

## CUPROPAGONITE FROM HALL'S VALLEY, PARK COUNTY, COLORADO

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### ABSTRACT

Cupropavonite from Hall's Valley, Park County, Colorado (USNM 92902) has a cell with double the *c* edge of pavonite. The cell (*a* 13.40, *b* 4.00, *c* 2 x 16.435 Å,  $\beta$  93.87°) contains  $8[\text{Cu}_{1.01}\text{Ag}_{0.58}\text{Bi}_{2.51}\text{Pb}_{0.33}\text{S}_5] = 8[\text{Me}_{4.43}\text{S}_5]$  as compared with  $4[\text{Cu}_{0.33}\text{Ag}_{0.78}(\text{Bi}_{2.74}\text{Pb}_{0.22}\text{Sb}_{0.03})\text{S}_5] = 4[\text{Me}_{4.10}\text{S}_5]$ , on the average, for pavonite. The new occurrence has a lower Pb content than type cupropavonite; the ratio Cu:Pb (of the cations that substitute for Ag and Bi in pavonite) is 3:1, as compared with 3:2 in the type material and in average pavonite. The excess of about 0.4 atoms of metal *per* formula unit over the requirement of the ideal pavonite formula ( $\text{AgBi}_3\text{S}_5$ ) counters a deficiency of about 0.5 atoms of  $\text{Bi}^{3+}$ , which is presumed to be replaced by  $\text{Pb}^{2+}$  in type-locality cupropavonite and by  $\text{Pb}^{2+}$  and  $\text{Ag}^+$  in the new occurrence. The excess metal is judged to be Cu and to be accommodated interstitially in the structure, as has been proposed for the excess of 0.1 atom of metal in pavonite. The remaining Cu and Ag are present in about equal amounts. The distribution may be mainly ordered, causing the doubling of the periodicity along [001] and certain distinct differences in intensities compared with pavonite. A general formula based on these assumptions has the form  $4[\text{Cu}^{\geq 0.5y+z}\text{CuAgBi}_5(\text{Pb}_y\text{Ag}_z)\text{S}_{10}]$ , in which  $\text{Cu}^1$  is interstitial Cu.

**Keywords:** cupropavonite, Park County, Colorado, pavonite, structural formula, interstitial Cu.

### SOMMAIRE

Le paramètre *c* de la maille de la cupropavonite de Hall's Valley, comté de Park, Colorado, U.S.A. (USNM 92902) est double de celui de la pavonite. La maille (*a* 13.40, *b* 4.00, *c* 2 x 16.435 Å,  $\beta$  93.87°) contient  $8[\text{Cu}_{1.01}\text{Ag}_{0.58}\text{Bi}_{2.51}\text{Pb}_{0.33}\text{S}_5] = 8[\text{Me}_{4.43}\text{S}_5]$  au lieu de  $4[\text{Cu}_{0.33}\text{Ag}_{0.78}(\text{Bi}_{2.74}\text{Pb}_{0.22}\text{Sb}_{0.03})\text{S}_5] = 4[\text{Me}_{4.10}\text{S}_5]$  en moyenne pour la pavonite. Le contenu en Pb de cette nouvelle occurrence est inférieur à celui de la pavonite type. Le Cu et le Pb qui remplacent dans la cupropavonite une partie de l'Ag et du Bi de la pavonite sont ici présents dans un rapport Cu/Pb voisin de 3 alors que ce

