# Orthopyroxene from the Serra de Magé meteorite: A structure-refinement procedure for a *Pbca* phase coexisting with a *C2/c* exsolved phase

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#### ABSTRACT

An X-ray structure-refinement procedure was developed to characterize orthopyroxene from the Serra de Magé meteorite. In the studied sample, the orthorhombic phase coexists with exsolution lamellae of C2/c augite, parallel to (100), with  $a_{\text{Opx}}^* \equiv a_{\text{Aug}}^*$ ,  $[101]_{\text{Opx}}^* \equiv c_{\text{Aug}}^*$ ,  $b_{\text{Opx}}^* \equiv b_{\text{Aug}}^*$ . Diffraction maxima of the monoclinic phase overlap those of the orthorhombic phase with h+l=2n, yielding violations of the extinction conditions for space group Pbca and simulating the lower symmetry space group  $P2_1ca$ . Structural parameters of the Pbca phase together with the  $W_{\text{Cpx}}$  parameter, which expresses the fraction of the C2/c phase present, were refined first.  $F_c$  values were calculated using the equation  $F_c = \sqrt{(1-W_{\text{Cpx}})[F_c]_{\text{Opx}}^2 + W_{\text{Cpx}}[F_c]_{\text{Cpx}}^2}}$ , with  $[F_c]_{\text{Cpx}}$  values taken from the structure refinement of an augite. Observed  $F_o$  values were then corrected by subtracting the calculated contribution of the monoclinic phase with use of the equation  $F_o^{\text{corr}} = \sqrt{[F_o]^2 - [F_c]_{\text{Cpx}}^2}W_{\text{Cpx}}}$  and used for a further refinement of the orthorhombic phase. The final residual indices,  $R_{\text{obs}}$ , were 2.11 and 1.78% for two crystals with different augite contents. This refinement procedure confirms Pbca as the correct space group and provides more accurate structural parameters for the Serra de Magé orthopyroxene.

#### Introduction

In the orthopyroxene from the Serra de Magé meteorite, the presence of reflections 0kl with k = 2n + 1, in violation of the systematic absence for the b-glide plane in the Pbca space group, had led to the suggestion of the lower symmetry space group P2<sub>1</sub>ca (Harlow et al. 1979), as also proposed for other lunar and terrestrial pyroxenes (Smyth 1974, 1975; Steele 1975; Luo et al. 1992). However, the results of the refinements of two Serra de Magé orthopyroxene crystals according to P21ca (Domeneghetti et al. 1995b) did not support the assumption of this lower symmetry space group. Violations of extinction conditions for *Pbca* symmetry, and the monoclinic Laue symmetry 2/m (rather than orthorhombic symmetry mmm) exhibited by the reflections, were explained by the presence of an exsolved C2/c augite phase, as already proposed by Sasaki et al. (1984). TEM observations by Domeneghetti et al. (1995b) of microstructures of Serra de Magé orthopyroxene showed augite lamellae, parallel to (100), thicker than 500 Å, and irregularly spaced within the orthopyroxene matrix  $(a_{\text{Opx}}^* = a_{\text{Aug}}^*, [101]_{\text{Opx}}^* =$  $c_{\text{Aug}}^*, b_{\text{Opx}}^* \equiv b_{\text{Aug}}^*$ ). With this orientation of the monoclinic cell with respect to the orthorhombic cell, all diffraction effects of monoclinic C2/c augite are superimposed on those of the *Pbca* phase with h + l = 2n, thus explaining the monoclinic-like equivalence of the reflections. An estimate of the amount of the monoclinic phase present was obtained from the ratio  $100(F_{\rm o})^2/(2F_{\rm c})^2$ , where  $F_{\rm o}$  represents the observed structure factors of the forbidden reflections and  $F_{\rm c}$  represents the calculated structure factors of the corresponding reflections in the C2/c phase. The latter are from a refinement of an augite sample with about the same composition as the (100) lamellae. The proportion of augite determined for the two crystals was about 4 and 12%, showing a nonhomogeneous distribution of the exsolved phase in different orthopyroxene crystals.

