# An attempt to simulate high pressure structures of Mg-silicates by an energy minimization method

## MASAMICHI MIYAMOTO<sup>1</sup>

NASA, Johnson Space Center Houston, Texas 77058

### and Hiroshi Takeda

## Mineralogical Institute, Faculty of Science University of Tokyo, Hongo, Tokyo 113, Japan

#### Abstract

A computer program that refines unknown repulsive parameters of ions or structural parameters of a crystal by minimizing the potential energy has been applied to simulating the crystal structures of high pressure minerals. The program (WMIN, Busing, 1981) is written under the assumption of Born-Mayer type repulsive forces. The repulsive parameters, ionic radius A and ionic compressibility B for Mg<sup>2+</sup>, Si<sup>4+</sup>, and O<sup>2-</sup> ions obtained from the structural data of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> (forsterite), have been applied to simulating the crystal structures of  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub>,  $\gamma$ -Mg<sub>2</sub>SiO<sub>4</sub>, hypothetical Mg<sub>2</sub>SiO<sub>4</sub> with Sr<sub>2</sub>PbO<sub>4</sub> and CaFe<sub>2</sub>O<sub>4</sub> structures, and MgSiO<sub>3</sub> with perovskite and ilmenite structures. The results were compared with the observed structures refined by the X-ray diffraction data. We also tried to predict the crystal structures of the minerals *in situ* at high pressure under the constraint that the cell dimensions are fixed at the observed values at high pressure. Our approach appears to be moderately successful in constructing the "frameworks" of relatively complex crystal structures of high pressure minerals.

#### Introduction

High pressure minerals give us information about the structure and evolution of the Earth's mantle. It is important to study high pressure phase transformations of these minerals on the basis of more detailed crystal structures. Many workers have carried out crystal structure analyses of these minerals under high pressure as well as 1 atm, and given their crystal chemical interpretations (e.g., Hazen, 1976; Hazen and Prewitt, 1977; Hazen and Finger, 1978, 1979; Akimoto et al., 1976; Finger et al., 1979; Levien and Prewitt, 1981). These approaches have the advantage of observing directly the crystal structure changes. However, we have often encountered difficulty in growing large enough single crystals for the structure refinements or in carrying out high pressure experiments. It is necessary to develop a method to predict the crystal structures semi-theoretically in order to supplement the experimental work.

The distance least-squares (DLS) method (Meier und Villiger, 1969) has been successful in simulating crystal

structures from ionic radii data and other additional constraints. For example, the crystal structures of nine different polymorphs of  $Mg_2SiO_4$  were simulated with the DLS method by Baur (1972), and Ito and Matsui (1978) predicted the perovskite-type structure of  $MgSiO_3$ . Thus, the crystal structures of high pressure minerals have been shown to be simulated from the ionic radii data alone.

Carrying this one step further, an attempt has been made to estimate the repulsive parameters of some ions in silicates from their crystal-structure data and to predict the crystal structures of mantle minerals (Miyamoto et al., 1979; Takeda et al., 1979; Miyamoto and Takeda, 1980) by making use of the WMIN program (Busing, 1981). The computer program WMIN was written by Busing (1970) to compute the "potential energy" of a crystal by assuming the Born-Mayer type repulsive energy as modified by Gilbert (1968) (Born and Huang, 1954; Tosi, 1964). This program enabled us to obtain repulsive parameters of ions from the structural parameters of a known structure, or to adjust the structural parameters of a hypothetical structure by employing the refined repulsive parameters thus obtained so that they give a minimum potential energy. We have already shown that the repulsive parameters defined as ionic compressibility and ionic radius can be determined from the structural parameters of a rela-

<sup>&</sup>lt;sup>1</sup> On leave of absence from Department of Pure and Applied Sciences, College of General Education, University of Tokyo, Komaba, Tokyo 153, Japan.

tively complex crystal structure such as  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> (forsterite), and that a complex crystal structure can be predicted by energy minimization using the program WMIN (e.g., Miyamoto and Takeda, 1980; Takeda et al., 1979). In this simulation each ion in the crystal structure occupies the position where the calculated potential energy is minimum.

The repulsive parameters of  $Mg^{2+}$ ,  $Si^{4+}$ , and  $O^{2-}$  and some transition metals ( $Mn^{2+}$ ,  $Fe^{2+}$ ,  $Co^{2+}$ , and  $Ni^{2+}$ ) in their olivine structures and those of  $Ca^{2+}$  in diopside and  $Al^{3+}$  in spinel obtained with the WMIN program have been reported by Miyamoto and Takeda (1980) and Miyamoto et al. (1982).

In this paper, we report an attempt to simulate the crystal structures of high pressure minerals by the WMIN program on the basis of the repulsive parameters of  $Mg^{2+}$ ,  $Si^{4+}$ , and  $O^{2-}$  ions obtained from the crystal structure data of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> (forsterite). These predicted structures have been compared with those obtained by X-ray crystal structure analyses to examine the differences between the simulated and observed structures and the limit of the ionic model for the mineral structures. Then, the method was applied for prediction of the crystal structure at high pressure by employing the cell dimensions measured *in situ* at high pressure.

#### Estimation of repulsive parameters

The expression for the potential energy used in this study is

$$W = \frac{1}{2} \sum_{i j} \left[ \frac{Q_i Q_j}{r_{ij}} + P_{ij} \cdot f_0 \cdot (B_i + B_j) \\ \cdot \exp\left(\frac{A_i + A_j - r_{ij}}{B_i + B_j}\right) \right], \quad (1)$$

where  $r_{ij}$  is the distance between the two ions;  $Q_i$  is ionic charge (Busing, 1970, 1981). The second term in square brackets represents the repulsive energy.  $A_i$  is an ionic radius and  $B_i$  is an ionic compressibility (softness parameter). A constant  $f_0$  is a standard force arbitrarily chosen to have a numerical value of one and dimensions kcal/ (mol  $\cdot$  Å) (Busing, 1970). This expression for the repulsive term is discussed in more detail in Gilbert (1968), Busing (1970, 1981), and Ida (1976).  $P_{ij}$  is a Pauling coefficient (Pauling, 1960), which is expressed as,

$$P_{ij} = 1 + \frac{Q_i}{N_i} + \frac{Q_j}{N_j},$$

where  $N_i$  is the number of outer electrons contained in the ith ion. When the crystal is a stable phase,  $\partial W/\partial X_j = 0$  for all structural parameters  $X_j$ 's (i.e.,  $X_j$ 's are the cell dimensions and coordinates of ions) is expected. Therefore, we can determine the repulsive parameters by minimizing the following function D by the nonlinear least-squares method

$$D = \sum_{j} \left\{ \frac{\partial W}{\partial X_{j}} \right\}^{2}$$
(2)

The details of the procedure are described in Busing (1970, 1981) and Miyamoto and Takeda (1980). We assumed complete ionicity in this study, and thus the ionic charges for Mg, Si, and O are taken as +2, +4, and -2, respectively. We neglected the effect of the thermal and van der Waals energies to simplify the problems for the interpretation of the ionic structures of the minerals because these energies might cause additional sources of ambiguity at the present stage of development (Ida, 1976).

