NORDSTRANDITE FROM MONT ST-HILAIRE, QUEBEC

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ABSTRACT

Nordstrandite Al(OH)₃ occurs as aggregates of tabular and blocky crystals in association with natrolite, tetranatrolite, analcime, sodalite, albite and microcline in pegmatites, miarolitic cavities and xenoliths in nepheline syenite at Mont St-Hilaire, Quebec. Single-crystal X-ray-diffraction studies confirmed the triclinic symmetry of the mineral. The reduced cell has a 6.148(2), b 6.936(2), c 5.074(1) Å, α 95.76(2), β 99.06(2), γ 83.30(2)° and Z=4. Crystal morphology indicates the presence of a centre of symmetry. The mineral is colorless, white and, rarely, beige, pink and pale green, transparent to translucent. The lustre is vitreous, slightly pearly on the perfect $\{110\}$ cleavage. Hardness ~3, $D(\text{meas.}) = 2.42 \text{ g/cm}^3$. Optically, the mineral is biaxial positive, α 1.580(1). β 1.583(1), γ 1.602(1), 2V 24°. Elongation (-), dispersion moderate, r < v. The orientation of principal vibration axes is defined by the spherical coordinates $X(-59^{\circ}, 37^{\circ})$, $\overline{Y}(74^{\circ}, 64^{\circ})$, $Z(174^{\circ}, 67^{\circ})$, Electron-microprobe analyses of the blocky and tabular crystals gave Al₂O₃ 65.6, 64.9; SiO₂ 0.06, 0.98; CaO 0.00, 0.00; Na₂O 0.01, 0.02; FeO 0.00, 0.05, totals 65.67, 65.95 wt. %, respectively.

Keywords: nordstrandite, nepheline syenite, Mont St-Hilaire, Quebec, triclinic, physical and optical properties, electron-microprobe analyses.

SOMMAIRE

On trouve la nordstrandite Al(OH)₃ en agrégats de cristaux tabulaires et massifs associée aux phases natrolite, tétranatrolite, analcime, sodalite, albite et microcline dans les pegmatites, les miaroles et les xénolithes de la syénite néphélinique au mont St-Hilaire (Québec). La diffraction sur monocristal confirme la symétrie triclinique de cette espèce. Les paramètres de la maille réduite sont a 6.148(2), b 6.936(2), c 5.074(1) Å, α 95.76(2), β 99.06(2), γ 83.30(2)°, $Z \cong 4$. La morphologie des cristaux indique la présence d'un centre de symétrie. La nordstrandite est incolore, blanche ou, plus rarement, beige, rose ou vert pâle, et transparente à translucide. Son éclat est vitreux, légèrement nacré sur le clivage parfait {110}. Dureté ~ 3, D(mes.) 2.42. Elle est optiquement biaxe positive. a 1.580(1), β 1.583(1), γ 1.602(1), 2V 24° avec allongement négatif et dispersion modérée, r < v. L'orientation des axes principaux de vibration se définit en coordonnées sphériques: X (-59°,

37°), \overline{Y} (74°, 64°), Z (174°, 67°). L'analyse des cristaux massifs et d'aspect tabulaire à la microsonde électronique donne, respectivement, Al₂O₃ 65.6, 64.9; SiO₂ 0.06, 0.98; CaO 0.00, 0.00; Na₂O 0.01, 0.02; FeO 0.00, 0.05; total 65.67, 65.95% par poids.

(Traduit par la Rédaction)

Mots-clés: nordstrandite, syénite néphélinique, mont St-Hilaire, Québec, triclinique, propriétés physiques et optiques, analyse à la microsonde électronique.

INTRODUCTION

A quarter of a century ago Van Nordstrand et al. (1956) synthesized a new form of Al (OH)₃; it was later named nordstrandite by Papée et al. (1958). Many natural occurrences have since been reported. According to Milton et al. (1975), these may be classified into three distinct types: (1) most commonly as a weathering product in bauxitic soils derived from limestones, e.g., in Sarawak, Borneo (Wall et al. 1962), Guam (Hathaway & Schlanger 1962, 1965), Montenegro (Tertian 1966), Hungary (Náray-Szabó & Péter 1967), Croatia (Marić 1968), Jamaica (Davis & Hill 1973, as reported in Milton et al. 1975) and in the Sokolovsko-Sarbay magnetite mines, U.S.S.R. (Kulikova et al. 1974); (2) as a vein or fissure-filling mineral in dolomitic oil shale in the Green River Formation, northwestern Colorado (Milton et al. 1975); and (3) as an alteration product of dawsonite and alumohydrocalcite in New South Wales, Australia (Goldberry & Loughnan 1970, 1977). More recently, another type of occurrence of the mineral, in alkalic igneous rocks, was reported: Petersen et al. (1976) found nordstrandite as a late mineral in pegmatitic pockets in nepheline syenite at Narssârssuk, Greenland, and Sabina (1977) reported it as a vug mineral in the sodalite-bearing rocks at the Princess sodalite mine near Bancroft, Ontario.

