

Radiometric geochronology in the Eastern Alps: results and problems

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ABSTRACT. — The Eastern Alps are one of the European regions in which geological research has benefitted by the highest number of radiometric data and the most significant. The contributions of radiometric geochronology to the reconstruction of the Alpine, Hercynian, « Caledonian » and older events are critically discussed. All radiometric data available in the literature up to 1984 are listed in the attached tables. Some specific problems involving geochronology in the Eastern Alps are discussed in detail. They include: 1) the radiometric age of the « Caledonian » metamorphism, which is confirmed to be close to 460 Ma, as the first result by BORSI et al. (1973) indicates; 2) the hypothesis of a Paleozoic Megacycle, to which all geological processes developed during the Paleozoic (including the Hercynian events) should be referred. This hypothesis is rejected and a two-event Paleozoic history is suggested (« Caledonian » plus Hercynian events); 3) the existence of Caledonian regional uplifting, for which new evidence is shown; 3) the hypothesis that the presumed « Caledonian » processes are Cadomian (or Assynthian), a hypothesis which is rejected on the basis of radiometric data and geologic reasoning.

Key words: Radiometric Geochronology, Eastern Alps, « Caledonian » processes, Hercynian processes, Alpine processes.

RIASSUNTO. — Le Alpi Orientali sono una delle regioni europee in cui la ricerca geologica ha potuto disporre di un elevatissimo numero di dati radiometrici di elevata significatività. In questa nota vengono criticamente discussi i contributi dati dalla geocronologia radiometrica alla ricostruzione degli eventi Alpino, Ercinico, « Caledoniano » ed anteriori. Tutti i dati radiometrici disponibili in letteratura fino al 1984 sono elencati nelle tabelle indicate. Vengono discussi in dettaglio alcuni specifici problemi che coinvolgono la geocronologia, e precisamente: 1) l'età radiometrica del metamor-

fismo « Caledoniano », che è confermata prossima a 460 Ma circa, in accordo con i primi risultati di BORSI et al. (1973); 2) l'ipotesi del Megaciclo Paleozoico, al quale apparrebbero tutti i processi geologici verificatisi nel Paleozoico, inclusi quelli Ercinici: tale ipotesi viene respinta, a favore di una articolazione della storia paleozoica su due eventi ben distinti, il « Caledoniano » e l'Ercinico; 3) l'esistenza di un sollevamento regionale « Caledoniano », per il quale viene messa in luce una nuova prova; 4) l'ipotesi che i processi ritenuti « Caledoniani » siano in realtà Cadomiani (o Assynthici), ipotesi che viene respinta sulla base di dati radiometrici e considerazioni geologiche.

Parole chiave: Geocronologia radiometrica, Alpi Orientali, Processi « Caledoniani », Processi Ercinici, Processi Alpini.

Introduction

Following the title closely, the present report consists of two different sections.

The first summarizes the most important radiometric results, selected among those which contributed most to the construction or improvement of the presently available regional geologic picture.

The second section is a discussion of some problems related to radiometry in the Eastern Alps, crucial for an understanding of pre-Alpine geology in this region. The specific items proposed for discussion were taken into consideration because, in the authors' opinion, they represent the main geological problems in which the type of geological thinking is crucially controlled by the type of consideration given to the geochronological data.

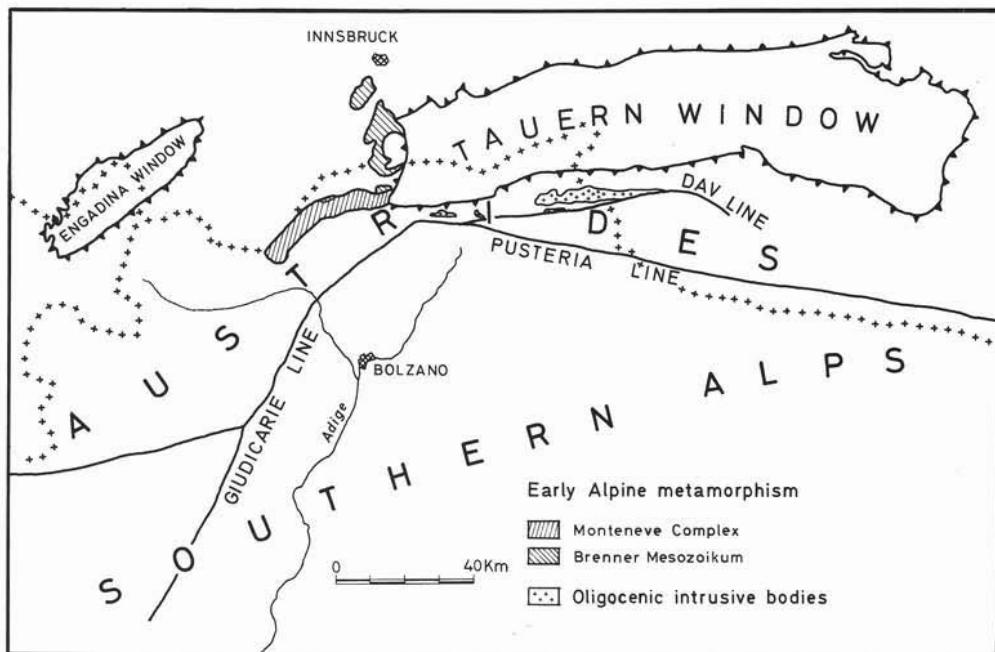


Fig. 1. — Distribution of the early Alpine (Cretaceous) metamorphic effects and location of the Oligocene intra-Austroalidic alignment Rensen - Vedrette di Ries.

1. Results

The Eastern Alps are one of the European regions in which geological research has benefitted by the highest number of radiometric data, and the most meaningful. These top-level results are mainly due to the systematic joint efforts carried out by authors from Pisa and Padova, as well as the important activity developed by scientists from Bern, Hannover, Leeds and Vienna.

The history of radiometric geochronologic research in the Eastern Alps is short, only about fifteen years. Indeed, the first significant isotopic datations go back to the late 1960s (BORSI et al., 1966, 1968, 1969; BORSI & FERRARA, 1967; MILLER et al., 1967; SCHMIDT et al., 1967; GRAUERT, 1969), and most of the presently available age values were produced in the second half of the 1970s and during the present decade.

The most important age results will be mentioned below, following a chronological classification.

1.1. ALPINE EVENTS

1.1.1. Early Alpine regional metamorphism

The Cretaceous age of the metamorphism in the Brenner Mesozoikum (fig. 1) was demonstrated by MILLER et al. (1967) by means of some Rb-Sr data obtained on biotites from the Raibl beds. The regional importance of this Early Alpine metamorphism and its implication in the metamorphic evolution of the pre-Mesozoic basement soon began to come to light: in fact SCHMIDT et al. (1967) supplied some Rb-Sr data demonstrating that biotite in the Monteneve Complex (Schneebergerzug; fig. 1) has the same Cretaceous age as the biotite in the Brenner Mesozoikum.

This very important result, however, was not sufficient to avoid the development of a widespread controversy concerning the age of the main metamorphism in the Monteneve Complex: some authors attributed a pre-Alpine age (e.g., PURTSCHELLER, SCHMIDT)

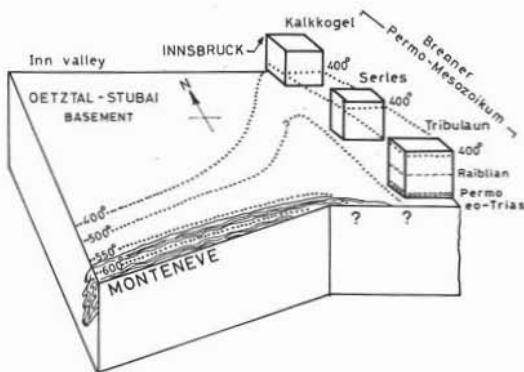


Fig. 2. — Reconstruction of the northern part of the early Alpine thermal structure (taken from SASSI et al., 1980).

to this metamorphism, others an Alpine age (e.g., JUSTIN-VISENTIN & ZANETTIN, SASSI). This controversy has only died down in the last few years, when the bulk of data was so rich and clear (see summarizing pictures in SASSI et al., 1980; DEL MORO et al., 1982; THÖNI, 1983) that no objections could be raised.

These data, also based on different lines of evidence including Rb-Sr and K-Ar isotopic interpretations, demonstrated that the Cretaceous metamorphic temperature was sufficiently high to produce the crystallization of kyanite + staurolite in the Monteneve Al-rich metapelites. Furthermore, the regional character of the related thermal anomaly was clearly shown: a NE-SW-stretched, dome-like, Cretaceous thermal structure affected a large part of the Austroalpine block to the west of the Tauern Window (Oetztal-Breoni area plus Monteneve Complex plus Merano-Mules basement) as sketched in figs. 2 and 3). All these data revealed that the pre-Alpine interpretation was erroneous.

1.1.2. Alpine regional metamorphism

While the Early Alpine metamorphism was revealing all its importance in the above-mentioned areas by means of the determining support of radiometric geochronology, the same isotopic methods were important for determining the age and distribution of the Alpine metamorphism.

This event developed close to the Eocene-Oligocene boundary, approx. 40 Ma ago. The related regional cooling began approx. 30 Ma ago, finishing approx. 12-20 Ma ago. These results are mainly due to BESANG et al. (1968), JÄGER et al. (1969), KREUZER et al. (1973), SATIR (1975), BORSI et al. (1978 a) and CLIFF (1981).

The most important result concerning the Alpine metamorphism in the Eastern Alps is related to its regional distribution. To the same extent that all authors were, rightly, convinced that the typical Alpine domain is represented by the Pennines (Tauern Window), almost all of them also firmly believed that the Austroalpine were not involved in the Alpine metamorphism. On the contrary, BORSI et al. (1973, 1978 a) demonstrated that large parts of the Austroalpine south of the Tauern Window were significantly affected by the Alpine recrystallization (fig. 4). More precisely, their Rb-Sr data on micas demonstrated that the northernmost Austroalpine block — lying between the Pennine-Austroalpine boundary to the north and the Defereggener-Anterselva-Valles tectonic line to the south — was completely affected by Alpine recrystallization, the mineral effects of which were described by CONTINI & SASSI (1980).

Therefore, radiometric data were able to display the unexpected importance of a regional tectonic line, which had been known for a long time (DAL PIAZ, 1936) but which had been consigned to oblivion by almost all authors.

1.1.3. Post-Permian magmatism

Other important results of radiometric research are related to the intra-Austroalpine Periadriatic Alignment of granitoid bodies (fig. 1): Rensen - M. Alto - Rio Vena - Cima di Vila - Vedrette di Ries (+ Pollard, in the surroundings of Lienz). Their Alpine age (30 Ma) was demonstrated by BORSI et al. (1978 a, 1978 b, 1979). Furthermore, the same Rb-Sr isotopic data suggested a complex origin for these acidic melts, in which Sr/Sr ratios are significantly lower than those of the typically crustal acidic melts.

As regards the contribution of radiometric research to the reconstruction of the post-Hercynian geology of the Eastern Alps, we must also mention here the important re-

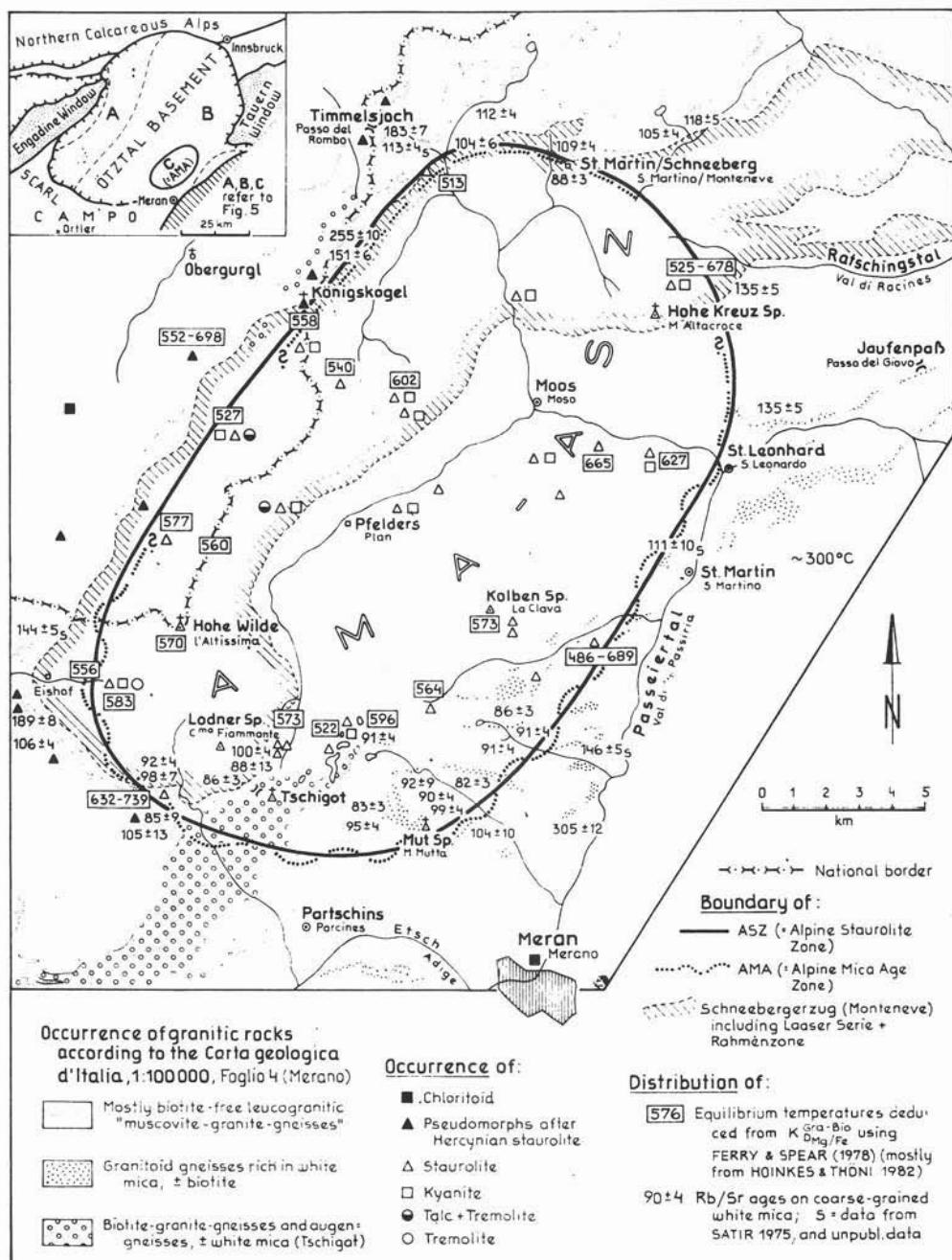


Fig. 3. — Distribution of white mica Rb/Sr ages and equilibrium temperatures in Schneebergerzug (Monte Neve) area and the bordering Altkristallin (taken from THÖNI, 1983).

sults regarding Triassic magmatism (BORSI & FERRARA, 1967; BORSI et al., 1968; FERRARA

& INNOCENTI, 1974) and Tertiary volcanism (BORSI et al., 1969) in the Southern Alps.

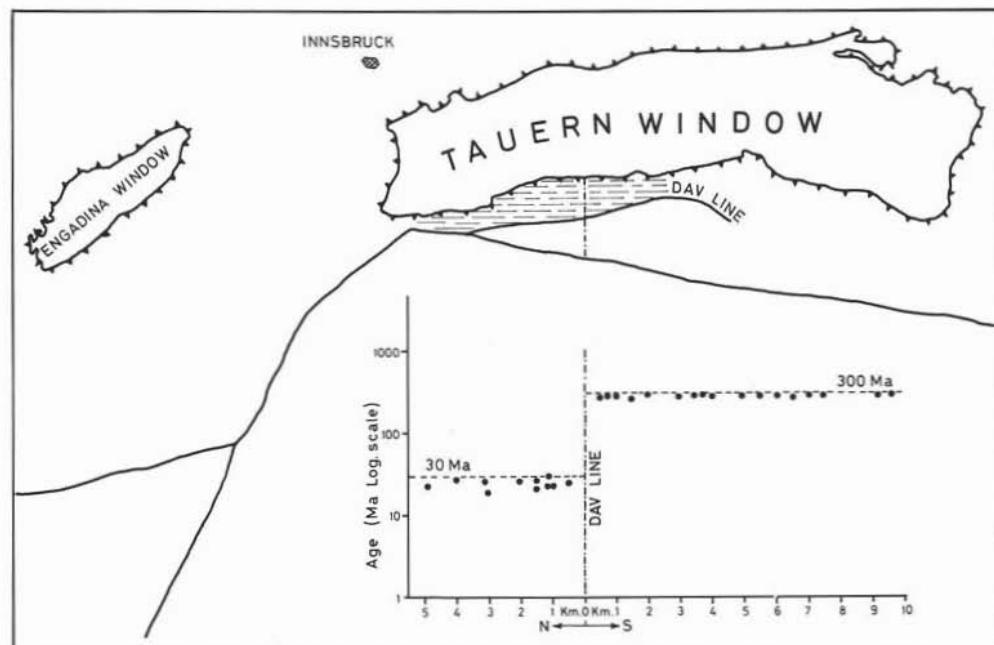


Fig. 4. — Distribution of the Alpine recrystallization within the Austridic basement to the south of the Tauern Window (taken from BORSI et al., 1973, 1978 a).

1.2. HERCYNIAN EVENTS

1.2.1. Hercynian granitoid plutonism

Let us begin from the Hercynian granitoids occurring within the Southern Alps, where many radiometric data finally managed to extinguish the age-old debate concerning the Alpine versus Hercynian age of the Periadriatic Plutons (fig. 5).

Some authors maintained their Paleozoic age, mostly giving reliability and confidence to the very rare and ambiguous finding of granitoid pebbles in the Permian conglomerate (e.g., SANDER, 1923). Instead, other authors (e.g., DAL PIAZ, 1942), considering the unquestionably Alpine age of the Adamello massif (which produced contact metamorphism on Mesozoic rocks) and the particular localization of all these granitoid bodies along very important tectonic lines of Alpine age (Giudicarie and Insubric Lines), supported an Alpine age for the whole granitoid alignment.

Radiometric data demonstrated the Upper Paleozoic age of these granitoids (age range 280-290 Ma: BORSI et al., 1966, 1972; DEL MORO & VISONÀ, 1982, and unpubl. data),

showing that in geology a very good line of reasoning (such as that used by the Alpine party: DAL PIAZ, 1942) sometimes turns out to be wrong because a completely unexpected new « truth » comes to light. In the case in point, the new unexpected truth was the existence of a Hercynian paleo-Insubric Line, which clearly controlled the injection of the Upper Paleozoic melts, and was then significantly reactivated during the Alpine orogeny. It is worth pointing out here that this important result of large-scale structural geology was obtained by means of radiometric geochronologic research.

The above-mentioned age of the Periadriatic granitoids is not very different from that obtained for the Cima d'Asta massif (fig. 5: 274 Ma, BORSI et al., 1974). Similar age values were also obtained for the Hercynian part of the Eisenkappel massif, located along the Insubric Line in Austria, across the boundary with Yugoslavia.

All the reported age values indicate that the Hercynian granitoid plutonism developed in the Southern Alps and along the Periadriatic Lineament during the Upper Carboniferous and Lower Permian, and that it

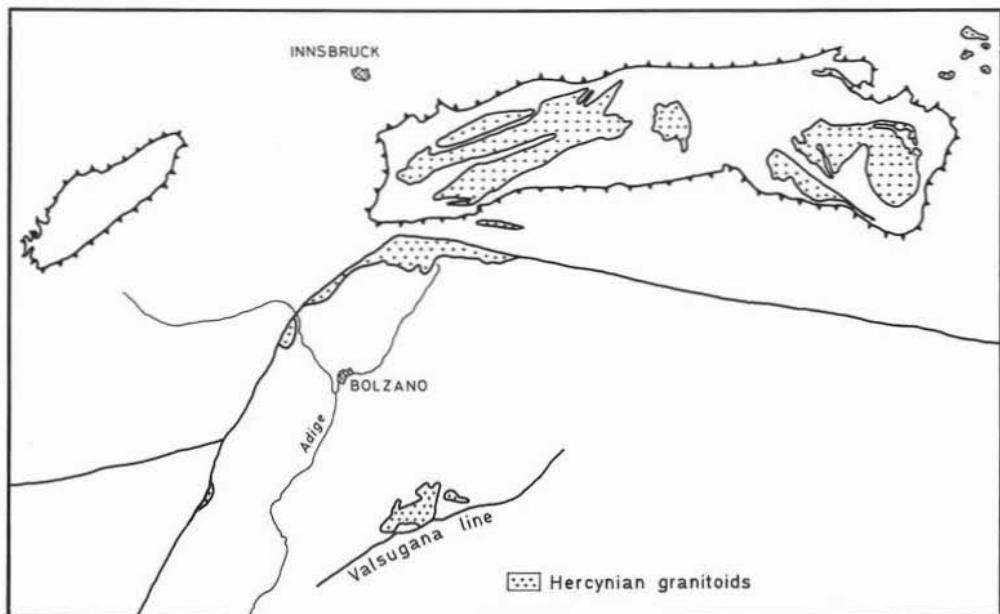


Fig. 5. — Location of the Hercynian intrusive bodies.

was preceded by the emplacement of basic to intermediate rocks, dated at 282 ± 14 Ma (DEL MORO & VISONÀ, 1982) in the Bressanone district.

