Hydrothermal synthesis of nukundamite and its crystal structure

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Abstract

Crystals corresponding to nukundamite were obtained by hydrothermal synthesis in a temperature gradient. The empirical formula of the synthetic mineral analyzed by electron microprobe is $Cu_{3,39}Fe_{0.61}S_4$. Its crystal data are: trigonal, space group $P\overline{3}m1$, a = 3.7830(2), c = 11.1950(8)Å; Z = 1. The crystal structure has been determined using counter-measured intensities. The characteristic S-S bond recognized in minerals such as covelline and pyrite is confirmed in the structure, and its bond length in the S₂ group is 2.062(20)Å. The metal atoms are in two non-equivalent sites and are tetrahedrally coordinated with sulfur.

Introduction

Nukundamite was recently described as a new mineral from the Undu mine, Nukundamu, Fiji by Rice *et al.* (1979). It had been referred to as idaite or Cu_5FeS_6 , a mineral which corresponds to a synthetic Cu–Fe sulfide with a composition of $Cu_{55}FeS_{65}$.

Idaite was originally described as a supergene sulfide mineral from the Ida mine, southwest Africa, by Frenzel (1958, 1959). He suggested that the composition of idaite was Cu_5FeS_6 , because the optical properties of the mineral were similar to those of a synthetic phase synthesized by himself and by Merwin and Lombard (1937), and the X-ray powder data of a 0003-004X/81/0304-0398\$02.00 synthetic and a natural specimen from Javis township, Ontario were similar. He also proposed that the synthetic phase has a hexagonal cell a = 3.90 and c =16.95Å. Later, Yund (1963), on the basis of synthetic experiments under dry conditions, found that the chemical composition of the synthetic phase is $Cu_{5.5x}Fe_xS_{6.5x}(=Cu_{3.38}Fe_{0.62}S_4)$ and its cell parameters are a = 3.77 and c = 11.18Å from the results indexed as a hexagonal cell. From the electron probe microanalyses, Levy (1967), Sillitoe and Clark (1969), and Constantinou (1975) determined that the chemical composition of naturally-occurring supergene idaite is Cu_3FeS_4 . On the other hand, both Frenzel and Ottemann (1967) and Clark (1970) have reported a

		Nutrient materials							Heating		
Run	No.	Cu	ight pe Fe	S	Assemblages	Mineralizer	(°C)	(kg/cm ²)	(days)	Products	
HHH	087	55.00	12.00	33.00	bnss+nk+py	5m NH4Cl	450	340	13	nk+bnss+cp	
HHH	085	55.00	12.00	33.00	bnss+nk+py	11	400	300	18	nk+bnss+cp	
HHH	027	56.50	14.50	29.00	bnss+cp+py	6m NH4Cl	350	300	9	nk+cp+py	
HHH	040	56.50	14.50	29.00	bnss+cp+py	5m NH4C1	350	300	11	nk+bnss+cp	
ннн	042	50.00	19.00	31.00	bnss+cp+py	11	350	310	14	nk+cp+bnss	
ннн	049	65.00	3.00	32.00	cv+bnss+nk	11	350	315	22	nk+cv+bnss	
ннн	053	35.00	28.00	37.00	bnss+cp+py	19	350	285	21	nk+cp+py	
ннн	059	65.00	3.00	32.00	cv+bnss+nk	и	300	320	34	nk+cv+bnss	
HHH	062	55.00	12.00	33.00	bnss+nk+py	11	300	290	23	nk+bnss+cp	
ннн	066	50.00	13.00	37.00	cv+nk+py	17	300	290	20	nk+cv+py	
ннн	086	55.00	12.00	33.00	bnss+nk+py	11	250	290	36	nk+dg+cp	

Table 1. Experimental results for the hydrothermal synthesis of nukundamite

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hypogene sulfide with the composition of Cu_5FeS_6 and with X-ray powder data in agreement with those of the synthetic phase reported by Yund (1963). Therefore, Sillitoe and Clark (1969) and Clark (1970) considered this to be a new mineral, which differed from supergene idaite originally described by Frenzel (1958).