The present paper describes the occurrence and properties of nordstrandite from the nepheline syenite at Mont St-Hilaire, Quebec. The mineral was first identified from this locality in 1971 and was documented by Chao & Baker (1979).

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OCCURRENCE

At Mont St-Hilaire, nordstrandite occurs in pegmatites, miarolitic cavities and biotite-rich xenoliths in the nepheline syenite. The mineral is commonly associated with natrolite, tetranatrolite, analcime, sodalite, albite and microcline. At least one of the carbonate minerals calcite, siderite, dolomite, ancylite and dawsonite is always present in small amounts, except in xenoliths where apatite takes the place of carbonates. Other minor minerals that may be present are aegirine, pyrite, zircon, rutile, fluorite, an unidentified mineral UK43 (Chao & Baker 1979), a nordstrandite-like phase of composition Al(OH)₃ and an iron aluminum member of the serpentine group. The mineral commonly forms tabular rhombic microcrystals in random aggregates in cavities and interstices of natrolite, analcime or albite (Figs. 1a, 1b). The mineral also forms globular aggregates and radiating clusters commonly with a core of

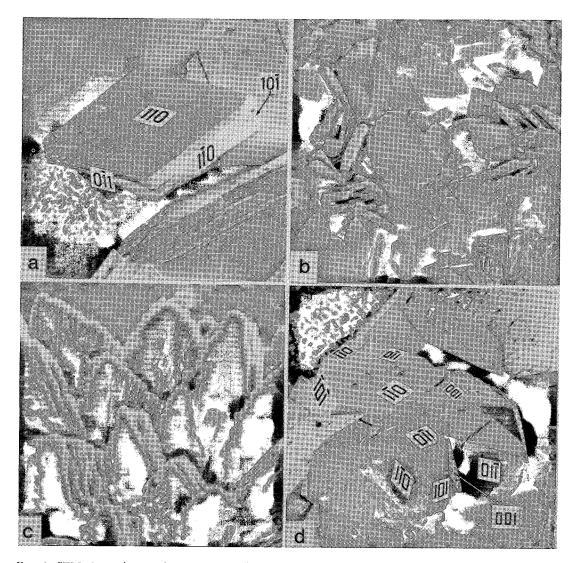


FIG. 1. SEM photomicrographs of nordstrandite, Mont St-Hilaire, Quebec. (a) Tabular rhombic crystal on natrolite (massive) and tetranatrolite (fibrous); long dimension of crystal approximately 0.5 mm. (b) Random aggregates of tabular crystals; average size 0.2 mm. (c) Parallel growths of tabular rhombic crystals in a globular aggregate. (d) Blocky crystals; largest crystal approximately 1.5 mm long.

analcime or siderite. In some aggregates the individuals are not recognizable even with high magnification under the scanning electron microscope. In others, "leaves" formed by subparallel to parallel growths of small, tabular, rhombic crystals are apparent (Fig. 1c). Nordstrandite less commonly forms relatively larger (1-2 mm) crystals of a distinct blocky habit (Fig. 1d) in natrolite vugs. These crystals are translucent to opaque with a transparent outer zone. The opaque parts of the crystals contain numerous minute needle-shaped inclusions of dawsonite, randomly arranged.

Textural relationships indicate that the order of crystallization of the essential minerals in the nordstrandite associations is microcline, albite, natrolite or analcime or sodalite, tetranatrolite, carbonates and nordstrandite. Pyrite and serpentine are the only minerals formed later than nordstrandite. Thus, the occurrence and paragenesis of nordstrandite at Mont St-Hilaire are similar to those at Narssârssuk, Greenland (Petersen et al. 1976).