Instead, as regards the Hercynian granitoids *within the Austroalps*, they seem to be very rare. In fact, these rocks can be recognized only by means of radiometric age determinations and not by field observation or petrographic analyses, and only a few acidic bodies turn out to be Hercynian in age (fig. 5; 262 Ma by BORSI et al., 1980; 289 Ma by KÖHLER, fide SÖLLNER et al., 1982; 240 Ma by CAVAZZINI, 1983; 248 ± 29 Ma by JÄGER & METZ, 1971).

The predominance of relatively young age values (in comparison with those found in the Southern Alps) may be explained by the fact the dated Austroalpine rocks are mostly aplites and pegmatites.

The extremely small amount of Hercynian granitoids within the Austroalps demonstrates that the crustal level rich in Hercynian granitoids was not yet reached by post-Alpine plus post-Hercynian erosion activities: only the uppermost part of the Hercynian metamorphic domain was split open — the part

in which only a few, very mobile, aplitic and pegmatitic veins were injected from the hotter (and granitoid-richer) underpart.

As regards the Hercynian granitoids *within the Pennines*, we may say that a very important Hercynian plutonism is well recorded. In fact, the core of the Tauern Window consists entirely of Hercynian granitoids (fig. 5), altered into gneisses (the so-called Central Gneisses) by the Alpine metamorphism. The age of this plutonism is certainly Upper Paleozoic. However, the radiometric chronologic picture is not so clear as initially believed, and complicated problems have recently been pointed out (CLIFF, 1981). Considering that the whole of this very important set of problems concerns Austrian (and not Italian) territories, we will not give it further space here, but we recommend our readers to become aware of these problems by reading CLIFF (1981) and references quoted in it.

1.2.2. Hercynian regional metamorphism

Its two-phase development in the Southern Alps was pointed out nearly twenty years

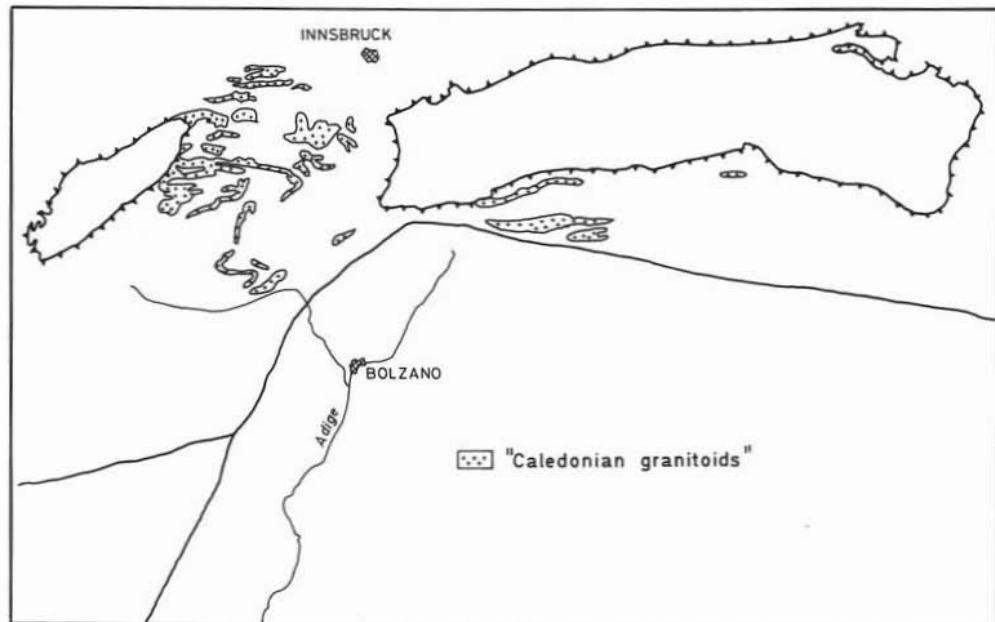


Fig. 6. — Location of the main acidic magmatic bodies of « Caledonian » age.

ago (e.g., SASSI & ZIRPOLI, 1968). Recently, radiometric geochronologic data have confirmed this two-phase development and assigned an age value of 350 Ma (corresponding to the Viséan according to the time-scale proposed by ODIN, 1982) to the older phase, and 320 Ma (corresponding to the Viséan-Namurian boundary according to the same time-scale) to the younger one (DEL MORO et al., 1980, 1984).

It is worth pointing out that these age values obtained in the Southern Alps agree perfectly with those obtained from the Lower Austridic rocks to the north of the Tauern Window by SATIR & MORTEANI (1979). Furthermore, some age values obtained from micas in Austria fall in this range, showing on one hand the significant consistency of the Hercynian metamorphism, and on the other, the synchronism of the Hercynian metamorphic processes on a regional scale.

Numerous Rb-Sr data from biotites indicate that the Hercynian regional cooling took place in the time range 310-300 Ma. The fact that these age values have systematically been found throughout the Eastern Alps in

the districts which escaped Alpine reworking shows that no parts of the pre-Permian basement are known in the Eastern Alps (Pennides, Austrides, Southern Alps) which escaped the Hercynian metamorphism. This metamorphism however was very low-grade in western Carnia and vanishes eastwards.

1.3. « CALEDONIAN » EVENTS

The contribution provided by radiometric geochronology to the unravelling of pre-Hercynian events is very important. However, in every case this contribution was the datation of a geological process which had previously been detected by means of geologic and petrologic research.

1.3.1. « Caledonian » acidic magmatism

During the first stage of this research, geologic and petrologic criteria indicated the existence within the Austrides of numerous old granitoids, the emplacement of which must have been pre-Carboniferous, i.e., such that Viséan metamorphism (350 Ma old) could have affected them. Later, radiometric geochronology demonstrated the Upper Or-

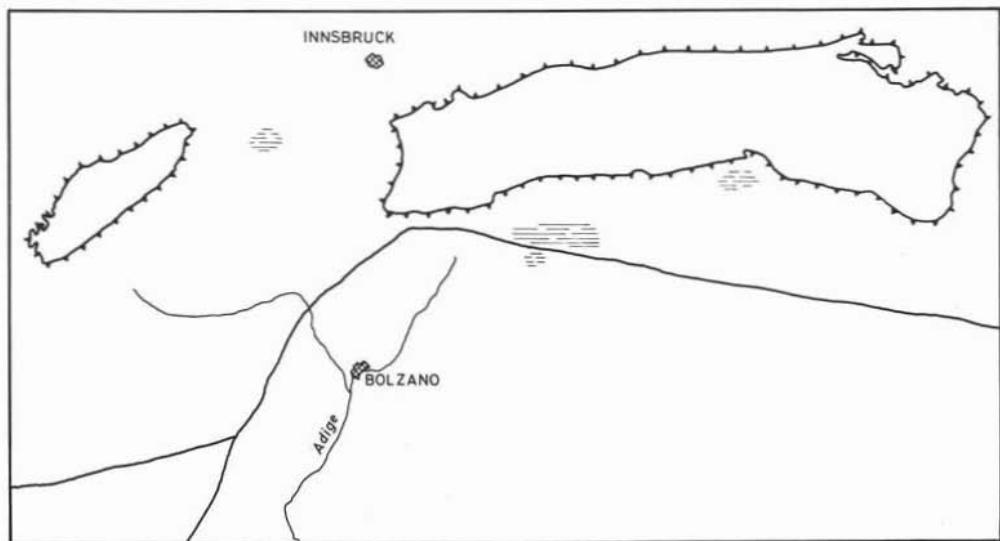


Fig. 7. — Approximate source areas of metasedimentary rock samples which supplied Rb/Sr « Caledonian » age values.

dovian age of this widespread plutonism — the most important recorded in the Austroalps of the Eastern Alps. The age range is 450-420 Ma (HARRE et al., 1968; BORSI et al., 1973, 1980; SATIR, 1975; mostly Rb-Sr isochrons on whole rocks), and 450-440 Ma if a more restrictive selection of the data to be included in the statistical analysis is carried out (SASSI & ZIRPOLI, 1979).

A similar age range was obtained from the acidic sheet-like gneisses, sometimes characterized by an augen texture (SCHMIDT et al., 1967; HARRE et al., 1968; SATIR, 1975; BRACK, 1977; BORSI et al., 1980; HAMMERSCHMIDT, 1981; SÖLLNER & SCHMIDT, 1981; mostly Rb-Sr whole-rock isochrons). These rocks probably represent old volcanics (HERITSCH & TEICH, 1976; SASSI & ZIRPOLI, 1979; HEINISCH & SCHMIDT, 1982), i.e., outpoured products of the same acidic magmatism which also produced the above-mentioned abundant plutons.

If the amount of all these melts (emplaced either in plutonic or volcanic conditions) is evaluated (fig. 6) and the coeval South-Alpine and Austroalpic porphyroids are also taken into consideration (BELLINI & SASSI, 1981; HEINISH, 1981), this Ordovician magmatic cycle — called « Caledonian » acidic magmatism or, better, « Upper Ordovician granite-

rhyolite association » — turns out to be extremely important, the most important magmatic activity recorded in the presently outcropping crustal levels in the whole of the Eastern Alps. In this respect, an intriguing piece of information is that a granitoid having a similar age (446 ± 18 Ma) has also been found underneath the Po Plain, in actual fact about 5000 meters below Venice (Assunta hole, drilled by AGIP in 1976).

1.3.2. « Caledonian » regional metamorphism

Field data SASSI & ZANFERRARI, 1972; BORSI et al., 1973; PURTSCHELLER & SASSI, 1975) indicate that the above-mentioned plutonic bodies were emplaced in previously metamorphosed foliated rocks. Furthermore, it was pointed out (SASSI et al., 1974) that the abundant production of Ordovician acidic melts cannot have been unrelated to a regional thermic anomaly, which must necessarily also have produced a regional metamorphism. Therefore, field data and theoretical expectations agree in maintaining the existence of a « Caledonian » regional metamorphism, and in assigning to it a slightly older age than that of the Caledonian acidic plutonism.

Radiometric geochronology has confirmed these expectations. BORSI et al. (1973) obtained the first significant although not perfectly defined (see discussion below, in the second section of this paper) — age value: 497 ± 38 Ma (Rb-Sr whole-rock isochron on paragneisses). Similar values have been obtained in the neighbouring Austrian areas by BRACK (1977; 466 ± 166 Ma; Rb-Sr whole-rock on paragneisses); SÖLLNER & SCHMIDT (1981; 460 ± 30 and 463 ± 37 Ma; Rb-Sr whole-rock isochrons on paragneisses and migmatites respectively; ca. 460 Ma: U/Pb on zircons). Finally, DEL MORO et al. (1984) obtained an identical value (463 ± 18 Ma) from an albite-muscovite Rb-Sr isochron from the Pusteria phyllites (Southern Alps).

All these data (the source area of which is sketched in fig. 7) are extremely consistent in indicating that the age of the «Caledonian» regional metamorphism in approx. 460 Ma.

1.4. THE OLDEST (PRE-CAMBRIAN) EVENTS

No systematic radiometric data exist, for the Italian part of the Eastern Alps, concerning pre-Cambrian events. However, it is useful to summarize here briefly, for the sake of completeness, the presently available data concerning the whole of the Eastern and Central Alps.

In the Eastern Alps, the pre-Cambrian stage is only indicated by a few radiometric age values and U/Pb data on detrital zircons contained in the metasediments (paragneisses) from the Silvretta Kristallin. Integrating these data with other zircon age values from the Central Alps, major pre-Cambrian events producing zircons may be admitted at 2000, 1500 and 700 Ma.

However, these pre-Cambrian ages have only been obtained from isolated zircon grains and not from their host-rocks, which in all cases turn out to be younger — Paleozoic — in their magmatic or metamorphic ages. Therefore, neither pre-Cambrian magmatic rocks nor metamorphic rocks have been found in the Eastern Alps.

As regards sedimentation, the only pre-Cambrian age which may at present be supposed refers to the sedimentation of the protoliths of the paragneisses and micaschists making up the bulk of the Austridic ba-

sediment: a maximum age value of 800 ± 125 Ma has been estimated by SCHMIDT & SÖLLNER (1982), by means of an integrated interpretation of all isotopic data available for the Eastern Alps.

2. Discussion of some « Caledonian » geochronological problems

Let us now consider in detail some geochronological problems concerning the « Caledonian » event in the Eastern Alps. This topic is proposed here because, in the authors' opinion, it represents the major problem in the Alps in which geochronology has been used, improperly and on a pretext, by some authors for stocking the fires of unnecessary controversies.

2.1. WHAT IS AGE OF THE « CALEDONIAN » METAMORPHISM?

The controversy on the age of the « Caledonian » metamorphism will first be summarized and analysed.

As mentioned above, BORSI et al. (1973) published a Rb-Sr isochron on the Pusteria paragneisses, specifying that the data points fit a fan of parallel isochrons (related to an age value in the range 450-500 Ma) better than the calculated isochron value of 497 ± 38 Ma.

SATIR (1975) criticized this age result, and reinterpreted BORSI et al.'s isotopic data, classifying them into two sample subgroups. However, this classification is not based on any geological or petrographic reasoning. Furthermore, Satir's age data calculated from each of these two isochrons (350 ± 70 and 370 ± 134 Ma) are at least as meaningless as the age value proposed by BORSI et al. (1973), both from the radiometric (compare the spread of the errors) and from the geological viewpoints.

It is astonishing to realize that the type of defect which stimulated Satir's criticism (i.e., the scattering of data points along the isochron) affects most of the published and uncriticized isochrons, and very often the spread of the scatter is larger than that which aroused Satir's sensibilities. If the criteria used in Satir's criticism are applied to all isochrons published in the world, Rb-Sr geochronology would be reduced to very

few exceptionally good results, while the majority of them would be accused of being unreliable and involved in Byzantine and bizarre reinterpretations.

The conclusions are:

(i) Satir's criticism is methodologically incorrect;

(ii) the age value suggested by BORSI et al.'s (1973) isochron on paragneisses maintains all its validity as regards the dating of the «Caledonian» metamorphism in the range 500-450 Ma.

The above statements are supported by the following further radiometric results: 1) Rb-Sr whole-rock isochron on the Winnebach paragneisses by SÖLLNER & SCHMIDT (1981: 460 ± 30 Ma); 2) Rb-Sr whole-rock isochron on the Winnebach migmatites by SÖLLNER & SCHMIDT (1981: 463 ± 37 Ma); 3) U/Pb data on zircons from the same area by SÖLLNER & SCHMIDT (1981: ca. 460 Ma); 4) Rb-Sr muscovite-albite isochron from South Alpine phyllites by DEL MORO et al. (1984: 463 ± 18 Ma); 5) lastly, if we calculate a three-point Rb-Sr isochron on whole rock utilizing the paragneiss data points published by FRANK et al. (1976) for constructing a mixed isochron based on para- and ortho-derivates, an age value of 456 ± 24 Ma is obtained (see SCHMIDT & SÖLLNER, 1982).

All these data indicate that the age of the «Caledonian» metamorphism (ca. 460 Ma) falls close to the age value first reported by BORSI et al. twelve years ago.

2.2. ARE WE DEALING WITH A MEGACYCLE?

Another important problem related to geochronology is ascertaining whether the pre-Hercynian Paleozoic processes represent a single event (i.e., the «Caledonian» Event) or parts of a long-lasting megacycle (i.e., a Paleozoic Megacycle).

The latter interpretation was proposed by HEINISCH & SCHMIDT (1976) and SCHMIDT (1977). After a period of latency, it seems to have found new supporters in that Schmidt demonstrated instead that he was more inclined towards accepting the «Caledonian» event (SASSI & SCHMIDT, 1982).

The megacycle idea was based on a statistical analysis of all radiometric data available in the literature: they cover the whole Paleozoic time range almost continuously. This

result, however, is only obtained if all data are included in the statistical analysis, whatever the type of isotopic system (i.e., K-Ar, Rb-Sr, etc.) and whatever the type of analysed body (i.e., biotite, zircon, whole-rock, etc.).

If this approach is extended to post-Paleozoic times, we obtain a really sensational result: the Megacycle which began in the early Paleozoic and operated throughout the Paleozoic prolonged its activity during the Mesozoic and Tertiary!

Joking apart, as a matter of fact the grounds for the megacycle interpretation vanish completely if the statistical analysis is carried out rigorously, i.e., only putting strictly homogeneous radiometric data (that is, those having identical significance) in the same sample population. SASSI & ZIRPOLI (1979) demonstrated that, following this rigorous approach, a clear two-modal distribution of the age values is obtained, consisting of a «Caledonian» cluster of age values and a Hercynian one. The age data published after SASSI & ZIRPOLI's statistical analysis confirm this result, increasing the statistical weight of both clusters and making the interposed empty space more significant.

Therefore, if our reasoning is kept very close to the facts, the megacycle does not exist at all or, if it does exist, data for detecting it are not yet available in the Eastern Alps. Currently available radiometric data indicate clearly that the Paleozoic thermal history (magmatism and metamorphism) in the Eastern Alps is a two-event history, consisting of a «Caledonian» event (magmatism and metamorphism) of Ordovician age (SASSI & SCHMIDT, 1982) and a Hercynian event (magmatism and metamorphism) of Carboniferous to Lower Permian age.

2.3. CHRONOLOGICAL EVIDENCE FOR A «CALEDONIAN» REGIONAL UPLIFT

A few lines above, we wrote: Paleozoic thermal history. However, readers should not believe that the «Caledonian» event was only thermal. Indeed, dynamic activity related to the «Caledonian» event was reported ten years ago (SASSI et al., 1974), and more recently radiometric geochronology (U/Pb method) has been used for demonstrating tectonic activity in the age range

450-420 Ma (CLIFF, 1980). SASSI & SCHMIDT (1982) agreed in defining the « Caledonian » event in the Eastern Alps as a « tectonic-thermic » event, thus stressing the importance of both its components, tectonic and thermic.

Radiometric geochronology supplies an indirect tool for detecting an effect of an important regional uplift of « Caledonian » age. In fact, radiometric geochronology in the Austrides indicate that:

(i) Alpine granitoids only occur in the Rensen-Vedrette di Ries alignment, the emplacement of which was strictly controlled by a tectonic line (fig. 1). This means that only Alpine granitoids injected into a very high (shallow) crustal level outcrop, and that post-Alpine erosion was not deep enough to affect the level rich in Alpine granitoids;

(ii) Hercynian granitoids outcrop only locally in the Austrides (fig. 5), and they are mostly aplites and pegmatites. In this case too, only rare and shallow granitoids outcrop, indicating that the post-Hercynian erosion of the Austrides was not deep enough to affect the crustal level rich in Hercynian granitoids;

(iii) « Caledonian » granitoids outcrop abundantly in the Austrides (fig. 6). As a matter of fact, the Ordovician crust outcropping below the Upper Ordovician to Silurian

metasediments (i.e., the Austridic phyllitic complexes) is abundantly riddled with Ordovician granitoids (compare the distribution of acidic magmatic rocks within the Austrides in figs. 5 and 6).

Therefore, the Ordovician crust was deeply eroded, and this erosion cannot be related either to the post-Hercynian or to the post-Alpine uplifts, for the above reasons.

These considerations are graphically represented in fig. 8, which shows the present erosion level and the different levels rich in variously aged granitoids. The erosion which abundantly split open the Ordovician plutonic bodies must have taken place definitely after their emplacement (440-420 Ma ago) but, significantly, before the Hercynian metamorphism (Viséan, 350 Ma ago): in fact, this erosion stage must certainly have been followed by a burial stage, necessary for these rocks to have been involved in the Viséan metamorphism, as radiometric data demonstrate.

