

Lead-tellurium oxysalts from Otto Mountain near Baker, California: X. Bairdite, $\text{Pb}_2\text{Cu}^{2+}\text{Te}^{6+}\text{O}_{10}(\text{OH})_2(\text{SO}_4)(\text{H}_2\text{O})$, a new mineral with thick HCP layers

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ABSTRACT

Bairdite, $\text{Pb}_2\text{Cu}^{2+}\text{Te}^{6+}\text{O}_{10}(\text{OH})_2(\text{SO}_4)(\text{H}_2\text{O})$, is a new tellurate-sulfate from Otto Mountain near Baker, California, U.S.A. It occurs in vugs in quartz associated with khinite, cerussite, goethite, and hematite. It is interpreted as having formed from the partial oxidation of primary sulfides and tellurides during or following brecciation of quartz veins. Bairdite is monoclinic, space group $P2_1/c$, with unit-cell dimensions $a = 14.3126(10)$, $b = 5.2267(3)$, $c = 9.4878(5)$ Å, $\beta = 106.815(7)$ °, $V = 679.41(7)$ Å³, and $Z = 2$. Bairdite occurs as diamond-shaped tabular crystals up to about 250 µm long and 5 µm thick, in subparallel and fan-shaped aggregates. The color is lime green, the streak is pale lime green, and the luster is adamantine. The Mohs hardness is estimated at between 2 and 3. Bairdite is brittle with an irregular fracture and one perfect cleavage on {100}. The calculated density based on the empirical formula is 6.062 g/cm³. Bairdite is biaxial (+), with calculated indices of refraction of $\alpha = 1.953$, $\beta = 1.966$, and $\gamma = 2.039$. The measured $2V$ is 47(2)°, dispersion is $r < v$, strong and the optical orientation is $Y = \mathbf{b}$; $Z \wedge \mathbf{a} = 34^\circ$ in obtuse angle β . The pleochroism is strong: Z (pale green) << X (green) < Y (green). Electron microprobe analyses (average of 4) provided: PbO 34.22, CaO 0.06, CuO 23.80, TeO₃ 26.34, SO₃ 5.74, H₂O 2.81 (structure), total 92.97 wt%. The empirical formula (based on 17 O atoms pfu) is: $\text{Pb}_{2.05}\text{Ca}_{0.01}\text{Cu}_{3.99}^{2+}\text{Te}_{2.00}^{6+}\text{S}_{0.96}\text{O}_{17.00}\text{H}_{4.16}$. The eight strongest powder X-ray diffraction lines are [d_{obs} in Å (hkl) I]: 4.77 (110, $\bar{1}02$) 50, 4.522 (002,011, $\bar{1}11$) 66, 3.48 (multiple) 62, 2.999 (311, $\bar{4}11$) 97, 2.701 (502, $\bar{1}13$, $\bar{2}13$) 79, 2.614 (013,020) 100, 1.727 (multiple) 65, and 1.509 (911,033,324) 83. The crystal structure of bairdite ($R_1 = 0.072$ for 1406 reflections with $F_o > 4\sigma F$) contains edge-sharing chains of Te⁶⁺O₆ and Cu²⁺O₆ octahedra parallel to \mathbf{b} that are joined by corner-sharing in the \mathbf{a} direction, forming thick stair-step-like hexagonal close packed layers parallel to {100}. The polyhedral sheet has similarities to those in the structures of timroseite and paratimroseite. The thick interlayer region contains PbO₁₀ polyhedra and half-occupied SO₄ groups. Raman and infrared spectral data are presented.

