

tine, lavendulan, schoepite, vandendriesscheite, kahlerite and metakahlerite are confirmed for the first time from the British Isles.

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References

- Breithaupt, J. F. A. (1837) *J. prakt. Chem.* **10**, 505.
 Clark, A. M., Couper, A. G., Embrey, P. G. and Fejer, E. E. (1986) *Mineral. Mag.* **50**, 731–3.
 Collins, J. H. (1881) *Catalogue of the Minerals in the Museum of the Royal Institution of Cornwall*, **2**, 32.
 Foshag, W. F. (1924) *Am. Mineral.* **9**, 30–1.
 Goldsmith (1877) *Proc. Acad. Nat. Sci. Philadelphia*, 192.
 Guillemin, C. (1956) *Bull. Soc. franc. Min. Crist.* **79**, 7–95.
 Harrison, R. K., Tresham, A. E., Young, B. R. and Lawson, R. I. (1975) *Bull. Geol. Survey Great Britain*, **52**, 1–26.
 Macpherson, H. G. and Livingstone, A. (1982) *Glossary of Scottish Mineral Species 1981. Scottish Journal of Geology*.
 Miller, J. M. and Taylor, K. (1966) *Bull. Geol. Survey Great Britain*, **25**, 1–8.
 Weisbach, A. (1871) *Neues Jahrb. Min.* 869–70.
 — (1877) *Ibid.*, 1.
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KEYWORDS: supergene alteration, walpurgite, eulytine, lavendulan, schoepite, vandendriesscheite, kahlerite, metakahlerite, Dalbeattie, Scotland.

*Chemistry Department, University of
 Manchester Institute of Science and
 Technology, Manchester M60 1QD*

*10 Lynton Drive, Burnage,
 Manchester M19 2LQ*

R. S. W. BRAITHWAITE

J. R. KNIGHT

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Occurrences of grandidierite, serendibite and tourmaline near Ihosy, southern Madagascar

TOURMALINE is a nearly ubiquitous accessory mineral in metapelites at low and medium grades of metamorphism, whereas the high-temperature borosilicates kornerupine, grandidierite, and serendibite are very rare and occur at the high-grade conditions of the granulite and pyroxene-hornfels facies (de Roever and Kieft, 1976; Grew, 1988; Lonker, 1988, etc.). Grandidierite and serendibite have been found near Ihosy, in southern Madagascar, the former in a seven-phase anatectic gneiss, the latter in a calcsilicate gneiss.

Grandidierite in anatectic gneiss

In the Precambrian Ihosy formation, banded cordierite–garnet gneisses, leptynites, two-pyroxene metabasic granulites, calcsilicate gneisses and

marbles indicate conditions of the hornblende intermediate-pressure granulite facies. A widespread anatectic event produced migmatitic gneisses with seven phases: quartz, K-feldspar, plagioclase, garnet, cordierite, sillimanite, and biotite (with green spinel included in cordierite and garnet) and aluminous residues rich in biotite (biotite–cordierite–garnet or sillimanite, \pm quartz) and sometimes containing sapphirine and kornerupine (Mégérin, 1968; von Knorring *et al.*, 1969; Nicollet, 1985, 1988, 1990). Monazite and zircon from a granodiorite dyke produced by the anatectic event yielded a U–Pb age of 561 ± 12 Ma (Andriamarofahatra *et al.*, in preparation).

Dark green tourmaline and blue grandidierite (included in cordierite) occur as very rare crystals in the anatectic gneisses. However, these two

borosilicates are never found in the same thin section. In a quarry, 1 km from Ihosy, grandierite is somewhat more abundant (up to ten crystals per thin section) in coarse-grained quartzo-feldspathic segregations. In these lenses, the same minerals as in the seven-phase gneisses are 1 cm or more in dimension. Grandierite is included in cordierite and garnet. Within the cordierite, small anhedral grains (≈ 0.1 mm) of the borosilicate are in sharp contact with biotite, sillimanite, green spinel, and opaques; it has been observed by other authors that grandierite and sillimanite are closely associated (e.g. van Bergen, 1980; Grew, 1983). The garnet may contain large polyminerals inclusions composed of slender prisms and needles of grandierite (≈ 0.5 mm long) with hexagonal basal sections, K-feldspar, plagioclase, quartz, green biotite, and apatite (Fig. 1). The borosilicate shows the usual blue pleochroism. With a X_{Fe} [$\text{Fe}/(\text{Fe} + \text{Mg})$] between 0.25 and 0.29 (Table 1), the mineral is moderately Fe rich (cf.