2.4. « CALEDONIAN » OR CADOMIAN?

A final remark deals with the possibility that the processes considered above as « Caledonian » are instead Cadomian (or Assynthian).

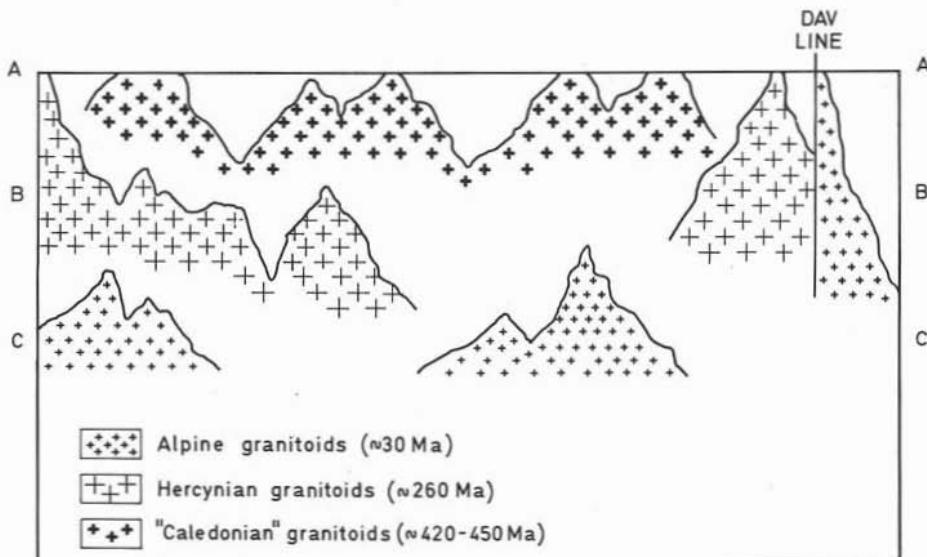


Fig. 8. — Schematic representation of the variously-aged granitoids in the Austrides. - AA: present erosion level; BB: level rich in Hercynian granitoid bodies; CC: level rich in Alpine granitoids.

First of all, it should be noted that the use of the term « Caledonian » only aims at indicating a given time range—that in which true Caledonian event took place in the Caledonides. Any causal or spatial relationships between Ordovician evolution in the Eastern Alps and the Caledonian event in the Caledonides is not implicit in our use of the term « Caledonian » inserted, as it is here, within quotation marks.

VAI (1975) maintains that a Cadomian hypothesis is more well-founded than a « Caledonian » one. However, a detailed comparison of the age values from the Eastern Alps with the true Cadomian and the true Caledonian values led BORSI et al. (1975) to reject Vai's view and to confirm the « Caledonian » interpretation.

More recently, discussing the « Caledonian » age values from Central Europe, DORNSIEPEN (1979) considered them as Cadomian, partially rejuvenated by the Hercynian event.

Two main objections are to be made against Dornsiepen's hypothesis, after having recalled that: (i) a partial (i.e., incomplete) rejuvenation is expected to produce random scatter of data points in the isochron diagram below the original isochron line, i.e., within the field between the original line (isochron of 650 Ma in our case) and the isochron line corresponding to the younger rejuvenating event (350 Ma, in our case); (ii) a new alignment according to a pseudo-isochron could only appear occasionally, as a random and rare *lusus*.

With these premises, the objections are:

- 1) Why should the assumed partial rejuvenation have produced new alignments of

the data points in the isochron diagram everywhere, i.e., the less probable result?

- 2) Why do the unexpected new alignments not correspond to several, random age values in the range 650-350 Ma, but in every case fit isochron lines systematically corresponding to Caledonian age values?

These objections are sufficient to discredit the Cadomian hypothesis definitely. We could also add the observation that geological evidence for a Cadomian event is completely lacking in the Eastern Alps: therefore a Cadomian hypothesis in this region is completely arbitrary, because it has no direct support, either radiometric or geologic. On the contrary, the « Caledonian » hypothesis is not only based on a rich set of well-defined radiometric data, but also on numerous geological indications. Furthermore, the Upper Ordovician age of the « rhyolite-granite association » is not only based on radiometric data, but also on strictly consistent biostratigraphic datations, which indicate that volcanism certainly took place in the Upper Ordovician (FLAJS & SCHÖNLAUB, 1976), and not in pre-Cambrian times, as the Cadomian hypothesis would imply.

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PRESENTATION OF THE TABLES

In order to obtain directly comparable data, three types of standard tables were compiled, one for each of the three radiometric methods. When necessary, isotopic parameters and age values were recalculated by using the recommended decay constants (STEIGER and JÄGER, 1977). However, such recalculation was not possible in some cases, when the used values of the decay constants were not mentioned in the papers.

The numerous Rb/Sr and K/Ar data were assembled following different criteria. Rb/Sr data firstly were divided according to the structural geological units to which they belong: three large groups of data (Southern Alps, Austrides and Pennides) were thus obtained. Inside each of these three groups, data were further assembled considering both the geographical location and the age range. As regards K/Ar tables, these were compiled on geographical base only.

TABLE 1
U/Pb analytical data

Rock type	Sample n.	Mat.	U ppm	Th ppm	Pb _{tot} ppm	$\frac{206\text{Pb}}{204\text{Pb}}$	$\frac{207\text{Pb}}{204\text{Pb}}$	$\frac{208\text{Pb}}{204\text{Pb}}$	AGE Ma				Ref.
									$\frac{206\text{Pb}}{238\text{U}}$	$\frac{207\text{Pb}}{235\text{U}}$	$\frac{208\text{Pb}}{232\text{Th}}$	$\frac{207\text{Pb}}{206\text{Pb}}$	
<u>SE Tauern Window</u>													
Tonalite	C 6716	Zi	854	--	--	4077	--	--	--	--	--	--	[314±7
		Zi	355	--	--	2019	--	--	--	--	--	--	
Biotite augen gneiss	G 5447	Zi	881	--	--	1818	--	--	--	--	--	--	343
		Zi	3268	--	--	498	--	--	--	--	--	--	
		Zi	3313	--	--	824	--	--	--	--	--	--	
		Zi	2857	--	--	961	--	--	--	--	--	--	
		Zi	2990	--	--	1064	--	--	--	--	--	--	
		Zi	3929	--	--	657	--	--	--	--	--	--	
		Zi	4082	--	--	799	--	--	--	--	--	--	[313±10
		Zi	3550	--	--	1218	--	--	--	--	--	--	
		Zi	4203	--	--	1948	--	--	--	--	--	--	
Foliated granodiorite	F T4	Zi	3178	--	--	1237	--	--	--	--	--	--	
		Zi	1654	--	--	2478	--	--	--	--	--	--	
		Zi	1830	--	--	3138	--	--	--	--	--	--	
Grandioritic gneiss	F T7	Zi	1872	--	--	1600	--	--	--	--	--	--	
		Zi	2287	--	--	2086	--	--	--	--	--	--	
Metamorphic tonalite	Ep	Ep	8,47	--	50.09	18.60	15.60	41.91	--	--	--	--	[19
		Sph	93.27	--	11.24	45.90	16.99	49.66	[215	--	--	--	
		WR	10	--	12.85	18.49	15.56	38.98		--	--	--	
		Clinoz.	6.94	--	43.47	18.50	15.59	38.59		--	--	--	
		Clinoz.	7.29	--	46.48	18.53	15.63	38.71	[22	--	--	--	
		Pl	.11	--	7.08	18.45	15.60	38.34		--	--	--	
		KF	.20	--	46.12	18.48	15.61	38.40		--	--	--	
		Ap	18.42	--	4.87	18.60	15.62	38.44		--	--	--	
<u>Detztal-Monteneve</u>													
Two-mica granitic gneiss	KAW 280	Zi	1118	--	84.5	1265.6	85.29	76.71	483±13	480±15	--	476±20	[34
Tonalitic gneiss	B 3/4	Zi	730.5	--	60.2	1057.1	72.63	169.25	474±15	486±19	--	540±25	

TABLE 2
K/Ar analytical data

Rock type	Sample n.	Material	K%	^{40}Ar rd ml/g	^{40}Ar rd%	AGE Ma	Ref.
<u>Euganean Hills</u>							
Basalt	10	WR	1.50	$1.68 \cdot 10^{-4}$	65	42±1.5	6
Lattice	29	Bi+Px	3.48	$1.44 \cdot 10^{-4}$	68	35±1	
Trachyte	8	Bi	6.48	$1.20 \cdot 10^{-4}$	72	30±1	
Trachytic dyke	9	Bi	6.07	$1.29 \cdot 10^{-4}$	73	32±1	
Trachytic dyke	11	Bi	6.79	$1.13 \cdot 10^{-4}$	52	28±1	
Trachyte	14	WR	3.79	$1.30 \cdot 10^{-4}$	71	32±1	
Trachyte	26	Bi	5.96	$1.31 \cdot 10^{-4}$	68	33±1	
Trachyte	27	Bi	6.47	$1.20 \cdot 10^{-4}$	79	30±1	
Rhyolite	12	WR	4.98	$1.36 \cdot 10^{-4}$	90	34±1	
Rhyolite	21	WR	4.25	$1.33 \cdot 10^{-4}$	91	33±1	
Rhyolite	25	Bi	6.26	$1.33 \cdot 10^{-4}$	69	33±1	
Rhyolitic dyke	28	WR	3.98	$1.31 \cdot 10^{-4}$	81	33±1	
<u>Adamello</u>							
Medium grained granodiorite	A/4	Bi	7.58	$11.8 \cdot 10^{-5}$	--	30±2	4
<u>Mules</u>							
Tonalitic granodiorite	AA7	Bi	2.69	$1.76 \cdot 10^{-4}$	13	41±1	7
<u>Vedrette di Ries</u>							
Muscovitic pegmatite	AA74-15	Mu	9.01	$1.14 \cdot 10^{-4}$	48	28±1	12
Muscovitic pegmatite	AA74-16	Mu	11.55	$1.07 \cdot 10^{-4}$	52	27±1	
Tonalite	AA74-46	Bi	8.04	$1.06 \cdot 10^{-4}$	62	26±1	
Tonalite	AA74-60	Bi	7.96	$1.08 \cdot 10^{-4}$	67	27±1	
---	D7401	Bi(Chl)	5.83	$7.27 \cdot 10^{-6}$	75.5	30.9	
---	D7402	Bi	7.29	$8.63 \cdot 10^{-6}$	75.8	29.4	
		Hbl	.821	$1.07 \cdot 10^{-6}$	37.7	32.4	
				$1.23 \cdot 10^{-6}$	57.8	37.1	

TABLE 2 (continued)

K/Ar

Rock type	Sample n.	Material	K%	^{40}Ar rd ml/g	^{40}Ar rd%	AGE Ma	Ref.
<u>Predazzo</u>							
Monzonite	P.te Cast. 65	WR	3.57	$9.66 \cdot 10^{-4}$	44	230	3
Monzonite	Malgola 49c	WR	3.02	$7.44 \cdot 10^{-4}$	45	179	
Monzonite	Cenzoccoli	WR	5.89	$9.94 \cdot 10^{-4}$	89	236	
Vulcanite	Vulcanite 34	WR	3.67	$9.92 \cdot 10^{-4}$	75	212	
Vulcanite	Vulcanite 60	WR	3.67	$9.72 \cdot 10^{-4}$	86	231	
Vulcanite	Vulcanite 103	WR	3.86	$9.13 \cdot 10^{-4}$	92	217	
<u>Monzoni</u>							
---	FT	Phl	7.91	$9.66 \cdot 10^{-4}$	87	229+8	5
---	FP	Phl	7.61	$9.49 \cdot 10^{-4}$	90	225+8	
<u>Bressanone-Ivigna-Mt. Croce</u>							
Granodioritic granite	89	Bi	5.19	$1.22 \cdot 10^{-3}$	83	228+8	7
Granodioritic granite	B13	Bi	4.94	$1.20 \cdot 10^{-3}$	94	230+8	
Granodioritic granite	AA9	Bi	5.60	$1.19 \cdot 10^{-3}$	95	230+8	
Granite-Granodiorite	122	Bi	4.54	$1.19 \cdot 10^{-3}$	78	230+8	
Granodiorite	123	Bi	1.45	$8.71 \cdot 10^{-4}$	34	197+7	
Granite-Granodiorite	126	Bi	5.32	$1.17 \cdot 10^{-3}$	85	229+8	
Granite or Granodiorite	I XIV	Bi	6.24	$9.70 \cdot 10^{-4}$	84	210+8	
Granite-Granodiorite	I19	Bi	7.13	$7.31 \cdot 10^{-4}$	82	167+6	
Fayalite granite	333	Bi	3.49	$1.10 \cdot 10^{-3}$	96	250+3	23
Fayalite granite	79/7	Bi	3	$1.20 \cdot 10^{-3}$	94	279+3	
<u>Mt. Sabion</u>							
Medium grained granodiorite	A/6	Bi	7.44	$80.2 \cdot 10^{-5}$	--	192+7	4
Fine grained granodiorite	A/7	Bi	4.64	$75.8 \cdot 10^{-5}$	--	182+6	
Biotitic pegmatite	A/8	Bi	3.50	$74 \cdot 10^{-5}$	--	178+6	
Muscovitic pegmatite	A/10	Mu	7.54	$63.6 \cdot 10^{-5}$	--	154+5	
<u>SE Tauern Window</u>							
Mu-Bi-Augengneiss	Gr	Bi	6.47	$4.93 \cdot 10^{-6}$	50	19 + .5	20
		Mu	7.87	$7.44 \cdot 10^{-6}$	74	23.5+ .5	
Augengneiss	Tr	Bi	6.75	$6.83 \cdot 10^{-6}$	84	25.2+ .5	
		Mu	8.70	$7.51 \cdot 10^{-6}$	67	20.9+ .5	
Amphibolite	A346	Bi	7.59	$14.91 \cdot 10^{-6}$	85	48.6+1	
Bi-Augengneiss	C308	Bi	5.94	$8.92 \cdot 10^{-6}$	75	37 + 2	
Coarse Augengneiss	C213	Bi	6.69	$5.96 \cdot 10^{-6}$	84	22 + .5	
Bi-Gneiss	Z162	Bi	6.51	$5.70 \cdot 10^{-6}$	72	21.9+ .5	
Bi-Gneiss	Z160	Bi	7.08	$6.09 \cdot 10^{-6}$	76	21.5+ .4	
Aplitic Gneiss	C18	Bi	7.12	$5.53 \cdot 10^{-6}$	82	19.5+1	
Bi-Pegmatite	C212	Bi	7.10	$5.58 \cdot 10^{-6}$	81	19.5+ .5	
Bi-Gneiss	Z150	Bi	7.41	$5.17 \cdot 10^{-6}$	79	17.5+ .2	
Bi-Gneiss	Z151	Bi	7.19	$4.62 \cdot 10^{-6}$	72	16.1+ .2	
Bi-Gneiss	Z159	Bi	7.23	$4.61 \cdot 10^{-6}$	72	16 + .3	
Mu-Schist	C422	Mu	8.48	$7.46 \cdot 10^{-6}$	58	22 + 1	
Mu-Schist	C55	Mu	7.99	$7.04 \cdot 10^{-6}$	79	22 + .5	
Mu-Qz-Schist	C495	Mu	8.74	$7.55 \cdot 10^{-6}$	51	21.5+1	
Mu-Dol-Quartzite	O VI	Mu	7.67	$6.46 \cdot 10^{-6}$	82	21 + .5	
Micaschist	C315	Mu	8.49	$7.21 \cdot 10^{-6}$	61	21 + 1	
Ga-Mu-Schist	C31	Mu	6.97	$4.86 \cdot 10^{-6}$	78	17.5+ .5	
Bi-Amphibolite	A346	Amph(40-80)	.312	$7.89 \cdot 10^{-6}$	94	597 +12	
		Amph(80-120)	.365	$8.74 \cdot 10^{-6}$	86	546 +10	
Amphibolite	A233	Amph(80-120)	.21	$1.61 \cdot 10^{-6}$	64	450 +9	
Amphibolite	A300	Amph(80-120)	.853	$3.44 \cdot 10^{-6}$	88	98.3+2.1	
Amphibolite	A255	Amph(80-120)	.796	$1.77 \cdot 10^{-6}$	69	54.9+1.3	
Amphibolite	C250	Amph(80-120)	.27	$.46 \cdot 10^{-6}$	33	42.9+2.2	
Amphibolite	C170	Amph(80-120)	.244	$.32 \cdot 10^{-6}$	21	32.7+2.8	
Kyanite schist	C523	WR	4.02	$3 \cdot 10^{-6}$	24	18.5+2	20
Phyllite	L122	WR	1.68	$1.22 \cdot 10^{-6}$	53	18.2+ .5	
Calc-phyllite	L52	WR	3.47	$1.63 \cdot 10^{-6}$	69	11.7+ .2	

TABLE 2 (continued)

