Re-examination and Crystal Structure Analysis of Litidionite

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Abstract

Litidionite, CuNaKSi₄O₁₀, occurs with tridymite in lapilli strongly modified by fumarolic activity in the Vesuvius crater. It is triclinic, space group $P\overline{1}$; the lattice parameters are: $a = 9.80(1) \ b = 8.01(1), \ c = 6.97(1) \ \text{Å}, \ \alpha = 114.12(8)^{\circ}, \ \beta = 99.52(6)^{\circ}, \ \gamma = 105.59(8)^{\circ}, \ Z = 2$. Its crystal structure—determined by Patterson and Fourier methods, using single crystal diffractometer X-ray data, and refined to an R index of 0.032 for 1450 reflections with $F_o > \sigma F_o$ —is the same as for fenaksite, FeNaKSi₄O₁₀. Tubular chains of silica tetrahedra, formed by the condensation of vlasovite-type chains, are interconnected by copper and sodium atoms which have a five-fold pyramidal coordination. Potassium atoms occur in the large cavities within the chains.

Introduction

The name litidionite was given by E. Scacchi in 1880 (Zambonini, 1935) to very tiny blue crystals associated with glass of the same color in lapilli found in the Vesuvius crater. Zambonini (1935) had difficulty measuring the interfacial angles of litidionite and could only conclude it to be triclinic or monoclinic sphenoidal. He also determined $2V_x = 56^\circ$, and indices $\alpha_D = 1.548$, $\beta_D = 1.574$. The chemical analysis he carried out, which yielded the formula (K, Na)₂Cu(Si₃O₇)₂, is very doubtful owing to the difficulty in obtaining pure litidionite.

No further data or publication appeared on this mineral other than an unindexed powder pattern (JCPDs card No. 18-713). In the text books litidionite is considered as an inadequately characterized and doubtful species. Accordingly we have re-examined litidionite and determined its crystal structure. Material used in this study was provided by Professor Carobbi from specimens that had been previously studied by himself and Zambonini.

Occurrence

Litidionite occurs as tiny plates in lapilli strongly modified by the fumarolic activity subsequent to the Vesuvius eruption of 1873. The dimension of the lapilli range from a few millimeters up to 25-30 mm. The nuclei of the lapilli are small fragments of rock or well-formed augite crystals ejected by the volcano. The fumarolic action coated the nuclei with a layer of white microcrystalline tridymite (considered as opal by Zambonini) which is overlaid by a blue glassy crust whose cavities are coated by litidionite, tridymite, and minor wollastonite. Litidionite is the unique copper mineral occurring in these lapilli, although several other copper minerals (covellite, tenorite, chlorothionite, cyanochroite, chalcocyanite, and others) were produced by the fumarolic activity in the Vesuvius crater.

X-Ray Crystallography

The lattice parameters of litidionite (Table 1) were measured with a Philips PW 1100 computer-controlled diffractometer. Its X-ray powder pattern, obtained with a Philips PW 1050 diffractometer using CuK α radiation and NaF as internal standard, is compared in Table 2 with the data on the JCPDS card No. 18-713. The three strong lines of the JCPDS data—d =4.31, d = 4.10, and d = 3.82 Å—correspond to the three strongest lines of tridymite (JCPDS card No. 18-1170).

The powder pattern and the lattice parameters of litidionite are very close to those of fenaksite, FeNaKSi₄O₁₀ (Golovachev *et al*, 1971).

Chemical Formula

Carobbi (Zambonini, 1935) carried out one chemical analysis of litidionite and gave the tentative formula $Na_2K_2CuSi_2O_7$. Zambonini (1935) considers

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	litidionite	fenaksite	
a	9 80(1) Å	10.00 Å	
ь С	8.01 "	8.18	
α	114 12°(8)	6.98	
β	99.52°(6)	100.7°	
Y Charai 1 C 1	105.59°(8)	105.0°	
Cnemical formula	CuNaKSi4010	FeNaKSi4010	
space group Z	P1 2	P1 2	
Dobs.	2.75		

TABLE 1. Unit Cell Data of Litidionite and Fenaksite

TABLE 3. Atomic Coordinates* and Equivalent Isotropic Temperature Factors**

The unit cell parameters given by Golovachev <u>et al.</u> did not correspond to the reduced cell and were transformed to those given in this Table through the following matrix (0 1 $1/\overline{1}$ 0/1 0 0).

