X-RAY EXAMINATIONS OF CHROME ORES: (I) LAT-
TICE DIMENSIONS; (II) THEORETICAL DENSITIES

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(I) LATTICE DIMENSIONS

As a part of an investigation of certain properties of chrome ores, the lattice dimensions of five chromites from various localities ranging in composition from $33 \%$ to $53 \% \mathrm{Cr}_{2} \mathrm{O}_{3}$ were determined.

The diffraction patterns were obtained using a multiple unit apparatus, made by the General Electric Company. The radiations employed were $K \alpha$ doublet of molybdenum of the wave length $0.712 \AA$.U. The spacings were standardized with sodium chloride.

The length of the edge of the unit cube $a_{0}$ for the five ores, given in the table II is calculated from the two following equations of $x$-ray spectroscopy ${ }^{1,2}$

$$
\begin{align*}
& \frac{d}{n}=\frac{\lambda}{2 \sin \theta}  \tag{1}\\
& a_{0}=\frac{d_{h k l}}{n} \sqrt{\left(h^{2}+k^{2}+l^{2}\right) n^{2}} \tag{2}
\end{align*}
$$

$d / n$, the spacing measurements of the diffraction lines, recorded in the third column of the table II are for the chrome ore A. In the fourth column are given the estimated relative intensities for the same ore, assigning the brightest line $a$ value of ten.

In the cubic system the ratios of $\sin ^{2} \theta$, (the latter being computed from the equation I) should give values which are integers and usually small. These values given in column five for A alone, diverge within the limits of the experimental error from the corresponding whole numbers given in the second column of the table II. The whole numbers represent the sum of $\left(h^{2}+k^{2}+l^{2}\right) n^{2}$ and aid in assigning the indices of the planes producing the diffraction pattern.

The ores examined were:
A. Rhodesian African lump chrome ore
B. Transvaal African lump chrome ore
C. Grecian lump chrome ore
D. Cuban lump chrome ore
E. A high grade chrome ore, locality not known.

The ores for which no analyses were supplied were analyzed. In the case of the ores for which only partial (factory) analyses were available, the undetermined constituents were estimated in order to note roughly any obvious relationship existing between the lattice dimension and the chemical composition. The estimates were based upon the analyses of the typical ores found in the localities whence the ores were obtained. Publications on the mineral resources of U.S. ${ }^{3}$ and the records of the Geological Surveys were consulted in this connection. Two of the typical analyses of the Rhodesian and Grecian ores given by McDowell ${ }^{4}$ are:

|  | Rhodesian | Grecian |
| :--- | :---: | :---: |
| $\mathrm{Cr}_{2} \mathrm{O}_{3}$ | 48.4 | 39.5 |
| FeO | 14.4 | 15.8 |
| $\mathrm{Al}_{2} \mathrm{O}_{3}$ | 12.1 | 26.2 |
| $\mathrm{MgO}_{2 i O}^{2}$ | 14.2 | 15.8 |
| $\mathrm{SiO}_{2}$ | 5.9 | 3.0 |
| CaO | - | - |

The estimates of the undetermined oxides of the Rhodesiam (A) and Grecian (D) ores were partly based upon these. For Transvaal ore the estimates were based upon Wagner's ${ }^{5}$ work. Parantheses are placed around the figures which are estimated. The analyses of the ores are given in the table I.

Table I
Analyses of the Chrome Ores

|  | $\mathrm{A}$ <br> Rhodesian | B <br> Transvaal | C <br> Cuban | D <br> Grecian | E | Hungarian (Todoky ${ }^{10}$ ) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{Cr}_{2} \mathrm{O}_{3}$ | 47.14 | 42.68 | 33.83 | 37.74 | 53.08 | 54.22 |
| $\mathrm{SiO}_{2}$ | 7.75 | 3.34 | 3.11 | 6.09 | 5.11 | 6.16 |
| FeO | (15.00) | (27.00) | 13.07 | (15.00) | 14.00 | 25.94 |
| CaO |  |  | Trace |  | 2.59 |  |
| MgO | (15.00) | (10.00) | 17.61 | (15.00) | 8.95 | 9.77 |
| $\mathrm{Al}_{2} \mathrm{O}_{3}$ | (13.00) | (15.00) | 31.11 | (26.00) | 16.11 | 2.29 |
| Loss on Ignition |  |  | 1.12 |  |  |  |
| MnO |  |  |  |  |  | 0.94 |
| $\mathrm{TiO}_{2}$ |  |  |  |  |  | 0.68 |
| $\mathrm{H}_{2} \mathrm{O}$ |  |  |  |  |  | 0.09 |
| Total | 97.89 | 98.02 | 99.85 | 99.83 | 99.84 | 100.09 |

The results seem to indicate that the values of $a_{0}$ for the different ores are inversely proportional to the $\mathrm{Al}_{2} \mathrm{O}_{3}$ content; or the higher the amount of aluminates present in the isomorphous mixture of spinels, the lower is the length of the edge of the unit cube (See Fig. 1).