Table 1: Electron microprobe analyses

	1	2	3	4
SiO ₂	20.56	35.43	24.91	35.83
Al ₂ O ₃	51.96	33.87	36.36	32.41
FeO _T	7.17	6.14	1.76	0.65
MgO	0.15	0.02	0.11	0.02
MnO	9.81	7.67	14.91	12.91
CaO	-	2.27	15.49	4.67
Na ₂ O	-	1.41	0.31	0.51
K ₂ O	0.02	0.05	-	0.03
TiO ₂	0.02	0.16	0.02	-
Total	89.69	87.02	93.86	87.02
Oxygens	15	24.5	18.5	24.5
Si	2.006	5.733	2.976	5.689
Al	5.975	6.461	5.122	6.067
Fe ²⁺	0.585	0.831	0.176	0.086
Mn	0.012	0.003	0.011	0.003
Mg	1.426	1.850	2.656	3.055
Ca	-	0.394	1.953	0.795
Na	-	0.442	0.072	0.157
K	0.006	0.010	-	0.006
Ti	0.002	0.020	0.002	-
XMg	0.71	0.69	0.94	0.97

1 - Grandierite, seven-phase anatectic gneiss (234)

2 - Tourmaline, seven-phase anatectic gneiss (Ih 5)

3 - Serendibite, calcsilicate gneiss (230)

4 - Tourmaline, calcsilicate gneiss (230).

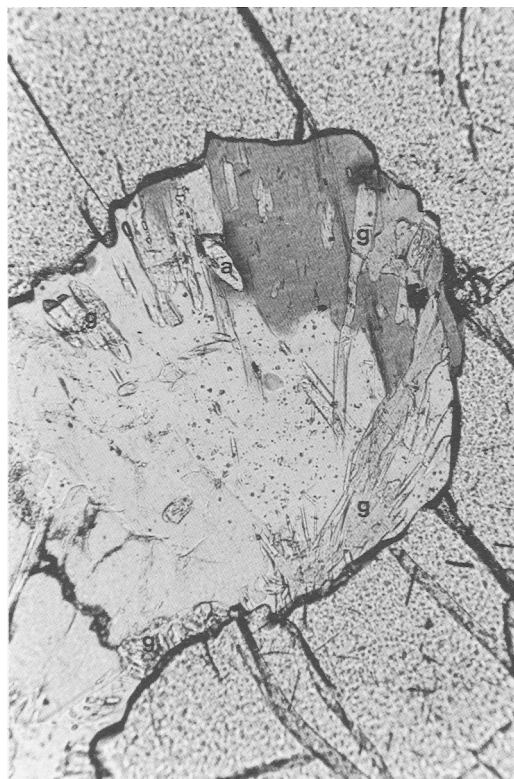
B₂O₃ not determined; all Fe as FeO.

FIG. 1. Sheaf-like aggregate of grandierite (g) with quartz, plagioclase, K-feldspar, green biotite, apatite (a) included in garnet. Note that prisms of the borosilicate grow on the garnet. The width of the photomicrograph is approximately 1.0 mm; plane-polarized light.

Huijsmans *et al.*, 1982; de Roever and Kieft, 1976; Semroud *et al.*, 1976; van Bergen, 1980; Haslam, 1980; Grew, 1983). Tourmaline in the banded gneisses has a composition intermediate between dravite and schörlite (Table 1); X_{Fe} is similar to that of the grandierite. Green biotite included in garnet (Fig. 1) is less titaniferous and magnesium rich than the brown mica in the matrix, as a result of equilibration with adjacent garnet during cooling.

Metamorphic pressures and temperatures are estimated in the seven-phase gneisses to have been near 700°C and 4–5 kbar and the application of the seven-phase thermodynamic model of Lee and Holdaway (1977) indicates that $P_{\text{H}_2\text{O}}$ was 0.35–0.4 P_t (total pressure) (Nicollet, 1985, 1988). Relations between grandierite-bearing rocks and partial melting have often been noted (e.g. de Roever and Kieft, 1976; van Bergen, 1980). The grandierite gneiss from Ihosy and the P – T conditions of its formation are quite similar to those described by Lonker in Canada (1988). At both localities, it may be suggested that the grandierite was generated through low $P_{\text{H}_2\text{O}}$ partial melting of tourmaline-bearing pelitic rocks.