K/Ar

Rock type	Sample n.	Material (ν)	K%	^{40}Ar rd ml/g	^{40}Ar rd%	AGE Ma	Ref.
<u>SW Tauern Window</u>							
Tonalitic granite	76	Bi	7.62	$113.4 \cdot 10^{-6}$	95	343 \pm 3	2
		Hb1	.764	$104.9 \cdot 10^{-7}$	85	318 \pm 3	
			.749	$94.7 \cdot 10^{-7}$	85	294 \pm 3	
Tonalitic granite	77	Bi	7.66	$105.5 \cdot 10^{-6}$	90	319 \pm 3	
		Hb1	.592	$93 \cdot 10^{-7}$	90	360 \pm 3	
			.750	$94 \cdot 10^{-7}$	85	293 \pm 3	
Tonalitic granite	78	Bi	7.56	$87.6 \cdot 10^{-7}$	75	28.9 \pm .3	
Tonalitic granite	79	Bi	7.65	$90.3 \cdot 10^{-7}$	70	29.6 \pm .3	
Tonalitic granite	83	Bi	7.30	$58.7 \cdot 10^{-7}$	65	20.2 \pm .2	
		Mu	7.82	$74.5 \cdot 10^{-7}$	75	23.9 \pm .2	
Tonalitic granite	82	Bi	7.55	$80.5 \cdot 10^{-7}$	80	26.7 \pm .2	
Hornblendite	403	Bi	7.02	$74.7 \cdot 10^{-7}$	80	26.6 \pm .2	
Hornblendite	404	Bi	7.13	$75.7 \cdot 10^{-7}$	85	26.6 \pm .2	
Tonalitized Augen & Flasergneiss	84	Bi	7.04	$53.9 \cdot 10^{-7}$	65	19.3 \pm .2	
Augen & Flasergneiss	86	Bi	7.13	$63.4 \cdot 10^{-7}$	70	22.3 \pm .2	
Augen & Flasergneiss	87	Bi	7.10	$66.2 \cdot 10^{-7}$	70	23.3 \pm .2	
Aplitic granite	80	Bi	7.26	$60.6 \cdot 10^{-7}$	60	20.9 \pm .2	
Aplitic granite	80a	Bi	5.91	$50.4 \cdot 10^{-7}$	55	21.4 \pm .2	
Aplitic granite	81	Mu	8.64	$73.2 \cdot 10^{-7}$	70	21.3 \pm .2	
Aplitic granite	85	Bi	6.84	$57.5 \cdot 10^{-7}$	55	21.1 \pm .2	
<u>W Tauern Window</u>							
Granitic gneiss	KAW 816	Ph	8.25	$5.06 \cdot 10^{-6}$	74.65	14 \pm 1	33
Granitic gneiss	KAW 811	Ph	8.98	$5.39 \cdot 10^{-6}$	73.79	15 \pm 1	
		Bi (96%)	6.61	$3.85 \cdot 10^{-6}$	62.35	14.6 \pm 1	
			6.88	$3.99 \cdot 10^{-6}$	--	12	
		Chl(90%)	.81	$.88 \cdot 10^{-6}$	29.70	27 \pm 4	
			.14	$.54 \cdot 10^{-6}$	--	97	
		KF	12.28	$19.85 \cdot 10^{-6}$	80.35	40.8 \pm 2	
Granitic gneiss	KAW 824	Ph	8.87	$6.66 \cdot 10^{-6}$	73.73	19 \pm 1	
		Bi	5.16	$3.35 \cdot 10^{-6}$	61.9	16 \pm 1	
Granitic gneiss	KAW 422	Ph	8.24	$7.55 \cdot 10^{-6}$	79.44	23 \pm 1	
Calcareous micaschist	KAW 1318	Mu	6.42	$8.03 \cdot 10^{-6}$	76.14	31 \pm 2	
		Bi	6.28	$4.42 \cdot 10^{-6}$	61.63	17.7 \pm 1	
<u>Merano-Mules-Anterselva basement</u>							
Paragneiss	KAW 1319	Mu	6.61	$9.15 \cdot 10^{-6}$	88.24	34 \pm 2	33
Biotite gneiss	KAW 1320	Bi	7.14	$5.48 \cdot 10^{-6}$	69.14	19 \pm 1	
Leucogranitic gneiss	T 1066	WM(>.43)	8.91	$67.18 \cdot 10^{-6}$	81.75	184 \pm 9	35
Granitoid gneiss	T 1246	WM(>.43)	9	$55.13 \cdot 10^{-6}$	95.65	151 \pm 6	
Granitoid gneiss	T 1247	WM(.15-.43)	8.91	$34.82 \cdot 10^{-6}$	94.89	98 \pm 4	
Granitoid gneiss	T 1285	WM(>.43)	8.94	$28.03 \cdot 10^{-6}$	86.77	79 \pm 3.6	
Leucogranitic gneiss	T 1298	WM(>.43)	8.94	$27.24 \cdot 10^{-6}$	75.17	77 \pm 4	
Leucogranitic gneiss	T 1299	WM(>.43)	9.03	$28.07 \cdot 10^{-6}$	77.45	78 \pm 4	
Leucogranitic gneiss	T 1307	WM(>.43)	8.74	$99.50 \cdot 10^{-6}$	96.02	271 \pm 11	
Leucogranitic gneiss	T 1316	WM(.15-.43)	8.69	$29.55 \cdot 10^{-6}$	94.04	85 \pm 3.6	
Leucogranitic gneiss	T 1319	WM(>.43)	9.03	$28.58 \cdot 10^{-6}$	87.23	80 \pm 3.7	
Leucogranitic gneiss	T 1325	WM(>.43)	8.73	$27.98 \cdot 10^{-6}$	86.37	81 \pm 3.7	
Granite gneiss	T 1349	KF	8.84	$19.69 \cdot 10^{-6}$	44.22	56.4 \pm 5	35
		KF(augen)	--	$19.49 \cdot 10^{-6}$	84.85	55.8 \pm 2.6	
Muscovite-pegmatite	T 1364	WM .15-.43 mm	8.77	$27.62 \cdot 10^{-6}$	65.20	79 \pm 5	
Leucogranitic gneiss	T 1373	WM >.30 mm	8.89	$29.42 \cdot 10^{-6}$	93.35	83 \pm 3.5	
Augengneiss	1554	WM(80-100)	9.42	$21.05 \cdot 10^{-6}$	92.55	56.3 \pm 1.8	28
		WM(100-200)	8.97	$18.98 \cdot 10^{-6}$	63.70	53.3 \pm 2.5	
Augengneiss	1555	WM(100-200)	9.05	$21.62 \cdot 10^{-6}$	88.81	60.1 \pm 2	
Augengneiss	1556	WM(60-80)	9.12	$29.48 \cdot 10^{-6}$	92.92	80.8 \pm 2.6	
		WM(80-100)	9.19	$24.40 \cdot 10^{-6}$	91.83	66.7 \pm 2.2	
		WM(100-200)	9.31	$19.63 \cdot 10^{-6}$	92.08	53.1 \pm 1.7	
Augengneiss	1796	Bi(80-100)	7.97	$8.24 \cdot 10^{-6}$	79.45	26.2 \pm .9	
Augengneiss	1799	Bi(80-100)	7.79	$6.76 \cdot 10^{-6}$	71.01	22 \pm .9	
		WM(100-150)	9.10	$10.98 \cdot 10^{-6}$	75.41	30.5 \pm 1.2	

TABLE 2 (continued)

K/Ar

Rock type	Sample n.	Material	K%	^{40}Ar rd ml/g	^{40}Ar rd%	AGE Ma	Ref.
Augengneiss lamella	1557	WM(50-60) WM(60-80)	9.21 9.08	$18.49 \cdot 10^{-6}$ $17.96 \cdot 10^{-6}$	92.24 91.02	50.6 \pm 1.6 49.9 \pm 1.6	
Pegmatitic orthogneiss	AA14	Mu	9.08	$4.65 \cdot 10^{-4}$	87	114 \pm 3	8
Pegmatitic orthogneiss	AA15	Mu	9.08	$4.23 \cdot 10^{-4}$	72	104 \pm 3	
Orthogneiss	D7353	Bl	7.82	$74.48 \cdot 10^{-6}$	96.4	233	21
Orthogneiss	D7382	Bl	7.72	$84.79 \cdot 10^{-6}$	97.1	257	
Orthogneiss	D7391	Bl	6.73	$63.78 \cdot 10^{-6}$	94.5	224	
		Mu	8.47	$108.69 \cdot 10^{-6}$	96.1	297	
	D7420	Hbl	.342	$4.91 \cdot 10^{-6}$	86.9	329	
<u>Monteneve</u>							
Orthogneiss	KW1135	Ph	9.02	$30.3 \cdot 10^{-6}$	92.5	82.5 \pm 4	33
		Bl	6.87	$21.6 \cdot 10^{-6}$	90.7	77.5 \pm 3	
Muscovite gneiss	KAW 1208	Mu	8.815	$31.5 \cdot 10^{-6}$	64.8	89 \pm 5	33
Garnet micaschist	KAW 1207	Mu	7.46	$26.1 \cdot 10^{-6}$	90.3	87 \pm 4	
		Mu	7.46	$25.8 \cdot 10^{-6}$	86.2	86 \pm 4	
		Bl	6.995	$21.25 \cdot 10^{-6}$	93.6	75.5 \pm 3.0	
Garnet micaschist	KAW 1150	Mu	7.94	$28.86 \cdot 10^{-6}$	68.6	90 \pm 4	
		Bl	7.33	$23.75 \cdot 10^{-6}$	91.1	81 \pm 3	
Garnet micaschist	KAW 1151	Mu	7.53	$25.3 \cdot 10^{-6}$	91.95	82.5 \pm 4.0	
		Bl	6.23	$20.09 \cdot 10^{-6}$	89.1	79.5 \pm 4.0	
Pegmatite	KAW 1143	Mu	8.79	$34.0 \cdot 10^{-6}$	93.0	96 \pm 4	
		Mu	8.79	$33.7 \cdot 10^{-6}$	90.1	95 \pm 4	
		Mu	8.84	$29.6 \cdot 10^{-6}$	85.1	82.5 \pm 4.0	
Paragneiss	KAW 1136	Mu	7.875	$29.1 \cdot 10^{-6}$	89.7	91 \pm 4	
		Bl	6.97	$28.7 \cdot 10^{-6}$	95.4	102 \pm 4	
		Bl	6.97	$28.2 \cdot 10^{-6}$	94.5	101 \pm 4	
Bi-Mu-schist	KAW 1134	Mu	6.77	$23.9 \cdot 10^{-6}$	86.0	87 \pm 4	
		Bl	7.23	$23.5 \cdot 10^{-6}$	91.7	81 \pm 4	
Augengneiss	KAW 421	Ph	9.09	$31.0 \cdot 10^{-6}$	89.5	85 \pm 4	
		Bl	7.11	$23.2 \cdot 10^{-6}$	95.4	81 \pm 3	
Orthogneiss	KAW 1138	Mu	9.06	$111.2 \cdot 10^{-6}$	95.5	287 \pm 12	
		Bl	7.42	$42.4 \cdot 10^{-6}$	96.3	139 \pm 6	
		Bl	7.42	$40.8 \cdot 10^{-6}$	98.2	134 \pm 5	
Orthogneiss	KAW 1141	Mu	9.16	$28.6 \cdot 10^{-6}$	91.7	77.5 \pm 3.0	
		Bl	8.04	$24.1 \cdot 10^{-6}$	93.7	74.5 \pm 3.0	
Orthogneiss	KAW 1127	Mu	8.47	$29.0 \cdot 10^{-6}$	91.7	85 \pm 4	
		Bl	8.03	$28.7 \cdot 10^{-6}$	93.4	89 \pm 4	
Paragneiss	KAW 1126	Mu	8.07	$44.3 \cdot 10^{-6}$	89.0	133 \pm 6	
		Bl	7.36	$53.6 \cdot 10^{-6}$	94.9	175 \pm 7	
<u>Oetztal</u>							
Biotite-plagioclase gneiss	169	Bl(600-300) Bl(600-300)	6.62 6.73	$24.56 \cdot 10^{-6}$ $24.61 \cdot 10^{-6}$	89.9 90.2	90.7 \pm 1.1	29
Biotite-plagioclase gneiss	170	Bl(600-400) Mu(600-400) Mu(400-200)	6.97 2.69 7.05	$24.92 \cdot 10^{-6}$ $33.91 \cdot 10^{-6}$ $37.83 \cdot 10^{-6}$	86.3 83.5 90.5	88.1 \pm 1.1 29.4 \pm 4 129 \pm 2	
Granitized biotite-plagioclase gneiss	171	Bl(400-200) Bl(400-200)	7 7.05	$108.1 \cdot 10^{-6}$ $109.3 \cdot 10^{-6}$	94.9 92.4	355 \pm 4	
Biotite-plagioclase gneiss	172	Bl(600-400) Bl(400-200) Mu(600-400) Mu(400-200)	6.10 6.15 5.05 6.35	$37.93 \cdot 10^{-6}$ $35.68 \cdot 10^{-6}$ $53.24 \cdot 10^{-6}$ $62.25 \cdot 10^{-6}$	86.6 86 92 93.6	151 \pm 2 141 \pm 2 249 \pm 3 232 \pm 3	
Muscovite granitic gneiss	173	Mu(600-400) Mu(400-200)	7.74 8.59	$59.26 \cdot 10^{-6}$ $54.50 \cdot 10^{-6}$	92.7 83.1	184 \pm 2 154 \pm 2	
Diabase	174	WR(630-400) WR(400-200) WR(200-100) WR(100-63) WR(<63)	.576 .558 .542 .520 .584	$12.79 \cdot 10^{-6}$ $12.39 \cdot 10^{-6}$ $12.13 \cdot 10^{-6}$ $11.38 \cdot 10^{-6}$ $12.28 \cdot 10^{-6}$	85.3 77.6 81.6 75.5 63.7	490 \pm 6 491 \pm 6 494 \pm 6 484 \pm 6 468 \pm 6	
Diabase	375	WR(630-500) WR(630-500) WR(630-500) Bl(630-500) Bl(630-500)	3.144 3.136 3.125 6.22 6.29	$12.05 \cdot 10^{-6}$ $12.01 \cdot 10^{-6}$ $12.05 \cdot 10^{-6}$ $22.14 \cdot 10^{-6}$ $22.56 \cdot 10^{-6}$	95.8 95.7 95.8 93.7 94.5	94.3 \pm 1.2 88.1 \pm 1.1	

TABLE 2 (continued)

K/Ar

Rock type	Sample n.	Material (ν)	K%	^{40}Ar rd ml/g	^{40}Ar rd%	AGE Ma	Ref.
Diabase	376	WR(400-315)	.2,065	$8.50 \cdot 10^{-6}$	92.3	101 \pm 1.3	
		WR(200-100)	1.583	$6.62 \cdot 10^{-6}$	91.5	102.6 \pm 1.3	
		WR(630-500)	.866	$15.60 \cdot 10^{-6}$	95	407 \pm 5	
		WR(400-315)	.840	$15.54 \cdot 10^{-6}$	94.8		
		WR(400-315)	.860	$15.59 \cdot 10^{-6}$	94.8	413 \pm 5	
		WR(400-315)		$15.62 \cdot 10^{-6}$	94.6		
Biotite-plagioclase gneiss	175	WR(200-100)	.802	$14.57 \cdot 10^{-6}$	93.9	410 \pm 5	
		Bi(600-400)	5.84	$56.19 \cdot 10^{-6}$	92.0	228 \pm 3	29
		Bi(400-200)	6.68	$65.74 \cdot 10^{-6}$	89.3	233 \pm 3	
		Mu(400-200)	6.54	$48.31 \cdot 10^{-6}$	88.8		
		Mu(400-200)	6.71	$48.74 \cdot 10^{-6}$	91.7	176 \pm 2	
		Mu(400-200)	6.73				
Granodiorite	176	Bi(400-200)	7.46	$101.3 \cdot 10^{-6}$	92.6	314 \pm 4	
Augengneiss	177	Bi(600-400)	6.54	$45.28 \cdot 10^{-6}$	83.8	167 \pm 2	
		Bi(400-200)	6.61	$41.83 \cdot 10^{-6}$	84.8	154 \pm 2	
		Mu(400-200)	8.32	$101.2 \cdot 10^{-6}$	90.0		
		Mu(400-200)	8.43	$103.5 \cdot 10^{-6}$	92.7	286 \pm 4	
Augengneiss	178	Bi(400-200)	7.28	$62.61 \cdot 10^{-6}$	88.4	202 \pm 3	
		Bi(600-400)	7.33	$67.54 \cdot 10^{-6}$	91.9	218 \pm 3	
		Mu(600-400)	8.10	$104.6 \cdot 10^{-6}$	92.1	300 \pm 4	
		Mu(400-200)	8.40	$62.4 \cdot 10^{-7}$	70.0	179 \pm 2	
Diabase	372	WR(600-500)	.507	$46.15 \cdot 10^{-7}$	90.5		
		WR(600-500)	.516	$46.21 \cdot 10^{-7}$	90.4	215 \pm 3	
		WR(400-315)	.516	$46.26 \cdot 10^{-7}$	90.3	214 \pm 3	
		WR(200-100)	.591	$54.35 \cdot 10^{-7}$	87.2	219 \pm 3	
Diabase	373	WR(600-500)	.366	$39.81 \cdot 10^{-7}$	93.7		
		WR(600-500)	.358	$39.64 \cdot 10^{-7}$	93.7	257 \pm 3	
		WR(400-315)	.348	$38.93 \cdot 10^{-7}$	89.6	265 \pm 3	
		WR(200-100)	.455	$50.10 \cdot 10^{-7}$	89.2	259 \pm 3	
Diabase	374A	WR(630-500)	1.487	$73.48 \cdot 10^{-7}$	93.8	120.7 \pm 1.5	
		WR(400-315)	1.452	$72.33 \cdot 10^{-7}$	92.9	121.7 \pm 1.5	
		WR(200-100)	1.340	$68.34 \cdot 10^{-7}$	89.2	124.4 \pm 1.6	
Diabase	374B	WR(630-400)	1.292	$50.36 \cdot 10^{-7}$	91.0	95.8 \pm 1.2	
		WR(400-315)	1.284	$49.54 \cdot 10^{-7}$	90.5	94.9 \pm 1.2	
		WR(200-100)	1.238	$49.12 \cdot 10^{-7}$	82.6	97.4 \pm 2	

TABLE 3
Rb/Sr analytical data - SOUTHERN ALPS

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{m}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
<u>Euganean, Lessinian, Marosticano Hills</u>									
Mugearite	CE 86	WR	50	850	.166	.7028	.7028		
Latitbasalt	CE 92	WR	60	1140	.149	.7033	.7032		
Latite	CE 42	WR	45	670	.190	.7036	.7035		
Trachyte	CE 119	WR	180	580	.877	.7038	.7033		
Trachyte	CE 116	WR	141	252	1.580	.7043	.7035		
Trachyte	CE 113	P1	19	1530	.035	.7032	.7032		
Trachyte	CE 113	WR	120	450	.753	.7038	.7034		
Alkali Rhyolite	CE 114	P1	15	894	.047	.7035	.7035		
Alkali Rhyolite	CE 52	WR	180	400	1.271	.7041	.7035		
Persilicic alkali-rhyolite	CE 57	KF	396	58	19.280	.7125	.7031		
Nephelinitic basalt	L 1	WR	25	1150	.061	.7028	.7028		
Ankaratritic basalt	M 103	WR	30	920	.092	.7037	.7036		

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_m$	$(^{87}\text{Sr}/^{86}\text{Sr})_t$	AGE Ma	Ref.
<u>Predazzo</u>									
Granite	B-GE-44	WR	459	20.3	66.3	.875		181 \pm 9	3
Granite	B-GT-1B	WR	329	102.2	9.3	.726			
Granite	B-GT-36	WR	546	11.8	139.6	1.161		230 \pm 7	
Granite	B-RS	Mu	11117	11.0	38952	126.760		227 \pm 7	
Granite	B-GE-9	WR	291	56.9	14.9	.756		244 \pm 12	
		Bi	1516	5.3	1156	4.750		245 \pm 7	
Granite	B-GT-10B	WR	365	138.4	7.6	.734			
		Bi	1050	26.0	121.5	1.120		238 \pm 7	
Granite	B-GT-30	WR	372	122.2	8.8	.725			
		Bi	970	33.7	85.5	.997		241 \pm 7	
		Feld.	674	96.6	20.3	.756		179 \pm 9	
Granite	Pr-6 bis	Bi	1685	14.0	393	2.010		233 \pm 7	
Granite	Pr-7	Feld.	436	29.5	43.4	.843		225 \pm 11	
Monzonite	B Canzoccoli 5	Bi	567	32.8	50.7	.873		234 \pm 12	
Monzonite	B Malgola 49C	Bi	599	32.5	54.1	.874		222 \pm 11	
		Bi	602	31.4	56.6	.888		228 \pm 7	
Monzonite	B Pente Cast. 65	Bi	463	40.8	33.1	.815		235 \pm 12	
		Bi	504	28.4	52.2	.880		236 \pm 12	
Monzonite	B-151	Bi	504	26.0	57.1	.911		241 \pm 7	
Monzonite	L-3	Bi	440	23.4	55.3	.885		230 \pm 11	
Syenite	L-1	Bi	728	14.8	148.9	1.210		237 \pm 7	
Syenite	B-114	Bi	449	40.2	32.6	.816		241 \pm 12	
Syenite	B-30-J	Bi	424	76.0	16.2	.758		235 \pm 12	
Granite	Pr-6	WR	378	121.4	9.0	.736			
Granite	Pr-6	Bi	1435	16.5	274	1.640		239 \pm 7	3
		Feld.	667	87.9	22.1	.777		234 \pm 12	
Granite	Pr-8	WR	286	50.5	16.5	.757		227 \pm 11	
		Bi	1446	13.1	358.4	1.912		236 \pm 7	
		Feld.	387	48.3	23.4	.783		237 \pm 12	
Granite	B-GE-14	WR	387	32.4	34.9	.816		226 \pm 11	
Granite	B-GT-19	WR	422	18.0	69.3	.914		213 \pm 6	
Granite	B-GT-22A	WR	378	15.3	72.8	.932		219 \pm 7	
Granite	B-GT-36	WR	546	8.6	195.3	1.352		233 \pm 7	
Granite	B-GT-147	WR	389	15.3	75.6	.948		227 \pm 7	
Camptonite	Pr-2	WR			.40	.7041		.7028	
Camptonite	Pr-3	WR			.20	.7055		.7048	
Granite	B-GT-30	WR			8.59	.7360		.7072	
Monzonite	B-cp-18	WR			.34	.7034		.7023	
Monzonite	B-151	WR			.59	.7053		.7033	
Syenite	B-180	WR			1.32	.7100		.7057	
Syenite	B-170	WR			1.59	.7094		.7040	
Monzonite	P.Castellani	WR			.30	.7080		.7070	
Monzonite	B Canzoccoli	WR			.30	.7062		.7050	
Volcanic	B-60	WR			.42	.7059		.7046	
<u>Monzoni</u>									
Monzodiorite	X2	Bi	429	24.7	51.07	.8712		229 \pm 8	5
Monzodiorite	X1	Bi	503	24.9	59.63	.9099		241 \pm 8	
Gabbro	48	Bi	380	63.4	17.45	.7630		234 \pm 8	5
Gabbro	28	Bi	446	75.9	17.09	.7642		243 \pm 8	
Diorite	121	Bi	434	49.4	25.62	.7948		246 \pm 8	
Diorite	274	Bi	617	31.4	57.95	.8910		226 \pm 8	
Syenitic dyke	F-1	Bi	530	33.3	46.74	.8680		245 \pm 8	
Gabbro	GT	WR	13	.875	.04	.7037			
Gabbro	48	WR	34	1395	.07	.7034			
Gabbro	268	WR	12	802	.04	.7039			
Gabbro	320	WR	40	1238	.09	.7052			