The repulsive parameters A and B for  $Mg^{2+}$ ,  $Si^{4+}$ , and  $O^{2-}$  ions have been estimated by employing the structural parameters of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> (forsterite) obtained from single-crystal X-ray diffraction data (Smyth and Hazen, 1973). We designate A for Mg and B for Mg as  $A_{Mg}$  and  $B_{Mg}$ , respectively, and so on. The reported repulsive parameters A and B of Mg, Si, and O ions (Miyamoto and Takeda, 1980) were initially obtained by fixing  $A_{Mg}$  and  $B_{Mg}$  at the values 0.97 and 0.065, respectively, which were obtained for MgCl<sub>2</sub> by Busing (1970) with the WMIN program, and by varying the remaining 4 parameters ( $A_{Si}$ ,  $B_{Si}$ ,  $A_O$ , and  $B_O$ ) simultaneously to convergence (Table 1, solution N1).

Using solution N1 as initial values, all 6 repulsive parameters for Mg, Si, and O are varied simultaneously to converge on the basis of the structural data of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub>. The weight matrix calculated from the standard errors of the observed structural parameters (Smyth and Hazen, 1973) is included for the nonlinear least-squares (Busing, 1981). The results for the repulsive parameters are listed in Table 1 for both cases when the Pauling coefficient P<sub>ij</sub> is included (solution WP) and when P<sub>ij</sub> is excluded (i.e., P<sub>ij</sub> = 1; solution WN). The damping constant for the parameter change in the nonlinear least-squares method is 1.0 throughout our calculations.

# Prediction of crystal structures of high pressure minerals

Once the repulsive parameters in equation (1) are estimated for all kinds of ions contained in the crystal, we can simulate the crystal structure so that potential energy W in equation (1) gives the minimum value by adjusting structural parameters  $X_j$ 's. The method is described in more detail by Busing (1970, 1981).

Before attempting to simulate the structures of high pressure minerals, we have reconstructed the structure of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> (forsterite) with the WMIN method by using the repulsive parameters obtained from the forsterite structure in the preceding section to check our procedures. Table 2a shows the atomic coordinates of the reconstructed structures of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> compared with those obtained from the single-crystal X-ray diffraction data (Smyth and Hazen, 1973). The reconstructed structures are calculated for solutions, WN, WP,and N1.

713

WP+ WN N1 Mg<sup>2+</sup> 0.97\* 0.788 0.749 A<sub>i</sub> <sup>B</sup>i 0.0284 0.065\* 0.0218 Si4 A<sub>i</sub> 0.608 0.528 0.382 Bi 0.0172 0.0149 0.0065 0<sup>2-</sup> 1.770 1.789 1.848 A<sub>i</sub> Bi 0.105 0.104 0.102 P<sub>ij</sub> With P<sub>ij</sub> Without f<sub>0</sub>= 1 kcal/(mol·Å) A<sub>i</sub>: Ionic radius (Å); B<sub>i</sub>: Ionic compressibility (Å) Weighted; \* Fixed, Busing (1970)

Table 1. Repulsive parameters of  $Mg^{2+}$ ,  $Si^{4+}$ , and  $O^{2-}$  for

solutions N1, WN, and WP

P<sub>ii</sub>:Pauling coefficient

There are no significant differences in the atomic coordinates between the reconstructed structures by using WN and that by WP. The average value of deviations of the atomic coordinates by N1 from the observed ones is the largest among the three reconstructed structures. Also, the maximum value of the deviations of bond lengths by N1 from the observed values is the largest among the three reconstructed structures. However, the mean value of bond lengths in coordination polyhedra in the N1 structure agrees best with the observed one among the three reconstructed structures (Table 2b).

Table 2a. Calculated positional parameters ( $\times$  10<sup>4</sup>) of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> and comparison with the observed ones

		WN	WP	N1	N7*	Obs.**	120kb <sup>+</sup>
M2	х	9826[-89]	9827[-88]	9814[-101]	9782[-133]	9915(2)	9784
	У	2766[- 8]	2765[- 9]	2800[ 26]	2819[ 45]	2774(1)	2754
Si	х	4265[ 3]	4265[ 3]	4344[ 82]	4325[ 63]	4262(1)	4294
	У	0954[ 14]	0953[ 13]	0974[ 34]	0970[ 30]	0940(1)	0993
01	x	7602[-55]	7607[-50]	7717[ 60]	7728[ 71]	7657(3)	7743
	У	0880[-33]	0879[-34]	0863[- 50]	0868[- 45]	0913(2)	0874
02	x	2143[-72]	2137[-78]	2121[- 94]	2127[- 88]	2215(4)	2134
	У	4485[ 11]	4486[ 12]	4492[ 18]	4492[ 18]	4474(2)	4444
03	х	2851[ 74]	2855[ 78]	2880[ 103]	2883[ 106]	2777(2)	2743
	у	1618[-10]	1620[- 8]	1635[ 7]	1622[- 6]	1628(1)	1652
	z	0332[ 1]	0332[ 1]	0278[- 53]	0286[- 45]	0331(2)	0244
Aver	age	of deviation	s ( calc	obs.[/n)			
		0034	0034	0057	0059		
Cell	dim	ensions (Å)					
	а	4.777[21]	4.776[20]	4.799[ 43]	4.756*	4.756	4.652
	b	10.199[-8] 1	0.199[-8] 1	0.141[-66] 1	0.207*	10.207	9.781
	с	5.986[ 6]	5.987[7]	5.911[-69]	5.980*	5.980	5.795
Cell	vol	ume (Å <sup>3</sup> )					
		291.6[13]	291.6[13]	287.7[-26]		290.3	263.7

Cell dimensions are fixed at values calculated by Kumazawa and Anderson (1969); Repulsive parameters from N1 ]: Deviation from the observed value

Deviation and standard error refer to the last digit.