CRYSTALLOGRAPHY

Considerable confusion regarding the unit-cell data of nordstrandite exists in the literature, as shown in Table 1. From powderdiffraction data of synthetic nordstrandite, Lippens (1961) derived a monoclinic cell that resembles the gibbsite cell. However, not all the powder-diffraction lines could be indexed satisfactorily on the basis of Lippens's cell. The triclinic nature of nordstrandite was first recognized by Saalfeld & Mehrotra (1966) from a Weissenberg single-crystal X-ray study of the natural material from Sarawak. Their proposed cell, doubly primitive but not bodycentred, was adopted by Roberts et al. (1974) in the Encyclopedia of Minerals, but the symmetry of the mineral was erroneously recorded as monoclinic. A body-centred triclinic cell was later established by Saalfeld & Jarchow (1968) in their structure determination; this cell has since been referred to by several authors (e.g., Peterson et al. 1976). In an independent attempt to determine the crystal structure of the mineral, Bosmans (1970) derived another cell from powder-diffraction data of a synthetic material. The cell was considered to be the reduced cell despite a doubling of the c axis indicated by electron diffraction. The reduced cell of Bosmans was adopted in the Powder Diffraction File, but it is now apparent that it is only a subcell.

Much of the confusion was apparently caused by the lack of good single crystals; hence the cell data were obtained mostly from The powder-diffraction work. well-formed crystals of nordstrandite from Mont St-Hilaire made it possible to re-examine its cell geo-

	1	2	3	4	5	6
α(Å) b α(°) β	8.63(2) 5.01(2) 19.12(3) 92.00	8.893 5.004 10.237 92.93 110.38 90.53	8.752(9) 5.069(5) 10.244(13) 109.326(88) 97.662(95)	5.082 5.127 4.980 93.67 118.92 70.27	5.114 5.082 5.127 70.27 74.00 58.47	6.148(2) 6.936(2) 5.074(1) 95.76(2) 99.06(2) 83.30(2)
Υ ν(Å)	826.17	90.53 426.30	88.340(84) 424.97	106.13	106.09	211.47(7)
Z	(16)	8	8	2	2	4
ransformation atrix to the resent cell	?	- ¹ 2- ¹ 2 ¹ 2 0 1 ¹ 2 12	يد يد يد يد يد 0 1-0	-1 1-1 0 1 1 1 0 0	-101 1-11 010	100 010 001
Inverse matrix	?	1]] -12 12-32 12-12	1-10 00-1 111	1 0 0 يز يز يز يز-يز ا	يد يد يد 1 0 0 2	100 010 001
Cell type	?	Doubly primitive, not I-cell	Doubly primitive, I-cell	Sub-cell	Tri-acute sub-cell	Reduced cell

TABLE 1. CELL DIMENSIONS OF NORDSTRANDITE

Lippens (1961), for synthetic nordstrandite. The powder patterns cannot be indexed on the basis of this cell.

2. Saalfeld & Mehrotra (1966) . 3. Saalfeld & Jarchow (1968) .

Bosmans (1970). The cell was thought to be the reduced cell but electron diffraction revealed the doubling of σ . Bosmans (1970). The cell may be transformed to 4 by the matrix 010/001/110 4.

5.

6. This study. metry by single-crystal X-ray diffraction using a precession camera. Six crystals were studied and numerous photographs taken with various important zone axes as the precession axis. The triclinic body-centred cell of Saalfeld & Jarchow (1968) was duly confirmed, and the true reduced cell with Z=4 was established. The relationship of this reduced cell to the cells reported by other authors is given in the form of transformation matrices in Table 1. The reduced cell of nordstrandite is related to the monoclinic cell of gibbsite by the matrix $1\overline{10}/00\overline{1}/110$. This relationship is supported by the almost identical nature of the hk0 reciprocal lattice net of nordstrandite and the hol net of gibbsite.

To compare the cell parameters of nordstrandite from different localities, we indexed the published powder-diffraction data sets with reference to the reduced cell and used them to refine the cell parameters by a least-squares method. The results are presented in Table 2. The comparison shows a significantly larger variation in a than in b and c. It is also interesting to note that the cell volume of all natural nordstrandite is slightly smaller than that of the synthetic material, a reflection, perhaps, of the substitution of Si for Al in the natural mineral.

The X-ray powder-diffraction data for nordstrandite from Mont St-Hilaire are given in Table 3. They are essentially identical to those of the synthetic material, the mineral from Guam and other localities.

Morphologically, the rhombic tabular crystals are bounded by a well-developed {110}, moderately developed $\{1\overline{1}0\}$ and $\{0\overline{1}1\}$, with or without a poorly developed {101} (Fig. 1a). In contrast, the most prominent forms on the blocky crystals (Fig. 1d) are $\{0\overline{1}1\}$ and $\{1\overline{1}0\}$, followed by $\{101\}$ and $\{001\}$, with $\{110\}$ being the least prominent form. A poorly developed {010} is present on some crystals. The form $\{1\overline{1}0\}$ is invariably striated parallel to $\{110\}$. Crystals of both habits appear to possess a centre of symmetry, consistent with the results of the crystal-structure analyses by Saalfeld & Jarchow (1968).