To obtain a more accurate estimate of the exsolved augite and a more accurate structural determination of the orthopyroxene host we designed a procedure for refining the *Pbca* phase in the presence of a *C2/c* exsolved phase. We present here the strategy and the results of this refinement procedure applied to the same orthopyroxene samples studied by Domeneghetti et al. (1995b).

#### **EXPERIMENTAL METHODS**

#### Structure-refinement procedures

We used for this study the same set of X-ray single-crystal diffraction data collected by Domeneghetti et al. (1995b) for two Serra de Magé samples, Opx SDM N.13, an untreated crystal, and Opx SDM N.7, disordered by heating at 900 °C for 47 h, 40 min. The intensity data

were obtained using a Philips PW 1100 four-circle automated diffractometer and graphite-monochromatized  $MoK\alpha$  radiation ( $\lambda = 0.71073$  Å). Net X-ray diffraction intensities were determined by measuring step-scan profiles and analyzing them by the Lehmann and Larsen (1974)  $\sigma_I/I$  method as modified by Blessing et al. (1974). The reflections hkl,  $\hbar kl$ ,  $h\bar{k}l$ , and  $\bar{h}kl$  were measured up to  $\theta \leq 30^{\circ}$  in the  $\omega$ -scan mode. The intensities were corrected for absorption using the semiempirical method of North et al. (1968). The cell dimensions were determined using a locally improved version (Cannillo et al. 1983) of the Philips LAT routine and are reported in Table 1. The values of equivalent pairs in 2/m Laue symmetry (hkl and  $h\overline{k}l$ ,  $h\overline{k}l$  and  $h\overline{k}l$ ) were averaged. We considered as extinct those reflections that are extinct in the Pbca space group and are either out of register with nodes of the monoclinic net or in register with nodes absent for the extinction conditions in the C2/c space group. Considering that all monoclinic reflections overlap orthorhombic reflections with h + l = 2n, the indices of the extinct reflections can be found by taking into account the extinction conditions in both Pbca and C2/c space groups, together with the transformation matrix  $hkl_{Phea} \rightarrow hkl_{C2/c}$ :  $\frac{1}{2}0 - \frac{1}{2}$ , 0 1 0, 0 0 1. The resulting extinction conditions are as follows: h0l with l = 2n + 1, hk0 with h = 2n + 11, 0kl with k = 2n + 1 and l = 2n + 1, 0kl with k = 2n+ 1 and l = 4n. This agrees with the presence of reflections 0kl with k = 2n + 1 and l = 4n + 2, which are forbidden in the Pbca space group and occur in Serra de Magé orthopyroxene. Reflections with  $I \geq 3\sigma_I$  were considered as observed and used with unit weights in the structure refinement, which was performed without chemical constraints using a locally modified version of the least-squares program ORFLS (Busing et al. 1962). The program allows the assignment of two scattering curves,  $f_1$  and  $f_2$ , to each site and refinement of the occupancy factors X(f) with the constraint  $X(f_1) + X(f_2) =$ 1. The sum  $X(f_1)N_1 + X(f_2)N_2$ , where  $N_1$  and  $N_2$  are the atomic numbers referring to  $f_1$  and  $f_2$ , provides the mean atomic number at the site. The atomic scattering curves are from the International Tables for X-ray Crystallography (Ibers and Hamilton 1974) and from Tokonami (1965). The ionization states used were 2.0+ for Si, Mg, and Fe and 1.5- for O.

The refinement was performed in two steps. First, starting from the values obtained in the Pbca refinement by Domeneghetti et al. (1995b), atomic positions, anisotropic displacement parameters, and occupancy factors at M1 and M2 sites of the orthorhombic phase were refined simultaneously with the fraction  $W_{\rm Cpx}$  of the C2/c exsolved phase. In this procedure, the  $F_{\rm c}$  values were calculated by summing the contribution of both phases with use of the equation  $F_{\rm c} = \sqrt{(1-W_{\rm Cpx})[F_{\rm c}]_{\rm Cpx}^2} + W_{\rm Cpx}[F_{\rm c}]_{\rm Cpx}^2}$ . The  $[F_{\rm c}]_{\rm Cpx}$  values were obtained from the atomic positional and anisotropic displacement parameters of an augite crystal with composition close to that of the (100) lamellae (see Table 3 in Domeneghetti et al. 1995b). These parameters were previously transformed into an appro-