rapport est voisin de 3/2 dans la cupropavonite type et dans la pavonite moyenne. L'addition d'environ 0.4 atomes de métal par formule par rapport à la formule idéale de la pavonite  $\text{AgBi}_3\text{S}_5$  compense le retrait d'environ 0.5 atomes de  $\text{Bi}^{3+}$ . On suppose que  $\text{Bi}^{3+}$  est remplacé par  $\text{Pb}^{2+}$  dans la cupropavonite type et par  $\text{Pb}^{2+}$  et  $\text{Ag}^+$  dans la présente occurrence. On estime que le métal en excédent est Cu et qu'il est présent sous forme d'atomes interstitiels dans la structure comme on l'a déjà proposé pour l'excès de 0.1 atome par formule pour la pavonite type. Le Cu restant et l'Ag sont présents en quantités à peu près égales. Leur distribution pouvait être ordonnée le long de [001] causant ainsi le doublement de *c* et des différences d'intensité par rapport à la pavonite. Une formule générale basée sur ces suppositions est:  $4[\text{Cu}^{\geq 0.5y+z}\text{CuAgBi}_5(\text{Pb}_y\text{Ag}_z)]$  où  $\text{Cu}^1$  représente le Cu interstitiel.

(Traduit par la Rédaction)

**Mots-clés:** cupropavonite, comté de Park, Colorado, pavonite, formule structurale, cuivre interstitiel.

### INTRODUCTION

Karup-Møller (1972) discovered that "alaskaite" from the Alaska mine, Colorado, which is known to be a mixture, consists in part of pavonite containing exsolution lamellae of a pavonite variety. He could not distinguish the host pavonite from the lamellae by X-ray diffraction, and detected only one lattice on Weissenberg films, which he interpreted as having a distorted (triclinic) cell with a doubled *c* axis (2 x 16.2 Å). Guinier films failed to reveal either the triclinic geometry or the doubling of the *c* axis. However, microprobe analyses showed that the lamellar phase differs significantly in chemical composition from its host. "Normal" pavonite is a Ag-Bi sulfosalt mineral with subordinate Cu. In the lamellae, Cu proved to be in excess of Ag.

A re-examination of the specimens led Karup-Møller & Makovicky (1979) to revise the chemical data for the Cu-rich lamellae (Table 1, nos. 1, 2) and their host (Table 2, no. 6). They described the lamellae as a new mineral, cupropavonite, having pavonite-like cell dimensions

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TABLE 1. MICROPROBE ANALYSES AND ATOMIC PROPORTIONS OF CUPROPAGONITE

Analysis	WEIGHT PERCENT						Total
	Cu	Ag	Bi	Pb	Sb	S	
1	5.7	6.2	56.4	13.5	0.1	18.0	99.9
2	5.9	6.1	56.8	13.4	0.3	17.6	100.1
3	7.26	7.01	59.32	7.65	--	18.10	99.34

  

Analysis	ATOMIC PROPORTIONS (calc. on a basis of S = 5)					Total Metal
	Cu	Ag	Bi	Pb	Bi+Pb+Sb	
1	0.80	0.51	2.40	0.58	2.98	4.29
2	0.85	0.52	2.48	0.59	3.07	4.44
3	1.01	0.58	2.51	0.33	2.84	4.43

1, 2. Alaska Mine, Colorado; repeat analyses of lamellae within "normal" pavonite. Sample 1124 in Karup-Møller & Makovicky (1979). The atomic proportions deduced here differ slightly from the values calculated by the authors. 3. Hall's Valley, Park County, Colorado (USNM 92902); analyst E. Gasparini. Data reduction effected by EMPADR VII (Rucklidge & Gasparini 1969).

TABLE 2. CATION PROPORTIONS (S = 5) FOR MICROPROBE ANALYSES OF PAVONITE

CATIONS	1	2	3	4	5	6	7
Cu	0.23	0.33	0.37	0.32	0.43	0.27	0.33
Ag	0.90	0.83	0.73	0.75	0.69	0.77	0.78
Bi	2.70	2.85	2.68	2.74	2.71	2.75	2.74
Pb	0.15	0.15	0.21	0.23	0.27	0.32	0.22
Sb	--	0.03	0.07	0.04	0.02	0.01	0.03
Cu+Ag	1.13	1.16	1.10	1.07	1.12	1.04	1.11
Bi+Pb+Sb	2.85	3.03	2.96	3.01	2.98	3.07	2.96
Total Metal	3.98	4.19	4.06	4.08	4.10	4.11	4.07