The interfacial angles of a blocky crystal were measured using a two-circle optical goniometer. The observed (ϕ, ρ) angles are in good agreement with the angles calculated from the cell parameters (Table 4). No goniometry measurements were made for the tabular crystals, as the only suitable crystal was broken when it was detached from the matrix. The forms present on the tabular crystal were, therefore, determined by X-ray precession goniometry by shooting X-rays along several zone axes of the broken crystal.

PHYSICAL AND OPTICAL PROPERTIES

Nordstrandite from Mont St-Hilaire is colorless, white and, rarely, pink, beige and pale green. The color of the pale green variety has been shown by scanning clectron microscopy and X-ray diffraction to be due to a thin coating of an iron aluminosilicate mineral of the serpentine group. The lustre of the mineral is vitreous and is somewhat pearly on the perfect {110} cleavage. The Mohs hardness is close to 3. The density, determined by the flotation method, is 2.42 g/cm³, in agreement with the

. <u> </u>	1	2	3	4	5	6	7	8	9
α(Å) b α(°) β Υ	6.125(4) 6.923(3) 5.082(2) 95.67(4) 98.88(4) 83.53(4)	6.137(1) 6.927(2) 5.075(1) 95.82(2) 98.93(2) 83.35(2)	6.148(2) 6.936(2) 5.074(1) 95.76(2) 99.06(2) 83.30(2)	6.148(2) 6.933(2) 5.079(1) 95.80(2) 98.93(2) 83.22(2)	6.156(3) 6.928(6) 5.078(2) 95.76(5) 98.97(4) 83.38(6)	6.149(2) 6.934(2) 5.078(1) 95.79(3) 98.92(3) 83.30(3)	6.141(8) 6.932(5) 5.087(6) 95.93(7) 98.62(8) 83.33(9)	6.167 6.924 5.082 95.62 99.08 83.27	6.163(2) 6.931(3) 5.081(2 95.70(3) 98.92(3) 83.32(3)
₽(Å ³)	210.8(1)	210.93(6)	211.47(7)	211.61(6)	211.7(2)	211.72(7)	211.8(2)	212.11	212.24(9)

TABLE 2. DIMENSIONS OF THE REDUCED CELL OF NORDSTRANDITE*

* Refined by a least-squares method using x-ray powder diffraction data from the following sources:

1. Sokolovsko-Sarbay magnetite mines, USSR (Kulikova et al. 1974).

2. Southeast Dinarides, Croatia (Maric 1968).

3. 4.

- 5.
- Southeast Dinarrows, croatia (marte 1990). Mont St-Hilaire, Quebec (this study). Guam (Hathaway & Schlanger 1962, 1965). West Sarawak, Borneo (Saalfeld & Mehrotra 1966). West Sarawak, Borneo (Wall *et al.* 1962). 6.
- 7.
- Sydney Basin, Australia (Goldberry & Loughnan 1977). Narssårssuk, Greenland (Petersen *et al.* 1976). Calculated from parameters of the I-cell, a = 8.715(4), b = 5.082(5), $\sigma = 10.248(6)$ Å, $\alpha = 109.56(6)$, $\beta = 97.34(5)$, $\gamma = 88.06(3)^\circ$. 8.

9. Synthetic nordstrandite (Bosmans 1970).

TABLE'4. FORMS AND ANGLES FOR NORDSTRANDITE FROM MONT ST-HILAIRE, QUEBEC

Forms	[¢] obs	[¢] calc	^ρ obs	^P calc	^A calc	^B calc	Ccalc
001	62.3	62.02	10.1	10.26	81,52	85.21	
010		0		90.00	95.90		85.21
100		95.90		90.00		95,90	81.52
110	52.0	52.0	90.0	90.00	43,90	52.00	79.90
TTO	134.6	134.65	90.0	90.00	38.75	134.65	86.95
0T1	167.0	166.31	34.0	33.96	79.25	122.87	37.66
101	89.8	90.09	44.1	44.86	45.43	90.06	36.06

A=hkl A 100; B=hkl A 010; C=hkl A 001.

values reported for the mineral from other localities $(2.41-2.43 \text{ g/cm}^3)$.