Table	2b.	Comparison of bond lengths between calculated an	nd
		observed structures of $\alpha$ -Mg <sub>2</sub> SiO <sub>4</sub>	

	WN	N1	N1*	Obs.**	120kb(N]) <sup>+</sup>
Si-01	1.595[-20]	1.623[ 8]	1.622[ 7]	1.615	1.608
-03[x2]	1.612[-23]	1.634[- 1]	1.633[- 2]	1.635	1.626
-02	1.643[-10]	1.660[ 7]	1.659[ 6]	1.653	1.655
Mean	1.616[-19]	1.638[ 3]	1.637[ 2]	1.635	1.629{~ 6}
02-03[x2]	2.534[-19]	2.539[- 14]	2.545[- 8]	2.553	2.526
03-03	2.595[ 1]	2.627[ 33]	2.648[ 54]	2.594	2.614
01-02	2.676[-67]	2.707[ 36]	2.702[- 41]	2.743	2.665
01-03[x2]	2.720[-37]	2.780[ 23]	2.767[ 10]	2.757	2.775
Mean	2.630[-30]	2.662[ 2]	2.662[ 2]	2.660	2.647{- 13}
		Octahedron	MI		
M1-01[x2]	2.087[ 2]	2.037[- 48]	2.046[- 39]	2.085	1.983
-02[x2]	2.093[ 25]	2.088[ 20]	2.091[ 23]	2.068	2.043
-03[x2]	2.149[ 17]	2.165[ 33]	2.156[ 24]	2.132	2.064
Mean	2.110[ 15]	2.097[ 2]	2.098[ 3]	2.095	2.030{- 65}
02-03[x2]s 01-02[x2]s 01-03[x2]s 01-02[x2] 01-02[x2] 01-03[x2] 02-03[x2] Mean	2.534[-19] 2.886[39] 2.922[68] 3.024[1] 3.068[-36] 3.402[66] 2.973[20]	2.539[- 14] 2.841[- 6] 2.780[- 74] 2.991[- 32] 3.032[- 72] 3.412[ 76] 2.933[- 20]	2.545[- 8] 2.821[- 26] 3.026[ 3] 3.026[ 3] 3.052[- 52] 3.400[ 64] 2.956[ 3]	2.553 2.847 2.854 3.023 3.104 3.336 2.953	2.526 2.764 2.775 2.928 2.947 3.238 2.863{- 90}
		Octahedron	M2		
M2-03[x2]	2.039[-28]	1.971[- 96]	1.979[- 88]	2.067	1.941
-02	2.073[ 22]	2.042[- 9]	2.040[- 11]	2.051	1.981
-01	2.197[ 15]	2.207[ 25]	2.218[ 36]	2.182	2.069
-03[x2]	2.267[ 53]	2.299[ 85]	2.328[ 114]	2.214	2.183
Mean	2.147[ 14]	2.132[- 1]	2.145[ 12]	2.133	2.050{- 83}
03-03 s	2.595[ 1]	2.627[ 33]	2.648[54]	2.594	2.614
01-03[x2]s	2.922[ 68]	2.912[ 58]	2.890[36]	2.854	2.775
02-03[x2]	2.888[-42]	2.854[-76]	2.854[-76]	2.930	2.802
03-03[x2]	3.016[ 19]	2.991[-6]	2.998[1]	2.997	2.871
01-03[x2]	3.066[ 38]	3.032[ 4]	3.057[29]	3.028	2.895
02-03[x2]	3.216[ 23]	3.202[ 9]	3.235[42]	3.193	3.040
03-03	3.391[ 5]	3.284[-102]	3.332[-54]	3.386	3.181
Mean	3.017[ 18]	2.991[- 8]	3.004[5]	2.999	2.880{-119}

Cell dimensions are fixed at observed values.

Smyth and Hazen (1973) Cell dimensions are fixed at values calculated by Kumazawa and Anderson (1969)

Shared

Deviation from the observed value Deviation from the observed value (at 1 atm)

Deviations refer to the last digit.

Table 3a shows the simulated structures (calculated structure or W-structure) for  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub> (modified spinel structure) with WMIN using the repulsive parameters of Mg, Si, and O ions obtained from the  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> structure. In these refinements except for the N1\* procedure, the cell dimensions have been varied simultaneously with the atomic coordinates. The W-structures are obtained for solutions of repulsive parameters, WN and N1, and are compared with the structure (X-structure) obtained from the single-crystal X-ray diffraction data (Horiuchi and Sawamoto, 1981). The initial values of atomic coordinates for  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub> were taken from those for  $\beta$ -Co<sub>2</sub>SiO<sub>4</sub> (Morimoto et al., 1974). Table 3b shows the comparison of bond lengths of coordination polyhedra between the W-structures and X-structure. Although the maximum deviation of bond lengths in the W-structure for N1 from those in X-structure is larger than that for WN, the mean bond-lengths of coordination polyhedra for N1 are closer to those of the X-structure than those for WN (Table 4). No significant difference was noted in the calculated structures and bond lengths between WN and WP. We will use repulsive parameter N1 for calculating W-structures hereafter, because the mean bond-lengths of coordination polyhedra agree best with those of the X-structure

Table 3a. Calculated positional parameters (× 10<sup>4</sup>) of  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub> and comparison with the observed ones

		WN	N1		N1*		Obs.**	120kb <sup>+</sup>	180kb <sup>+</sup>
Mg2	z	9841[ 140]	9906[	205]	9912[	211]	9701(4)	9910	9918
Mg3	У	1257[- 19]	1282[	6]	1280[	4]	1276(2)	1281	1281
Si	у	1220[ 22]	1207[	9]	1206[	8]	1198(2)	1200	1198
	z	6185[ 17]	6196[	28]	6191[	23]	6168(2)	6212	6218
01	z	2285[ 119]	2296[	130]	2283[	117]	2166(6)	2329	2353
02	z	7149[- 15]	7155[-	9]	7140[-	- 24]	7164(6)	7231	7270
03	У	9879[- 21]	9919[	19]	9918[	18]	9900(3)	9949	9959
	z	2586[ 28]	2573[	15]	2579[	21]	2558(5)	2544	2526
04	х	2651[ 36]	2602[-	13]	2606[-	- 9]	2615(4)	2585	2576
	У	1229[ 4]	1228[	3]	1226[	1]	1225(3)	1232	1235
	z	9912[- 13]	9921[-	4]	9919[-	- 6]	9925(3)	9924	9933
Aver	age (	of deviations 0039	( cald 0040	(	0040	)			
Ce11	dime	ensions (Å)							
	a	5.728[ 30]	5.686[-	12]	5.6983	k	5.6983(4)	5.5684	5.5034
	р.	11.484[ 46] 1	1.452[	14] 1	1.4380	k	11.4380(7)	11.0949	10.9233*
	с	8.212[-45]	8.182[-	75]	8.2566	ł	8.2566(8)	8.0882*	8.0039
Cell	volu	ume (Å <sup>3</sup> )							
		540.2[ 21]	532.8[-	·53]			538.14	499.70	481.16
Mg1 Mg3 * C ** O + C ( ] Devi	(0,0 with ell o bserv ell o 1975 : Dev ation	,0); Mg2, 01 x=z=1/4; Sp dimensions ar ved value, Ho dimensions ar ); Repulsive viation from n and standar	, 02 w ace gro riuchi re fixed parame the obs	ith x= oup: 1 l at c & Saw l at v eters served refe	0 and 1 Imma observed vamoto values o from N 1 value er to t	y=1/4 1 val (1981 estim } he la	; Si, O3 w ues. ); ( ): S ated from M st digit.	with x=0; Standard e Mizukami e	error et al.

among the three solutions of repulsive parameters. Table 5 shows the results of the W-structure refinements of  $\gamma$ -Mg<sub>2</sub>SiO<sub>4</sub> (spinel structure) (Matsui et al., 1979).