Serendibite in calcsilicate gneiss

Ten kilometres southeast of the preceding outcrop (Vohimena's hill; 46°15'30" E, 22°27'30" S), a serendibite-bearing calcsilicate gneiss is in contact with an olivine marble. The serendibite-bearing gneiss contains anorthite (An_{90-100}), scapolite

(Me_{75-83}), aluminous diopside, green amphibole, colourless spinel, large poikilitic blue tourmaline, and a little calcite. Tourmaline includes several minerals, especially clinopyroxene, spinel, and a few small crystals of serendibite. Scattered grains of the latter in one crystal of tourmaline are in optical continuity, suggesting that tourmaline is a breakdown product after serendibite. The serendibite is colourless to very light green. The mineral is much less coloured than the prussian blue crystals in a calcsilicate rock associated with clintonite clinopyroxenites from Ianapera, in SW Madagascar (Nicollet, 1988, 1990). The Ihosy serendibite has low birefringence and fine polysynthetic twinning. It may be mistaken for sapphirine, but it is distinguishable from the latter by a larger extinction angle and by its occurrence in calcic rocks. The ferromagnesian minerals in the rock are magnesium rich. Tourmaline (Table 1) is a magnesian uvite and the X_{Mg} ratio of the spinel is greater than 0.9. Hornblende and clinopyroxene are close to their Mg end-members ($X_{\text{Mg}} \approx 0.97$). The Fe–Mg partitioning between tourmaline and serendibite (K_d) is ≈ 0.4 as in the more iron rich pair from Ianapera; it is lower than the K_d of coexisting tourmaline and serendibite from Melville Peninsula, Canada ($K_d = 0.61$; Hutcheon *et al.*, 1977).

The P – T conditions of the crystallization of this rock are similar to those estimated for the neighbouring seven-phase anatectic gneisses (Nicollet, 1988).

KEYWORDS: grandidierite, serendibite, tourmaline, Ihosy, Madagascar, metamorphism.

Département de Géologie, UA 10,
5 Rue Kessler, Clermont Fd, 63038, France

CHRISTIAN NICOLLET

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References

- de Roever, W. F. and Kieft, C. (1976) *Am. Mineral.* **61**, 332–3.
 Grew, E. S. (1983) *Mineral. Mag.* **47**, 401–3.
 — (1988) *Am. Mineral.* **73**, 345–57.
 Haslam, H. W. (1980) *Mineral. Mag.* **43**, 822–3.
 Huijsmans, J. P. P., Barton, M., and van Bergen, M. J. (1982) *Neues Jahrb. Mineral. Abh.* **143**, 249–61.
 Hutcheon, I., Gunter, A. E. and Lecheminant, A. N. (1977) *Can. Mineral.* **15**, 108–12.
 Lee, S. M. and Holdaway, M. J. (1977) *Geophys. Monogr.* **20**, 79–94.
 Lonker, S. W. (1988) *Contrib. Mineral. Petrol.* **98**, 502–16.
 Mégerlin, N. (1968) *C.R. Sem. Géol. Madagascar*, 67–9.
 Nicollet, C. (1985) *Précamb. Res.* **28**, 175–85.
 — (1988) Thèse d'Etat. Clermont Fd, 350pp.
 — (1990) In *Granulites and Crustal Differentiation* (Vielzeuf, D. and Vidal, P., eds.), Kluwer Academic publishers, in press.
 Semroud, B., Fabriès, J., and Conquére, F. (1976) *Bull. Soc. fr. Mineral. Cristallogr.* **99**, 58–60.
 van Bergen, M. J. (1980) *Mineral. Mag.* **43**, 651–8.
 von Knorring, O. V., Sahama, T. G., and Lehtinen, M. (1969) *Bull. Geol. Soc. Finland*, **41**, 79–84.

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Serendibite from the northwest Adirondack Lowlands, in Russell, New York, USA

SERENDIBITE, to a first approximation, $\text{Ca}_2(\text{Mg}, \text{Fe}^{2+})_3(\text{Al}, \text{Fe}^{3+})_{4.5}\text{B}_{1.5}\text{Si}_3\text{O}_{20}$, has been reported from about eight localities worldwide (Deer *et al.*, 1978; Nicollet, 1988), including Johnsburg, New York, in the southern Adirondack Highlands (Larsen and Schaller, 1932; Grew *et al.*, in press). It is a mineral of high-temperature calc-silicate skarns, mostly in the granulite-facies. In this

paper we report a new occurrence from the northwest Adirondacks Lowlands. This occurrence has many mineralogical and chemical features in common with the serendibite–diopside rocks in the Johnsburg deposit, although located 130 km distant.

Serendibite occurs in core from hole 1872 drilled by St. Joe Resources Company (presently