Optical properties of the mineral were determined on a spindle stage. Crystals previously oriented by X-ray goniometry were reoriented with the aid of extinctions and interference figures to rotate about the optic normal. Indices of refraction and 2V were measured in sodium light, and all immersion liquids were checked with a refractometer. In white light a moderate, asymmetrical dispersion r < v was observed. The orientation of the optical indicatrix with respect to the reduced cell is shown in Figure 2, and is defined by the spherical coordinates $X(-59^{\circ}, 37^{\circ}), \overline{Y}(74^{\circ}, 64^{\circ}), Z(174^{\circ}, 67^{\circ}).$ The physical and optical properties of nordstrandite from Mont St-Hilaire are compared with those of the mineral from other localities in Table 5.

The sign of elongation of the nordstrandite crystals from Mont St-Hilaire is negative (lengthfast), consistent with that of the mineral from Guam (Hathaway & Schlanger 1965), but opposite that of the acicular crystals from Greenland (Petersen et al. 1976). Our examination of the nordstrandite from Guam confirmed the report by Hathaway & Schlanger. Moreover, nordstrandite from Bancroft, Ontario, and in specimens from Narssârssuk, Greenland (41217 and 41251 in the collection at the National Museum of Natural Sciences, Ottawa) also gave negative elongation. Thus, there is little doubt that the acicular crystals studied by Petersen et al. (1976) are elongate along a different crystallographic direction, perhaps the b axis of the reduced cell. The angle $Z \wedge b$, estimated from the stereographic projection (Fig. 2), is about 23°. close to the maximum extinction angle of about 20° observed by these authors for their material.

ELECTRON-MICROPROBE ANALYSES

Nordstrandite from Mont St-Hilaire was analyzed with an electron microprobe using a defocused beam. Natural gibbsite (for Al), kyanite (Si), anorthite (Ca), albite (Na) and olivine (Fe) were used as standards with jadeite

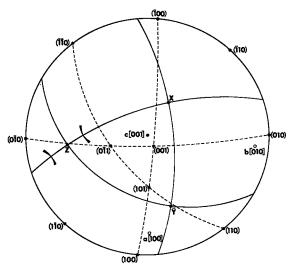


FIG. 2. Stereographic projection of crystallographic and optical elements of nordstrandite.

and the nordstrandite from Guam as internal standards. The analyses for the blocky and tabular crystals are, respectively, Al_2O_3 65.6, 64.9; SiO₂ 0.06, 0.98; CaO 0.00, 0.00; Na₂O 0.01, 0.02; FeO 0.00, 0.05; totals 65.67, 65.95 wt. %. The H₂O contents were not determined because of the small amount of sample available. Both analyses compare favorably with the theoretical Al_2O_3 content of 65.36% for Al (OH)₃. The higher SiO₂ content of the tabular crystal is probably not related to crystal habit because the nordstrandite from Guam, which has a similar tabular habit, has only 0.02% SiO₂.

POLYMORPHISM IN Al(OH)₃

It is commonly accepted that Al(OH)₃ exists in three structural forms, namely, bayerite, gibbsite and nordstrandite. They differ essentially in the stacking of the layers of hydroxyl ions. Bayerite is believed to have the brucite structure with the stacking sequence ABAB typical of the hexagonal closest packing (Yamaguchi & Sakamoto 1958, Lippens 1961). The structure of gibbsite, as determined by Megaw (1934) and refined by Saalfeld & Wedde (1974), is based on the stacking sequence ABBA. Speculations on the basis of X-ray powder-diffraction data suggested that nordstrandite has a structure with mixed bayerite and gibbsite layers of ABABBABA, but structure determination using single-crystal X-raydiffraction intensity data (Saalfeld & Jarchow