TABLE 1. Crystal and refinement data

	Opx SDM N.13, natural	Opx SDM N.7, heated
a (Å)	18.316(13)	18.320(18)
b (Å)	8.907(7)	8.917(7)
c (Å)	5.218(6)	5.219(6)
V (Å <sup>3</sup> )	851.2	852.6
Crystal dimensions (µm)	$170 \times 180 \times 200$	180 × 310 × 330
Radiation	ΜοΚα	$MoK\alpha$
g range (°)	2-30	2-30
No. of measured refl.	5438	5625
R <sub>sym</sub> ( <i>mmm</i> ) (%)	6.3	7.1
No. of ind. refl. (mmm)	1247	1252
R <sub>sym</sub> (2/m) (%)	4.4	1.6
No. of ind. refl. (2/m)	2743	2808
No. of Ions (in Pbca)	851	897
Robs (%)	2.11	1.78
R <sub>all</sub> (%)	4.24	3.78

Note:  $R_{\text{sym}}(mmm) = \Sigma_{\text{hol}}(I_{\text{hol}} - \bar{I})/\Sigma_{\text{hol}}\bar{I}$ , where  $\bar{I} = (I_{\text{nol}} + I_{\bar{\text{hol}}} + I_{\bar{\text{hol}}} + I_{\bar{\text{hol}}})/4$ ;  $R_{\text{sym}}(2/m) = \Sigma_{\text{hol}}I_{\text{hol}} - \bar{I})/\Sigma_{\text{hol}}\bar{I}$ , where  $\bar{I} = (I_{\text{hol}} + I_{\bar{\text{hol}}})/2$ ;  $I_{\text{cos}} = \text{reflections}$  with  $I \geq 3\sigma_i$ ; and  $R_{\text{cos}}$  and  $R_{\text{all}}$  are the final conventional discrepancy indices expressed as  $\Sigma \|F_{\circ}\| - \|F_{\circ}\|/\Sigma \|F_{\circ}\|$ .

priate form for the orthorhombic cell of the orthopyroxene by means of the unit-cell transformation matrix 2 0 1, 0 1 0, 0 0 1; the unit-cell origin was also moved by  $\frac{1}{4}a$  and  $\frac{1}{4}b$  to ensure the coincidence of the c-glide plane and the  $2_1$  screw axes in the Pbca and C2/c unit cells. After a few least-squares cycles, the first step of the refinement reached convergence. For the two crystals, Opx SDM N.13 and Opx SDM N.7, the residual indices  $R_{\text{obs}}$  were, respectively, 3.20% (for 1725 observed reflections) and 2.25% (for 1918 observed reflections). The calculated fractions of the monoclinic phase were 2.80 and 11.27%, respectively.

In the second step of the refinement procedure, the observed  $F_{\rm o}$  values were corrected by subtracting the calculated contribution of the monoclinic phase with use of the equation  $F_{\rm o}^{\rm corr} = \sqrt{[F_{\rm o}]^2 - [F_{\rm c}]_{\rm Cpx}^2} W_{\rm Cpx}$ . The equivalent pairs in mmm Laue symmetry (hkl and  $\hbar kl$ ) were then averaged and used for a further refinement of the Pbca phase. For the two crystals the refinement converged to residual indices  $R_{\rm obs}$  of 2.11% (for 851 observed reflections) and 1.78% (for 897 observed reflections), respectively.

An attempt to refine separately the monoclinic phase was unsatisfactory because its low concentration resulted in weak intensity of the augite reflections. For the same reason, changes in the augite composition (even 10% in Ca and Fe content) do not significantly affect the values of  $F_o^{\rm cor}$  or the final result of the refinement of the *Pbca* phase.

Selected interatomic distances and angles, and mean atomic numbers at M1 and M2 sites, are reported in Table 2. The final positional and displacement parameters are reported in Table 3. The observed and calculated structure factors (Tables 4 and 5) have been deposited.<sup>1</sup>

<sup>&</sup>lt;sup>1</sup> A copy of Tables 4 and 5 may be ordered as Document AM-96-618 from the Business Office, Mineralogical Society of America, 1015 Eighteenth Street NW, Suite 601, Washington, DC 20036. Please remit \$5.00 in advance for the microfiche.