Table 6 shows the data of the W-structure of a hypothetical structure which is isostructural with Sr<sub>2</sub>PbO<sub>4</sub>. This high pressure polymorph of Mg<sub>2</sub>SiO<sub>4</sub> is proposed to be isomorphous with  $\delta$ -Mn<sub>2</sub>GeO<sub>4</sub> (Morimoto et al., 1972). The initial values of atomic coordinates for this Wstructure were taken from those of  $\delta$ -Mn<sub>2</sub>GeO<sub>4</sub> (Morimoto et al., 1972). Table 6 also shows the data for another hypothetical structure isomorphous with CaFe<sub>2</sub>O<sub>4</sub>. The initial values of atomic coordinates were taken from Prof. Y. Matsui (pers. comm., 1980).

We tried to simulate the structures of high pressure phases of MgSiO<sub>3</sub> on the basis of the repulsive parameters of Mg, Si, and O ions obtained from the forsterite structure. Table 7 shows the data of the W-structure of MgSiO<sub>3</sub>-ilmenite. The results are compared with those obtained by the single-crystal X-ray diffraction method (Horiuchi et al., 1982). As initial values for MgSiO3ilmenite ideal atomic coordinates of FeTiO3 were employed. Table 8 shows the data of the W-structure of MgSiO<sub>3</sub>-perovskite compared with those obtained from the X-ray powder diffraction data (Ito and Matsui, 1978; Matsui, 1982). The initial values of atomic coordinates for MgSiO<sub>3</sub>-perovskite were the same as those obtained by Ito and Matsui (1978).

Table 3b. Comparison of bond lengths between calculated and observed structures of B-Mg2SiO4

	NI	N1 *	0bs.**	120kb(N1) <sup>+</sup>	180kb(N1)+
Si -04[x2]	1.642[ 11]	1.643[ 12]	1.631(2)	1.629	1.621
-03	1.637[- 1]	1.638[ 0]	1.638(4)	1.624	1.615
-02	1.676[- 26]	1.675[- 27]	1.702(3)	1.661	1.653
Mean	1.649[- 2]	1.650[- 1]	1.651	1.636{- 15}	1.628{- 23}
02 -04[x2]	2.620[- 16]	2.622[- 14]	2.636(4)	2.612	2.607
03 -04[x2]	2.697[- 9]	2.703[- 3]	2.706(4)	2.689	2.681
04 -04	2.727[ 9]	2.728[ 10]	2.718(4)	2.690	2.668
02 -03	2.779[ 25]	2.775[ 21]	2.754(3)	2.723	2.691
Mean	2.690[- 3]	2.692[- 1]	2.693	2.669{- 24}	2.656{- 37}
		Octahedron	Mg1		
Mg1-04[x4]	2.042[- 4]	2.044[- 2]	2.046(3)	1.986	1.958
-03[x2]	2.108[- 7]	2.131[ 16]	2.115(4)	2.058	2.023
Mean	2.064[- 5]	2.073[ 4]	2.069	2.010{- 59}	1.980{- 89}
04 -04[x2]	2.816[ 11]	2.808[3]	2.805(6)	2.737	2.700
03 -04[x4]	2.843[ 1]	2.859[17]	2.842(4)	2.788	2.754
04 -04[x2]e	2.959[- 22]	2.970[-11]	2.981(4)	2.878	2.836
03 -04[x4]e	3.024[- 16]	3.045[5]	3.040(4)	2.931	2.874
Mean	2.918[- 7]	2.931[6]	2.925	2.842{- 83}	2.799{-126}
		Octahedron	Mg2		
Mg2-01	1.955[- 80]	1.958[- 77]	2.035(6)	1.956	1.949
-04[x4]	2.076[- 17]	2.081[- 12]	2.093(3)	2.012	1.980
-02	2.251[ 156]	2.289[ 194]	2.095(6)	2.167	2.119
Mean	2.085[ 1]	2.095[ 11]	2.084	2.029{- 55}	1.998{- 86}
01 -04[x4]e	2.844[ 56]	2.853[ 65]	2.788(4)	2.799	2.770
04 -04[x2]	2.913[- 4]	2.914[- 3]	2.917(6)	2.813	2.764
04 -04[x2]e	2.959[- 22]	2.970[- 11]	2.981(4)	2.878	2.836
02 -04[x4]	3.071[- 18]	3.097[ 8]	3.089(4)	2.965	2.909
Mean	2.950[ 8]	2.964[ 22]	2.942	2.870{- 72}	2.826{-116}
		Octahedron	Mg3		
Mg3-01[x2]	1.999[- 18]	2.002[- 15]	2.017(2)	1.946	1.919
-03[x2]	2.112[- 11]	2.112[- 11]	2.123(3)	2.030	1.994
-04[x2]	2.112[- 17]	2.133[ 4]	2.129(2)	2.085	2.056
Mean	2.074[- 16]	2.082[- 8]	2.090	2.020{- 70}	1.990{-100}
01 -04[x2]e	2.844[ 56]	2.853[ 65]	2.788(4)	2.799	2.770
03 -03	2.846[ 5]	2.852[ 1]	2.851(1)	2.785	2.752
01 -01 e	2.863[ 39]	2.872[- 30]	2.902(2)	2.798	2.762
03 -04[x2]	2.883[ 26]	2.893[- 16]	2.909(4)	2.834	2.803
01 -03[x2]e	2.965[ 27]	2.964[- 28]	2.992(3)	2.836	2.779
03 -04[x2]e	3.024[ 16]	3.045[ 5]	3.040(4)	2.931	2.874
03 -04[x2]	3.028[ 93]	3.053[- 68]	3.121(5)	2.954	2.900
Mean	2.933[ 21]	2.945[- 9]	2.954	2.858{- 96}	2.814 {-140}
<pre>* Cell dime + Cell dime (1075);</pre>	nsions are fi nsions are fi	xed at observ xed at values	ed values. estimated	from Mizuka	ni <u>et al.</u>

Horiuchi & Sawamoto (1981); ( ): Sta ]: Deviation from the observed value; ): Standard error

e: Edge shared

Deviation and standard error refer to the last digit.