the analysis as very doubtful owing to the difficulty in separating the mineral from associated tridymite and blue glass. In effect, if it is assumed, on the basis of the close resemblance of the crystallographic data, that litidionite and fenaksite are isostructural, the

TABLE 2. X-ray Powder Data of Litidionite

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Atom	x/a	у/Ъ	z/c	$B_{\rm H}({\rm \AA}^2)$
Cu	0.4122(1)	0.1287(1)	0.1624(1)	0.74
Si(1)	0.8593(1)	0.1742(2)	0.8119(2)	
Si(3) Si(4)	0.7222(1) 0.2110(1) 0.7256(1)	0.3545(2) 0.2877(2) 0.3825(2)	0.5664(2) 0.9582(2) 0.1557(2)	0.66 0.59 0.68
0(1)	0.5682(3)	0.2072(4)	0.0323(4)	1.25
0(2)	0.2716(3)	0.1161(4)	0.9166(5)	1.04
0(3)	0.8535(3)	0.3320(4)	0.0452(5)	1.22
0(4)	0.0302(3)	0.1851(4)	0.8549(5)	1.64
0(5)	0.2521(3)	0.0432(4)	0.2803(5)	1.34
0(6)	0.7943(3)	0.4314(4)	0.4102(4)	1.20
0(7)	0.5561(3)	0.2017(4)	0.4459(5)	1.30
0(8)	0.2562(4)	0.4471(5)	0.2155(5)	1.76
0(9)	0.8331(3)	0.2535(4)	0.6323(5)	1.30
0(10)	0.2716(3)	0.4072(4)	0.8286(5)	1.61
K	0.0038(1)	0.2033(1)	0.3348(2)	1.58
Na	0.4090(2)	0.1295(3)	0.6633(4)	

* With standard deviations in parentheses.

** After Hamilton (1959).

formula should be $CuNaKSi_4O_{10}$. A chemical formula very close to this can be obtained from Carobbi's analysis if we assume that the material analyzed contained about 40 percent tridymite.

The crystal structure analysis fully confirms that

	This p	aper		JCPDS 18-71:		
d _{obs.}	d _{calc} .	hk1	I	d	I	
6.75	6.76	ī10	35	6.73	50	
				6.08	10	
				5.16	10	
4.485	4.483	200	12			
				4.31	60	trid
				4.10	60	trid.
4.046	4.048	111	15			
				3.82	40	trid
3.652	3.664	211	18	3.66	10	
3.372	3.379	220	100	3.36	100	
3.223	3.226	210	75	3.22	60	
3.046	3.042	112	10			
3.027	\$3.033	111	10	2.01		
5.027	3.020	002	10	3.04	30	
	2.977	311				
2.976	2.976	212	13	2.96	60	
	(2.973	202				
2.835	2.840	120	18	2.82	30	
2.675	2.672	311	37	2.65	30	
2.567	2.563	302	13			
2.460	2.457	032	7			
2.409	2.413	230	85	2.40	80	
		20		2.23	10	
1.980	1.982	122	10			
1.924	1.920	042	10			
1.916	1.919	510	10	1.91	10	
		-	12	1.84	10	
1.817	1.810	342	5	1.80	10	
1,789	1.793	500	15			

TABLE 4. Analysis of the Anisotropic Thermal Parameters in Litidionite*

Atom	r.m.s.	U _i a	0 _i b	U _i c	Atom	r.m.s.	U _i a	U _i b	U _i c
Cu	0.084(1) 0.089(1) 0.114(1)	64 38 64	57 106 145	155 67 100	0(5)	0.103(6) 0.129(5) 0.154(5)	128 111 134	125 54 54	50 60 126
Si(1)	0.073(3) 0.091(2) 0.105(2)	160 71 82	92 127 143	63 29 100	0(6)	0.082(7) 0.112(6) 0.163(5)	90 46 137	63 65 38	171 96 84
Si(2)	0.067(3) 0.090(3) 0.112(2)	80 85 169	65 37 65	178 88 91	0(7)	0.098(6) 0.107(6) 0.169(5)	133 64 53	118 99 151	46 44 91
Si(3)	0.078(3) 0.086(3) 0.095(2)	123 97 146	62 29 97	53 143 94	0(8)	0.100(7) 0.134(6) 0.197(5)	81 113 155	95 10 98	149 118 77
Si(4)	0.064(3) 0.093(3) 0.113(2)	75 76 170	58 52 54	168 78 92	0(9)	0.087(7) 0.124(6) 0.162(5)	85 35 124	56 134 117	170 99 94
0(1)	0.093(7) 0.109(6) 0.165(5)	107 51 136	86 56 34	141 129 90	0(10)	0.084(7) 0.121(6) 0.199(5)	87 44 133	164 106 89	53 132 115
0(2)	0.085(7) 0.119(5) 0.134(5)	51 103 138	155 85 115	67 29 73	К	0.112(2) 0.149(2) 0.160(2)	95 27 117	147 96 58	87 126 144
0(3)	0.105(6) 0.121(6) 0.144(5)	99 40 52	123 78 145	112 138 56	Na	0.109(3) 0.135(3) 0.172(3)	55 52 57	78 78 162	155 72 73
0(4)	0.079(8) 0.118(6) 0.205(5)	163 97 75	74 30 65	96 88 174					