Ottemann and Frenzel (1971) analyzed idaite from various localities, including the type specimen from the Ida mine, and concluded that naturally-occurring idaites have chemical compositions ranging from Cu_3FeS_4 to Cu_5FeS_6 . Similar analytical data were obtained by Goble and Smith (1973) and Briggs *et al.* (1977). However, Wang (1976) concluded that the X- ray powder data for supergene idaite described by Frenzel (1959) are different from those of the synthetic phase or hypogene idaite, and that the cell of the supergene phase is tetragonal, as originally suggested by Levy (1967). Ugarte and Burkin (1977) obtained idaite of composition Cu_3FeS_4 as leaching products of bornite in ferric sulphate solutions at $40^{\circ}C$ and determined the cell dimensions to be a =5.26 and c = 10.46Å, similar to those of chalcopyrite.

Thus there are two kinds of minerals described as idaite; one is supergene idaite with a composition of Cu_3FeS_4 , and another is a hypogene sulfide with a composition of Cu_3FeS_6 . Recently, Rice *et al.* (1979) studied both minerals and concluded that the hypo-



Fig. 1. Scanning electron micrographs for hydrothermally-synthesized nukundamite. Length of scale bar is 150 µm.

gene mineral differs from supergene idaite, and named it nukundamite as a new mineral.

Our specimen was produced by a hydrothermal method and corresponds to the nukundamite. The synthetic crystals are larger than the crystal of natural nukundamite studied by Rice *et al.* (1979) and therefore were suitable for a X-ray single crystal study. In this paper we describe the chemical composition, crystal data and structure of synthetic nukundamite.

Hydrothermal synthesis of nukundamite

Nukundamite was synthesized under hydrothermal conditions using a thermal gradient transport method similar to that described by Scott and Barnes (1971), Scott (1975), and Sugaki et al. (1975, 1976). Hydrothermal syntheses were performed in a gold tube (4 mm inside diameter, 50 mm long), with an aqueous chloride solution as the transporting agent. The results are given in Table 1 together with the compositions of the nutrient materials, the experimental conditions, and the synthesized phases. As described previously (Sugaki et al., 1975), nukundamite ("idaite") coexists with pyrite covelline, bornite, and chalcopyrite at 300° and 350°C. The assemblage nukundamite + chalcopyrite is observed in the temperature range from 250° to 450°C. This assemblage is in conflict with the bornite + pyrite assemblage produced in dry synthesis experiments (Yund and Kullerud, 1966) and has been discussed in detail in Sugaki et al. (1975). The hydrothermally-synthesized nukundamite occurs as thin, platy, reddishbrown, hexagonal crystals up to 2 mm in diameter (Fig. 1).

Under the ore microscope, it has very strong pleochroism changing from reddish-orange to yellowish-gray and it also shows very strong anisotropism in greenish-gray under crossed polars. The reflectance of synthetic nukundamite was measured with a Leitz microphotometer MPV II in air and oil. The values for R_e and R_o in air are as shown in Table 2. These values are close to those of natural nukundamite of Rice *et al.* (1979). Twinning with a twin plane of (001) and a perfect cleavage on the basal plane were observed.

