Derivative structures based on the alunite octahedral sheet: mitridatite and englishite

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SUMMARY. It is proposed that mitridatic and englishite contain octahedral sheets of the alunite type and that their formulae are $[Ca_6(H_2O)_3(PO_4)_2][Fe_3(OH)_6(PO_4)(PO_3OH)]_3$ and $[K_2Na(H_2O)_4Ca_{4.5}(PO_3OH)_3][Al_3(OH)_6(PO_4)(PO_3OH)]_3$ respectively.

DURING the course of compiling new single crystal data on basic iron and aluminium phosphates, a striking relationship was noted between two very complex structures mitridatite and englishite—and the relatively simple alunite structure type. The alunite structure is based on a network of corner-linked Me³⁺–oxygen octahedra whose cation centres reside on the nodes of the familiar Kagomé net. An orthohexagonal unit of this net includes six octahedra (fig. 1). Many compounds, both natural and synthetic, are based on the alunite structure, including crandallite, Ca[Al₃ (OH)₆(PO₄)(PO₃OH)], a species recently investigated in detail by Blount (1974).

The relationship was noted during a three-dimensional crystal structure analysis on mitridatite. Despite the large cell with an estimated 40 atoms in the asymmetric unit, it was hoped that the chemistry of this important phase could be rationalized. Presently, we have located the Fe³⁺ atoms, the associated PO₄ tetrahedra, and the hydroxyl groups; these define the components of an alunite-like sheet, $[Fe_3^{3+}(OH)_6$ $(PO_4)(PO_3OH)]^{2-}$, parallel to {100} and located at intervals x = 1/4 and 3/4. Unfortunately, progress on the rest of this structure is slow owing to difficulty in assigning correct phases associated with weak superstructure reflexions. A single crystal study on englishite shows that reflexions in the b^*c^* plane are related to those of mitridatite and that both compounds reveal a strong pseudohexagonal intensity distribution in this plane (cf. Moore, 1974).

Cell parameters for englishite, mitridatite, and crandallite appear in Table I. The data for englishite, heretofore unpublished, were obtained from a single crystal fragment of Fairfield, Utah, material. Owing to the large cell, it was not possible to refine the full three-dimensional cell based on the powder data since ambiguities exist in indexing. The mineral exhibits perfect micaceous cleavage parallel to $\{100\}$; preferential orientation of flakes crushed with glass and placed on a slide afforded the enhanced hoo reflexions, which were utilized to refine the *a*-axial dimension (Table II). The remaining data were obtained from calibrated precession photographs and probably have associated errors of 0.3 %. Streaks appeared on the photographs indicating

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FIG. 1. The Kagomé net whose nodes are the octahedral centres in sheets of the alunite type. Associated tetrahedra above (+) or below (—) the sheets are indicated. The axes a_1 and a_2 refer to the orthohexagonal unit in Table IV; b_E and c_E , and b_M and c_M are the crystal cell coordinates for englishite and mitridatite in Table I.

TABLE I. Cell data for englishite, mitridatite, and crandallite

		Englishite*	Mitridatite [†]	$Crandallite^+$	* Errors in $\underline{b},\ \underline{c},\ and\ \beta$ are uncertain since the results					
a		38.43(2) Å	17.52(2)	7.00(1)	were computed from precession films					
<u>b</u>		11-86 Å	19.35(4)		† Data from Moore (1974)					
c		20.67 Å	11.25(2)	16.19	+ Data from Blount (1974)					
β		111° 16	95°55′		Observed densities are from Larsen and Shannon (1930)					
Space group		A2/a or Aa	A2/a or Aa	R3m	for englishite and crandallite, and from Moore (1974)					
Density	obs -	2.65 g.cm ⁻³	3.24	2.95	for mitridatite.					
	calc.	2.67	3.27	3.00						
z		8	4	1						

a small degree of disorder in the crystal. Powder data, obtained from a single crystal Gandolfi camera, are presented in Table III. No attempt was made to index these data since a full three-dimensional single crystal data set is not yet at hand.

The crucial information appears in Table IV. Here, the cell data have been transformed into the orthonexagonal equivalent of the alunite cell, with $a_1 = a_{\text{hex}}$, $a_2 = a_{\text{hex}}\sqrt{3}$ and $c' = c_{\text{hex}}$. It is noted that a, c/3, and b/3; $a\sqrt{3}$, b, and c; and $c, \frac{1}{2}a\sin\beta$,