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TABLE 3. SELECTED X-RAY POWDER DIFFRACTION DATA FOR NORDSTRANDITE

	1 Mont St-Hilaire Quebec		ire	2 Guam		3 Sara Bor		4 Sydney Austra	Basin	5 Croatia		6 Synthetic	
hkl	doalo	d _{obs}	I _{obs}	d _{obs}	Iobs	dobs	Iobs	d _{obs}	I _{obs}	d _{obs}	I _{obs}	dobs	Iobs
1 <u>1</u> 0	4.785	4.779	100	4.789	100	4.78	vs	4.79	100	4.782	100	4.791	
1 <u>To</u>	4.320	4.317	30	4.322	12	4.33	S	4.34	20	4.318	40	4.320	100 27
01 <u>T</u> 10T	4.208 4.161	4.223 4.168	20 5	4.207	10	4.206	S	4.20	12	4.207	40	4.205	18 12
011	3.886	3.890	10	4.156 3.887	7 4	4.153 3.886	s m	4.15 3.865	12 8	4.162 3.883	40 30	4.160 3.888	12
101	3.597	3.603	10	3.600	4	3,600	m	3.600	6	3.597	30	3.609	11 8 5
<u>0</u> 20	3.432	3.428	5	3.429	4	3.425	m	3.435	3	3.432	20	3.427	5
111 120	3.356 3.125					3.140		3.320	2				
200	3.020	3.019	5	3.022	2	3.023	VW W	3.040	3	3.021	20	3.028	3
121	2.852	2.852	5	2.850	ž	2.849	w	2.860	2	2.848	20	2.848	3
211	2 705	2 700	•	0 704	-	0 704	D	2.785	3			0 710	
210	2.705 2.665	2.708	3	2.704 2.663	1 <1	2.704 2.637	Bw w					2.710	3
002	2.496			2.497	~i	2.037	-					(2,501)	1
<u>1</u> 21	2.482	2.484	10	2.480	3	2.479	Bm	2.484	3	2.479	30	2,480	12
211 220	2.453 2.392)	2.443	5	2.450	1	2.445	W	2.445	1	2.454	20	2.455	8
11 2	2.392	2.392	40	2.392	9	2.392	s	2.393	10	2.392	70	2.393	27
211	2.322	2.323	1	2.349	<1	2.330	VW					2.333	5
211	2.261	2.263	40	2.263	15	2.261	s	2.269	9	2.261	70	2.271	29
13 <u>0</u> 031	2.216 2.149	2.219 2.147	1 5	2.148	1	2.225	W VW	2.152	2			2.217 2.146	3
122	2.114	2.112	ĭ	2.140		2.097	VW	2.152	2			2.113	3 3 2
12 <u>2</u> 212	2.0741	2.073	3	2.074	1	2.073	VW					2.074	3
112	2.070	2.075	5			2.0/5						2.074	5
031	2.017	2.017	30	2.033 2.016	<1 8	2.016	s	2.018	10	2.015	50	2.016	24
310	1.988	1.983	ĩ	2.010	Ū	1.982	vw	2.010		2.010	30	(1.991)	2
301	1.971		-			1.959	VW					1.975	1
022	1.943 1.902	1.944 1.899	3 30	1.9434 1.9008	<1 8	1.939 1.899	VW	1.902	10	1.898	50	1.945 1.902	10
222 310 320	1.880	1.033	30	1.8797	<1	1.099	s	1.902	10	1.090	50	(1.877)	2
320	1.820											1.821	24 2 1 5 19 2 2 3
20 <u>2</u> 132	1.798 1.781 ເ			1.8017	<1					1.799	40	1.804	3
301	1.778	1.779	15	1.7807	5	1.779	ms	1.783	7			1.784	13
04 <u>0</u>	1.716	1.716	1	1.7152	<1	1.715	VW					1.715	1
312	1.702 1.697	1.698	2	1.7016	<1	1.704	w					1.704	2
311 141	1.671	1.670	7	1.6706	2	1.672	W					1.668	4
141 231 013	1.654		,							1 650			
013	1.649			1.6518	<1	1.653	VW			1.652	20	1.653	3
231	1.632 1.620)											1.632	2
321 321 132	1.618											1.616	1
132	1.614											1.010	
222 222 132 12 <u>3</u> 213	1.594 }	1.594	7	1.5948	2	1.593	wm	1.599	2	1.593	20	1.598	5
132	1.592 ⁽ 1.574 (11050		11050	·
123	1.572	1.571	2	1.5722	<1	1.574	W	1.575	1	1.572	20	1.572	4
213	1.564												
240 132	1.563 1.556											1.560	1
141	1.556	1.550	5	1.5495	r	1.549	WM	1.547	٦	1.547	20	1.547	6
<u>41</u> T	1.5141												
T41	1.513	1.513	10	1.5134	4	1.513	m	1.515	2	1.511	30	1.517	7
21 <u>3</u> 332	1.478 1.465	1.478 1.464	10 1	1.4773	3	1.477	m	1.475	2	1.476	30	1.479	6 2
312	1.440)	1.404	•	1.4638	<1	1.465	VW					1.465	2
330	1.440	1.438	15	1.4395	5	1.440	me					1 441	14
T23	1.439	1.430	15	1.4355	5	1.440	ms					1.441	14
123 2 <u>4</u> 1	1.438) 1.430 (
240	1.430}			1.4273	1	1.429	wm					1.430	5
03 <u>3</u> 303	1.403	1.403	2	1.4037	٦	1.405	W					1.404	3
303	1.387	1.386	1	1.3866	<1	1.387	VW					1.388	3
150 042	1.369	1.370	1			1.373	VW					1.370	1
322	1.362 } 1.360 }											1.365	1
32 <u>2</u> 323 213	1.354												
213	1.354	1.353	3	1.3534	<1	1.353	w						
22 <u>3</u> 42 <u>2</u>	1.353 1.353		-										
251	1.305	1.305	5	1.3051	1	1.305	m						
			-		•								