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{m}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
Gabbro	28	WR	79	992	.23	.7058			
Gabbro	84	WR	41	1177	.10	.7073			
Monzogabbro	64	WR	87	973	.26	.7048			
Monzodiorite	X1	WR	124	873	.40	.7062			
Monzodiorite	X2	WR	96	1347	.20	.7060			
Monzodiorite	301	WR	208	710	.84	.7074			
Diorite	274	WR	102	1118	.26	.7055			
Syenitic dyke	F-1	WR	225	520	1.25	.7084			
Syenitic dyke	F-100	WR	240	286	2.43	.7125			
Pyroxenite	PT	WR	21	144	.40	.7083			
Ladinian limestone	CL	WR	4	127	.08	.7094			
Monzodiorite	MD	WR	144	830	.50	.7091			
Xenolith in MD	IMD	WR	210	253	2.40	.7261			
Phyllite	SM-153	WR	128	117	3.20	.7326			
<u>Mt. Sabion</u>									
Granodiorite	A/6	Bi	522	4.61	377.07	2.2481		285+ 9	4
Granodiorite	A/7	Bi	454	4.03	374.33	2.2190		283+ 9	
Biotitic pegmatite	A/8	Bi	650	4.33	523.49	2.8031		278+ 9	
Muscovitic pegmatite	A/10	Mu	1392	4.11	1592.26	7.0993		282+ 9	
<u>Mt. Croce</u>									
Leucogranite	18	WR	198	125	4.59	.7264			7
Pegmatite	18p	WR	248	30	23.60	.8046			
Granodioritic granite	AA9	WR	144	213	1.96	.7196			
Granodioritic granite	MC14	WR	144	198	2.10	.7197			
Granodioritic granite	MC I	WR	150	176	2.47	.7201			
Granitic porphyrite	MC II	WR	163	135	3.49	.7229			
Basic dyke	MC VI	WR	145	296	1.42	.7151			
Leucogranite	MC VII	WR	243	36	19.70	.7919			
Pegmatite	MC VIII	WR	179	116	4.48	.7294			
Aplitic granite	MC IX	WR	202	68	8.60	.7460			
Leucogranite	18	Bi	740	5.39	465.56	2.4748	(.7101)	266+ 5	
Granodiorite	AA9	Bi	604	5.29	380.00	2.2345	(.7101)	282+ 6	
Granodiorite	MC14	Bi	544	11.14	149.94	1.3351	(.7101)	292+10	
Granodiorite	MC I	Bi	475	3.66	443.44	2.5261	(.7101)	287+ 5	
<u>Ivigna</u>									
Granite-Granodiorite	I21	WR	155	187	2.39	.7189			
Granite-Granodiorite	I22	WR	153	208	2.13	.7183			
Microgranite	I22a	WR	211	31	19.30	.7904			
Granodiorite	I23	WR	154	228	1.90	.7171			7
Granite-Granodiorite	I26	WR	155	178	2.52	.7171			
Granite-Granodiorite	I19	WR	134	255	1.50	.7142			
Lamprophyric dyke	I XIII	WR	108	272	1.15	.7116		.7079+4	301+ 2
Pegmatitic dyke	I XIV	WR	204	107	5.52	.7333			
Aplitic	I XVIIa	WR	172	114	4.36	.7260			
Granite or Granodiorite	I XVIig	WR	208	141	4.26	.7258			
Pegmatite	I XVIII	WR	293	21	40.50	.8819			
Granite-Granodiorite	I19	Bi	591	2.85	712.50	2.6187	(.7079)	188+ 4	
Granite-Granodiorite	I22	Bi	469	10.81	131.75	1.2105	(.7079)	267+11	
Granodiorite	I23	Bi	110	27.91	11.45	.7500	(.7079)	258+98	
Granite-Granodiorite	I26	Bi	389	11.50	101.40	1.0791	(.7079)	257+13	
Granite-Granodiorite	I XIV	Bi	435	5.61	242.04	1.5049	(.7079)	231+12	
<u>Bressanone</u>									
Granodioritic granite	AA3	WR	160	179	2.58	.7212			
Granodioritic granite	B10	WR	155	162	2.78	.7205			
Granodioritic granite	B9	WR	179	142	3.65	.7255			
Granodioritic granite	B13	WR	166	160	3.02	.7222			
Granodioritic granite	B12	WR	155	181	2.48	.7204		.7091+ 5	291+ 6
Granodioritic granite	30v	WR	146	188	2.24	.7177			

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{f}}$	AGE Ma	Ref.
Granodioritic granite	50v	WR	147	192	2.21	.7161			
Microgranite	B XIX	WR	153	139	3.19	.7217			
Aplitic or Aplitic granite	B XXI A	WR	187	50	10.80	.7535			
Aplitic	B XXII	WR	188	39	13.90	.7658			
Aplitic	B XXIV	WR	178	45	11.50	.7579			7
Granodioritic granite	B XXV	WR	164	173	2.75	.7195			
Granodiorite	AA3	Bi	391	7.75	154.99	1.3360	(.7091)	284 \pm 11	
Granodiorite	B10	Bi	546	4.82	376.19	2.2266	*	283 \pm 6	
Granodiorite	B9	Bi	522	9.40	171.65	1.4030	*	284 \pm 9	
Granodiorite	B13	Bi	394	24.17	48.11	.9120	*	296 \pm 25	
Granodiorite	B XXV	Bi	512	4.00	435.76	2.5043	*	289 \pm 5	
	80-11	WR	82	209	1.14	.7121 \pm 2	.7076 \pm 2		23
	80-12	WR	24	213	.33	.7096 \pm 2	.7083		
Gabbro-norites & Diorites	80-9	WR	123	319	1.11	.7138 \pm 2	.7094		
	80-10	WR	203	340	1.73	.7177 \pm 3	.7108		
	80-42	WR	42	255	.48	.7100 \pm 2	.7081		
	1230	WR	193	168	3.33	.7226 \pm 4			
	293a	WR	175	134	3.78	.7241 \pm 3			
	269	WR	168	241	2.02	.7182 \pm 3			
	345	WR	174	177	2.85	.7213 \pm 4	.7102 \pm 4		
	264	WR	156	187	2.42	.7190 \pm 4			
Granodiorites & Granites	510	WR	113	239	1.37	.7141 \pm 5	.7096 \pm 6	282 \pm 14	
	340	WR	167	180	2.68	.7200 \pm 4	.7095 \pm 5		
	237	WR	184	137	3.89	.7249 \pm 2	.7096 \pm 4		
	229	WR	196	180	3.15	.7226 \pm 2	.7102 \pm 4		
	274	WR	206	130	4.59	.7285 \pm 4			
	360	WR	184	172	3.10	.7221 \pm 1			
Fayalite granites	311/4	WR	158	80.4	5.69	.7324 \pm 2			
	311/3	WR	185	60	8.99	.7455 \pm 2			
	315	WR	229	71	9.31	.7460 \pm 4			
	79/7	WR	168	33.5	14.63	.7676 \pm 3			23
	316	WR	245	39	18.53	.7826 \pm 3			
Fayalite granites	1291	WR	186	24.3	22.35	.7972 \pm 5	.7101 \pm 2	275 \pm 1	
	290Fg	WR	207	19.7	30.88	.8316 \pm 2	.7085 \pm 3	266 \pm 1	
	305/c	WR	217	17.6	36.23	.8516 \pm 3			
	6	WR	169	77	6.37	.7327 \pm 9			
	527	WR	216	44	14.27	.7622 \pm 6			
	311/8	WR	301	12.1	74.24	.9882 \pm 8			
	333	WR	352	12.4	84.37	1.0286 \pm 7			
Fayalite granite	6	Bi	330	22.3	43.59	.8749 \pm 8	.7084 \pm 10	269 \pm 5	
		Bi	398	20.0	58.98	.9340 \pm 21	.7083 \pm 10	269 \pm 5	
		KF	363	86	12.29	.7480 \pm 5			
		P1	33	116	.83	.7207 \pm 3			
		Fa	51	7.5	19.78	.7850 \pm 7			
Fayalite granite	79-7	Bi	293	9.3	94.21	1.0729 \pm 14	.7101 \pm 14	271 \pm 5	
		KF	393	41	28.28	.8117 \pm 4			
		P1	11	16.5	1.98	.7291 \pm 3			
		Fa	30	9.0	9.76	.7369 \pm 8			
Fayalite granite	290Fg	Bi	843	9.5	285.41	1.7893 \pm 16	.7153 \pm 26	264 \pm 4	
Fayalite granite	305/c	Bi	705	4.9	492.03	2.5532 \pm 62	.7163 \pm 3	262 \pm 4	
Fayalite granite	1291	Bi	671	4.4	535.92	2.7820 \pm 71	.7108 \pm 20	272 \pm 4	
Fayalite granite	333	Bi	802	3.1	1032.15	4.4726 \pm 145	.7220 \pm 70	255 \pm 4	
Fayalite granite	315	Bi	645	18.0	107.76	1.1055 \pm 22	.7121 \pm 9	257 \pm 5	
Fayalite granite	311/4	Bi	494	16.6	88.77	1.0465 \pm 18	.7109 \pm 5	266 \pm 5	

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_m$	$(^{87}\text{Sr}/^{86}\text{Sr})_f$	AGE Ma	Ref.
Two-mica cordierite granites	321	WR	227	87	7.59	.7408± 5			23
	329	WR	217	155	4.05	.7282± 7			
	344	WR	238	92	7.51	.7413± 7	.7106± 2	286± 5	
	245	WR	225	81	8.06	.7435± 3			
Two-mica cordierite granite	329	P1	619	3.5	650.34	3.2909± 46	.7121± 7	279± 4	
		P1	119	449	.77	.7137± 2	.7107		
Two-mica cordierite granite	344	Bi	1089	6.3	624.16	3.1907± 144	.7115± 9	279± 5	
Gabbro-norite	80-11	Bi	493	6.8	228.05	1.6032± 11			276± 4
Gabbro-norite	237	Bi	317	9.4	101.21	1.1078± 42			277± 5
Gabbro-norite	345	Bi	497	3.8	444.95	2.4479± 125			274± 5
Granodiorites & Granites	229	Bi	416	9.0	140.56	1.2646± 33			277± 5
	340	Bi	537	3.6	514.13	2.7306± 100			276± 4
	1230	Bi	509	4.1	422.07	2.3588± 32			275± 4
<u>Cima d'Asta</u>									
Granito di Cima d'Asta	1	Bi	863	10.49	263.23	1.7900			282± 8
Granito di Cima d'Asta	2	Bi	940	3.78	1004.36	4.7697			281± 8
Granito di Caoria	3	Bi	1604	6.20	1066.67	5.0405			285± 8
Granodiorite	4	Bi	555	4.49	443.51	3.1592			290± 10
Granite	CDA46	WR	205	97	6.12	.7345			9
Granite	CDA47	WR	290	28	30.42	.8297			
Granite	CDA48	WR	275	26	30.12	.8264			9
Pegmatite	CDA49	WR	328	11	85.93	1.0441	.7102± 5	273± 1	
Granodiorite	CDA54	WR	120	245	1.41	.7154			
Granite	CDA62	WR	182	135	3.89	.7248			
Qz-Diorite	CDA52	WR	68	265	.74	.7096	.7067± 13		
Qz-Norite	CDA61	WR	45	253	.52	.7087	.7067± 8		
Granite	CDA33	WR	146	190	2.21	.7154			
Granite	CDA34	WR	180	196	2.65	.7187			
Aplice-Pegmatite	CDA35	WR	218	179	3.52	.7224			
Granodiorite	CDA36	WR	154	210	2.13	.7183			
Aplice-Pegmatite	CDA37	WR	274	52	15.25	.7676			
Aplice	CDA38	WR	224	75	8.61	.7433			
Tonalite	CDA40	WR	123	241	1.47	.7149			
Granite	CDA13	WR	176	184	2.77	.7214			
Tonalite	CDA15	WR	129	220	1.69	.7167			
Aplice	CDA18	WR	115	124	2.68	.7216			
Granite	CDA19	WR	197	199	2.85	.7233			
Porphyrite	CDA20	WR	186	267	2.02	.7172			
Porphyrite	CDA28	WR	54	45	3.49	.7244			
Granite	CDA47	Bi	1250	10.07	415.81	2.3240			272± 5
Granodiorite	CDA63	Bi	554	4.80	372.51	2.1644			268± 4
Granite	CDA34	Bi	839	5.42	539.50	2.8029			272± 5
Tonalite	CDA40	Bi	557	5.21	350.38	2.0546			269± 6
Granodiorite	83	Bi	742	3.43	805.82	3.9583			277± 5
Granite	116	Bi	914	4.41	758.59	3.7173			273± 5
Granite	CDA19	Bi	491	3.53	473.59	2.5220			269± 5
Qz-Diorite	CDA52	Bi	402	13.95	86.15	1.0459			276± 14.5
Granite	CDA19	P1	104	211	1.42	.7167	.7113± 17		
Granodiorite	CDA36	P1	17	401	.12	.7092	.7087± 13		
<u>Basement</u>									
Porphyroids	AA 76.2	WR	70	232	.88	.7165± 4			24
	AA 76.3	WR	74	95	2.25	.7226± 4			
	AA 76.20	WR	84	85	2.89	.7260± 4			
	AA 76.34	WR	142	33	12.58	.7745± 6	.7118± 2	350± 3	
	AA 76.35	WR	145	32	13.34	.7777± 7			
	AA 76.37	WR	161	31	14.98	.7870± 3			

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$\{\text{Rb}_{\text{S}}/\text{Rb}_{\text{M}}\}_{\text{M}}$	$\{\text{Sr}_{\text{S}}/\text{Sr}_{\text{M}}\}_{\text{S}}$	AGE Ma	Ref.
Metavolcanoclastics	AA 76.1	WR	50	25	5.87	.7385± 6			
	AA 76.19	WR	66	67	2.87	.7252± 6	.7119± 9	317± 6	
	AA 76.36	WR	48	41	3.43	.7272± 2			
Phyllite	AA 78.22	Bi	419	9.5	134.88	1.3165± 10			
		WR	62	199	.91	.7175± 2		314± 5	
Phyllite	AA 77.13	Mu	489	111	12.78	.7769± 3			
		Pt+Qz	14	57	.71	.7233± 3		312± 6	
Metavolcanoclastic Phyllite	AA 76.1	Mu	323	10.3	94.26	1.1336± 23			
Metavolcanoclastic Phyllite	AA 78.23	WR	99	4	5.29	.7488± 4			
		Mu	349	103	9.73	.7702± 2		354± 10	
Acidic gneisses		Ga	8	7	3.29	.7380± 1			26
	AA 80-13	WR	121	163	2.15	.7230± 2			
	AA 80-14	WR	112	113	2.87	.7260± 2			
	AA 80-15	WR	119	153	2.26	.7237± 2	.7155± 5	257± 17	
	AA 80-22	WR	132	154	2.49	.7249± 3			
	AA 80-27	WR	95	167	1.65	.7216± 1			
		KF	165	249	1.92	.7227± 4			
Acidic gneiss	AA 80-13	KF	165	249	1.92	.7227± 4			26
Acidic gneiss	AA 80-14	KF	173	186	2.71	.7255± 3			
Acidic gneiss	AA 80-15	KF	189	212	2.58	.7250± 2	.7160± 11	248± 33	
Acidic gneiss	AA 80-22	KF	170	219	2.25	.7238± 2			
Acidic gneiss	AA 80-27	KF	165	203	2.36	.7247± 3			
Acidic gneiss	AA 80-13	WR	121	163	2.15	.7230± 2			
Acidic gneiss		KF	165	249	1.92	.7227± 4			309± 24
Acidic gneiss		Ab	7.7	35.9	.62	.7164± 3			
	AA 80-14	WR	112	113	2.87	.7260± 2			
		KF	173	186	2.71	.7255± 3			293± 6
Acidic gneiss		Ab	6.0	28.9	.60	.7166± 4			
	AA 80-15	WR	119	153	2.26	.7237± 2			
		KF	189	212	2.58	.7250± 2			302± 4
Acidic gneiss	AA 80-22	WR	132	154	2.49	.7249± 3			
		KF	170	219	2.25	.7238± 2			
		Ab	6.5	26.1	.72	.7170± 3			326± 3
Acidic gneiss	AA 80-27	WR	95	167	1.65	.7216± 1			
		KF	165	203	2.36	.7247± 3			
		Ab	11.5	40.3	.83	.7179± 5			321± 1
Acidic gneiss		Bi	445	15.0	89.16	1.1292± 23			
	AA 78-28	WR	301	78	11.26	.7869± 6			
		Ab	5.4	2.6	6.17	.7634± 3			414± 97
Acidic gneiss	AA 80-21	Mu	390	96	11.90	.8012± 6	.7227	463± 18	
		WR	203	104	5.71	.7561± 3			
Atesino Porphyry Plateau		Mu	353	189	5.44	.7559± 8			
Rhyolitic ignimbrites	DAV 1	WR	251	59	12.32	.7537± 5			22
	DAV 2	WR	224	91	7.12	.7361± 4			
	DAV 3	WR	226	80	8.22	.7385± 6			
	DAV 5	WR	258	74	10.16	.7443± 6			
	DAV 6	WR	276	75	10.70	.7448± 10			
	DAV 7	WR	272	58	13.56	.7564± 5			
	DAV 8	WR	202	97	6.05	.7315± 4			
	DAV 9	WR	222	98	6.58	.7323± 4	.7103± 6	239± 8	
	VA 719	WR	208	138	4.36	.7258± 3			
	VA 721	WR	189	101	5.42	.7273± 2			
	VA 731	WR	192	150	3.69	.7234± 2			
	VA 739	WR	192	149	3.73	.7233± 2			
	VA 779	WR	199	98	5.88	.7296± 2			
	VA 780	WR	196	103	5.51	.7288± 2			
	VA 1016	WR	210	120	5.05	.7266± 3			

TABLE 3 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
Rhyolitic ignimbrite	VA 719	Bi	416	14.5	85.7	1.0419±3		272±4	
Rhyolitic ignimbrite	VA 739	Bi	425	10.8	119.7	1.1770±6		274±4	
Rhyolitic ignimbrite	VA 746	Bi	528	4.9	353.7	2.0676±57		270±4	
Rhyolitic ignimbrite	VA 757	Bi	736	5.3	471.7	2.5053±39		268±4	
Rhyolitic ignimbrite	VA 779	Bi	609	6.6	294.2	1.8125±42		263±4	
Rhyolitic ignimbrite	VA 780	Bi	577	7.7	234.6	1.5983±8		266±4	
Rhyolitic ignimbrite	DAV 1	Bi	605	3.7	571.9	2.9234±41		272±4	
Rhyolitic ignimbrite	DAV 2	Bi	624	9.9	195.0	1.4601±11		270±4	
Rhyolitic ignimbrite	DAV 9	Chl	102	36.3	8.1	.7407±5		(265±7)	
Rhyolitic ignimbrite	VA 719	Bi	416	14.5	85.7	1.0419±3		22	
Rhyolitic ignimbrite	VA 739	Bi	425	10.8	119.7	1.1770±6			
Rhyolitic ignimbrite	VA 746	Bi	528	4.9	353.7	2.0676±57			
Rhyolitic ignimbrite	VA 757	Bi	736	5.3	471.7	2.5053±39			
Rhyolitic ignimbrite	VA 779	Bi	609	6.6	294.2	1.8125±42	.7171±52	267±2	
Rhyolitic ignimbrite	VA 780	Bi	577	7.7	234.6	1.5983±8			
Rhyolitic ignimbrite	DAV 1	Bi	605	3.7	571.9	2.9234±41			
Rhyolitic ignimbrite	DAV 2	Bi	624	9.9	195.0	1.4601±11			
Rhyolitic ignimbrite	DAV 1	KF	278	229	3.52	.7209±2			
Rhyolitic ignimbrite	DAV 9	KF	357	204	5.06	.7186±3			
Rhyolitic ignimbrite	DAV 9	P1+Qz	128	171	2.17	.7170±3	.7102±4	226±8	
Rhyolitic ignimbrite	VA 739	KF	364	231	4.56	.7248±1			
Rhyolitic ignimbrite	VA 780	KF	443	179	7.17	.7337±3			
Rhyolitic ignimbrite	VA 780	P1+Qz	19	18	3.07	.7204±3			