TABLE IIEnglishite; the hoo reflexions

TABLE IIIEnglishite: X-ray powder data

I	d	<u>hk1</u>	<u>a</u> sin β	1 1	<u>d</u>	I	<u>d</u>	I	<u>d</u>	I	<u>d</u>
75	17.89 Å	200	35.78 \$	9	17.66Å	1	3.18 Å	4	2.34 1	1	1.62 Å
100	8.96	400	35.84	10	8.94	3	3.08	1	2.22	3	1.58
45	5.965	600	35.79	1	7.01	6	2.96	1	2.19	2	1.42
40	3.581	10.0.0	35.81	2	5.86	8	2.84	2	2.09	3	1.36
75	2.983	12.0.0	35.80	6	5.57	2	2.77	2	1.97	2	1.16
70	2.558	14.0.0	35.81	1	5.25	6	2.68	2	1.93	2	1.13
20	2.240	16.0.0	35.84	4	3.85	2	2.62	2	1.84	3	1.08
15	1.9897	18.0.0	35.82	1	3.58	-2	2.58	5	1.71	2	1.03
5	1.7911	20.0.0	35.82	1	3.40	3	2.54	2	1.68	3	0.98
15	1.6295	22.0.0	35.85	2	3.35	1	2.46	1	1.65	-	
Flakes	crushed wit	th glass; S	i internal	114	.6 mm diam	• camei	ra, single-	crystal	Gandolfi	techni	que.
standa	rd (<u>a</u> 5.430:	1 A). Chart	diffractogram,	Fe-	Kα radiati	on, Mn	filter	•			,
Cu− <u>K</u> α	radiation,	° per min.		• •	_						

TABLE IV. Alunite sheet structures: orthohexagonal unit

	Alunite	Englishite	Mitridatite	Crandallite	
<u>a</u> 1	<u>a</u> 6.97 Å	<u>c</u> /3 6.89 Å	<u>b</u> /3 6.45 Å	<u>a</u> 7.00 Å	
<u>a</u> 2	a y 3 12.07	<u>b</u> 11.86	<u>c</u> 11.25	a√ 3 12.11	
<u>c</u> '	<u>c</u> 17.38	$\frac{1}{2a}$ sin β 17.90	<u>a</u> sin β 17.43	<u>c</u> 16.19	
Vol. per 0 ²⁻	17.4 Å ³	18.9 Å ³	17.9 A ³	16.2 Å ³	
No. of sheets	3	2	3	3	

 $a \sin \beta$ for alunite, englishite, and mitridatite respectively are related according to the orthohexagonal unit. Crandallite is included for purposes of comparison since it is a phosphate and since 'acid' phosphate tetrahedra occur in the structure.

To relate their crystal chemistry further, it is assumed that mitridatite contains $[Fe_3^{3+}(OH)_6(PO_4)(PO_3OH)]^{2-}$ sheets, and englishite (and crandallite) $[Al_3^{3+}(OH)_6(PO_4)(PO_3OH)]^{2-}$ sheets of the alunite type. The chemical analysis for englishite in Larsen and Shannon (1930) and for mitridatite in Moore (1974) were recomputed on the basis of this interpretation. It is seen that in the *c*'-repeat, alunite and crandallite possess three sheets, and englishite and mitridatite two sheets each. The remaining ions are assumed to reside between these sheets. In this manner, the distributions between the sheets and in the sheets can be assigned (Table V). Presumably, these complexes of ions between the sheets contribute to the complexity of the mitridatite and englishite

crystal cells. Further analysis of their structures should provide an answer for the distribution of these inter-sheet ions.

Further support for this interpretation comes from calculations on the oxygen packing. The volumes per oxygen are 17.4, 18.9, 17.9, and 16.2 Å³ for alunite, englishite, mitridatite, and crandallite respectively (Table IV). Finally, the calculated compositions for englishite and mitridatite are in good agreement with the published analyses (Table VI) as are the computed densities and observed specific gravities (Table I).

 TABLE V. Proposed crystal-chemical formulae for crandallite, mitridatite, and englishite

	Between sheets	In sheets
Crandallite Mitridatite Englishite	$\begin{array}{l} [Ca_3]^{6+} \\ [Ca_6(H_2O)_3(PO_4)_2]^{6+} \\ [K_2Na(H_2O)_4Ca_{4\cdot5}(PO_3OH)_3]^{6+} \end{array}$	$\begin{array}{l} [Al_{3}(OH)_{6}(PO_{4})(PO_{3}OH)]_{3}^{6}-\\ [Fe_{3}(OH)_{6}(PO_{4})(PO_{3}OH)]_{3}^{6}-\\ [Al_{3}(OH)_{6}(PO_{4})(PO_{3}OH)]_{3}^{6}-\end{array}$

 TABLE VI. Chemical analyses of mitridatite and englishite, compared with theory for the formulae of Table IV

	Englishite		Mitridatite		
	1	2	3	4	1. Larsen and Shannon, 1930
к ₂ 0	5.4	5.34	-	-	2. Theoretical composition
Na ₂ 0	1.6	1.75	-	-	3. Ito anal in Moone 1974
Ca0	14.1	14.31	17.4	18.03	4. Theoretical composition
A1203	24.7	26.03	-	-	it metretrear composition
$Fe_2^0_3$	-	-	38.3*	38.51	* Includes 2.7% Mn ₂ 03
P205	37.8	36.24	31.5	30.44	2.5
н ₂ 0	16.5	16.33	12.8	13.02	
Sum	100.1	100.00	100.0	100.00	

These interpretations shed new light on the complex mineral chemistry of mitridatite and englishite. The common appearance of mitridatite as a low-temperature product in phosphate-bearing pegmatites has been noted by Moore (1974) and englishite occurs associated with crandallite and related phosphates (Larsen and Shannon, 1930). As the alunite sheet is a stable moiety over a wide range of conditions, it is not unreasonable that its octahedral basis can be utilized in a variety of more complicated structures.

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