Fig. 1
Tests on synthetic spinels ${ }^{6}$ had shown that the structures of the chromites of $\mathrm{Zn}, \mathrm{Mg}, \mathrm{Fe}$ and Mn are generally intermediate between the structures of their ferrites and their aluminates. The lattice dimensions of the aluminates ranged from 8.062 to 8.271 Å.U.; those of the chromites varied from 8.296 to 8.436 , the ferrites had the values between the limits of 8.423 to 8.457 . Therefore, when either $\mathrm{Al}_{2} \mathrm{O}_{3}$ or $\mathrm{Fe}_{2} \mathrm{O}_{3}$ replaces $\mathrm{Cr}_{2} \mathrm{O}_{3}$ from the chromite molecule the value of $a_{0}$ should be lowered or raised in proportion to the replacement, in accordance with the additive law of Vegard. ${ }^{7}$

Table II
Powder Diffraction Data from Chrome Ores


Table III

| Synthetic <br> Chromite $^{6}$ | Rhode- <br> sian $A$ | Trans- <br> vaal $B$ | $E$ | Grecian <br> $D$ | Cuban <br> $C$ | Ural $^{9}$ | Hunga- <br> rian $^{10}$ |
| :--- | :--- | :--- | :---: | :---: | :---: | :---: | :---: |
| $8.344 \pm$ | $8.283 \pm$ | $8.267 \pm$ | $8.241 \pm$ | $8.216 \pm$ | $8.179 \pm$ | $8.355 \pm$ | 8.05 |
| 0.003 | 0.001 | 0.001 | 0.001 | 0.001 | 0.002 | $0.004 ;$ <br> $8.358 \pm$ <br> 0.004 |  |

In this investigation it appears that $\mathrm{Cr}_{2} \mathrm{O}_{3}$ is replaced by $\mathrm{Al}_{2} \mathrm{O}_{3}$ alone since the values of $a_{0}$ gradually decline from the synthetic $\mathrm{FeCr}_{2} \mathrm{O}_{4}$, which contains no $\mathrm{Al}_{2} \mathrm{O}_{3}$ to the Cuban ore which contains $31 \% \mathrm{Al}_{2} \mathrm{O}_{3}$. This seems to be further substantiated by the fact that the 5 ores examined are nonmagnetic while all the known ferrites, except $\mathrm{ZnFe}_{2} \mathrm{O}_{4}$, are ferromagnetic and zinc seldom occ̣urs in chromite. ${ }^{8,12}$

It may be inferred that iron which is generally analyzed as $\mathrm{Fe}_{2} \mathrm{O}_{3}$ is present in these ores wholly in ferrous form (ferrous chromite, $\mathrm{FeC}_{2} \mathrm{O}_{4}$; ferrous aluminate, $\mathrm{FeAl}_{2} \mathrm{O}_{4}$ ) and not as ferrite.

No characteristic diffraction lines were observed for the minerals of $\mathrm{SiO}_{2}$ and CaO , although they are not minerals of the spinel series. Their presence in the isomorphous spinel molecule whether interstitial or substitutional does not seem to have any marked effect on the lattice constants.

The $x$-ray data obtained by other investigators on chrome ores are given in the table III, together with those obtained in this investigation. Todoky's value of $8.05 \AA . \mathrm{U}$. appears to be low considering the fact that it is a high grade chrome ore as revealed by its analysis (see Table I)

The quality of the chrome ore, it is evident, can not be definitely ascertained by means of $x$-ray examination; however, it can be roughly surmised as to its $\mathrm{Al}_{2} \mathrm{O}_{3}$ content.
(II) THEORETICAL DENSITIES

An attempt was made to compute the theoretical densities of the ores by employing the well known relation:

$$
\begin{aligned}
\rho & =\frac{n M}{a_{0}{ }^{3}} \\
\text { in which } \rho & =\text { density } \\
n & =\text { number of molecules per unit cell; } \\
& 8 \text { in the case of spinels } \\
M & =\text { molecular weight of the substance } \\
a_{0} & =\text { length of the edge of the unit cube. }
\end{aligned}
$$

The above equality holds true in the case of a substance composed only of a single kind of crystal. The ore is regarded as such, consisting of an isomorphous mixture of spinels.