In order to examine the crystal structure change with pressure, we tried to simulate the structures at high pressure. Because the cell dimensions are easier to measure in situ at high pressure than the single-crystal X-ray diffraction intensities, the compressibility of cell dimen-

Table 4. Maximum values of deviations between calculated and observed structures

	α−Mg	Si04	B-Mg	Si04
	WN	N1	WN	N1
Fractional coordinate	0.0089	0.0103	0.0140	0.0205
Cell dimensions (Å)	0.021 (0.44)	0.069 (1.2)	0.046 (0.40)	0.075 (0.91)
Bond length of polyhedron	(Å)			
Si-O	0.023 (1.4)	0.008 (0.50)	0.032 (1.9)	0.026 (1.5)
Mg-0	0.053 (2.4)	0.096 (4.6)	0.116* (5.5)	0.156* (7.4)
0 -0	0.068 (2.4)	0.102 (3.0)	0.082 (2.6)	0.093 (3.0)
Mean	0.030	0.020 (0.68)	0.039 (1.3)	0.021 (0.71)

): The ratio in percent to the corresponding observed value

Table	5.	Calculate	d crystal	data	and	bond	lengths	for	Y
ľ	Mg <sub>2</sub>	SiO <sub>4</sub> and	comparis	on with	their	obser	ved value	es	

Table 7. Calculated crystal data and bond lengths for MgSiO<sub>3</sub>ilmenite and comparison with observed values

	N1	N1*	Obs.**	180kb(N1) <sup>+</sup>	240kb(N1) <sup>+</sup>
u	0.3687[2]	0.3685[0]	0.3685(6)	0.3705	0.3711
a (Å)	8.042[-23]	8.0649*	8.0649(1)	7.8396+	7.7645
V (Å <sup>3</sup> )	520.1[45]	524.56	524.56	481.82	468.10
Bond lengt	:h (Å)				
Si-0[x2] 0 -0[x6]	1.653[- 2] 2.700[- 3]	1.655[0] 2.703[0]	1.655(4) 2.703	1.636{- 19} 2.671{- 32}	1.628{- 27} 2.659{- 44}
Mg-0[x6]	2.063[- 7]	2.070[0]	2.070(4)	1.996{- 74}	1.972{- 98}
0 -0[x6]s 0 -0[x6]	2.845[- 8] 2.987[-13]	2.853[0] 3.000[0]	2.853 3.000	2.773{- 80} 2.872{-128}	2.746{-107} 2.832{-168}
Space grou * Cell di Previou + Cell di (1975); ** Observe [ ]: Devi Deviation	p: Fd3m imensions are sly reported imensions are { }: Devia d value, Sass ation from the and standard	fixed at ob: by Miyamoto fixed at va ation from t aki <u>et al</u> .(1' he observed error refer	served value and Takeda lues estimat he observed 982); (): value; s: S to the last	s. (1980) ed from Mizuka value (at 1 af Standard erro hared digit.	ami <u>et al</u> . cm) or

sions is often available. We predicted the crystal structures in situ at high pressure by the WMIN method under the constraint that the cell dimensions are fixed at the observed values at high pressure. The estimated fractional coordinates and bond lengths are shown in Tables 2, 3, 5, 7, and 8. The cell dimensions of  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub> at high pressure were calculated from the elastic constants obtained by Kumazawa and Anderson (1969). The cell

Table 6. Calculated crystal data of two hypothetical polymorphs of Mg<sub>2</sub>SiO<sub>4</sub>

	Sr <sub>2</sub> I	b0 <sub>4</sub> -type	
	×	у	z
Mg	0.0225 (0.0819)	0.2978 (0.3329)	1/2
01	0.2061 (0.2164)	0.0756 (0.0448)	1/2
02	0.2990 (0.3925)	0.3384 (0.3087)	0
Cell dimensions (Å)	a	b	с
	4.862 (4.98)	9.241 (8.85)	2.806 (2.75)
Cell volume (Å <sup>3</sup> )	126.1 (121.2)		
Space group	Pbam, z=	=2; Siat (O	,0,0)
	Cal	<sup>Fe</sup> 2 <sup>0</sup> 4-type	
	×	У	z
Mg(VIII)	0.5709	0.7619	3/4
Mg(VI)	0.4274	0.5757	1/4
Si	0.8379	0.5931	1/4
01	0.4127	0.7506	3/4
02	0.1045	0.5805	3/4
03	0.2188	0.3966	1/4
	0.4304	0.4107	3/4
04			
O4 Cell dimensions (Å)	a	ь	с
O4 Cell dimensions (Å)	a 13.785	ь 8.196	с 2.894
O4 Cell dimensions (Å) Cell volume (Å <sup>3</sup> )	a 13.785 327.0	b 8.196	с 2.894

	N1	N1*	Obs.**	180kb(N1) <sup>+</sup>	240kb(N1)*
Mg z	3555[- 42]	3547[- 50]	3597(1)	3534	3531
Si z	1460[-117]	1495[- 82]	1577(1)	1523	1530
0 x	3331[ 117]	3305[ 91]	3214(5)	3279	3274
У	0313[- 48]	0286[- 75]	0361(4)	0246	0234
z	2413[ 5]	2415[ 7]	2408(1)	2428	2433
Average o	of deviations (	calc obs	.//n)		
	0066	0061			
Cell dime	ensions (Å)				
a	4.797[ 69]	4.7284*	4.7284(4)	4.6382	4.6081
b	4.797[ 69]	4.7284*	4.7284(4)	4.6382	4.6081
С	13.735[ 186]	13.5591*	13.5591(16)	13.0441+	12.8725
۷	273.7[ 112]		262.54	240.02	236.72
Bond leng	th (Å)	Si0,	octahedron		
Si-0[x3] -0[x3] Mean	2.012[ 182] 1.700[- 68] 1.856[ 57]	1.950[ 120] 1.710[- 58] 1.830[ 31]	1.830(2) 1.768(2) 1.799	1.883 1.708 1.796{- 3}	1.864 1.705 1.785{- 14}
0 -0[x3]e -0[x3]f -0[x3] -0[x3] Mean	2.427[ 96] 2.647[ 150] 2.650[- 3] 2.696[ 14] 2.605[ 64]	2.405[74] 2.598[101] 2.634[-19] 2.658[-24] 2.574[33]	2.331(4) 2.497(4) 2.653(3) 2.682(3) 2.541	2.370 2.541 2.597 2.609 2.529{- 12}	2.359 2.525 2.576 2.598 2.515{- 26}
		Mg0	octahedron		
Mg-0[x3] -0[x3] Mean	2.190[ 27] 1.995[ 5] 2.093[ 16]	2.146[- 17] 1.966[- 24] 2.056[- 21]	2.163(2) 1.990(2) 2.077	2.058 1.909 1.984{- 93}	2.031 1.888 1.960{-117}
0 -0[x3]f -0[x3] -0[x3]e -0[x3] Mean	2.647[150] 2.954[13] 3.030[-22] 3.084[8] 2.929[37]	2.598[101] 2.919[- 22] 2.976[- 76] 3.026[- 50] 2.880[- 12]	2.497(4) 2.941(3) 3.052(3) 3.076(4) 2.892	2.541 2.801 2.883 2.898 2.781{-111}	2.525 2.762 2.836 2.871 2.749{-143]
Mg, Si wi * Cell d	th x=y=0; Spanimensions are	ce group: R3	Positional erved values.	parameter: x	104

Observed value, Horiuchi <u>et al</u>.(1982); e:Edge shared; f:Face shared ] ]. Deviation from the observed value; (): Standard error }: Deviation from the observed value (at 1 atm)

Deviation and standard error refer to the last digit.

dimensions of  $\beta$ - and  $\gamma$ -Mg<sub>2</sub>SiO<sub>4</sub> at high pressure were estimated by extrapolating the values obtained in situ at about 90 kbar by Mizukami et al. (1975). By a similar extrapolation of the values after Yagi et al. (1982), the cell dimensions of MgSiO<sub>3</sub>-perovskite were estimated. The cell dimensions of MgSiO<sub>3</sub>-ilmenite were estimated by extrapolating the values obtained for ZnSiO<sub>3</sub>-ilmenite in situ at high pressure by Sato et al. (1977).