TABLE 4
Rb/Sr analytical data - AUSTRIDES

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
<u>Mules</u>									
Tonalitic granodiorite	AA7	WR	50	370	.39	.7083			7
" "	AA7a	WR	53	319	.48	.7102			
" "	AA7b	WR	46	383	.35	.7110			
" "	AA7c	WR	38	327	.34	.7083			
" "	AA12	WR	45	326	.40	.7096			
" "	AA7	Bi	324	7.84	120.19	.7625	(.7080)	32.1±10.4	
<u>Merano-Mules-Anterselva basement</u>									
Epimetamorphic dyke	AA76-33	Bi	507	5.7	257.68	.7976±30		24.2±2.1	10
		WR	97	277	1.02	.7091±4		31.4±17.4	
		Mu	296	82	10.45	.7133±8			
Epimetamorphic dyke	AA77-30	WR	98	311	.91	.7080±2		17.6±20	
		Bi	448	14.8	87.61	.7297±9			
		Mu	227	161	4.07	.7092±4			
<u>Rensen</u>									
Granites	AA76-4	WR	149	259	1.66	.7109±4			11
	AA76-5	WR	142	260	1.58	.7106±7			
	AA76-8	WR	134	207	1.87	.7106±3			
	AA76-9	WR	183	196	2.70	.7114±2			
	AA76-10	WR	130	249	1.50	.7108±5			
	AA76-12	WR	140	123	3.27	.7117±8			
	AA76-6	WR	69	307	.65	.7079±5			
	AA76-7	WR	66	347	.55	.7078±6			
	AA76-11	WR	75	311	.70	.7083±3			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{m}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{f}}$	AGE Ma	Ref.
Granodiorites & tonalites	AA76-13	WR	78	367	.62	.7079± 4			11
	AA100	WR	84	338	.72	.7083± 3			
	AA101	WR	72	348	.60	.7077± 5			
	AA102	WR	79	356	.64	.7070± 6			
Granite	AA76-4	Bi	683	4.1	479.02	.8299±27		17.5±1.0	
		Mu	469	32	42.50	.7281±17		29.6±8.1	
Granite	AA76-5	Bi	650	7.6	246.43	.7748±15		18.4±1.2	
		Mu	514	20.6	50.28	.7314± 3		30.8±2.9	
Granodiorite	AA76-11	Bi	471	2.2	630.89	.8710±64		18.2±1.7	
Granodiorite	AA33	Bi	492	3.6	395.47	.8095	(.709)	18.0±4.0	8
<u>Vedrette di Ries</u>									
Granites, acid dykes, tonalites and granodiorites	AA75-1	WR	191	86	6.40	.7125± 9			12
	AA75-2	WR	181	113	4.66	.7116± 7			
	AA75-3	WR	118	227	1.50	.7102± 3			
	AA75-5	WR	215	67	9.26	.7135± 4			
	AA75-8	WR	146	342	1.24	.7097± 5	.7096± 1	31 ±3	
	AA75-9	WR	159	263	1.76	.7105± 3			
	AA75-10	WR	188	224	2.42	.7109± 3			
	AA75-11	WR	223	49	13.20	.7154±10			
	AA75-12	WR	181	85	6.17	.7125± 4			
	AA73-57a	WR	120	275	1.26	.7116± 3			
	AA74-45	WR	30	431	.20	.7110± 4			
Tonalites	AA74-46	WR	101	318	.92	.7112± 3	.7109± 1	24 ±12	
	AA74-47	WR	123	331	1.07	.7113± 5			
	AA74-48	WR	97	272	1.03	.7115± 5			
	AA74-60	WR	106	243	1.27	.7114± 5			
	AA74-61	WR	104	286	1.05	.7110± 4			12
Tonalite	AA30 a	WR	159	220	2.10	.7116± 3			
	AA30 b	WR	111	270	1.19	.7112± 4			
	AA30	Bi	497	2.6	575.27	.9417	(.709)	28 ± 3	8
Tonalite	AA31	Bi	522	2.9	539.47	.9219	(.709)	28 ± 3	
Granites	AA73-81	WR	172	203	2.45	.7101± 2			12
	AA74-53	WR	197	147	3.86	.7108± 4			
	AA74-67	WR	170	317	1.55	.7097± 4			
	AA74-68	WR	163	388	1.21	.7097± 3	.7089± 1	33 ± 7	
	AA74-69	WR	167	376	1.28	.7095± 3			
	AA74-70	WR	176	348	1.46	.7095± 4			
	AA73-84	WR	153	315	1.40	.7107± 4			
	AA74-49	WR	132	377	1.02	.7105± 3			
	AA74-52	WR	126	373	.98	.7105± 4			
	AA74-55	WR	112	325	1.00	.7102± 3	.7098± 4	43.5±44.5	
	AA74-71	WR	139	345	1.16	.7107± 3			
	AA74-72	WR	145	323	1.30	.7105± 3			
Microgranites, granodiorites and aplites	AA75-4	WR	119	349	.99	.7105± 5			
	AA73-57b	WR	145	97	4.30	.7225± 4			
	AA73-57c	WR	137	145	2.75	.7219± 4	.7207± 5	28 ± 22	
	AA73-59	WR	154	130	3.45	.7219± 4			
	AA73-58	WR	148	126	3.39	.7289± 9			
	AA73-64	WR	99	201	1.43	.7131± 4			
	AA74-44	WR	130	178	2.11	.7158± 4			
	AA74-65	WR	110	222	1.44	.7126± 5			
	AA74-73	WR	152	190	2.31	.7159± 3			
	AA73-43	WR	143	589	.70	.7082± 4			12
Mt. Alto	AA73-43	Bi	712	4.5	457.77	.9068±32		30.5±1.1	
	AA74-50	WR	145	612	.69	.7079± 4			
	AA32	WR	122	579	.61	.7111± 3			
Vedrette di Ries	AA32	Bi	586	3.9	446.24	.8669	(.7090)	25.0±3.0	8
	AA32 b	WR	94	475	.57	.7115± 4			12
	AA32 c	Bi	524	3.6	432.12	.8577±17		24.0±0.7	
	AA73-58	WR	146	125	3.40	.7285± 4			
	AA73-58	KF	313	202	4.49	.7295± 2			
	AA73-58	Bi	760	1.2	2021.24	1.5360±27			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{J}}$	AGE Ma	Ref.
Tonalite	AA73-98	WR	95	291	.95	.7115+ 3			
	AA73-98	P1	51	543	.27	.7111+ 5	.7111+ 1	29.9+1.3	
	AA73-98	Bi	445	23	56.14	.7349+25			
	AA74-9	Bi	579	4.2	401.25	.8629+13	(.7100)	26.9+3.0	
	AA74-44	WR	130	178	2.11	.7158+ 4			
	AA74-44	KF	268	263	.96	.7166+ 3		28.1+3.6	
	AA74-44	P1	41	317	.37	.7146+ 2			
	AA74-44	Bi	588	3	587.35	.9469+ 6	.7145+3	27.8+0.1	
	AA74-46	WR	101	318	.92	.7112+ 3			
	AA74-46	P1	10	497	.06	.7110+16		26.9+1.0	
Tonalite	AA74-46	Bi	496	2.5	568.19	.9238+44			
	AA74-60	WR	106	243	1.27	.7114+ 5			
	AA74-60	P1	5.7	555	.03	.7109+ 5			
	AA74-60	KF	49	204	.69	.7113+ 2	.7110+ 1	28.1+0.5	
	AA74-60	Ho	17	35	1.44	.7104+ 5			
Merano-Mules-Anterselva basement	AA74-60	Ga	10	14	2.12	.7117+ 4			12
	AA74-60	Bi	433	4.6	270.10	.8189+ 9			
	AA74-65	WR	110	222	1.44	.7126+ 5			
	AA74-65	P1	9	428	.06	.7119+ 3	.7121+ 1	27.1+ .7	
	AA74-65	KF	146	296	1.43	.7128+ 3			
	AA74-65	Bi	553	1.8	869.05	1.0461+14			
Pegmatite	KAW 302	Mu	406.3	32.02	37.21	.84498		257+30	30
Pegmatitic orthogneiss	AA 12	Bi	771	4.92	485.53	1.4233	(.7200)	101+ 3	8
Pegmatitic orthogneiss	AA 13	Mu	1095	4.35	991.26	4.4086	(.7200)	261+ 5	
Pegmatitic orthogneiss	AA 14	Mu	555	10.32	164.06	1.2635	(.7200)	233+ 9	
Pegmatitic orthogneiss	AA 15	Mu	513	13.02	118.89	1.1493	(.7200)	254+12	
Pegmatitic orthogneiss	AA 16	Mu	527	10.03	160.70	1.2959	(.7200)	252+ 9	
Pegmatitic orthogneiss	AA73-17	Mu	1108	4.07	1069.89	4.3666	(.7200)	239+ 3	
Pegmatitic orthogneiss	AA73-50	Mu	1603	132.96	35.18	.7943	(.7200)	148+30	
Pegmatite	KAW 1143	WR	205.9	38.3	15.65	.7598			
		Mu	530.3	8.7	185.2	1.184		176+ 7	
		Mu	707.7	9.2	229.9	1.050		95+ 4	
Pegmatitic orthogneiss	AA73-103	Bi	2013	6.1	959.56	1.1173+21	(.7100)	30+ 1.4	10
		Mu	1029	40.0	74.52	.8851+16		166+16	
Aplitic orthogneisses	AA73-32	WR	206	47	12.7	.7647+ 4			
	AA73-33	WR	247	14.7	49.7	.9038+ 8			
	AA73-34	WR	210	35	17.7	.7798+ 4			
	AA73-55	WR	260	16.7	45.8	.9071+ 7			
	AA78-43	WR	86	210	1.2	.7207+ 4			
	AA78-46	WR	101	219	1.3	.7208+ 3			
	AA78-47	WR	286	80	10.3	.7520+ 2			
	AA78-48	WR	69	171	1.2	.7260+ 4	.7157+ 4	262+ 5	
	AA78-53	WR	149	114	3.8	.7288+ 5			
	AA78-54	WR	145	61	6.9	.7422+ 4			
Pegmatite	AA78-55	WR	120	54	6.4	.7379+ 4			
	AA78-56	WR	162	47	10.0	.7554+ 8			
	AA78-57	WR	151	28	15.9	.7739+12			
	AA73-56	WR	241	13.2	53.8	.9228+ 5			
		Mu	1239	2.5	2294.9	7.1046+284		194+ 6	
	D 7558	Mu	1187	3.09	1439	3.705		146	21
		KF	30.40	79.43	1.107	.7203			
Pegmatitic orthogneiss	AA73-100	KF	262	216	3.52	.7357+ 2			
		Mu	301	13.3	67.13	.9546+ 21		242+ 4	
Aplitic orthogneiss	AA77-11	WR	165	39.3	12.2	.7641+ 2			
		Mu	739	4.4	576.64	2.5127+ 50		218+ 3	
Pegmatitic orthogneiss	VDB0-1	WR	107	85	3.66	.7322+ 2			
		KF	263	138	5.52	.7322+ 5		240+ 4	
		Mu	517	92	177.7	1.3199+ 9			
<u>Ötztal</u>									
Muscovitic orthogneiss	KAW 149	Mu	1343	8.66	548.91	2.9907		292+11	32
		WR	417	34.32	35.91	.9282		429+35	

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{SM}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{S}}$	AGE Ma	Ref.
Tonalitic gneiss	B 3/4	Bi	512.3	2.66	712.60	3.5724		283±11	34
Two mica-granitic gneiss	KAM 200	Bi	1526.3	2.05	4371.14	11.1873		169±7	
		Mu	766.7	5.67	469.78	2.7576		307±12	
		WR	262.2	50.05	15.29	.7977		430±83	
Biotite plagioclase gneiss	TR 170	WR	93.1	189.34	1.43±2	.7274±36			29
Biotite microcline plagioclase gneiss	TR 171	WR	167.3	59.02	8.25±14	.7662±46			
Biotite plagioclase gneiss	TR 172	WR	111.1	332.4	.70±10	.7213±43	.7179±24	424±9	
Muscovitic gneiss	TR 173	WR	336.5	36.28	27.27±38	.8744±26			
Augengneiss	TR 177	WR	336.0	22.22	45.02±59	1.0020±60			
Augengneiss	TR 178	WR	206.5	53.24	11.33±15	.7909±40			
Muscovitic gneiss	173	WR	336.5	36.28	27.27±38	.8744±26			
		Mu	1240.8	7.7	582.7±8.7	3.27±11			
		Mu	1236.6	--	--	--	.7573±56	302±11	
		Mu	1248.2	7.9	565.9±9.6	3.20±9			
		Mu	--	--	575.5±8.0	3.23±9			
Granodiorite	176	Bi	532.8	6.4	267.5±3.2	1.810±45			
		Bi	528.5	--	--	--	.7204±27	287±8	
		Bi	--	6.4	266.9±2.7	1.808±45			
		Bi	533.5	--	267.1±2.0	1.809±32			
Augengneiss	177	WR	336.0	22.22	45.02±59	1.0020±60			
		Bi	2014.2	2.8	4316±78	11.63±30			
		Bi	2017.2	--	--	--	.8918±68	172±4	
		Bi	--	2.6	4852±97	12.79±36			
Augengneiss	177	Bi	2029.4	--	4584±70	12.11±23			29
Augengneiss	178	Bi	1154.6	3.5	1519±26	6.88±16			
		Bi	1157.1	--	--	--			
		Bi	--	3.4	1584±28	7.08±18		283±7	
		Bi	1150.4	--	1150±22	6.97±12	.7439±4	293±7	
		WR	206.5	53.24	11.33±15	.7909±40			
		Mu	579.8	5.8	333±8	2.21±6			
		Mu	581.9	--	--	--		310±16	
<u>Merano-Mules-Anterselva basement</u>	AA1	WR	218	71	8.86	.7662			8
	AA2	WR	341	13.5	76.86	1.1951			
	AA3	WR	204	85	6.99	.7572			
Anterselva Orthogneisses	AA4	WR	438	13.1	103.19	1.3789			
	AA5	WR	331	18.6	53.30	1.0534			
	AA6	WR	316	10.6	91.68	1.3037	.7080±34	449±4	
	AA7	WR	469	18.5	78.75	1.1951			
	AA8	WR	226	63	10.37	.7700			
	AA9	WR	144	85	4.94	.7404			
Casies Orthogneisses	AA10	WR	107	204	1.51	.7175			
	AA11	WR	112	179	1.80	.7206			
Anterselva Orthogneiss	AA1	Bi	957	3.30	1328.75	6.6509			310±4
		Mu	586	7.26	262.87	1.9810			319±8
Anterselva Orthogneiss	AA2	Bi	1701	3.54	3306.46	14.7680			296±4
		Mu	1125	3.14	1929.54	9.4825			316±4
Anterselva Orthogneiss	AA3	Bi	1022	2.53	2292.23	10.5059			303±4
		Mu	582	7.93	235.69	1.8381			313±8
Anterselva Orthogneiss	AA4	Bi	2578	4.70	4698.00	20.7544			298±3
		Mu	1524	4.78	1532.28	7.4629			305±4
Anterselva Orthogneiss	AA 8	Bi	1196	5.36	894.16	4.6248			301±4
Casies Orthogneiss	AA 10	Bi	422	4.02	351.56	2.3136			304±6
	AA 20	WR	159	68	6.84	.7513			
	AA 21	WR	96	241	1.16	.7187			
	AA 22	WR	120	248	1.40	.7207			
Augengneisses	AA 23	WR	179	68	7.64	.7583	.7121	(409)	
	AA 24	WR	97	234	1.20	.7208			
	AA 25	WR	175	72	7.07	.7522			
	AA 26	WR	118	212	1.61	.7201			
Augengneiss	AA 21	Bi	532	13.24	117.02	.7727			18±8
Augengneiss	AA 22	Bi	605	5.16	344.42	.8577			24±4

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{m}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
Augengneiss	AA 24	Bi	441	6.43	200.06	.7907		17+ 5	
Augengneiss	AA 20	Mu	503	9.94	149.02	.8861		67+ 9	
Augengneiss	AA 23	Mu	568	10.15	164.91	.9004		67+ 8	
Augengneiss	AA 25	Mu	537	10.12	155.91	.8616		53+ 8	
Augengneiss	KAW 421	WR	291.3	27.7	31.01	.9083		451+18	33
		Ph	1026.4	4.7	733.3	2.399		149+ 5	
		Bi	1646.4	2.3	2770	3.927		78+ 3	
Orthogneiss	KAW 1138	WR	238.3	47.8	14.56	.8047		451+18	
		Mu	767.0	5.7	474.1	2.829		310+12	
		Mu	764.2	5.9	445.0	2.690		307+12	
		Bi	1369.8	4.9	1037	3.481		184+ 6	
		Bi	1351.0	5.2	935.6	3.266		188+ 7	
Orthogneiss	KAW 1141	WR	174.4	158.7	3.186	.7280		451+18	
		Mu	439.9	15.3	84.40	.8601		115+10	
		Bi	809.8	2.6	979.7	1.782		76+ 3	
Orthogneiss	KAW 1127	WR	84.4	158.0	1.549	.7231		451+18	
Orthogneiss	KAW 1127	Mu	254.9	44.8	16.50	.7419		88+47	33
		Bi	513.0	4.7	325.1	1.080		78+ 3	
	KAW 1135	WR	305.4	32.5	27.63	.8869			
	KAW 421	WR	291.3	27.7	31.01	.9083			
Orthogneisses	KAW 1138	WR	238.3	47.8	14.56	.8047			
	KAW 1141	WR	174.4	158.7	3.186	.7280	.7102+14	451+18	
	KAW 1127	WR	84.4	158.0	1.549	.7231			
Orthogneiss	KAW 1317	WR	460.7	11.6	123.7	1.511			
		WR	460.7	11.6	123.7	1.511		451+18	
		Mu	1545.7	4.0	2250	11.17		319+12	
		Mu	1574.3	3.9	2413	11.43		304+12	
		Bi	2639.5	4.2	8031	35.94		306+12	
Orthogneiss	KAW 1135	WR	305.4	32.5	27.63	.8869		451+18	
		Ph	1140.8	5.1	725.6	2.046		117+ 4	
		Bi	1623.8	3.7	1489	2.490		78+ 3	
Augengneiss	AA73-63	WR	227	32.9	19.98	.8365+ 18			10
		Mu	866	6.3	398	1.1849+ 35		65+ 1.4	
Anterselva granitic Orthogneiss	AA 5	WR	331	19	53.30	1.0534+ 54			
		Bi	1849	.9	5898.13	26.3400+520		304+10	
<u>Ötztal</u>	[AA74-34	WR	269	23.6	33.77	.9236+ 12			13
	AA74-35	WR	102	159	1.86	.7202+ 4			
	AA74-37	WR	91	191	1.43	.7152+ 5			
Orthogneisses	AA74-38	WR	107	118	762.94	.7247+ 4	.7073+ 5	448+14	
	AA74-39	WR	630	6.9	1.37	2.6445+ 49			
	AA74-40	WR	89	125	2.63	.7208+ 4			
	AA74-36	WR	95	193	312.91	.7154+ 5	.7076+ 8	434+12	
		Bi	563	2.8	2.08	3.9060+457		294+10	
<u>Stubai</u>	[AA76-21	WR	256	38	19.70	.8098+ 3			13
	AA76-23	WR	262	39	19.84	.8200+ 4			
Orthogneisses	AA76-24	WR	419	11.3	114.61	1.3909+ 8	.7091+135	425+43	
	AA76-25	WR	355	15	71.65	1.1415+ 7			
	AA76-2	WR	287	48	17.55	.8259+ 5			
	AA76-27	WR	235	42	16.46	.8194+ 4			
Orthogneiss	Bi	1348	4.1	1084.94	2.0751+ 41			82.7+ 2.5	
	Mu	811	5.6	504.25	2.8838+ 71			297+ 9	
<u>Merano-Mules-Anterselva basement</u>	KAW 1554	WR	145.4	62.89	6.716	.7494			28
	KAW 1555	WR	143.9	107.94	3.867	.7333			
	KAW 1556	WR	158.4	142.55	3.220	.7278			
Augengneisses	KAW 1796	WR	109.5	132.65	2.381	.7215	.7078+ 13	445+24	
	KAW 1799	WR	127.5	170.81	2.162	.7224			
	KAW 1797	WR	149.1	89.47	4.837	.7395			
	KAW 1798	WR	163.8	58.35	8.157	.7561			
Augengneiss	KAW 1554	Wm (80-100)	476.1	9.02	155.3	.8903		66.7+ 2.0	
		Wm (100-200)	477.2	9.53	147.3	.8803		65.4+ 2.0	
Augengneiss	KAW 1555	Wm (100-200)	488.7	12.29	117.1	.8838		93.6+ 3.0	
Augengneiss	KAW 1556	Wm (60-80)	469.3	15.17	91.86	.9828		195.8+ 5.0	
		Wm (80-100)	462.3	13.86	98.44	.9117		131.8+ 4.0	
		Wm (100-200)	451.1	17.87	73.88	.8335		101.1+ 3.0	