The small amount of the mineral impurities in the ore (possibly due to the improper separation of chromite from gangue), some of these being visible under the microscope, yet giving no evidence of their presence in the diffraction patterns, were regarded, in order to
simplify the "average molecular weight" ( $M$ ) computations, as being present within the chromiferrous spinelled in solid solution. The value for $M$ was determined from the mineral analysis of the ore. The following rules were observed in calculating the latter.

1. All $\mathrm{Cr}_{2} \mathrm{O}_{3}$ or FeO was regarded as entering the chromite molecule.
2. $\mathrm{SiO}_{2}$ was disposed of as magnesium olivine- $2 \mathrm{MgO} \cdot \mathrm{SiO}_{2}$; serpentine$2 \mathrm{H}_{2} \mathrm{O} \cdot 3 \mathrm{MgO} \cdot 2 \mathrm{SiO}_{2}$; or as uvarovite $-3 \mathrm{CaO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3} \cdot 3 \mathrm{SiO}_{2}$; microscopic study determining the choice.
3. The remaining oxides were distributed to form the other members of the spinel series- $\mathrm{MgO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3}, \mathrm{MgO} \cdot \mathrm{Al}_{2} \mathrm{O}_{3}, \mathrm{FeO} \cdot \mathrm{Al}_{2} \mathrm{O}_{3}$. The excess of either acidic or basic oxides, which practically all chrome ore analyses indicate, were reported as such-periclase ( MgO ) and corundum $\left(\mathrm{Al}_{2} \mathrm{O}_{3}\right)$.
4. Ignition loss reported in the Cuban ore, C , was regarded as $\mathrm{CO}_{2}$ and was disposed of as magnesite, $\mathrm{MgCO}_{3}$.

Taking, as an illustration, chrome ore C, for which a complete analysis was available, the percentages of the various oxides were divided by their molecular weights, and these were later combined to form different minerals, as is illustrated below.

$$
\text { Mols. of various oxides }=\frac{\text { Percentage }}{\text { Molecular Weight }}
$$

|  | $\begin{gathered} \mathrm{Cr}_{2} \mathrm{O}_{3} \\ 0.2225 \end{gathered}$ | $\begin{gathered} \mathrm{FeO} \\ 0.1821 \end{gathered}$ | $\begin{gathered} \mathrm{SiO}_{2} \\ 0.0513 \end{gathered}$ | $\begin{gathered} \mathrm{MgO} \\ 0.4371 \end{gathered}$ | $\begin{gathered} \mathrm{Al}_{2} \mathrm{O}_{3} \\ 0.3045 \end{gathered}$ | $\quad \mathrm{CO}_{2}$ (Ignition loss) 0.0254 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $0.1821 \mathrm{FeO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3}$ <br> $0.05132 \mathrm{MgO} \cdot \mathrm{SiO}_{2}$ <br> Magnesium olivine <br> $0.0404 \mathrm{MgO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3}$ <br> $0.0254 \mathrm{MgCO}_{3}$ <br> Magnesite <br> $0.2686 \mathrm{MgO} \cdot \mathrm{Al}_{2} \mathrm{O}_{3}$ <br> spinel <br> $0.0359 \mathrm{Al}_{2} \mathrm{O}_{3}$ <br> Corundum <br> 0.6037 Total | $\begin{aligned} & 0.1821 \\ & 0.0404 \end{aligned}$ | 0.1821 | 0.0513 | $\begin{aligned} & 0.1026 \\ & \\ & 0.0404 \\ & 0.0254 \\ & 0.2686 \end{aligned}$ | $\begin{aligned} & 0.2686 \\ & 0.0359 \end{aligned}$ | 0.0254 |

0.6037 is the combined mols. of different minerals present in 100 gms . of the ore. Therefore, the average molecular weight of the ore is $100 / 0.6037=165.6=M$, and the density

$$
\rho=\frac{8 \times 165.6 \times 1.649 \times 10^{-24}}{\left(8.179 \times 10^{-8}\right)^{3}}=3.989 .
$$

The mineral analyses, theoretical densities, and the true specific gravities of the chrome ores and the per cent difference between the theoretical and the actual values are given in the Table III.

It should be noted that the combined mols. of different minerals in each ore were uniformly divided into 100 , to obtain the average molecular weight of the ore, regardless of the fact that the analyses slightly deviated from this sum.