## Discussion

The main purpose of this paper is not to predict the "precise" crystal structures of high pressure minerals by the WMIN method. The precise structures should be determined by experimental studies (e.g., single-crystal X-ray diffraction). The purposes are: (1) to simulate the "framework" of the crystal structure of a relatively complex crystal by the WMIN method where each ion is located at the position which gives the minimum value of "potential energy" W in equation (1), (2) to find the mean deviations of the atomic coordinates thus obtained from those observed, (3) to reveal the limitations of this simple classical ionic model for the silicate minerals, and (4) to predict the structure in situ at high pressure, for cases in which the cell dimensions measured at high pressure are known.

Our WMIN studies appear to be successful in constructing the "framework" of the crystal structure, and show that each ion is located at the position which gives the minimum value of W in equation (1), even in a relatively complex crystal structure. As was shown in Miyamoto et al. (1982), repulsive parameter A is related to the traditional ionic radius (Shannon and Prewitt, 1969; Shannon, 1976).

Because the Si-O bond in silicate minerals is considered to show covalency (e.g., Matsui and Matsumoto, 1982), the fully ionized model used in this study may be unreasonable. Matsui and Matsumoto (1982) have pro-

Table 8. Calculated crystal data and bond lengths for MgSiO<sub>3</sub>perovskite and comparison with observed values

		11		N1*	Obs.**	240kb(N1)*	300kb(N1)*
Mg	x	5094[-	- 86]	5101[- 79]	5180(40)	5168	5186
	У	5253[·	-337]	5441[-149]	5590(30)	5599	5632
01	x	1053[	123]	1087[ 157]	0930(40)	1178	1200
	У	4727[	37]	4668[- 22]	4690(70)	4621	4612
02	x	1983[-	- 7]	1917[- 73]	1990(40)	1871	1860
	У	2021[	11]	1968[- 42]	2010(40)	1935	1928
	z	5522[-	- 58]	5550[- 30]	5580(30)	5591	5600
Aver	age of	deviat	ions (	calc obs.	/n)		
		0094		0079			
Cell	dimen	sions ()	<u>()</u>				
	a	4.848[	73]	4.7754*	4.7754(3)	4.6195	4.5806
	b	5.032[	103]	4.9292*	4.9292(4)	4.8074	4.7769
	c	7.097[	200]	6.8969*	6.8969(5)	6.7512	6.7148 <sup>+</sup>
	۷	173.1[	108]		162.35	149.93	146.93
Bond	lengt	h (Å)		S106	octahedron		
Si-0	1[x2]	1.851[	64]	1.808[ 21]	1.787	1.783	1.776
-0	2[x2]	1.819	201	1.793[- 6]	1.799	1.754	1.755
Mean	1	1.830[	38]	1.801 9]	1.792	1.767{- 25}	1.758{- 34}
02-0	][x2]s	2.576[	127]	2.535[ 86]	2.449	2.486	2.474
02-0	2[x2]s	2.581	161	2.520[- 32]	2.552	2.4/5	2.463
Mean		2.579	57]	2.538[ 16]	2.522	2.488{- 34}	2.475{- 47}
02-0	2[x2]	2.565	53]	2.527[ 15]	2.512	2.473	2.459
02-0	1[x2]	2.615	6	2.573[- 36]	2.609	2.527	2.519
Mean	1	2.597[	50]	2.557[ 10]	2.547	2.510{- 37}	2.498{- 49}
				Mg0 <sub>8</sub> -	square ant	iprism	
Mg-0 -0	1	2.319	2301	1.955[-122]	2.077	1.902	1.890
-0	2[x2]	2.030[-	- 5]	1.993[- 42]	2.035	1.940	1.926
Mean		2.089[	30]	2.025[- 34]	2.059	1.953{-106}	1.937{-122}
Mg~0	2[x2]	2.421	190]	2.338[107]	2.231	2.222	2.195
Mean	for 8	2.283[	79]	2.205 1]	2.204	2.127 (- 77)	2.108(- 96)
02-0	2[x2]	2.565[	53]	2.527[ 15]	2.512	2.473	2.459
02-0	02	2.807	1591	2.691 43	2.648	2.578	2.552
02-0	2[x2]	2.896	36]	2.774[- 86]	2.860	2.662	2.635
01-0	2[x2]	3.318	27]	3.196[- 95]	3.291	3.023	2.982
St (	1/2.0	1/2): 1	VJ. 01	with z=1/4.	Space group	· Dhom	2.313
1. 1		., - /	33 01	HIGH 2-1/49	share Aroch	• F D101	

xed at observed values.

Cell dimensions are fixed at observed values. Cell dimensions are fixed at values estimated from Yagi <u>et al</u>.(1982) Observed value, Powder X-ray data, Matsui (1982); s: Shared (): Standard error; Positional parameter: x 10' Standard errors of bond lengths are estimated to be less than 0.05Å ]: Deviation from the observed value At a standard error profers to the last digit

Deviation and standard error refer to the last digit

posed a model for olivine structures where the SiO4 ion is treated as a rigid body by a similar potential energy calculation. The deviations of the bond lengths in the simulated structure from those in the observed structure are almost the same in our model and that of Matsui and Matsumoto (1982). Our model may be able to simulate the length of the Si-O bond regardless of the specific nature of the bond. The structures of MgSiO<sub>3</sub>-ilmenite and MgSiO<sub>3</sub>-perovskite, where the Si atom is octahedrally coordinated are simulated approximately on the basis of the repulsive parameters of Si obtained from the olivine structure, where the Si atom is tetrahedrally coordinated.

The reasons why we use the expression of the repulsive parameters defined as in equation (1) are that (1) the repulsive parameters obtained in the WMIN program are directly applicable to the molecular dynamics (MD) calculations for the silicate structures (e.g., Matsui and Kawamura, 1980), and that (2) it may be possible to examine the change of the coordination number (of the Si atom from 4 to 6) in high pressure phase. This is because the repulsive parameters are defined as those proper to an individual ion. After we have examined in detail the limitations of the simple fully ionized model and of the formula of the repulsive force used in this study, further studies based on more precise models for minerals are required in the future.