Keywords: Bairdite, new mineral, tellurate, crystal structure, Raman spectroscopy, infrared spectroscopy, HCP layers, timroseite, paratimroseite, Otto Mountain, California

INTRODUCTION

Bairdite, the new mineral described here, is the eleventh new Pb-Te oxysalt mineral (Table 1) to be described from the remarkable secondary mineral assemblage at Otto Mountain, near Baker, California, U.S.A. (Kampf et al. 2010a; Housley et al. 2011). Bairdite is named for Jerry A. Baird (b. 1940) of Lake Havasu City, Arizona. Baird, a mineral collector for 45 years, has collected extensively at Otto Mountain and has provided numerous samples for research. He provided one of the two co-type specimens of bairdite and provided one of the two co-types of the recently described mineral fuettererite (Kampf et al. 2013a). Baird has agreed to the naming of the mineral in his honor.

The new mineral and name have been approved by the Commission on New Minerals, Nomenclature and Classification of the International Mineralogical Association (IMA2012-061).

TABLE 1. New minerals described from Otto Mountain

Mineral	Ideal formula	Reference
Ottite	$\text{Pb}_2\text{Te}^{6+}\text{O}_5$	Kampf et al. (2010a)
Housleyite	$\text{Pb}_6\text{Cu}^{2+}\text{Te}_4^{6+}\text{O}_{18}(\text{OH})_2$	Kampf et al. (2010b)
Thorneite	$\text{Pb}_6(\text{Te}^{6+}\text{O}_{10})(\text{CO}_3)\text{Cl}_2(\text{H}_2\text{O})$	Kampf et al. (2010c)
Markcooperite	$\text{Pb}_2(\text{UO}_2)\text{Te}^{4+}\text{O}_6$	Kampf et al. (2010d)
Timroseite	$\text{Pb}_2\text{Cu}^{2+}(\text{Te}^{6+}\text{O}_6)_2(\text{OH})_2$	Kampf et al. (2010e)
Paratimroseite	$\text{Pb}_2\text{Cu}^{2+}_2(\text{Te}^{6+}\text{O}_6)_2(\text{H}_2\text{O})_2$	Kampf et al. (2010e)
Telluroperite	$\text{Pb}_2\text{Te}^{4+}\text{O}_4\text{Cl}_2$	Kampf et al. (2010f)
Chromschieffelite	$\text{Pb}_{10}\text{Te}_6^{6+}\text{O}_{20}(\text{CrO}_4)(\text{H}_2\text{O})_5$	Kampf et al. (2012)
Fuettererite	$\text{Pb}_3\text{Cu}^{2+}\text{Te}^{6+}\text{O}_6(\text{OH})_7\text{Cl}_5$	Kampf et al. (2013a)
Agaite	$\text{Pb}_3\text{Cu}^{2+}\text{Te}^{6+}\text{O}_6(\text{OH})_7(\text{CO}_3)$	Kampf et al. (2013b)
Bairdite	$\text{Pb}_2\text{Cu}^{2+}\text{Te}_2^{6+}\text{O}_{10}(\text{OH})_2(\text{SO}_4)(\text{H}_2\text{O})$	This study

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Two co-type specimens, the second collected by one of the authors (B.T.), are deposited in the Natural History Museum of Los Angeles County, catalog numbers 64000 and 64001.

OCCURRENCE

Bairdite was found in the Bird Nest drift (35.27677°N , $116.09927^{\circ}\text{W}$) on the southwest flank of Otto Mountain, 0.4 miles northwest of the Aga mine, which is 1 mile northwest of Baker, San Bernardino County, California, U.S.A. Bairdite was also found in a quartz vein (designated NE3) northeast of the Bird Nest drift.