**TABLE 2.** Selected interatomic distances (Å) and angles (°); mean atomic numbers (m.a.n.) at M1 and M2 sites

	Opx SDM N.13, natural	Opx SDM N.7, heated
SiA-O1A	1.606(2)	1.606(2)
SiA-O2A	1.597(2)	1.594(2)
SiA-O3A	1.636(2)	1.638(2)
SiA-O3A	1.658(2)	1.657(2)
(SiA-O)	1.624(2)	1.624(2)
O3A-O3A-O3A	165.51	164.58
SiB-01B	1.620(2)	1.618(2)
SiB-O2B	1.596(2)	1.595(2)
SiB-O3B	1.666(2)	1.667(2)
SiB-O3B	1.669(2)	1.669(2)
(SiB-O)	1.638(2)	1.637(2)
O3B-O3B-O3B	143.86	143.13
M1-O1A	2.037(2)	2.046(2)
M1-O1A	2.151(2)	2.161(2)
M1-O1B	2.173(2)	2.178(2)
M1-01B	2.066(2)	2.076(2)
M1-02A	2.038(2)	2.041(2)
M1-02B	2.070(2)	2.073(2)
(M1-O)	2.089(2)	2.096(2)
M2-O1A	2.159(2)	2.149(2)
M2-O1B	2.117(2)	2.109(2)
M2-02A	2.051(2)	2.050(2)
M2-02B	2.004(2)	1.998(2)
M2-O3A	2.371(2)	2.367(2)
M2-O3B	2.528(2)	2.526(2)
(M2-O)	2.205(2)	2.200(2)
m.a.n.(M1)	13.4(1)	15.4(1)
m.a.n.(M2)	22.5(1)	20.9(1)

Note: Standard deviations are in parentheses.

### Electron microprobe analysis

New electron microprobe analyses (EMPA) were collected for orthopyroxene grains of a polished thin section of the Serra de Magé meteorite; this thin section was obtained from the same fragment of the meteorite that provided the two single crystals used in the X-ray analysis.

A CAMECA-CAMEBAX electron microprobe, operating with a focused beam ( $\sim 1~\mu m$ ) in wavelength-dispersive (WDS) mode at 15 kV and 15 nA sample current, was used. The counting time for peak and background was 20 s. Synthetic end-member mineral standards were used (Domeneghetti et al. 1995b). X-ray counts were converted to oxide weight percentages using the ZAF correction program. Precision of the analyses is  $\pm 1\%$  for major elements and 3–5% for minor elements.

Ca concentration profiles along directions perpendicular to exsolved lamellae in the orthopyroxene grains allowed us to select Ca-poor areas suitable for spot analyses. The chemical composition reported in Table 6 represents the average of 15 spot analyses in which the oxide total was  $100 \pm 1\%$ . The Fe³+ content was zero, consistent with the highly reduced state of the meteorite; total cation content was  $4.000 \pm 0.005$  atoms on the basis of six O atoms; the charge balance 3<sup>[4]</sup>Al + Na - 3<sup>[6]</sup>Al - 4Ti - 3Cr, where the elements are expressed in terms of atoms per formula unit (apfu), was  $\le 0.005$ . The mean Ca content measured was 0.036 apfu ( $\sigma = 0.005$ ), with a minimum value of 0.029.

#### **Determination of the site population**

Site populations were calculated using the results of both the structure refinements and microprobe analyses. The best value of the atomic fraction  $X_i^i$  for each element i at site j was determined with a minimization procedure using the program MINUIT (James and Ross 1975).