There are some differences between the repulsive parameters which are calculated including the Pauling coefficient  $P_{ij}$  (WP) and those excluding  $P_{ij}$  (WN) (Table 1). The calculated coordinates from repulsive parameter WP are, however, almost identical with those from WN (Table 2).

The differences (maximum values) in  $\alpha$ - and  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub> between the W-structures (WN and N1) and the observed structures are summarized in Table 4. The deviation of bond-length Mg2–O2 in  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub> for two W-structures in extraordinarily large, but the difference from the observed value is 5.5% (WN) or 7.4% (N1). Thus, the calculated structures may be reproducible within an error of about 5 percent. The maximum value of the deviations of atomic coordinates from the observed ones is about 0.02 for all W-structures (in Table 4).

The sequence of the bond-lengths in the coordination polyhedra is almost always reproduced in the W-structures (Table 2b and 3b). Only the four pairs: M1-O1 and M1–O2, M2–O3 and M2–O2, and O1–O3 and O2–O3 in  $\alpha$ -Mg<sub>2</sub>SiO<sub>4</sub>, and Si-O4 and Si-O3 in  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub>, are in the reverse order in W-structures from the corresponding sequence in the observed structure.

Since Tokonami et al. (1972) have discussed the stability relations between the  $\alpha$ ,  $\beta$ , and  $\gamma$  polymorphs of Co<sub>2</sub>SiO<sub>4</sub> on the basis of their refined crystal structures at 1 atm and experimental results of phase equilibria, we will not discuss the stability relations of the three polymorphs of Mg<sub>2</sub>SiO<sub>4</sub> until the precise crystal structures in situ at high pressure have been obtained by single-crystal X-ray diffraction methods.

717

	α-Mg <sub>2</sub> SiO <sub>4</sub>	β-Mg <sub>2</sub> SiO <sub>4</sub>	γ-Mg <sub>2</sub> SiO <sub>4</sub>	MgSiO <sub>3</sub> ilmenite	MgSiO <sub>3</sub> perovskite
\$i-0**	0.031	0.082	0.082	0.10*	0.085*
Ma-0**	0.29	0.25	0.20	0.20	0.15+

 Table 9. The trend of shortening of interionic distances of Mg-silicates with pressure

\* Six-fold coordination; + Eight-fold coordination \*\* Mean

On the basis of the results of simulated structures by

WMIN, Matsui (1982) has discussed the stability relation of a hypothetical  $Mg_2SiO_4$  with the  $Sr_2PbO_4$  structure under high pressure and has also shown that the Wstructure of  $MgSiO_3$ -perovskite is consistent with his powder X-ray diffraction data. The cell volumes of the Wstructures for  $MgSiO_3$ -ilmenite and -perovskite at 1 atm have a tendency to show larger values than for the observed ones.

The changes of bond lengths with increase of pressure are summarized as follows (e.g., Hazen and Prewitt, 1977): (1) In general, degrees of shortening of larger bonds are larger than those of shorter bonds. More specifically, (2) in Mg-silicates, the Mg-O polyhedron shrinks more than the Si-O polyhedron in a structure; (3) Unshared edges become shorter at a higher rate than shared edges; and (4) The shape of a polyhedron becomes regular, because unshared edges of cation polyhedra tend to approach the lengths of the shared edges. In some cases, the shortest shared edge remains almost constant (Takeda et al., 1979).

One of the notable exceptions to the above principle is the longest tetrahedral lateral edge (O1-O3) in the olivine structure (Table 2b). Although this is the largest among the tetrahedral O-O edges, this length does not shorten as do the others.

Although it is well understood that the transformation of Mg<sub>2</sub>SiO<sub>4</sub> or MgSiO<sub>3</sub> from the  $\alpha$ -phase to the  $\gamma$ -phase and to the ilmenite and perovskite structures achieves more densely packed structures, we view the change of the structures with pressure in terms of bond distances, which are the subjects of crystal-chemical interest. According to the general principles of structural change or bond distance variation with pressure, the rates of shortening of the Mg-O distances with high pressure become small from the  $\alpha$ -phase to the perovskite-type structure but not for that of the ilmenite-type (Table 9). At each transition, the mean length of the Si-O tetrahedral bond shows a discontinous increase in going from a lower pressure phase to a higher pressure phase (Tables 2, 3, and 5). However, for each of the phases this bond length decreases at increasing pressure until the next transition. Furthermore, the rate at which the bond length decreases with pressure is significantly higher for the high pressure  $\gamma$ -phase compared to the low pressure  $\alpha$ -phase (Table 9). Finally at lower mantle pressure, the rates of shortening

of both the Mg–O and Si–O distances tend to approach each other (Table 9).

In conclusion, the WMIN methods enables us to simulate the "framework" of the crystal structures of minerals, and may give us some useful pieces of information about the crystal chemistry of high pressure minerals at high pressure, until the refinements of crystal structures in situ at high pressure are available. For example, this method may be useful in predicting the crystal structures of these minerals on the basis of the cell dimensions obtained from the powder diffraction data in situ at high pressure, because it is, at present, difficult to carry out single-crystal diffraction experiments in situ at pressures higher than 100 kbar. However, any detailed discussion of the crystal structure should be based on the results of the refinement in situ at high pressure by single-crystal Xray structure analyses, since the ionic model for the crystal structure used in this study is a very simple one.

#### Acknowledgments

We are indebted to Dr. W. R. Busing for the WMIN program, to Prof. Y. Takéuchi, Prof. Y. Matsui, and Prof. Y. Ida for helpful suggestions and discussions, and to Prof. Y. Takano and Prof. K. Ito for discussions and computer time made available for this study. The calculations were performed on the HITAC M-280H computer at the Computer Centre, University of Tokyo. We thank Prof. C. T. Prewitt, Prof. S. Ghose, Dr. D. S. McKay, and an anonymous reviewer for critical reading of the manuscript.

#### References

- Akimoto, S., Matsui, Y. and Shono, Y. (1976) High pressure crystal chemistry of orthosilicates and the formation of the transition zone. In R. G. J. Strens, Ed., The Physics and Chemistry of Minerals and Rocks, p. 327-363. J. Wiley, London.
- Baur, W. H. (1972) Computer-simulated crystal structures of observed and hypothetical Mg<sub>2</sub>SiO<sub>4</sub> polymorphs of low and high density. American Mineralogist, 57, 709–731.
- Born, M. and Huang, K. (1954) Dynamical Theory of Crystal Lattices. Oxford University Press, London.
- Busing, W. R. (1970) An interpretation of the structures of alkaline earth chlorides in terms of interionic forces. Transactions of American Crystallographic Association, 6, 57–72.
- Busing, W. R. (1981) WMIN, a computer program to model molecules and crystals in terms of potential energy functions. Oak Ridge National Laboratory, Oak Ridge.
- Finger, L. W., Hazen, R. M. and Yagi, T. (1979) Crystal structures and electron densities of nickel and iron silicate spinels at elevated temperature or pressure. American Mineralogist, 64, 1002–1009.
- Gilbert, T. L. (1968) Soft-sphere model for closed-shell atoms and ions. Journal of Chemical Physics, 49, 2640–2642.
- Hazen, R. M. (1976) Effects of temperature and pressure on the crystal structure of forsterite. American Mineralogist, 61, 1280–1293.
- Hazen, R. M. and Finger, L. W. (1978) Crystal structures and compressibilities of pyrope and grossular to 60 kbar. American Mineralogist, 63, 297–303.
- Hazen, R. M. and Finger, L. W. (1979) Bulk modulus-volume

relationship for cation-anion polyhedra. Journal of Geophysical Research, 84, 6723-6728.