1. Harris & Chen (1975). 2 - 6. Karup-Møller & Makovicky (1979).  
 1. Porvenir Mine, Cerro Bonete, Bolivia (type specimen).  
 2. Silver Bell Mine, Red Mountain, Colorado (BSF 1821).  
 3. Old Lout, Ouray, Colorado (# 1139). 4. Gladiator Mine, Colorado (# 1140). 5. Manhattan, Nevada (ROM M21003). 6. Alaska Mine, Colorado (# 1124), host of cupropavonite (analyses 1 & 2, Table 1). Average of two. 7. Averages of 1 to 6.

but with a doubled *c* axis (*a* 13.45, *b* 4.02, *c* 33.06 Å,  $\beta$  93.50°, *C2/m* or *Cm*) and cell contents  $\text{Cu}_{7.2}\text{Ag}_4\text{Bi}_{20}\text{Pb}_{4.8}\text{S}_{40}$ .

This paper describes an occurrence of cupropavonite from the Missouri mine, Hall's Valley, Park County, Colorado, on the type specimen (USNM 92902) for cuprobismuthite and beryllite. The nature of the occurrence and the crystallographic properties of the mineral differ somewhat from those of the Alaska mine mineral.

### X-RAY-DIFFRACTION STUDIES

The material available for the present study consisted of a few, roughly platy crystals measuring less than 0.5 mm across their greatest dimension after removal from the specimen. Four crystals were oriented for rotation about *b* [010], which lies in the plane of the plates, and each crystal was used to obtain a set of

Weissenberg films. Also, the *hk0*\* and *0kl*\* layers were recorded on precession films with the best crystal.

The X-ray films reveal no departure from monoclinic symmetry and generally resemble the films obtained with Cerro Bonete, Bolivia (type) pavonite and with synthetic Cu- and Pb-free pavonite. However, there are obvious differences: (1) The *c* dimension is doubled by the presence of additional reflections, as reported by Karup-Møller (1972). Some reflections are as strong as the "normal" reflections. (2) The films show certain distinct differences in the relative intensities of reflections as compared with pavonite. (3) The [001]\* row and all reciprocal rows parallel to it are marked on the films by discontinuous streaks (although individual reflections in the rows are generally sharp), indicating some imperfections in the structure of these crystals parallel to (001).

The four sets of Weissenberg films show small variations in intensities. However, the differences appear to be due, not to intrinsic differences among the crystals, but to their poor quality and to the difficulty of mounting them for a uniform view of the X-ray beam during oscillation. The films lead to cell dimensions and to a cell volume that are essentially the same (except for the doubled *c* edge) as values recently obtained on natural and synthetic pavonite. This is in contrast to type cupropavonite, for which Karup-Møller & Makovicky (1979) reported a marked increase in the *a* and *c* dimensions (Table 3). The systematically absent reflections are compatible with the space group *C2/m*.

### MICROPROBE ANALYSES

The microprobe analysis (Table 1, no. 3) was performed on a polished section of one of the crystals previously used in the Weissenberg study. The polished area was of sufficient width (about 30 x 10 μm) to permit a test of the chemical homogeneity of the crystal. As far as can be determined with the microprobe, the Park County crystals consist of only one phase and the distribution of elements within the analyzed crystal is uniform. If some of the odd diffraction effects noted on the Weissenberg films are due to chemical or physical inhomogeneity, it is on a scale below the resolution of the microprobe.