The quantity minimized was  $M = \Sigma [R_w]^2 = \Sigma [(Q_{calc} - Q_{obs})/\sigma(Q_{obs})]^2$ , i.e., the weighted sum of the squares of the residuals (R) between the observed and the calculated values of chemical and structural (Q) parameters, with weights (w) inversely proportional to the relevant standard deviations,  $\sigma$ . The chemical parameters considered

TABLE 3. Atomic positional parameters and displacement parameters

	x	у	Z	$U_{11}$	$U_{22}$	$U_{33}$	U <sub>12</sub>	U <sub>13</sub>	$U_{23}$	$B_{\rm eq}$ (Å <sup>2</sup> )
				0	px SDM N.13	, natural				
SiA	0.22851(4)	0.65945(8)	0.5520(1)	0.0066(3)	0.0063(3)	0.0034(3)	-0.0006(3)	-0.0004(2)	0.0003(3)	0.43
SiB	0.47401(4)	0.33650(8)	0.7950(1)	0.0064(3)	0.0058(3)	0.0031(3)	0.0002(3)	-0.0004(2)	-0.0001(3)	0.40
01A	0.3161(1)	0.8380(2)	0.0440(3)	0.0077(8)	0.0086(8)	0.0043(7)	-0.0002(7)	0.0003(7)	0.0001(7)	0.54
01B	0.5625(1)	0.3370(2)	0.7957(4)	0.0057(7)	0.0088(8)	0.0045(7)	-0.0001(7)	-0.0018(7)	0.0011(7)	0.50
02A	0.3111(1)	0.5003(2)	0.0539(4)	0.0098(8)	0.0081(8)	0.0067(8)	-0.0016(7)	-0.0028(7)	0.0004(7)	0.65
02B	0.4341(1)	0.4852(2)	0.6976(4)	0.0112(8)	0.0087(8)	0.0071(8)	0.0017(7)	0.0004(7)	0.0017(7)	0.71
O3A	0.1976(1)	0.7686(2)	0.3243(4)	0.0096(8)	0.0124(9)	0.0061(8)	0.0016(7)	-0.0003(7)	0.0035(7)	0.74
O3B	0.4473(1)	0.2022(2)	0.5935(4)	0.0105(8)	0.0113(9)	0.0067(9)	-0.0006(7)	-0.0006(8)	-0.0030(7)	0.75
M1	0.37537(4)	0.65467(8)	0.8743(1)	0.0084(4)	0.0071(4)	0.0035(3)	-0.0001(3)	-0.0006(3)	0.0003(3)	0.50
M2	0.37799(3)	0.48326(5)	0.36791(8)	0.0103(2)	0.0096(2)	0.0042(2)	-0.0012(2)	-0.0022(2)	0.0000(2)	0.64
				C	px SDM N.7,	heated				
SiA	0.22838(3)	0.65955(7)	0.5514(1)	0.0051(3)	0.0056(3)	0.0041(2)	-0.0007(2)	-0.0008(2)	0.0001(2)	0.39
SiB	0.47386(3)	0.33637(7)	0.7946(1)	0.0050(3)	0.0049(3)	0.0043(3)	0.0006(2)	-0.0001(2)	0.0004(2)	0.37
01A	0.31602(9)	0.8387(2)	0.0412(3)	0.0064(7)	0.0078(7)	0.0063(7)	0.0001(6)	0.0001(6)	-0.0001(6)	0.54
01B	0.56216(9)	0.3379(2)	0.7951(3)	0.0047(7)	0.0087(7)	0.0068(7)	-0.0001(6)	-0.0007(6)	0.0003(6)	0.53
02A	0.31132(9)	0.4996(2)	0.0517(3)	0.0092(8)	0.0076(7)	0.0074(7)	-0.0017(6)	-0.0012(6)	0.0020(7)	0.64
02B	0.43368(9)	0.4840(2)	0.6951(3)	0.0084(7)	0.0090(8)	0.0077(7)	0.0011(6)	-0.0007(6)	0.0014(6)	0.66
03A	0.19772(9)	0.7698(2)	0.3246(3)	0.0077(7)	0.0114(8)	0.0074(7)	-0.0004(6)	-0.0012(7)	0.0035(6)	0.70
03B	0.44762(9)	0.2012(2)	0.5940(3)	0.0084(8)	0.0102(8)	0.0075(8)	-0.0009(6)	-0.0010(7)	-0.0030(6)	0.69
M1	0.37563(3)	0.65440(6)	0.8726(1)	0.0069(3)	0.0061(3)	0.0052(3)	-0.0002(2)	-0.0009(2)	0.0002(2)	0.48
M2	0.37780(3)	0.48378(5)	0.36644(8)	0.0091(2)	0.0096(2)	0.0068(2)	-0.0011(2)	-0.0025(2)	-0.0002(2)	0.67

Note: Standard deviations are in parentheses.