- Hazen, R. M. and Prewitt, C. T. (1977) Effects of temperature and pressure on interionic distances in oxygen-based minerals. American Mineralogist, 62, 309–315.
- Horiuchi, H. and Sawamoto, H. (1981)  $\beta$ -Mg<sub>2</sub>SiO<sub>4</sub>: Singlecrystal X-ray diffraction study. American Mineralogist, 66, 568–575.
- Horiuchi, H., Hirano, M., Ito, E., and Matsui, Y. (1982) MgSiO<sub>3</sub> (ilmenite-type): Single-crystal X-ray diffraction study. American Mineralogist, 67, 788–793.
- Ito, E. and Matsui, Y. (1978) Synthesis and crystal-chemical characterization of MgSiO<sub>3</sub> perovskite. Earth and Planetary Science Letters, 38, 443–450.
- Ida, Y. (1976) Interionic repulsive force and compressibility of ions. Physics of the Earth and Planetary Interiors, 13, 97–104.
- Kumazawa, M. and Anderson, O. L. (1969) Elastic moduli, pressure derivatives, and temperature derivatives of singlecrystal olivines and single-crystal forsterite. Journal of Geophysical Research, 74, 5961–5972.
- Levien, L. and Prewitt, C. T. (1981) High-pressure structural study of diopside. American Mineralogist, 66, 315–323.
- Matsui, M. and Matsumoto, T. (1982) An interionic potentialfunction model for Mg, Ca, and CaMg olivines. Acta Crystallographica, A38, 513–515.
- Matsui, Y. (1982) Computer simulation of structures of actual and hypothetical silicate crystals. In I. Sunagawa and K. Aoki, Eds., Collected Papers on "The Materials Science of the Earth's Interior," p. 11–17. The Japanese Association of Mineralogists, Petrologists and Economic Geologists, Tohoku University, Sendai.
- Matsui, Y. and Kawamura, K. (1980) Instantaneous structure of an MgSiO<sub>3</sub> melt simulated by molecular dynamics. Nature, 285, 648–649.
- Matsui, Y., Yamada, H., and Ito, E. (1979) Synthesis and crystal chemical characterization of high pressure minerals 2, Mg<sub>2</sub>SiO<sub>4</sub> and NaAlSiO<sub>4</sub>. (Abstr.) 20th High Pressure Conference, 81–82, Kobe.
- Meier, W. M. und Villiger, H. (1969) Die Methode der Abstandverfeinerung zur Bestimmung der Atomkoordinaten idealisierter Gerüststruckturen. Zeitschrift für Kristallographie, 129, 411–423.
- Miyamoto, M. and Takeda, H. (1980) An interpretation of the structures of mantle minerals at high pressure in terms of interionic forces. Geochemical Journal, 14, 243–248.
- Miyamoto, M., Takeda, H., and Matsui, Y. (1979) The prediction of crystal structures at high pressure by energy minimization method: Olivine and spinel. (Abstr.) Annual Meeting of Crystallographic Society of Japan, 3B-4, Nagoya.

- Miyamoto, M., Takeda, H., Fujino, K., and Takéuchi, Y. (1982) The ionic compressibilities and radii estimated for some transition metals in olivine structures. Mineralogical Journal, 11, 172–179.
- Mizukami, S., Ohtani, A., and Kawai, N. (1975) High-pressure X-ray diffraction studies on  $\beta$  and  $\gamma$ -Mg<sub>2</sub>SiO<sub>4</sub>. Physics of the Earth and Planetary Interiors, 10, 177–182.
- Morimoto, N., Tokonami, M., Koto, K., and Nakajima, S. (1972) Crystal structures of the high pressure polymorphs of Mn<sub>2</sub>GeO<sub>4</sub>. American Mineralogist, 57, 62–75.
- Morimoto, N., Tokonami, M., Watanabe, M., and Koto, K. (1974) Crystal structures of three polymorphs of Co<sub>2</sub>SiO<sub>4</sub>. American Mineralogist, 59, 475–485.
- Pauling, L. (1960) The Nature of Chemical Bond, 3rd ed. Cornell University Press, Ithaca.
- Sasaki, S., Prewitt, C. T., Sato, Y., and Ito, E. (1982) Singlecrystal X-ray study of  $\gamma$ -Mg<sub>2</sub>SiO<sub>4</sub>. Journal of Geophysical Research, 87, 7829–7832.
- Sato, Y., Ito, E., and Akimoto, S. (1977) Hydrostatic compression of ilmenite phase of ZnSiO<sub>3</sub> and MgGeO<sub>3</sub>. Physics and Chemistry of Minerals, 2, 171–176.
- Shannon, R. D. (1976) Revised effective ionic radii and systematic studies of interatomic distances in Halides and Chalcogenides. Acta Crystallographica, A32, 751–767.
- Shannon, R. D. and Prewitt, C. T. (1969) Effective ionic radii in oxides and fluorides. Acta Crystallographica, B25, 925–946.
- Smyth, J. R. and Hazen, R. M. (1973) The crystal structures of forsterite and hortonolite at several temperatures to 900°C. American Mineralogist, 58, 588–593.
- Takeda, H., Miyamoto, M., and Matsui, Y. (1979) The prediction of crystal structures at high pressure by energy minimization method: Stishovite. (Abstr.) Annual Meeting of Crystallographic Society of Japan, 3B-2, Nagoya.
- Tokonami, M., Morimoto, N., Akimoto, S., Syono, Y., and Takeda, H. (1972) Stability relation between the olivine, spinel and modified spinel. Earth and Planetary Science Letters, 14, 65–69.
- Tosi, M. P. (1964) Cohesion of ionic solids in the Born model. In F. Seitz and D. Turnbull, Eds., Solid State Physics, Vol. 16, p. 1–113. Academic Press, New York.
- Yagi, T., Mao, H. K., and Bell, P. M. (1982) Hydrostatic compression of perovskite-type MgSiO<sub>3</sub>. In S. Saxena, Ed., Advances in Physical Geochemistry, Vol. 2, p. 317–325. Springer-Verlag, New York.

Manuscript received, March 22, 1983; accepted for publication, January 27, 1984.