### THE COMPOSITION OF PAVONITE

The composition of pavonite has been dis-

cussed extensively by Harris & Chen (1975), Karup-Møller & Makovicky (1979) and others. The cation proportions derived from recent analyses and collected in Table 2 support the statement of Harris & Chen (1975) that the composition departs significantly from the ideal formula,  $\text{AgBi}_3\text{S}_5$ . Substantial amounts of Pb and Cu are reported in the analyses. The values listed in column 7 of Table 2 indicate that the average cell contents are  $4[\text{Ag}_{0.78}\text{Cu}_{0.33}\text{Bi}_{2.74}\text{Pb}_{0.22}\text{Sb}_{0.03}\text{S}_5]$  or, in terms of the ideal formula,  $4[\text{Ag}, \text{Cu}]_{1.1}\text{Bi}_{2.75}(\text{Pb}, \text{Sb})_{0.25}\text{S}_5$ . Thus, about one atom of Bi *per* unit cell is replaced, chiefly by Pb. The sum  $\text{Ag} + \text{Cu}$  is in excess of 1 in all the analyses. It averages 0.1 atoms *per* formula unit and represents an additional  $\frac{1}{2}$  atom *per* unit cell over that required by the ideal formula. The extra metal is a response to the ubiquitous presence of  $\text{Pb}^{2+}$  in the  $\text{Bi}^{3+}$  sites; presumably, it is accommodated interstitially and consists of Cu rather than the larger Ag atoms. The sum of the positive charges of the cations in column 7 is 9.86, assuming the valences  $\text{Ag}^+$ ,  $\text{Cu}^+$ ,  $\text{Pb}^{2+}$ ,  $\text{Sb}^{3+}$  and  $\text{Bi}^{3+}$ . The shortfall of positive charges required to balance the charge on S supports the suggestion of Harris & Chen (1975) that the interstitial Cu is present in the divalent state.

An idealized general formula that attempts to show the substitution for Ag and Bi and to separate substituting ( $\text{Cu}^0$ ) from interstitial ( $\text{Cu}^1$ ) copper can take the form  $4[\text{Cu}^{1.0y}(\text{Ag}_{1-y}\text{Cu}^0_y)\text{Bi}_{2.75}(\text{Pb}_y\text{Sb}_{z})_{0.25}\text{S}_5]$ . The values of  $y$  (the substitutional coefficient of Karup-Møller & Makovicky 1979), derived from column 7 in Table 2, are Cu 0.22, Ag 0.22, Pb 0.22 for this formula, giving a metal-sulfur composition for the cell of  $4[\text{Me}_{4.11}\text{S}_5]$ . These conclusions differ slightly from the composition of "typical" pavonite deduced by Karup-Møller & Makovicky (1979), although they are in agreement with the observation that the Cu:Pb ratio trends to 3:2.

#### THE COMPOSITION OF CUPROPAVONITE

Type cupropavonite differs from pavonite in having a substitutional coefficient for Ag and Bi of about 0.5 and, as a consequence, a higher metal content. However, the ratio Cu:Pb of the substituting cations remains near 3:2, as in pavonite. The average cell contents derived from the two analyses (Table 1, nos. 1, 2) are  $\text{Cu}_{0.64}\text{Ag}_{4.16}(\text{Bi}_{2.44}\text{Pb}_{0.59})_8\text{S}_{40} = \text{Me}_{35.04}\text{S}_{40}$ . These values differ somewhat from the values calculated from the analyses by Karup-Møller & Makovicky (1979).

The composition of the new occurrence

TABLE 3. CELL DIMENSIONS (Å) AND VOLUME (Å<sup>3</sup>), AND THE VALUE OF Z

	PAVONITE				CUPROPAVONITE		
	1	2	3	4	5	6	7
a	13.333	13.36	13.305	13.310	13.419	13.445	13.40
b	4.039	4.02	4.042	4.035	3.988	4.023	4.00
c	16.346	16.38	16.417	16.395	16.387	33.06	32.87
$\beta$	94.21°	94.18°	94.02°	93.91°	94.31°	93.50°	93.87°
V	877.9	877.4	880.7	878.5	874.5	1784.8	1757.8
Z	4	4	4	4	4	8	8