**TABLE 6.** Electron microprobe analysis of Serra de Magé orthopyroxene

Chemical composition (avg. of 15 spots)				
Oxid	e (wt%)	apfu (based on 6 O atoms)		
SiO <sub>2</sub>	52.42(30)	Si	1.987(7)	
Al <sub>2</sub> O <sub>3</sub>	0.36(6)	Al	0.016(3)	
FeO	24.96(22)	Fe <sup>2+</sup>	0.791(7)	
MgO	20.07(21)	Mg	1.134(12)	
MnO	0.80(8)	Mn	0.026(3)	
TiO,	0.15(3)	Ti	0.004(1)	
Cr <sub>2</sub> O <sub>3</sub>	0.15(4)	Cr	0.004(1)	
CaO	0.88(13)	Ca	0.036(5)	
Na <sub>2</sub> O	0.01(1)	Na	74	
Sum	99.80(40)	Sum	3.998(6)	

Note: Standard deviations are in parentheses.

were the atomic fractions of Si, Al, Ti, Cr, Mg, Fe<sup>2+</sup>, Mn, Ca, and Na. For each element, the residuals  $(R_1-R_9)$  are the difference between the sum of its atomic fraction at the structural sites and its total amount, in apfu, determined by the microprobe analysis. The structural parameters are the mean atomic numbers (m.a.n.) at the M1 and M2 sites and the mean bond distances M1-O and SiB-O, the experimental values of which were provided by the structure refinement. The relevant residuals  $(R_{10}-R_{13})$  were then calculated as follows:

$$\begin{split} R_{10} = & (12X_{\rm Mg}^{\rm M1} + 26X_{\rm Fe^2+}^{\rm M1} + 26X_{\rm Fe^3+}^{\rm M1} + 13X_{\rm [6]A1}^{\rm M1} \\ & + 22X_{\rm Ti}^{\rm M1} + 25X_{\rm Mn}^{\rm M1} + 24X_{\rm Cr}^{\rm M1}) - {\rm m.a.n.(M1)} \quad (1) \\ R_{11} = & [2.078X_{\rm Mg}^{\rm M1} + 2.135X_{\rm Fe^2+}^{\rm M1} + 2.030X_{\rm Fe^3+}^{\rm M1} \\ & + 1.929X_{\rm [6]A1}^{\rm M1} + 1.990X_{\rm Ti}^{\rm M1} + 2.010X_{\rm Cr}^{\rm M1} \\ & + 2.173X_{\rm Mn}^{\rm M1} + 0.006(X_{\rm Fe^2+}^{\rm M2} - X_{\rm Fe^2+}^{\rm M1})] \\ & - \langle {\rm M1-O} \rangle_{\rm obs} \end{split} \tag{2}$$

$$+25X_{Mn}^{M2}+11X_{Na}^{M2})-m.a.n.(M2)$$
 (3)

$$R_{13} = [(X_{[4]A1}^{T} + 15.4230)/9.4187] - \langle SiB-O \rangle_{obs}.$$
 (4)

In Equations 1 and 3 the coefficients of the atomic fractions  $X_i$  are the atomic numbers of the elements that enter the sites M1 and M2, respectively. In Equation 2 the coefficients of  $X_{Mg}^{M1}$  and  $X_{Fe^{2+}}^{M1}$  are the mean bond distances from enstatite and ferrosilite, and the coefficients of the other atomic fractions represent the cation-O mean bond distances for Fe<sup>3+</sup>, <sup>[6]</sup>Al, Ti, Cr, and Mn, respectively, as currently adopted for the M1 site in pyroxenes. The last term of the equation introduces an empirical correction for the influence of the M2-site population on the size of the M1 polyhedron, as in Equation 5 of Domeneghetti et al. (1995a). In Equation 4 the numerical values come from Equation 18 of Domeneghetti et al. (1995a).