Chemical composition and crystal data

The chemical composition of synthetic nukundamite crystals obtained by 11 runs shown in Table 1 was determined by electron probe microanalyses, using procedures similar to those described in Sugaki *et al.* (1975). The mean composition is Cu 56.8 ± 0.4 , Fe

Table 2. Reflectance values for synthetic nukundamite

λ (nm)	406	436	497	546	589	657
In air Ro	12.0	11.0	10.6	12.4	16.7	23.8
Re	21.3	21.3	22.0	22.1	22.2	25.1
In oil Ro	5.1	5.1	6.1	7.9	10.1	15.9
Re	9.8	9.8	11.1	11.2	11.5	15.3

9.0 \pm 0.3, S 33.8 \pm 0.2, total 99.6 weight percent. The empirical formula normalized to four sulfur atoms is Cu_{3,39 \pm 0.02}Fe_{0.61 \pm 0.02}S₄. This chemical formula is in good agreement with Cu_{5.5}FeS_{6.5}(=Cu_{3.38}Fe_{0.62}S₄) obtained by Yund (1963) in his dry experimental study and with Cu_{3.37}Fe_{0.66}S_{3.97} for natural nukundamite reported by Rice *et al.* (1979).

Cell parameters and the space group were determined from precession and Weissenberg photographs and intensity data measured with an automatic four-circle X-ray diffractometer using MoKa and $CuK\alpha$ radiations. The crystals are generally distorted and twinned with c as the twin axis, and the twinned crystals showed a hexagonal symmetry. Therefore, most of them were not suitable for singlecrystal investigation. However, after careful study a small crystal which exhibited almost no distortion in precession photographs was found and was used for single-crystal analysis. The Laue group of this crystal was determined to be $\overline{3}m$, because the principal axis shows three-fold symmetry and a set of mirror planes of {110} is found in the reciprocal lattice reproduced from three-dimensional diffraction data by the precession and Weissenberg method. These results differ from Rice et al. (1979), who considered the mineral to be hexagonal from the Laue photographs of inferior crystals. No systematic extinction was found. This indicates that the possible space groups of nukundamite are $P\overline{3}m1$, P321, or P3m1. The space group is uniquely determined to be P3m1 by the results of structural analysis.

X-ray powder diffraction data were measured by a Guinier camera with $CuK\alpha_1$ radiation. These results agree with those obtained by Yund (1963) and Rice *et al.* (1979). The cell dimensions calculated by the least-squares method from the powder data at 25°C are a = 3.7830(2), c = 11.1950(8)Å, V = 138.75(2)Å³. These values are in good agreement with a = 3.782(4) and c = 11.187(8)Å for natural nukundamite determined by Rice *et al.*

The density of nukundamite, measured by Berman density balance using carbon tetrachloride, is 4.49(3) g/cm³. This value is in good agreement with the calculated value of 4.53 g/cm³ for Z = 1.

Aton	ns	х	Y	Z	в(Å ²)
M(1)	Cu, Fe	2/3	1/3	0.1547(4)	0.87(7)
M(2)	Cu, Fe	2/3	1/3	0.5790(6)	1.92(13)
S(1)		0	0	0.0921(9)	1.08(14)
S(2)		2/3	1/3	0.3609(9)	0.96(11)

Table 3. Atomic parameters of nukundamite, Cu_{3.39}Fe_{0.61}S₄

Structural analysis

We selected the most suitable crystal for the structural analysis, but some distortion from handling was unavoidable. The single crystal of nukundamite for intensity measurements has a platy form with dimensions of $0.1 \times 0.15 \times 0.005$ mm. The mineral is so soft and deformable that it is difficult to obtain a distortion-free crystal. Intensities were measured on an automatic four-circle diffractometer (Philips PW 1100), using graphite monochromated MoK α radiation. The ω -2 θ scan mode was used to a 2 θ limit of 120° and scan widths were varied by ω angles. 898 independent reflections were measured, and 544 of them with values less than 2σ were omitted in the least-squares refinement. Corrections were made for Lorenz and polarization factors but not for absorption. The neutral atomic scattering factors were taken from the International Tables for X-ray Crystallography (1962).

A three-dimensional Patterson synthesis was made using observable F_o values from 354 reflections. Patterson vectors revealed on the map can be solved assuming metal and sulfur atoms to be located on (0, 0, z) and (2/3, 1/3, z). From the relative peak height and crystallochemical considerations, sulfur atoms were located on two sites of (0, 0, z) and (2/3, 1/3, z)and metal atoms were located on $(2/3, 1/3, z_1)$ and $(2/3, 1/3, z_2)$. Each z parameter was determined by a trial and error method using the Patterson peaks.