The unit cell parameters given in Table 1 were used angles (° for indexing

* Root mean square thermal vibrations along the ellipsoid axes (A) and angles (°) between the crystallographic axes and the principal axes (U₁) of the vibration ellipsoid.

Atoms		Distances	Atoms	Angles
	0(2)	1 623 Å	0(3) = Si(1) = 0(4)	106.0
51(1) -	- 0(3)	1.025 A	O(5) = 31(1) = O(4)	114 6
	0(4)	1.023	0(3)	109 /
	0(5)	1.5/2		100.4
	0(9)	1.633	O(4) = S1(1) = O(5)	110.6
		1 (12	O(9)	111 2
Average	2	1.613	0(3) = 31(1) = 0(3)	111.5
Si(2) -	- 0(6)	1.622	O(6) - Si(2) - O(7)	113.4
	0(7)	1.578	0(8)	105.0
	0(8)	1.618	0(9)	102.9
	0(9)	1.630	O(7) - Si(2) - O(8)	114.5
			0(9)	111.4
Average	2	1.612	O(8) - Si(2) - O(9)	109.0
s;(3) .	- 0(2)	1 580	O(2) = Si(3) = O(4)	107.1
01(3)	0(4)	1.626	0(8)	114.3
	0(4)	1 614	0(10)	110.7
	0(8)	1 622	O(4) = Si(3) = O(8)	109 2
	0(10)	1.032	0(4) = 31(3) = 0(0)	107.9
		1 (12	O(10) = O(10)	107.5
Averag	e	1.013	0(8) = 51(3) = 0(10)	107.0
Si(4)	-0(1)	1.582	O(1) - Si(4) - O(3)	112.8
	0(3)	1.632	0(6)	112.7
	0(6)	1.627	0(10)	115.3
	0(10)	1,635	O(3) - Si(4) - O(6)	103.2
			0(10)	106.7
Averag	e	1.619	O(6) - Si(4) - O(10)	105.1
Cu. = 0	(2)	1 961	Si(1) = O(3) = Si(4)	137.2
00 0	(5)	1 978	Si(1) = O(4) = Si(3)	157.0
0	(7)	1 081	Si(1) = O(9) = Si(2)	135.5
0	(1)	1 990	Si(2) = O(6) = Si(4)	131.8
0	(1)	2 5/9	Si(3) = O(8) = Si(2)	160.1
0		2.349	Si(3) = O(10) = Si(4)	133.2
Na - O	(7)	2.386		
	(2)	2,409	O(1) - Cu - O(2)	85.8
0	(1)	2.505	0(5)	177.9
0	(5)	2 555	0(7)	94.7
0	(7!)	2 581	$O(2) = C_{11} = O(5)$	92.6
0	(10)	2.501	0(2) $0(3)$	167.3
0	(10)	2.000	$O(5) = C_{11} = O(7)$	87.2
L	(9)	2.754	O(1) = Cu = O(1)	75.8
V 0/	(6)	2 667	O(1) = C1 + O(1)	93 5
R - 0(4	2.00/	0(5)	102 9
00	4)	2./34	0(3)	02.2
00	2)	2.022	0(7)	20.0
00	(9)	2.844		
0(3)	2.9/3		
00	(2)	3.049		
0((0)	3.069		
00	(8)	3.168		

fenaksite and litidionite are isostructural and that the

X-Ray Data Collection

The intensities were collected from a crystal frag-

ment (dimensions: $0.02 \times 0.06 \times 0.08$ mm) using the

single crystal automatic diffractometer and $MoK\alpha$

radiation monochromatized by a flat graphite crystal.