	1.2753	<1	1.276	vw
5	1.2499	1	1,250	wm
			1.242	vw
3	1.2257	1	1.226	wm
1			1.216	W
5	1.1969	1	^{1.197} }	Bm
	1.1936	1	1.192	5
2	1.1832	<1	1.183	wm
		_		VW
				wm
	1.1146	<].		WIII
				WIII
				wm
			1.010	m
	3	5 1.2499 3 1.2257 1 1.1969 5 1.1936	5 1.2499 1 3 1.2257 1 1 .1969 1 5 1.1936 1 2 1.1832 <1 1.1615 <1	$\begin{array}{c ccccccccccccccccccccccccccccccccccc$

This study. CuX α radiation, λ = 1.5418Å, 114.6mm Gandolfi camera, NBS Si as internal standard, visual intensities. Hathaway & Schlanger (1962, 1965). CuX α^1 radiation, λ = 1.54050Å, diffractometer cutoff 2°(20). Wall $e^{\pm} \alpha I$. (1962). Fe radiation, 114.8mm camera, s = strong, m = moderate, w = weak, v = very, B = broad. Goldberry &Loughnan (1977). Sample from the Glen Alice area near Mt. Canobola. Marić (1968).Diffractometer data, CuX α radiation. ٦.

2.

3.

4. 5.

6.

Bosmans (1970). Diffractometer data, Cux α radiation, λ = 1.5418Å, quartz as internal standard, net integrated intensities.

1968) showed the stacking in nordstrandite to be ABBCCA.

The polymorphism in Al(OH)₃, however, is apparently more complicated than is generally triclinic "hydrargillite" from believed. Α Schischimsk, Urals, with a 17.338, b 10.086, c 9.730 Å, α 94°10′, β 92°08′, γ 90°00′ and Z=32 was reported by Saalfeld (1960) as forming oriented intergrowths with gibbsite. In a private communication (1979), Saalfeld stated that many years later the same triclinic phase with the same oriented intergrowths with gibbsite was found again in a bauxite deposit in India. Despite the scanty data available, this triclinic phase is almost certainly not nordstrandite, as the powder pattern of nordstrandite cannot be indexed on the basis of the large triclinic cell.

Recently, an unidentified mineral, UK45, from Mont St-Hilaire (Chao & Baker 1979)