In the minimization procedure the following constraints were adopted: (1) Fe<sup>3+</sup> was considered to be equal to zero, as in the microprobe analysis; (2) <sup>[6]</sup>Al, Ti, and Cr were considered to be fully ordered at the M1 site

TABLE 7. Site population of Serra de Magé orthopyroxenes

Site	population	Opx SDM N.13, natural	Opx SDM N.7, heated
Т	Si	1.986	1.986
•	[4]A[	0.014	0.014
M1	Mg	0.897(4)	0.764(4)
	Fe <sup>2+</sup>	0.091(4)	0.220(4)
	[6]A[	0.000	0.000
	Mn	0.003	0.007
	Cr	0.004	0.004
	Fe <sup>3+</sup>	0.000	0.000
	Ti	0.005	0.005
M2	Mg	0.237(4)	0.364(4)
	Fe <sup>2+</sup>	0.704(4)	0.581(4)
	Mn	0.023	0.019
	Na	0.000	0.000
	Ca	0.036	0.036
	K <sub>D</sub>	0.034	0.180

Note: Standard deviations are in parentheses.

(Domeneghetti et al. 1985; Tazzoli and Domeneghetti 1987), Ca fully ordered at the M2 site, Mg and Fe<sup>2+</sup> distributed between the M1 and M2 sites, and Mn partitioned in the same way as Fe<sup>2+</sup> (Hawthorne and Ito 1978); (3) all the structural sites were considered to be fully occupied; (4) the charge balance in isomorphous replacements was constrained by  $X_{\rm I61_{Al}} + 2X_{\rm Ti^{4+}} + X_{\rm Cr} = X_{\rm Na} + X_{\rm I41_{AlCD}}$ .

The standard deviations used for weighting the residuals  $R_1$ – $R_9$  came directly from the microprobe analysis, and those for the residuals  $R_{10}$ – $R_{13}$  from the structure refinement. The standard deviations of the residuals  $R_{11}$  and  $R_{13}$  include both the error of the observed values, given by the refinement, and the error of the calculated values, which were estimated to be 0.01 and 0.001 for  $\langle \text{SiB-O} \rangle$  and  $\langle \text{M1-O} \rangle$ , respectively, according to the results of Domeneghetti et al. (1995a).

The calculated site populations for samples SDM N.13 and SDM N.7 are reported in Table 7.

# DISCUSSION

The validity of the refinement procedure adopted in this work for the two Serra de Magé orthopyroxene crystals is supported by the final values of the residual indices  $R_{\rm obs}$ , which must be compared with those obtained for the same samples by Domeneghetti et al. (1995b) using the traditional refinement in space group Pbca. The  $R_{\rm obs}$  indices decreased from 2.81 to 2.11% for Opx SDM N.13 and from 5.17 to 1.78% for Opx SDM N.7, which contains a higher fraction of the exsolved phase. The standard deviations of the atomic positional parameters, and consequently those of the interatomic distances, show a comparable improvement. In particular, for Opx SDM N.7 they are reduced to about one-third of the previous values. The displacement parameters also exhibit a general decrease, as is indicated by the new values of  $B_{\rm eq}$ .

The results of this work demonstrate that the correct space group for Serra de Magé orthopyroxene is indeed *Pbca* and that the violations of the extinction conditions

for Pbca space group are due to the exsolved augite lamellae. The adoption of a lower symmetry space group must definitely be rejected for this orthopyroxene and must be questioned also for the other reported  $P2_1ca$  orthopyroxenes from lunar and terrestrial occurrences. For these samples, if single crystals are available, the true symmetry could be verified using the method described in this paper.

The refinement procedure adopted here allowed us to determine the fraction of the exsolved C2/c phase, 2.80 and 11.27% for Opx SDM N.13 and Opx SDM N.7, respectively, in good agreement with the approximate values, 4 and 12%, calculated from the intensity of the "forbidden reflections" in Pbca space group (Domeneghetti et al. 1995b). Our refinement provided accurate values of the distribution coefficient  $K_D$  for both samples with use of a more correct determination of M1 and M2 site populations. The  $K_D$  values of 0.034 and 0.180 for Opx SDM N.13 and Opx SDM N.7, respectively, are lower in comparison with those of 0.042 and 0.205 obtained from the site population previously calculated by Domeneghetti et al. (1995b). This indicates that, if the diffraction effects of the exsolved phase are not subtracted from those of the host, the degree of order measured for the orthopyroxene can be lower than the actual value, leading to significant errors in the calculation of the cooling rates.

