C., are not easily distinguished optically from those of merwinite, but the X-ray diffraction



FIG. 2. Infra-red absorption spectra. Monticellite and β Ca₂SiO₄ obtained using 0.4 % concentration in KBr pellet, merwinite 0.2 % in KBr; γ -Ca₂SiO₄ obtained from 'Nujol' mull on NaCl plates.

patterns of the two phases are readily identified in mixtures, as has been demonstrated above.

Another diagnostic property of these phases was the infra-red absorption spectrum. Spectra were obtained on the Perkin-Elmer model 21 double-beam instrument, using both the 'Nujol' mull technique and pressed KBr pellets.¹ Typical spectra of γ Ca₂SiO₄, β -Ca₂SiO₄, merwinite and monticellite are shown in fig. 2. The diagnostic portion is mainly within the 9.5-12.5 micron range, although spectra were obtained from 2 to 15 microns. The four phases are seen to be quite distinct in their infra-red absorption spectra, and even structurally similar monticellite and γ Ca₂SiO₄ exhibit differences.

¹ Standard techniques of sample preparation for crystalline materials include the suspension of a few milligrams of the sample in a mull with a high-boiling paraffin oil such as 'Nujol', smeared between polished flat NaCl plates. The paraffin oil has a few sharp characteristic absorption bands, but an alternative method employs KBr, which is completely infra-red transparent, as a suspending medium. A small concentration of sample (0.2.2%) is mixed with finely ground KBr and pressed under a hydrostatic load of about 100 000 psi into a flat disk.

D. M. ROY ON MERWINITE

Although the KBr method generally gives better and more easily reproducible spectra, the pattern shown for γ -Ca₂SiO₄ is that obtained by the 'Nujol' method. The pattern obtained from the KBr pellet in this case was diffuse, which may be the result of a pressure-inversion. The spectra obtained by the 'Nujol' and KBr methods were comparable in the other three substances, although the KBr spectra were more sharp, more easily reproducible, and had no interfering absorption maxima.

Acknowledgements.—This work was done as part of a study of phase equilibria in steelplant refractories systems sponsored by the American Iron and Steel Institute, on the recommendation of the Research Sub-committee of the Refractories Committee. The author is indebted to Drs. E. F. Osborn and R. Roy for advice and discussion of the manuscript.

Conclusions.

Studies of phase relations among Ca_2SiO_4 , $Ca_3MgSi_2O_8$, and $CaMgSiO_4$ in the temperature range $600^{\circ}-900^{\circ}$ C. confirm the existence of each of these compositions as a unique compound in the system CaO--MgO-SiO₂. X-ray diffraction, optical and infra-red absorption data on crystalline material obtained in the hydrothermal runs all point to the conclusion that no appreciable solid solution exists among the phases in this temperature range.

References.

- BREDIG (M. A.), 1950. Journ. Amer. Ceram. Soc., vol. 33, p. 188.
- LARSEN (E. S.) and FOSHAG (W. F.), 1921. Amer. Min., vol. 6, p. 143 [M.A. 1–254]. NEWMAN (E. L.) and WELLS (L. S.), 1946. Journ. Res. Nat. Bur. Stand., vol. 36, p. 137.
- NURSE (R. W.), 1954. 'The Dicalcium Silicate Phase', Paper No. 3, pp. 56-90, Proceedings of the Third International Symposium on the Chemistry of Cement, 1952. Cement and Concrete Association, 52 Grosvenor Gardens, London, S.W. 1.
- O'DANIEL (H.) and TSCHEISCHWILI (L.), 1942. Zeits. Krist., vol. 104, p. 124 [M.A. 8–290].
- OSBORN (E. F.), 1943. Journ. Amer. Ceram. Soc., vol. 26, p. 321 [M.A. 10-108].
- PHEMISTER (JAMES), NURSE (R. W.), and BANNISTER (F. A.), 1942. Min. Mag., vol. 26, p. 225.
- RICKER (R. W.) and OSBORN (E. F.), 1954. Journ. Amer. Ceram. Soc., vol. 37, p. 133.
- Roy (D. M.), Roy (R.), and Osborn (E. F.), 1950. Ibid., vol. 33, p. 153.
- ROY (R.) and OSBORN (E. F.), 1952. Econ. Geol., vol. 47, p. 717.
- TRÖMEL (G.), 1949. Naturwiss., vol. 36, p. 88.
- ----- and Möller (H.), 1949. Fortschr. Min. Krist. Petr., vol. 28, p. 80 [M.A. 11-470].
- VAN VALKENBURG (A.) and MCMURDIE (H. F.), 1947. Journ. Res. Nat. Bur. Stand., vol. 38, p. 415.

194