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material (mesh)	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_\text{M}$	$(^{87}\text{Sr}/^{86}\text{Sr})_\text{i}$	AGE Ma	Ref.
Augengneiss	KAW 1796	Bi (35-60)	532.3	13.23	116.9	.7639		26 ± 1.1	
		Bi (80-100)	546.9	12.30	215	.8026		26.8 ± .7	
		Wm (60-100)	289.7	19.80	42.50	.7503		50.6 ± 2.9	
Augengneiss	KAW 1799	Bi (80-100)	652.7	12.20	155.6	.7709		22.2 ± .8	
		Wm(100-150)	352.5	10.96	93.44	.7545		24.7 ± 1.3	
Augengneiss	KAW 1797	Wm (35-60)	453.6	11.11	119.5	.8419		62.8 ± 1.6	28
		Wm (60-80)	438.4	19.27	66.35	.7886		56 ± 2.2	
		Wm (80-100)	449.4	10.80	121.7	.8332		56.4 ± 1.5	
		Wm(100-200)	438.8	11.49	111.5	.8082		45.3 ± 1.4	
		Wm(200-270)	436.9	15.92	79.97	.7817		39.5 ± 1.7	
Augengneiss	KAW 1798	Wm (60-80)	526.9	9.93	155.9	.8774		57.7 ± 1.5	
		Wm(100-200)	619.2	9.39	193.4	.8543		37.3 ± 1	
Augengneiss lamella	KAW 1557	WR	178.1	66.96	7.7377	.7618			
		Wm (35-60)	626.7	10.19	181.7	.9257		66.3 ± 2.0	
		Wm (60-80)	634.3	9.26	202.4	.9305		60.9 ± 1.3	
		Wm (80-100)	582.2	8.15	211.1	.9316		58.7 ± 2.3	
		Wm(100-200)	630.5	10.43	178.2	.9058		59.4 ± 2.4	
Orthogneiss	AA80-18	WR	66	61	3.17	.7376 ± 2			
		Bi	638	3.9	591.31	3.2485±57		300 ± 5	25
		Mu	213	195	3.16	.7381 ± 1			
Orthogneiss	AA80-19	WR	147	72	7.57	.7543 ± 4			
		Bi	1001	2.6	2142.45	9.90624±287		301 ± 5	
Granitic Orthogneiss	AA74-10	WR	273	30	26.68	.9035 ± 4			
		Mu	1032	4.7	766.6	2.7365±10		174 ± 3	15
Granitic Orthogneiss	AA77-4	WR	244	30.5	23.44	.8334 ± 5			
		Mu	816	5.3	469.2	1.2713 ± 20		69.1 ± 1	35
Leucogranitic gneiss	T 1023	Mu	1423.8	3.84	2169.8	11.001		328 ± 12	
		WR	397.1	18.53	65.02	1.1688		495 ± 20	
Leucogranitic gneiss	T 1066	Mu	1784.2	6.73	1137.5	5.555		293 ± 12	
		WR	476.6	29.89	47.69	1.0162		450 ± 25	
Granitoid gneiss	T 1246	Mu	941.2	4.03	760.2	1.9307		95 ± 3.8	
		WR	292.2	21.02	41.34	.9576		420 ± 28	
Granitoid gneiss	T 1246	WR	292.2	21.63	40.49	.9620		437 ± 29	35
		Mu	978.7	3.82	832.7	1.9075		82.7 ± 3.3	
Granitoid gneiss	T 1247	WR	293.6	20.83	42.01	.9758		444 ± 29	
		Mu	1310.8	5.88	718.6	1.8207		91.1 ± 3.6	
Granitoid gneiss	T 1285	Mu	403.1	34.04	35.15	.9368		453 ± 33	
		WR	405.2	34.10	35.26	.9349		447 ± 33	
Granitoid gneiss	T 1288	Mu	707.7	11.89	178.8	.9841		82.3 ± 5.4	
		WR	200.3	38.15	15.36	.7929		379 ± 70	
Leucogranitic gneiss	T 1298	Mu Ø > .43 mm	1158.8	9.01	399.1	1.3482		90.7 ± 3.6	
		Mu Ø .15-.071	1130.6	9.50	367.0	1.3032		90.5 ± 3.6	
		Feld+Qz	186.2	55.15	9.93	.8471		86.8 ± 10	
		WR	332.8	49.99	19.73	.8592		531 ± 54	
Leucogranitic gneiss	T 1299	Mu	1127.0	4.97	725.9	1.7561		85.7 ± 3.4	
		Feld+Qz	157.2	28.26	16.46	.8925		84.4 ± 7.4	
		WR	356.8	24.77	42.73	.9240		352 ± 28	
Leucogranitic gneiss	T 1307	Mu	1568.7	4.76	1639.2	7.954		305 ± 12	
		WR	438.5	19.34	68.56	1.1318		432 ± 19	
Granitoid gneiss	T 1311	Mu	1321.4	3.75	1429.1	4.684		184 ± 7.6	
		WR	362.8	17.82	61.38	1.0966		442 ± 21	
Leucogranitic gneiss	T 1316	Mu	759.6	13.86	164.5	1.045		104 ± 4	
		WR	246.6	69.23	10.45	.8163		712 ± 102	
Granitoid gneiss	T 1318	Mu	651.8	15.81	122.12	.9102		98 ± 7.7	
		WR	196.1	96.80	5.91	.7483		453 ± 173	
		WR	196.1	99.61	5.79	.7483		463 ± 477	
Leucogranitic gneiss	T 1319	Mu	1250.7	6.03	662.3	1.7364		92.5 ± 3.7	
		WR	323.3	28.01	34.18	.9108		412 ± 33	
Leucogranitic gneiss	T 1325	Mu Ø > .43 mm	957.5	51.73	54.44	.8371		105 ± 13	
		Mu Ø .07-.15	946.9	36.54	76.42	.8530		84.7 ± 9	35
Leucogranitic gneiss	T 1325	WR	261.1	44.09	17.31	.7819		291 ± 62	
		Mu Ø > .43 mm	1116.4	13.52	258.7	1.4989		189 ± 8	
		Mu Ø < .15 mm	1095.2	11.78	288.0	1.3027		106 ± 4	
		WR	287.9	14.70	58.25	.9565		297 ± 21	
Leucogranitic gneiss	T 1343	Mu Ø > .43 mm	957.5	51.73	54.44	.8371		310 ± 22	
		Mu Ø < .15 mm	286.9	15.12	56.45	.9593			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_\text{m}$	$(^{87}\text{Sr}/^{86}\text{Sr})_\text{f}$	AGE Ma	Ref.
Granitic gneiss	T 1345	Mu	565.6	20.86	80.47	.8473		104 ± 10	
		WR	190.1	133.02	4.16	.7348		417 ± 245	
Granitic gneiss	T 1347	Mu	833.8	7.48	341.90	1.2172		90.2 ± 3.6	
		WR	242.4	51.88	13.67	.7966		444 ± 78	
Granitic gneiss	T 1348	Mu	823.2	4.58	573.06	1.6529		99.1 ± 4	
		WR	268.2	30.98	25.56	.8822		472 ± 44	
Granitic gneiss	T 1349	Mu	397.5	11.33	105.90	.8675		91.7 ± 8.6	
		Bi	612.6	3.13	612.17	1.4061		77.5 ± 3.1	
		KF	176.7	211.1	2.45	.7322		195 ± 126	
		KF	196.8	206.1	2.75	.7324		230 ± 149	
Leucogranitic gneiss		WR	141.3	103.4	3.98	.7364		464 ± 257	
	T 1356	Mu	982.2	4.38	734.5	2.0082		112 ± 4.5	
		WR	267.1	33.07	23.82	.8718		476 ± 47	
	T 1357	Mu	678.	7.46	279.1	1.2797		135 ± 5.4	
Leucocratic layer in paragneiss	T 1359	Mu	319.7	42.89	21.73	.7480		83.1 ± 32	
		WR	87.3	251.6	1.01	.7235		926 ± 996	
Muscovite pegmatite	T 1364	Mu	1621.7	3.42	1874.8	4.384		135 ± 5.4	
		WR	385.1	27.93	40.98	.8673		270 ± 30	
Leucogranitic gneiss	T 1366	Mu	617.2	8.53	219.5	1.0885		100.4 ± 4	
		WR	225.8	43.83	15.08	.7967		403 ± 71	
Biotite granite gneiss	T 1367	Mu	392.9	16.08	71.64	.8154		88 ± 13	35
		Bi	650.1	3.51	570.2	1.3385		75.6 ± 3	
		WR	161.1	179.3	2.61	.7292		507 ± 386	
	T 1368	Mu	1091.7	7.98	421.7	1.3440		85.8 ± 3.4	
Leucogranitic gneiss		Feld+Qz	214.1	35.67	17.8	.8495		106 ± 12	
		WR	304.2	36.58	24.52	.8597		428 ± 45	
	T 1370	Mu Ø > 30 mm	1056.4	5.73	593.4	1.8213		109 ± 4.4	
		Mu Ø 15-30	1070.5	5.77	602.1	1.8345		109 ± 4.4	
Leucogranitic gneiss		WR	297.8	23.22	38.18	.9633		465 ± 31	
	T 1371	Mu	961.0	7.18	418.2	1.3925		87.8 ± 3.5	
		WR	284.4	21.26	39.64	.9201		372 ± 29	
	T 1373	Mu Ø > 30 mm	1420.3	12.24	386.1	2.1904		255 ± 10	
Leucogranitic gneiss		Mu Ø < 30 mm	1215.4	19.16	195.7	1.2501		151 ± 6	
		WR	259	29.16	26.23	.8855		469 ± 43	
	T 1378	Mu Ø > 30 mm	1713.5	2.57	2932.2	5.939		118 ± 4.8	
		Mu Ø 15-30	1724.1	2.42	3081.4	5.610		105 ± 4.2	
Leucogranitic gneiss		WR	461.1	21.10	66.06	1.1267		443 ± 20	
	T 1066	WR	476.6	29.89	47.69	1.0162			
	T 1246	WR	292.2	21.02	41.34	.9576			
	T 1247	WR	293.6	20.83	42.01	.9758			
Granitoid gneiss	T 1285	WR	403.1	34.04	35.15	.9368			
	T 1288	WR	200.3	38.15	15.36	.7929			
Leucogranitic gneiss	T 1307	WR	438.5	19.34	68.56	1.1318			
	T 1311	WR	362.8	17.82	61.38	1.0966			
Granitoid gneiss	T 1318	WR	196.1	96.80	5.91	.7483			
	T 1319	WR	323.3	28.01	34.18	.9108			
Granitic gneiss	T 1345	WR	190.1	133.02	4.16	.7348			
	T 1347	WR	242.4	51.88	13.67	.7966		7105 ± 3	35
Granite gneisses	T 1348	WR	268.2	30.98	25.56	.8822		441 ± 7	
	T 1349	WR	141.3	103.4	3.98	.7364			
Leucogranitic gneiss	T 1356	WR	267.1	33.07	23.82	.8718			
	T 1357	WR	217.3	79.24	7.80	.7608			
Leucogranitic gneiss	T 1366	WR	225.8	43.83	15.08	.7967			
	T 1367	WR	161.1	179.3	2.61	.7292			
Leucogranitic gneisses	T 1368	WR	304.2	36.58	24.52	.8597			
	T 1370	WR	297.8	23.22	38.18	.9633			
	T 1373	WR	259.0	29.16	26.23	.8855			
	T 1378	WR	461.1	21.10	66.06	1.1267			
Raibl Schist	KAW 420	Bi	392	1.89	644.51	1.441		80 ± 3	32
Monteneve		Bi	392	1.34	937.47	1.774		80 ± 3	
Two mica plagioclasic gneiss	KAW 303	Bi	392.2	3.51	335.51	1.082		80 ± 4	34
		Mu	277	--	--	--			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
Garnet quartz micaschist	KAW 304	Bi	353.3	5.12	204.36	.941		82+ 6	
		Mu	255.1	--	--	--			
Ore dyke	KAW 109	Bi	978.7	3.47	901.43	1.780		84+ 3	
		Bi	975.1	3.38	924.67	1.809		84+ 3	
Garnet micaschist	KAW 305	Bi	491.1	4.59	321.42	1.094		85+ 4	
		Mu	112.4	--	--	--			
Biotitic porphyroblastic gneiss	KAW 306	Bi	345.9	2.94	--	1.1117		80+ 4	
<u>Detzta</u>		Mu	112.4	--	--	--			
Plagioclastic Biotite-Muscovite gneiss	KAW 307	Bi	519.4	5.86	266.06	1.082		98+ 5	
		Mu	187.3	--	--	--			
	169	Bi	334.2	--	--	--			29
		Bi	332.8	--	--	--			
		Bi	334.2	12.6	--	--			85+39
		Bi	335.3	--	77 +85	.812+28			
Biotite plagioclase gneisses	170	WR	93.1	189.34	1.43+ 2	.727+36			
		Bi	420.4	7.7	161.8 + 1.4	.942+44			
		Bi	424	--	--	--			
		Bi	418	7.8	157.1 + 1.90	.902+39			
		Bi	--	--	160.1 + 1.25	.923+26			
		Mu	171	145.9	3.40+ 7	.723+ 6			
		Mu	171	--	--	--			
	175	Bi	466.7	13.4	102.2 + 2	.879+25			29
		Bi	462.1	--	--	--			
		Bi	--	13.7	99.16 + 1.15	.853+22			
		Bi	462.1	--	100.5 + 1.4	.863+14			
Biotite plagioclase gneisses	172	WR	111.1	332.4	.699+10	.721+43			
		Bi	463.5	15.4	87.7 + 9	.837+19			
		Bi	462.1	--	--	--			
		Bi	464.9	15.1	91.2 + 1	.837+20			
		Bi	--	--	89 + 1.6	.837+10			
		Mu	166.1	304.6	1.600+32	.728+29			
		Mu	171.4	--	--	--			
		Mu	166.1	305.9	1.573+25	.722+29			
		Mu	168.2	--	1.580+20	.725+10			
Biotite microcline plagioclase gneiss	171	WR	167.3	59.02	8.25 +14	.766+46			
		Bi	994.2	2.7	1463 +20	4.65 +16			
		Bi	995.6	--	--				
		Bi	1002.7	2.6	1538 +40	4.70 +16			
		Bi	--	--	1478 +40	4.67 +11			
Merano-Mules-Anterselva basement	AA73 19	WR	92	82	3.23	.7331			8
Paragneiss	AA73 22	WR	169	138	3.55	.7354			
Gneissic micaschist	AA73 23	WR	186	144	3.74	.7366			
Gneissic micaschist	AA73 25	WR	133	164	2.35	.7296			
	AA73 26	WR	55	218	.73	.7148			
Paragneisses	AA73 27	WR	71	253	.81	.7154			
	AA73 28	WR	139	241	1.67	.7200			
	AA73 24	WR	72	102	2.05	.7274			
Gneissic micaschist	AA 40	Bi	584	10.34	176.22	1.5053		313+ 9	8
	AA 41	Bi	464	4.69	327.84	2.1813		313+ 7	
	AA 42	Bi	391	8.35	136.46	.7768		29+ 9	
Paragneisses	AA 43	Bi	482	9.06	155.14	.7849		29+ 6	
	AA 44	Bi	461	2.95	460.49	.8929		26+ 3	
	AA 45	Bi	720	17.08	124.15	.8930		98+ 9	
Micaschist	AA 46	Bi	597	3.70	494.12	1.3151		85+ 3	
	AA 47	Bi	630	9.78	198.37	1.3600		227+ 7	
Paragneisses	AA 48	Bi	597	3.98	443.44	.9271		33+ 3	
	AA 49	Bi	407	5.60	231.18	1.7152		301+ 8	
Monteneve Garnet micaschist	KAW 1208	WR	128.5	77.7	4.795	.7246		--	33
		Mu	351.9	99.2	10.29	.7329		107+ 75	
Garnet micaschist	KAW 1207	WR	157.5	137.1	3.334	.7350		126+146	
		Mu	279	152.8	5.299	.7385		80+ 3	
		Bi	627.1	4.7	400.1	1.184			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{f}}$	AGE Ma	Ref.
Garnet micaschist	KAW 1150	WR	109.3	174	1.820	.7217		--	
		Mu	253.4	294.3	2.495	.7228		122 \pm 305	
		Bi	478.7	4.6	313.5	1.080		81 \pm 3	
Garnet micaschist	KAW 1151	WR	75.9	137.9	1.596	.7220		--	
		Mu	211.7	145.3	4.223	.7251		82 \pm 181	
		Bi	365.7	4	269.4	1.015		77 \pm 3	
Paragneiss	KAW 1136	WR	149.3	112.9	3.835	.7322		--	
		Mu	214.5	93.9	6.625	.7383		155 \pm 117	
		Bi	389.3	9.1	125.9	.8704		80 \pm 7	
Biotite Muscovite schist	KAW 1134	WR	140.8	82.4	4.961	.7393		--	
		Bi	717.6	3	743.1	1.537		76 \pm 3	
Paragneiss	KAW 1126	WR	159.3	141.5	3.265	.7293		--	
		Mu	164.1	86.1	5.533	.7380		271 \pm 140	
		Bi	399.2	4	300.5	1.057		78 \pm 3	
Paragneiss	KAW 1320	WR	90.2	321.4	0.81	.7095		--	
		Bi	442	21.8	59.09	.7724		76 \pm 13	
Paragneiss	AA73-66	WR		148	2.83	.7296 \pm 6			
		Bi	523	3.6	418.83	2.4279 \pm 30		287 \pm 1	8
Paragneiss	AA73-67	WR	96	218	1.28	.7161 \pm 4			
		Bi	440	9.8	129.71	1.2081 \pm 24		269 \pm 2	
Paragneiss	AA73-70	WR	109	112	2.82	.7237 \pm 5		21.9 \pm 4.3	
		Bi	505	3.7	390.18	.8447 \pm 99			
Paragneiss	AA73-71	WR	98	164	1.73	.7227 \pm 6		19.9 \pm 8.6	
		Bi	453	16.3	80.58	.7450 \pm 36			
Gneissic micaschist	AA73-77	WR	171	107	4.63	.7436 \pm 22		42.9 \pm 11.9	
		Bi	752	34.4	63.61	.7795 \pm 30			
Paragneiss	AA73-85	WR	193	195	2.86	.7398 \pm 6			
		Bi	757	6.8	331.51	1.0047 \pm 30		57.7 \pm 1.7	
Paragneiss	AA73-91	WR	182	91	5.79	.7381 \pm 21		208 \pm 23	
		Mu	323	25.4	37.28	.8315 \pm 38			
Paragneiss	AA73-91	Bi	611	12.3	152.21	1.3244 \pm 40		278 \pm 4	
		WR	136	123	3.22	.7351 \pm 22		313 \pm 69	
Paragneiss	AA73-92	WR	315	96	9.59	.7635 \pm 8			
		Bi	173	159	3.15	.7306 \pm 22		293 \pm 3	
Paragneiss	AA73-93	WR	160	335	1.38	.7192 \pm 5		298 \pm 1	
		Bi	499	5.5	262.52	1.8287 \pm 45			
Paragneiss	AA73-94	WR	130	229	1.31	.7181 \pm 5		289 \pm 2	
		Bi	501	5.4	270.96	1.8270 \pm 70			8
Paragneiss	AA73-95	WR	120	192	1.81	.7202 \pm 4		304 \pm 2	
		Bi	452	2.9	454.37	2.6785 \pm 135			
Paragneiss	AA74-1	WR	145	131	3.18	.7308 \pm 8		267 \pm 1	
		Bi	573	2.1	800.59	3.7665 \pm 145			
Paragneiss	AA74-2	WR	107	179	1.72	.7189 \pm 2		275 \pm 2	
		Bi	404	9	129.51	1.2190 \pm 24			
Paragneiss	AA74-7	WR	124	149	2.40	.7298 \pm 3		293 \pm 1	
		Bi	606	4.4	400.06	2.3896 \pm 10			
Paragneiss	AA74-8	WR	81	175	1.34	.7229 \pm 6		307 \pm 2	
		Bi	513	15.1	103.95	1.1713 \pm 9			
Paragneiss	AA74-13	Bi	605	3.8	481.84	1.1552 \pm 14		61.4 \pm .3	
		WR	177	123	4.18	.7388 \pm 2			
Paragneiss	AA74-18	WR	95	19	14.74	.7589 \pm 30		199 \pm 7	
		Bi	380	12.3	91.70	.9765 \pm 20			
Paragneiss	AA74-19	WR	169	116	4.23	.7372 \pm 22		203 \pm 2	
		Bi	578	6.7	271.01	1.5082 \pm 45			
Paragneiss	AA74-21	Bi	590	4.6	368	.8766 \pm 20		27.8 \pm 1.2	
		WR	148	124	3.46	.7326 \pm 5		95 \pm 34	
Gneissic micaschist	AA74-23	WR	116	30	11.27	.7214 \pm 6		49.9 \pm 1	
		Bi	532	4.8	321.84	.9417 \pm 19			
Paragneiss	AA74-26	WR	48	84	1.64	.7255 \pm 5		48.4 \pm 1.8	
		Bi	428	4.8	256.47	.9008 \pm 29			