1. Silver Bell Mine, Red Mountain, Colorado; Weissenberg method (Karup-Møller 1972). 2. Porvenir Mine, Cerro Bonete, Bolivia; precession method (Harris & Chen 1975). 3. Synthetic (Makovicky et al. 1977). 4. Synthetic; 4-circle single crystal diffractometer (E. W. N.). 5, 6. Alaska Mine, Colorado (#1124); Weissenberg method (Karup-Møller & Makovicky 1979). 7. Hall's Valley, Park Co., Colorado (USNM 92902); Weissenberg method (E. W. N.).

matches these features except that the principal metal increase occurs with Cu. The ratio of the substituting cations, Cu:Pb, is 3:1. Presumably the deficiency in Bi (compared with the ideal pavonite formula,  $\text{AgBi}_3\text{S}_5$ ) is made up by  $\text{Pb}^{2+}$  and  $\text{Ag}^+$  in the Park County crystals:  $\text{Cu}_{0.08}\text{Ag}_{3.36}(\text{Bi}_{2.51}\text{Pb}_{0.33}\text{Ag}_{0.16})_8\text{S}_{40} = \text{Me}_{35.44}\text{S}_{40}$ .

Despite the high metal content, the cell of the Park County cupropavonite shows no increase in volume, except for doubling, over natural and synthetic pavonite (Table 3). This factor and the similarity in cell dimensions of the two species offer evidence that the excess metal is accommodated interstitially, as in pavonite. The abundance of Cu in the cell and the small radius of its atom imply that the interstitial metal is Cu. When the cell contents are cast into a pavonite-like formula to bring  $\text{Cu}^0 + \text{Ag} = 1$ , the excess metal is seen to amount to about 0.4 atoms *per* pavonite formula unit:  $8[\text{Cu}^{0.43}(\text{Cu}^{0.58}\text{Ag}_{0.42})(\text{Bi}_{2.51}\text{Pb}_{0.33}\text{Ag}_{0.16})_8\text{S}_5]$ .

The sum of the charges on the cations derived for the Park County material (Table 1, no. 3) is 9.78, accepting the usual valences. It represents a greater shortfall in balancing the charge on S than that occurring in pavonite and supports the view (Karup-Møller & Makovicky 1979) that some, if not all, of the excess (interstitial) Cu in cupropavonite exists in the higher valence state,  $\text{Cu}^{2+}$ .

A generalized formula to account for the three available analyses of cupropavonite, and to incorporate the chemical characteristics noted above, has the form  $8[\text{Cu}^{1.0y+z}(\text{Cu}^0_y\text{Ag})_1\text{Bi}_{2.5}(\text{Pb}_y\text{Ag}_z)_{0.5}\text{S}_5]$ . If all the interstitial Cu is divalent, the value of  $\text{Cu}^1$  in this formula is exactly  $0.5y + z$ , and the metal-sulfur proportion can range from  $\text{Me}_{34}\text{S}_{40}$  to  $\text{Me}_{38}\text{S}_{40}$ , increasing in metal with increasing substitution of Ag for Bi. Table 4 shows the distribution of the cations determined in Table 1 in accordance

TABLE 4. DISTRIBUTION OF CATIONS IN CUPROPAVONITE ACCORDING TO  $\text{Cu}^i(\text{Cu}^s, \text{Ag})_1(\text{Bi}, \text{Pb}, \text{Ag}_z)_3\text{S}_5$

Analysis*	Bi	Pb <sub>y</sub>	Ag <sub>z</sub>	Ag	Cu <sup>s</sup>	Cu <sup>i</sup>	(.5y+z)	Total metal	Total charge <sup>†</sup>
1	2.40	0.58	0.02	0.49	0.51	0.29	0.31	34.3	9.96
2	2.48	0.59	--	0.52	0.48	0.37	0.31	35.5	10.36
3	2.51	0.33	0.16	0.42	0.58	0.43	0.33	35.4	10.21

\*From Table 1. †taking Cu<sup>i</sup> as Cu<sup>2+</sup>.