A unique set of data was collected out to $2\theta = 60^{\circ}$ by

the $\theta - 2\theta$ scan mode with a symmetric scan range of

1° in 2 θ from the calculated scattering angle. The scan rate was 0.05°/sec. Processing of the data was carried

chemical formula of the latter is CuNaKSi₄O₁₀.²

 TABLE 5. Interatomic Distances and Principal Bond Angles in Litidionite*

out in the manner described by Davies and Gatehouse (1973) to yield values of F_0 and σF_0 .

The intensities of 2677 independent reflections were measured; of these 1450 have $F_o > \sigma F_o$ and were used in subsequent calculations. Three standard reflections, monitorized at three-hour intervals, showed no variation in intensity greater than 5 percent. No absorption nor extinction corrections were applied.

Crystal Structure Analysis

The crystal structure analysis was carried out by Patterson and Fourier methods and confirmed that litidionite and fenaksite are isostructural. The least squares isotropic refinement, carried out on the structure amplitudes with the program ORFLS (Busing, Martin, and Levy, 1962), reduced the conventional Rindex to 0.047. Three successive least squares cycles performed with anisotropic thermal parameters led to an R index of 0.032 for the 1450 observed reflections. At this stage the refinement was stopped as the shifts of the atomic parameters were less than one tenth the standard deviations.

The final atomic parameters are given in Tables 3 and 4; bond distances and angles are listed in Table 5. The observed and computed structure factors are compared in Table $6.^3$

Description and Discussion of the Structure

The crystal structure analysis confirmed that the structure of litidionite is similar to that of fenaksite (Golovachev *et al*, 1971) and showed no deviation from the adopted chemical formula CuNaKSi₄O₁₀. The basic structural feature of both minerals is a tubular silicate radical Si₈O₂₀⁸⁻ (Figs. 1 and 2). This radical is made up by the condensation of two vlasovite type chains (Voronkov and Pyatenko, 1962) which are formed by rings of four tetrahedra. The tubular chains are parallel to *c* and are interconnected by Cu and Na atoms. Potassium occurs in the large cavities existing in the pipe-like tetrahedral chains with K-O distances ranging from 2.67 to 3.17 Å (Table 7).

The coordination polyhedron of copper is a fairly regular square pyramid whose base is formed by four oxygen atoms at distances ranging from 1.96 to 1.99 Å. An oxygen atom 2.55 Å from copper is the vertex

² P. H. Ribbe and T. D. Kurtz carried out a microprobe analysis The of litidionite and found only small amounts (less than 0.5 wt percent) of Ca, Ti, and Fe, besides the elements given in the formula.

³ To obtain a copy of Table 6, order Document AM-75-004 from The Mineralogical Society of America, Business Office, 1909 K St., N.W., Washington, D.C. 20006. Please remit in advance \$1.00 for a copy of the microfiche.

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FIG. 1. The crystal structure of litidionite projected along [010]. The bonds terminated by an arrow refer to oxygen atoms O(1) and O(5) occurring in chains which are not shown in the figure.

of the pyramid. The distance to the next closest oxygen is 3.24 Å. In fenaksite the oxygen atoms of the base have Fe-O distances ranging from 1.98 to 2.16 Å while the apical oxygen is 2.37 Å from Fe.

The coordination polyhedron around sodium atoms is similar to that around copper but is more distorted. Besides the five bonds with oxygen atoms forming the distorted pyramid, sodium has two longer bonds with O(9) and O(10), respectively 2.85 and 2.93 Å.

Each coordination pyramid around copper is connected by edge sharing to another copper pyramid and to one sodium pyramid in such a way that pairs of pyramids around copper alternate with pairs of pyramids around sodium forming serrate chains parallel to c.

The Si–O bond lengths show a well marked difference between the distances of Si from bridging oxygen atoms and the distances of Si from nonbridging oxygens, as has been pointed out by Cruickshank (1961). In the former group the mean Si–O distance is 1.578 Å and in the latter it is 1.626 Å.

The balance of electrostatic charges computed with the method of Brown and Shannon (1973) is satisfactory. The sum of the electrostatic charges reaching each oxygen atom ranges from 1.84 to 2.20 v.u. It should be noted that the non-bridging oxygen atoms are underbonded (from 1.84 to 1.94 v.u.) while the bridging ones are overbonded (from 2.03 to 2.20v.u.).

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FIG. 2. The crystal structure of litidionite projected along [100]. The bonds terminated by an arrow refer to oxygen atoms O(1) and O(7) occurring in chains which are not shown in the figure. The potassium atoms have been omitted for clarity.

Blacksburg, who carried out the microprobe analysis of litidionite.

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