TABLE 4 (continued)

Rb/Sr

Rock type	Sample n.	Material (mesh)	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{f}}$	AGE Ma	Ref.
Paragneiss	AA74-27	WR	196	204	2.78	.7297+ 3		297 \pm 3	8
		Bi	735	5	427.72	2.5251+208			
Paragneiss	AA74-29	WR	124	212	1.69	.7263+ 6		180 \pm 1	28
		Bi	625	5.2	347.29	1.6120+50			
Paragneiss	AA74-30	WR	191	317	1.74	.7303+ 2		144 \pm 1	25
		Bi	727	7.3	287.52	1.3134+40			
Hornblend-Garnet gneiss	KAW 1794	WR	82.4	218.24	1.092	.7169		24.3 \pm 0.8	28
		Bi (100-200)	380.7	5.82	190.7	.7823			
		WM (80-100)	188.3	167.04	3.266	.7245			
Paragneiss	AA79-9	WR	81	481	.48	.7129+ 1		41 \pm 1	25
		Bi	479	12.7	110.06	.7761+ 6			
Gneissic micaschist	AA80-1	WR	113	257	1.28	.7192+ 1		79 \pm 1	25
		Bi	402	6	196.72	.9397+10			
		Mu	259	192	3.90	.7224+ 1			
Gneissic micaschist	AA80-2	WR	103	122	2.44	.7302+ 2		241 \pm 4	326 \pm 23
		Bi	635	6.2	328.51	1.8501+32			
		Mu	232	176	3.82	.7366+ 2			
Gneissic micaschist	AA80-3	WR	120	321	1.08	.7185+ 2		78 \pm 1	93 \pm 12
		Bi	419	4.8	260.71	1.0074+19			
		Mu	196	127	4.48	.7230+ 5			
Gneissic micaschist	AA80-4	WR	125	79	4.61	.7464+ 1		80 \pm 1	75 \pm 18
		Bi	627	4.8	392.75	1.1871+36			
		Mu	191	202	2.74	.7444+ 4			
Gneissic micaschist	AA80-6	WR	68	374	.53	.7174+ 1		80 \pm 2	208 \pm 3
		Bi	387	3.1	375.71	1.1420+49			
		Mu	619	2.7	803.37	3.0842+89			
Gneiss	AA80-8	WR	140	267	1.52	.7160+ 4		208 \pm 3	189 \pm 3
		Bi	619	2.7	803.37	3.0842+89			
		Mu	670	3.5	647.96	2.4588+15			
Micaschist	AA80-16	WR	179	118	4.40	.7441+ 3		81 \pm 1	72 \pm 33
		Bi	740	3	782.33	1.6359+54			
		Mu	226	203	3.23	.7429+ 4			
Micaschist	AA80-17	WR	152	112	3.94	.7394+ 3		81 \pm 1	69 \pm 35
		Bi	616	2.9	652.88	1.4851+40			
		Mu	277	252	2.61	.7381+ 6			
Paragneiss	AA80-20	WR	152	228	1.93	.7227+ 9		104 \pm 2	274 \pm 22
		Bi	477	6.5	218.91	1.0421+ 5			
		Mu	120	215	1.62	.7206+ 2			
Micaschist	AA80-28	WR	430	4.3	303.76	1.1702+41		105 \pm 2	310 \pm 125
		Bi	159	208	2.21	.7232+10			
		Mu	146	284	1.49	.7194+ 3			
Gneissic micaschist	AA80-30	WR	503	4.7	318.68	1.0927+25		83 \pm 1	274 \pm 22
		Bi	220	228	2.80	.7245+ 2			
		Mu	42	162	.75	.7168+ 1			
Paragneiss	AA80-34	WR	363	12.9	82.41	.8122+ 8		82 \pm 1	25
		Bi	109	112	2.82	.7237+ 5			
		Mu	236	248	2.76	.7251+ 4			
Paragneiss	AA73-95	WR	120	192	1.81	.7202+ 4		394 \pm 174	321 \pm 8
		Bi	174	198	2.54	.7243+18			
		Mu	303	94.7	9.32	.7229+ 6			
Paragneiss	AA74-8	WR	138	214	1.87	.7255+ 2		218 \pm 81	325.5 \pm 15
		Bi	195	262	2.16	.7264+ 2			
		Mu	196	204	2.78	.7297+ 3			
Paragneiss	AA74-27	WR	288	156	5.37	.7417+ 2		112 \pm 8	25
		Bi	160	107	4.34	.7449+2			
		Mu	328	122	7.8	.7504+3			
Micaschist	AA78-27	WR	126	203	1.8	.7220+3		90.8 \pm 78	136 \pm 17
		Bi	265	363	2.11	.7224+2			
		Mu	240	214	3.26	.7234+8			
Micaschist	AA78-37	WR	183	116	4.57	.7432+3		50 \pm 36	35
		Bi	188	201	2.71	.7396+2			
		Mu	566.1	30.43	54.67	.7822			
Calcareous biotite phenegite schist	T 749							92.7 \pm 19	

TABLE 5
Rb/Sr analytical data - PENNIDES

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
<u>S.W. Tauern Window</u>									
Tonalitic granite	76	WR	84.8	270	.91	.7099±32			2
Tonalitic granite	77	WR	219	293	2.2	.7073±24			
Augengneisses	84	WR	66.8	338.2	.572	.7117±19			
	86	WR	134.3	154	2.522	.7184±19			
	87	WR	147.7	199.7	2.138	.7141±14			
Aplitic granites	80	WR	225.4	41.9	15.57	.7789±23			
	80a	WR	222.9	42.1	15.31	.7808±23	.7170±15	292 ± 5	
	81	WR	61.8	56.8	3.15	.7304±14			
Tonalitic granite	85	WR	282.3	23	35.46	.8650±16			
	76	Bi	432	8.3	150.8	.7772±38	.7095	31.6 ± .9	
	77	Bi	425	8.7	143	.7814±72	.7061	37.1 ± 3.7	
	78	Bi	650	7.8	243.5	.7997±56		26.7 ± 1.8	
	79	Bi	39	7.4	153	.7694±50		28 ± 2.8	
	82	Bi	542	6	262.3	.8016±56		25.1 ± 1.6	
	83	Bi	418	20.6	58.9	.7371±37		35.1 ± 5.1	
	84	Bi	563	13.6	120.6	.7438±48	.7115	18.8 ± 3	
	86	Bi	860	17.6	142.5	.7611±33	.7176	21.4 ± 1.9	
	87	Bi	779	19.3	117.3	.7506±34	.7134	22.3 ± 2.3	
Aplitic granite	80	Bi	1672	7.8	637.3	.9466±59	.7747	19 ± .7	
Aplitic granite	80a	Bi	1663	7	707.3	.9622±90	.7768	18.4 ± .9	
Aplitic granite	85	Bi	1622	13.3	363.8	.9946±54	.8510	27.8 ± 1.2	
Porphyric granite gneiss	KAW 111	WR	157.2	147.15	3.113	.7198			31
Biotite muscovite augengneiss	KAW 110	WR	266.4	57.87	13.47	.7575			
		WR		57.77			.7086	255	
Phengite gneiss	KAW 113	WR	108.8	153.11	2.066	.7162			
Biotite phengite gneiss	KAW 114	WR	71.4	--	.753	.7106			
Porphyric granite gneiss	KAW 111	Bi	851.5	3.82	662.9	.9318			
		Bi	844.4	3.75	669.5	.9358		22.9 ± 1.7	
Biotite muscovite augengneiss	KAW 110	Bi	1374.4	3.04	1377	1.1116		18.3 ± .9	
		Bi		2.99					
Phengite gneiss	KAW 113	Bi	692.5	3.44	596.7	.8808		19.5 ± 1.9	
		Bi	688.9	3.52	579.2	.8771		19.7 ± 1.9	
		Ph	455.8	8.30	160.8	.7734		25.4 ± 5.5	
Biotite phengite gneiss	KAW 114	Bi	515.8	3.25	467.7	.8378		19.2 ± 2.3	
		Ph	309.8	20.24	44.71	.7294		30 ± 19.7	
		Ph	311.3	20.34	44.68	.7292		30 ± 17.6	
<u>S.E. Tauern Window</u>									
Coarse Leucogranites	6702	WR	186.6	6.28	88.6	1.0220 ± 1			16
	6701	WR	163.8	5.74	85	1.0080 ± 2			
	6703	WR	173.1	6.86	75.1	.9922 ± 14			
		WR	--	--	--	.9932 ± 9			
		WR	168.7	6.91	72.6	.9883 ± 18			
	6712	WR	--	--	--	.8466 ± 2	.7141 ± 33	253 ± 6	
		WR	144.5	11	38.4	.8497 ± 16			
		WR	--	--	--	.8467 ± 17			
		WR	139.2	--	37.6	--			
	6707	WR	125.7	15.08	24.3	.8053 ± 16			
Coarse Leucogranites	6706	WR	159.2	29.08	15.9	.7700 ± 19			
		WR	--	--	--	.7697 ± 9			
	6713	WR	129.8	43	8.60	.7406 ± 14			
		WR	--	--	--	.7402 ± 9			
	6709	WR	152.5	67.93	6.52	.7316 ± 14			
		WR	--	--	--	.7320 ± 7			
	6711	WR	125.4	64.78	5.61	.7280 ± 14			
		WR	--	--	--	.7295 ± 10			16
		WR	115.3	--	--	.7275 ± 7			
		WR	120.5	--	--	--			
6708 L4		WR	--	--	--	--			
6708 L1		WR	185	163.88	3.28	.7230 ± 16			
		WR	--	--	--	.7209 ± 7			

TABLE 5 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{M}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{I}}$	Age Ma	Ref.
Fine grained leucocratic granites-gneisses	A 310	WR	270.5	102.54	7.67	.7344+36			
	A 312	WR	246.5	105.14	6.77	.7271+36			
	A 341	WR	211.2	162.04	3.77	.7190+36			
	A 337	WR	143.2	539.15	.75	.7116+35			
Fine leucogranites	A 414	WR	4.9	321.12	.04	.7097+10			
	A 417	WR	213	197.27	3.2	.7169+10			
Pegmatite	A 379	Mu	1832	.89	5932	21.5			
Leucogranite	C 6712	Mu	542	5.07	310	.8932			
	A11	--	--	--	--	.8244			
		WR	142	10.8	38	.8474			
Leucogranite	C 6711	Mu	393	17.9	63.5	.7447			
	Bi	644	14	132.7	.7564				
	Pl	17.7	55.9	.9	.7250				
		WR	120	64.6	5.61	.7284			
Tonalite	C 6710	Bi	418	37.2	32.5	.7169			
	Ap	.7	739	.002	.7105				
	Pl	17.7	55.9	.15	.7086				
Tonalite	C 6716	Bi	339	62.3	15.8	.7125			
	Ap	--	143	--	.7094				
	Ep	3.9	2201	.005	.7074				
	Pl	11	323	.1	.7096				
W Tauern Window									33
Porphyroblastic gneiss	KAW 913	WR	202.7	198.2	2.962	.7218			
		Ph	556.5	56.1	28.74	.7274			
		Bi	1008	17.1	171.6	.7556			
Micaceous gneiss	KAW 878	WR	135.6	104.3	3.765	.7198			
		Bi	1027.1	4.2	709.7	.8511			
Spotted granite	KAW 817	WR	132.1	275.2	1.386	.7122			
		Ph	516.9	27.7	54.50	.7517			
		Bi	1014	6.2	475.7	.7391			
Phengite arcasic gneiss	KAW 974	WR	199.5	805.1	.7177	.7144			
		Ph	646.5	86.4	21.67	.7214			
		Bi	781.5	47.3	47.90	.7226			
Meta-tonalite	KAW 1131	WR	106.1	282.3	1.088	.7113			
		Bi	466.4	18.3	73.99	.7280			
Meta-tonalite	KAW 1314	WR	68.1	360.4	.5593	.7084			
		Bi	405.2	16.5	71.12	.7233			
Granitic gneiss	KAW 422	WR	132.5	579.2	.6612	.7083			
		Bi	886.4	26.9	95.52	.7269			
Biotite schist	KAW 1142	WR	161.2	217.6	2.145	.7131			
		Bi	556.5	12.3	131	.7385			
Qz-biotite schist	KAW 1139	WR	164.1	48.6	9.782	.7128			
		Ph	302.8	21.6	40.69	.7273			
		Bi	429.3	2.9	433.1	.7868			
Calc-micaschist	KAW 1128	WR	77.6	168.7	1.331	.7120			
		Ph	347.8	3.1	327.5	.8067			
		Bi	604.2	1.4	1247	.9638			
Calc-micaschist	KAW 1147	WR	103.3	188.8	1.535	.7129			
		Bi	619.3	11.8	152.6	.7454			
									33
	KAW 1179	WR	144	27.1	15.46	.7656			
	KAW 816	WR	122.1	64.8	5.50	.7329			
		--	--	--	5.50	.7315			
	KAW 818	WR	157.2	32.5	14.15	.7617			
		--	--	--	14.15	.7618			
	KAW 1178	WR	171.7	25.3	19.74	.7818			
Granitic gneisses	KAW 856	WR	110.6	48.9	6.59	.7383	.7143+23	244+27	
		--	--	--	6.59	.7406			
	KAW 1176	WR	138.6	40.1	10.04	.7508			
	KAW 811	WR	163.9	24.4	19.70	.7867			
		--	--	--	19.70	.7845			
	KAW 824	WR	173.6	30.3	16.78	.7728			
		--	--	--	16.78	.7715			

TABLE 5 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{III}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{I}}$	AGE Ma	Ref.
Granitic gneiss	KAW 1179	WR	144	27.1	15.46	.7656		244 \pm 27	
		Ph	540.2	3	536.9	.9352		23 \pm 1.7	
		Bi	573.8	1.3	1280	.9754		11 \pm .7	
Granitic gneiss	KAW 816	WR	122.1	64.8	5.50	.7329		244 \pm 27	
		WR	--	--	5.50	.7315		--	
		Ph	520.4	20.6	73.64	.7872		56 \pm 11	
Granitic gneiss	KAW 818	WR	157.2	32.5	14.15	.7617		244 \pm 27	
		WR	--	--	14.15	.7618		--	
		Ph	682.6	3.8	532.4	.9071		20 \pm 1.7	
Granitic gneiss	KAW 1178	WR	902.7	5.4	492.5	.8141		7 \pm 1.7	
		Ph	171.7	25.3	19.74	.7818		244 \pm 27	
		Bi	703.8	3.4	625.4	1.037		30 \pm 1.6	
Granitic gneiss	KAW 856	WR	355.4	5.9	176.2	.8079		11 \pm 4.8	
		WR	110.6	48.9	6.59	.7383		244 \pm 27	
		WR	--	--	6.59	.7406		--	33
Granitic gneiss	KAW 856	Ph	659.6	8.3	234.6	.8138		24 \pm 3	
		Bi	820.4	2.6	961.2	.9018		12 \pm 1	
		WR	138.6	40.1	10.04	.7508		244 \pm 27	
Granitic gneiss	KAW 1176	Ph	608.4	6.4	275.9	.8183		18 \pm 3	
		Bi	884.7	3.3	800.6	.8769		11 \pm 1	
		WR	163.9	24.4	19.7	.7867		244 \pm 27	
Granitic gneiss	KAW 811	WR	--	--	19.7	.7845		--	
		Ph	735.9	3.4	644.2	.9443		18 \pm 1.4	
		Bi	1109.7	7.4	442.4	.8128		4 \pm 1.9	
		Bi	1155.3	--	454.5	.8171		4	
		Ch	135.7	5.4	73.26	.7645		-29 \pm 10	
		Ch	17.7	--	9.93	.7589		59	
		P1	21.8	12.9	4.94	.7810		27 \pm 41	
		KF	455.4	53.8	24.67	.7863		-6 \pm 41	
		Gr	15.3	18.1	2.46	.7744		50 \pm 41	
Granitic gneiss	KAW 824	WR	173.6	30.3	16.78	.7728		244 \pm 27	
		WR	--	--	16.78	.7715		--	
		Ph	714.4	4.6	488.6	1.036		39 \pm 2	
Pegmatite	KAW 827	Bi	904.8	23.1	113.9	.7487		-18 \pm 7	
		WR	142.7	175.6	2.36	.7189		--	
		Mu	1532.6	9.4	575.7	2.950		273 \pm 4	
<u>E Tauern Window</u>	FT 1	Ph	714.4	14	150.3	.9077		90 \pm 6	
		WR	271	61	12.88	.7646			17
		WR	281	64	12.71	.7642			
		WR	279	64	12.70	.7632			
		WR	265	69	11.14	.7579			
		WR	249	94	7.65	.7405			
Granitic gneisses	FT 8	WR	247	83	8.63	.7456			
		WR	251	70	10.47	.7581			
		WR	253	70	10.47	.7524			
		WR	229	87	7.64	.7438			
		WR	218	141	4.50	.7328			
		WR	216	139	4.51	.7338			
<u>S-W Tauern Window</u>	FT 9	WR	216	139	4.51	.7338			
		WR	144	205	2.03	.7126 \pm 17			
		Bi	880	5.5	462.34	.8225 \pm 21			
		WR	105	232	1.31	.7117 \pm 7			
		Bi	591	7.4	230.06	.7597 \pm 14			
		WR	108	274	1.15	.7128 \pm 4			
Granitic orthogneiss	AT 1	Bi	494	1.3	1139.07	.9753 \pm 21			
		WR	80	223	1.04	.7108 \pm 7			
Granitic orthogneiss	AT 2	Bi	629	3.3	553.20	.8435 \pm 12			
		WR	137	143	2.77	.7192 \pm 8			
Granitic orthogneiss	AT 6	Bi	976	3	951.95	.9219 \pm 49			
		WR	115	246	1.35	.7124 \pm 5			
Granitic orthogneiss	AT 9	Bi	535	1.4	1077.86	.9612 \pm 30			

TABLE 5 (continued)

Rb/Sr

Rock type	Sample n.	Material	Rb ppm	Sr ppm	$^{87}\text{Rb}/^{86}\text{Sr}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{m}}$	$(^{87}\text{Sr}/^{86}\text{Sr})_{\text{i}}$	AGE Ma	Ref.
Granitic orthogneiss	AT 10	WR	135	262	1.48	.7119 ⁺⁷		17.2 ^{+2.0}	
		Bi	593	3.5	488.12	.8310 ⁺⁵⁶			
Calcschist	AA73-4B	WR	40	239	.48	.7095 ⁺⁵		13 ^{+ 4.3}	
		Bi	704	15.5	131.71	.7338 ⁺²⁹			
S-E Tauern Window	MS 3520	Wm	539	54	29	.79498		214 ^{+ 8}	
Banded gneiss		Bi	892	28.5	90	.72978		17.4 ^{+ .3}	18
		Feld	66	813	.24	.70737			
Banded gneiss	MS 4000	Wm	340	53.8	18.4	.76756		226 ^{+ 8}	
		Bi	643	21.9	.84.4	.72077		16 ^{+ .3}	
		Feld	60	548	.32	.70962			

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