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#### REFERENCES CITED

- Blessing, R.H., Coppens, P., and Becker, P. (1974) Computer analysis of step-scanned X-ray data. Journal of Applied Crystallography, 7, 488– 492.
- Busing, W.R., Martin, K.O., and Levy, H.S. (1962) ORFLS, a Fortran crystallographic least-squares program. U.S. National Technical Information Service, ORNL-TM-305.
- Cannillo, E., Germani, G., and Mazzi, F. (1983) New crystallographic

- software for Philips PW 1100 single crystal diffractometer. CNR Centro di Studio per la Cristallografia, Internal Report 2.
- Domeneghetti, M.C., Molin, G.M., and Tazzoli, V. (1985) Crystal-chemical implications of the Mg<sup>2+</sup>-Fe<sup>2+</sup> distribution in orthopyroxenes. American Mineralogist, 70, 987–995.
- ———(1995a) A crystal-chemical model for *Pbca* orthopyroxene. American Mineralogist, 80, 253–267.
- Domeneghetti, M.C., Molin, G.M., Stimpfl, M., and Tribaudino, M. (1995b) Orthopyroxene from the Serra de Magé meteorite: Structure refinement and estimation of C2/c pyroxene contributions to apparent *Pbca* diffraction violations. American Mineralogist, 80, 923–929.
- Harlow, G.E., Nehru, C.E., Prinz, M., Taylor, G.J., and Keil, K. (1979) Pyroxenes in Serra de Magé: Cooling history in comparison with Moama and Moore County. Earth and Planetary Science Letters, 43, 173–181.
- Hawthorne, F.C., and Ito, J. (1978) Refinement of the crystal structures of (Mg<sub>0.778</sub>Co<sub>0.224</sub>)SiO<sub>3</sub>. Acta Crystallographica, B34, 891–893.
- Ibers, J.A., and Hamilton, W.C., Eds. (1974) International tables for X-ray crystallography, vol. 4, p. 99–101. Kynoch, Birmingham, U.K.
- James, F., and Ross, M. (1975) MINUIT, a system for function minimization and analysis of the parameter errors and correlations: CERN/DD, Internal Report 75/20. Computer Physics, 10, 343-347.
- Lehmann, M.S., and Larsen, F.K. (1974) A method for location of the peaks in step-scan measured Bragg reflections. Acta Crystallographica, A30, 580-584.
- Luo, G., Xue, J., Xu, H., Xu, H., and Hu, M. (1992) Confirmation of the terrestrial occurrence of orthopyroxene with space group P2<sub>1</sub>ca, American Mineralogist, 77, 115–120.
- North, A.C.T., Phillips, D.C., and Mathews, F.S. (1968) A semi-empirical method of absorption correction. Acta Crystallographica, A24, 351– 359
- Sasaki, S., Prewitt, C.T., and Harlow, G.E. (1984) Alternative interpretation of diffraction patterns attributed to low (P2<sub>1</sub>ca) orthopyroxene. American Mineralogist, 69, 1082–1089.
- Smyth, J.R. (1974) Low orthopyroxene from a lunar deep crustal rock: A new pyroxene polymorph of space group P2<sub>1</sub>ca. Geophysical Research Letters, 1, 27–29.
- ———(1975) Intracrystalline cation order in a lunar crustal troctolite. Proceedings of the 6th Lunar Science Conference, 821–832.
- Steele, I.M. (1975) Mineralogy of lunar norite 78235: Second lunar occurrence of P2<sub>1</sub>ca pyroxene from Apollo 17 soils. American Mineralogist, 60, 1086-1091.
- Tazzoli, V., and Domeneghetti, M.C. (1987) Crystal-chemistry of natural and heated aluminous orthopyroxenes. Physics and Chemistry of Minerals. 15, 131–139.
- Tokonami, M. (1965) Atomic scattering factors for O<sup>2-</sup>. Acta Crystallographica, 19, 486.

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