In the case of A , in order to form 0.0646 mols. of serpentine, the presence of which was confirmed by microscopic examination of the ore, $2.33 \% \mathrm{H}_{2} \mathrm{O}$ is necessary. This is not provided for in the analysis for A (Table) where only the three main unanalyzed constituents were estimated. However, when $2.33 \% \mathrm{H}_{2} \mathrm{O}$ is added to $97.89 \%$ already accounted for the sum approaches about $100 \%$, which, as mentioned above, was used in the average moecular weight calculation.

The sum of analyzed and unanalyzed oxides totaled $98.02 \%$ in the case of $B$ (See table $I$ ); the total number of mols. of different minerals was slightly smaller and the theoretical density and the "average molecular weight" were a little larger due to the fact that 100 was used as the basis instead of the actual value of $98.02 \%$. In this particular case the deviation between the theoretical and the actual densities was found to be the maximum, $5.93 \%$.

This method of computing the theoretical density of the ore by calculating an "average molecular weight" from the chemical composition should be applicable to all substances ${ }^{11}$ forming continuous solid solutions and obeying the additive law of Vegard.

## SUMMARY

Diffraction data by the "powder method" were obtained from 5 chrome ores found in various localities. The length of the edge of the unit cube ranged from $8.283 \pm 0.001$ to $8.179 \pm 0.002$, varying inversely with the $\mathrm{Al}_{2} \mathrm{O}_{3}$ content of the ore.

A method which has been applied to alloys for computing the theoretical density has been extended for calculating the "average molecular weight" of the chrome ore ; this method is generally applicable to substances, forming isomorphous mixtures and continuous substitutional solid solutions which obey the additive law of Vegard. The theoretical densities of the 5 ores obtained by this method had a maximum difference of $1.3 \%$ between the theoretical and the actual values in the çase of ores with complete analyses
Table III
The Mineral Analyses, Theoretical Densities, and the True Specific Gravities of the Chrome Ores

|  | A | B | C | D | E | Synthetic |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | Rhodesian | Transvaal | Cuban | Grecian |  | Chromite |
| $\mathrm{FeO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3}$ (Chromite). | 0.2089 | 0.2808 | 0.1821 | 0.2090 | 0.1951 |  |
| $\mathrm{MgO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3}$ (Magnochromite). | 0.1012 |  | 0.0404 | 0.0391 | 0.1384 |  |
| $\mathrm{MgO} \cdot \mathrm{Al}_{2} \mathrm{O}_{3}$ (Spinel) | 0.0772 | 0.0515 | 0.2686 | 0.2421 | 0.0065 |  |
| $2 \mathrm{MgO} \cdot \mathrm{SiO}_{2}$ (Magnesium olivine) |  | 0.0554 | 0.0513 | 0.1010 | . 0386 |  |
| $2 \mathrm{H}_{2} \mathrm{O} \cdot 3 \mathrm{MgO} \cdot 2 \mathrm{SiO}_{2}$ (Serpentine). | 0.0646 |  |  |  |  |  |
| $3 \mathrm{CaO} \cdot \mathrm{Cr}_{2} \mathrm{O}_{3} \cdot 3 \mathrm{SiO}_{3}$ (Uvarovite) .... |  |  |  |  | 0.0154 |  |
| $\mathrm{MgCO}_{3}$ (Magnesite) . . . . . . . . . . |  |  | 0.0254 |  |  |  |
| MgO (Periclase). |  | 0.0857 |  |  |  |  |
| $\mathrm{Al}_{2} \mathrm{O}_{3}$ (Corundum) | 0.0500 |  | 0.0359 | 0.0133 | 0.1511 |  |
| $\mathrm{FeO} \cdot \mathrm{Al}_{2} \mathrm{O}_{3}$ (Hercynite) |  | 0.0952 |  |  |  |  |
| Total No. of Mols. in 100 gms . of the ore. | 0.5019 | 0.5686 | 0.6037 | 0.6045 | 0.5451 |  |
| Average Molecular Weight. | 199.1 | 175.8 | 165.6 | 165.3 | 183.3 |  |
| Theoretical Density. | 4.619 | 4.145 | 3.989 | 3.934 | 4.245 | 5.085 |
| True Specific Gravity D ${ }^{20} / 20$ | 4.421 | 3.913 | 3.980 | 3.888 | 4.191 | 4.998 |
| Difference between theoretical and actual densities in per cent. . ......... | 4.79\% | 5.93\% | 0.226\% | 1.18\% | 1.29\% | 1.74\% |

available, and a maximum difference of $5.93 \%$ where some of the undetermined constituents were estimated.

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