Preliminary structural calculations assuming a center of symmetry and using z parameters derived from Patterson maps gave a R value of 0.30. Several cycles of least-squares calculations by the full-matrix, least-squares program (UNICS) were made, varying atomic positions and isotropic temperature factors

Table 4. Interatomic distances and bond angles for nukundamite

Interatomic distances				Bond angles			
M(1)-S(1)	x 3	2.294(13)	3	S(1)-M(1)-S(2)	107.8(5) °		
M(1)-S(2)		2.308(11)		S(1) - M(1) - S(1)	111.0(2)		
M(2)-S(2)		2.442(12)		S(2)-M(2)-S(2)	107.1(5)		
M(2)-S(2)	х3	2.285(13)		S(2) ~-M(2) -S(2) ~	111.7(2)		
M(2)-M(2)'		2.811(12)					
S(1)-S(1)'		2.062(20)					

for each atom in order to distinguish between copper and iron atoms. A final R value of 0.174 was obtained for 354 independent reflections. Three-dimensional Fourier and difference Fourier syntheses using the final atomic parameters shown in Table 3 were made. Because no unusual features were observed in the Fourier maps, the structure is believed to be correct and to justify the space group $P\overline{3}m1$. The Rvalue is higher than would be desirable which is probably a consequence of lack of corrections for absorption, anomalous dispersion, and site occupancy between copper and iron atoms.

Description of the crystal structure

Interatomic distances and bond angles are given in Table 4 and the structure is shown in Figure 2. The structure is characterized by pairs of S_2 and MS_4 tetrahedra. The S_2 group is surrounded by six metal atoms, and its S–S distance is 2.062(20)Å. This value agrees with the mean S–S distance (2.06Å) in orthorhombic sulfur (Caron and Donohue, 1965), but is slightly shorter than the S–S bond length of 2.071 or 2.086Å of covelline as given by Evans and Konnert (1976) or Ohmasa *et al.* (1977). S–S bonds are also found in other sulfide structures such as FeS₂, CoS₂,



Fig. 2. The crystal structure of nukundamite.

and NiS₂ (Elliott, 1960) and Cu₄Bi₄S₉ (Takéuchi and Ozawa, 1975). Sulfur is in tetrahedral arrangement around M(1) and M(2). Two tetrahedrons of M(1)- S_4 and $M(2)-S_4$ share a corner and their base S_3 triangles are oriented opposite to each other. The bond lengths of $M(1)-S(1)_3$ (2.294Å) and M(1)-S(2)(2.308Å) are slightly shorter than those for most common sulfide or sulfosalt minerals, 2.33-2.37Å (Ohmasa et al., 1977). The $M(2)-S(2)_3$ bond length (2.285\AA) is slightly shorter than M(1)-S(1) (2.294Å) and the bond length of M(2)-S(2) (2.442Å) is the longest of the M-S bond groups in the nukundamite structure. The $M(2)-S_4$ tetrahedron is distorted from the regular tetrahedron. Two $M(2)-S_4$ tetrahedra share an edge in the structure and the M(2)-M(2)distance is shortened to 2.811Å. A similar structure is found, for example, with respect to the Fe-Fe distance of 2.804Å in cubanite (CuFe₂S₃) (Szymanski, 1974).

Although the crystal structure of hydrothermallysynthesized nukundamite has been determined as mentioned above, we do not understand why the ratio of copper and iron atoms has a very constant value of 3.39 to 0.61, and so more detailed work on the structural analysis would be helpful. X-ray single-crystal study of supergene idaite has not yet been made. The X-ray powder data for idaite by Frenzel (1960) are distinctly different from our data for synthetic nukundamite, and more study on the composition and crystal structure of idaite is indicated.

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