Note:  $\text{Ag}_z = 3 - (\text{Bi} + \text{Pb})$ ,  $\text{Cu}^s = 1 - (\text{total Ag} - \text{Ag}_z)$ ,  $\text{Cu}^i = \text{total Cu} - \text{Cu}^s$ .

with the formula. For analyses 2 and 3, the value of Cu<sup>i</sup> is greater than the sum  $0.5y + z$ , and the total positive charge, taking Cu<sup>i</sup> to equal Cu<sup>2+</sup>, is higher than 10. This implies that only some of the interstitial Cu is divalent, and indicates that the total metal content is determined not only by the ratio of Pb to Ag in the Bi sites but also by the valence of the interstitial Cu. Hence, the quantity of Cu<sup>i</sup> in the formula is  $\geq 0.5y + z$ . These observations obviously make it difficult to conceive of an "ideal" composition for cupropavonite. The measured metal-sulfur proportions for the three analyses (Tables 1, 4) range from  $\text{Me}_{34.3}\text{S}_{40}$  to  $\text{Me}_{35.5}\text{S}_{40}$ .

The distribution of Cu in Table 4 implies that the ratio Cu<sup>s</sup>/Ag in the Ag sites is near 1. Cu and Ag exhibit only limited substitution for each other in sulfosalt minerals. Since the high Cu content is the principal chemical difference between cupropavonite and pavonite, it is likely that the two metals are mainly ordered in distinct positions in cupropavonite. This feature accounts for the doubling of the periodicity along [001] and promotes appreciable intensity differences between pavonite and cupropavonite. A general chemical formula for cupropavonite should, therefore, have the form  $4[\text{Cu}^i \geq (0.5y+z)]\text{CuAgBi}_3(\text{Pb}_y\text{Ag}_z)\text{S}_{10}$ , in which z has a maximum observed value of 0.32.

Table 1 shows that the principal difference between the two occurrences of cupropavonite

lies in the ratio Cu:Pb. The lower Pb content of the Park County crystals is presumed to result in the presence of Ag in the Bi sites. This multivalence occupation by Ag<sup>+</sup>, Pb<sup>3+</sup> and Bi<sup>3+</sup> of the Bi sites is the only apparent source of the disorder in the Park County crystals that is responsible for the streaking observed on the X-ray films.

#### ACKNOWLEDGEMENTS

I thank Dr. E. Gasparrini of Minmet Scientific, Toronto, for doing the microprobe analysis and Dr. George Switzer for the extended loan of a U.S. National Museum specimen (USNM 92902). I had the benefit of a discussion and correspondence with Dr. E. Makovicky concerning his work on type cupropavonite. This project was supported by a grant from the National Research Council.

#### REFERENCES

- HARRIS, D.C. & CHEN, T.T. (1975): Studies of type pavonite material. *Can. Mineral.* 13, 408-410.
- KARUP-MØLLER, S. (1972): New data on pavonite, gustavite and some related sulphosalt minerals. *Neues Jahrb. Mineral. Abh.* 117, 19-38.
- & MAKOVICKY, E. (1979): On pavonite, benjaminite and "oversubstituted" gustavite. *Bull. Minéral.* 102, 351-367.
- MAKOVICKY, E., MUMME, W.G. & WATTS, J.A. (1977): The crystal structure of synthetic pavonite,  $\text{AgBi}_3\text{S}_8$ , and the definition of the pavonite homologous series. *Can. Mineral.* 15, 339-348.
- RUCKLIDGE, J. & GASPARRINI, E.L. (1969): Electron microprobe analytical data reduction - EMPADR VII. *Dep. Geol. Univ. Toronto.*

Received August 1979, revised manuscript accepted January 1980.