] Mont	2	3	4	5	6	7
St-Hilaire Quebec	Guam	Rio Blanco Colorado	Sarawak Borneo	Narssårssuk Greenland	USSR	Croatia
Nepheline syenite	Limestone solution cavities	Fissure- fillings in dolomitic marlstone & oil shale	Limestone residual soil	Pockets in nepheline syenite	Carbonate rocks solution cavities	Limestone terra rossa
Tabular & blocky xls	Aggregates of bladed & flamboyant xls	Platy & fibrous aggregates	Rhombic xls, polycrystal- line aggre- gates	Spherules, sheaf-like aggregates of acicular xls	Stringers, aggregates of acicu- lar xls	Leaf-like aggregates
Colorless, white, pink,	Colorless, white	White	Colorless, pink	Cream, snow-white	White	White, pink
~ 3	3					
2.42	2.43		2.43	2.41		2.5
1.580(1)(Na) 1.583(1) 1.602(1)	1.580(4) 1.580(4) 1.596(4)	1.570	1.580 1.581 1.613	1.579(2)(Na) 1.579(2) 1.584(1)	1.58 1.58 1.596	1.565-1.570 1.565-1.570 1.585-1.590
24 ⁰ 21.9 ⁰	low 00		18 ⁰ 10.2 ⁰	o ^o	00	10-40 ⁰
(-)	(-)			(+)		
	Quebec Nepheline syenite Tabular & blocky xls Colorless, white, pink, pale green ~ 3 2.42 1.580(1)(Na) 1.602(1) 24 ⁰ 21.9 ⁰	Mont St-Hilaire QuebecGuamNepheline syeniteLimestone solution cavitiesTabular & blocky xlsAggregates of bladed & flamboyant xlsColorless, white, pink, pale green ~ 3Colorless, white2.422.431.580(1)(Na) 1.583(1)1.580(4) 1.580(4) 1.596(4)240 21.90low 00	Mont St-Hilaire QuebecGuamRis Blanco ColoradoNepheline syeniteLimestone solution cavitiesFissure- fillings fin dolomitic marlstone & oil shaleTabular & blocky xlsAggregates of bladed & flamboyant xlsPlaty & fibrous aggregates aggregatesColorless, white, pink, pale green ~ 3Colorless, whiteWhite2.422.431.580(4) 1.580(4)1.5701.602(1)1.580(4) 1.9901.570	Mont St-Hilaire QuebecGuamRio Blanco ColoradoSarawak BorneoNepheline syeniteLimestone solution cavitiesFissure- fillings in dolomitic marlstone & oil shaleLimestone residual soil marlstone & oil shaleLimestone residual soil polycrystal- line aggre- gatesTabular & blocky xlsAggregates of bladed & flamboyant xlsPlaty & fibrous aggregatesRhombic xls, polycrystal- line aggre- gatesColorless, white, pink, pale green ~ 3Colorless, whiteWhiteColorless, pink2.422.432.422.431.580 1.580(1) 1.580(4)1.5701.580 1.581 1.613240 21.90low 00180 10.20	Mont St-Hilaire QuebecGuamRio Blanco ColoradoSarawak BorneoNarssarssuk GreenlandNepheline syeniteLimestone solution cavitiesFissure- fillings to dolomitic marlstone & oil shaleLimestone residual soilPockets in nepheline soilTabular & blocky xlsAggregates of bladed & flamboyant xlsPlaty & fibrous aggregatesRhombic xls, polycrystal- line aggre- gatesSpherules, saggregates of actular xlsColorless, white, pink, pale green ~ 3Colorless, whiteWhiteColorless, pinkCream, snow-white2.422.432.432.432.411.580(1)(Na) 1.580(4)1.580(4) 1.596(4)1.5701.580 1.581 1.581 1.581 1.584(1)1.579(2)(Na) 1.584(1)240 21.9010w 00180 10.200000	Mont QuebecGuamRio Blanco ColoradoSarawak BorneoNarssarssuk GreenlandUSSRNepheline syeniteLimestone solution cavitiesFissure- fillings to dolomitic marlstone & oil shaleLimestone residual soilPockets in nepheline soilCarbonate rocks solution cavitiesTabular & blocky xlsAggregates of bladed & flamboyant xlsPlaty & fibrous aggregatesRhombic xls, polycrystal- jine aggre- gatesSpherules, seaf-like aggregates of actcular xlsStringers, aggregates of actcular time signedStringers, aggregates of actcular time signedStringers, aggregates of actcular time signedStringers, aggregates of actcular time signedColorless, white, pink, pale green ~ 3Colorless, whiteWhiteColorless, pinkCream, snow-whiteWhite1.580(1)(Na) 1.580(4)1.580(4) 1.596(4)1.5701.580 1.581 1.6131.579(2)(Na) 1.584(1)1.58 1.596240 21,90100180 10,20000

TABLE 5. OCCURRENCES AND PHYSICAL PROPERTIES OF NORSTRANDITE

This study Hathaway & Schlanger (1962, 1965) Milton $et \alpha I$. (1975)

Wall et al. (1962)

Petersen *et al.* (1976) Kulikova *et al.* (1974) Maric (1968) 5. 6.

has been established to be a new polymorph of $Al(OH)_3$. The mineral is triclinic with a 5.002, b 5.175, c 4.980 Å, α 97.50°, β 118.60°, γ 104.74° and Z=2, distinct from all known phases of Al(OH)₃ including the triclinic "hydrargillite" of Saalfeld. Moreover, a nordstrandite-like mineral has been found in association with nordstrandite on a specimen from Mont St-Hilaire. It occurs as powdery sprays and as spherulitic aggregates of extremely finegrained crystals. Its composition is close to Al(OH)₃, and its powder diffraction pattern is very similar to that of nordstrandite. The major difference in the X-ray powder pattern occurs in the shift of some moderately strong lines by up to 0.05 Å in either direction. Some lines are noticeably diffuse. This mineral, tentatively designated UK51, is thus most likely a new phase of Al(OH)₃, perhaps a phase with stacking disorder. A full description of these phases from Mont St-Hilaire will be presented in forthcoming papers.

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