Bairdite is very rare and has been confirmed to occur on only four specimens, two of which were used in this study and have been designated co-types. Bairdite crystals occur in vugs in quartz in association with khinite, cerussite, goethite, and hematite. Other minerals found elsewhere on the specimens include wulfenite and galena. Other species identified in the mineral assemblages at Otto Mountain include acanthite, agaite, anglesite, anatacamite, atacamite, boleite, brochantite, burckhardtite, calcite, caledonite, celestine, cerussite, chalcopyrite, Br-rich chlorargyrite, chromschieffelite, chrysocolla, devilline, diaboleite, eztelite, fluorite, fornicite, frankhawthorneite, fuetterite, gold, hessite, housleyite, iodargyrite, jarosite, khinite, kuranakhite, linarite, malachite, markcooperite, mattheddleite, mcalpineite, mimetite, mottramite, munakataite, murdochite, muscovite, ottoite, paratimroseite, perite, phosphohedyphane, plumbojarosite, plumbotsumite, pseudoboleite, pyrite, telluroperite, thorneite, timroseite, vanadinite, and vauquelinite.

Bairdite and most of the other secondary minerals in the quartz veins are interpreted as having formed from the partial oxidation of primary sulfides (e.g., galena and chalcopyrite) and tellurides (e.g., hessite) during or following brecciation of the quartz veins. Additional background on the occurrence is provided in Kampf et al. (2010a) and Housley et al. (2011).

PHYSICAL AND OPTICAL PROPERTIES

Bairdite occurs as diamond-shaped, tabular crystals up to about $250\ \mu\text{m}$ long and $5\ \mu\text{m}$ thick, in subparallel and fan-shaped aggregates (Figs. 1 and 2). Tablets are flattened on $\{100\}$ and bounded by the $\{011\}$ form (Fig. 3). No twinning was observed optically under crossed polars or based upon single-crystal X-ray diffraction. The color is lime green, the streak is pale lime green, and the luster is adamantine. Bairdite does not fluoresce under long-wave or short-wave ultraviolet light. The Mohs hardness could not be measured, but is estimated to be between 2 and 3, based upon the behavior of crystals when broken. The new mineral is brittle with irregular fracture and one perfect cleavage on $\{100\}$. The density could not be measured because it is greater than those of available high-density liquids and there is insufficient material for physical measurement. The calculated density based on the empirical formula and single-crystal cell is $6.062\ \text{g}/\text{cm}^3$. In cold, dilute HCl, bairdite crystals rapidly turn opaque white, and then dissolve slowly.

The indices of refraction could not be measured because of the small amount of material available and the difficulty in working with liquids of sufficiently high index of refraction using a spindle stage. We have endeavored to provide optical properties based upon a combination of measurements and

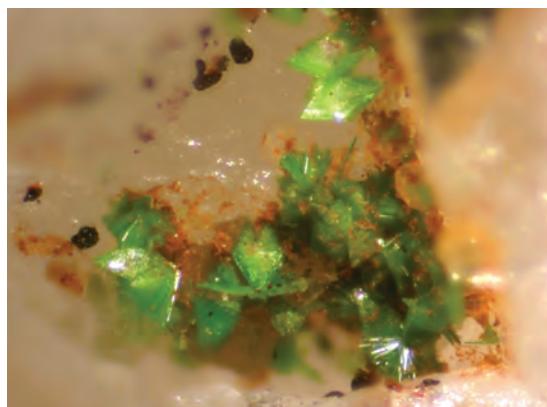


FIGURE 1. Crystals of bairdite on quartz on co-type specimen NHMLAC 64000, FOV 1.5 mm. (Jerry Baird image; color online).

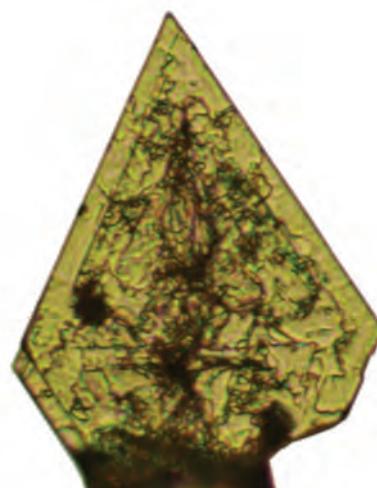


FIGURE 2. Bairdite crystal used in the morphological and optical studies (100 μm across; plane-polarized light; color online).

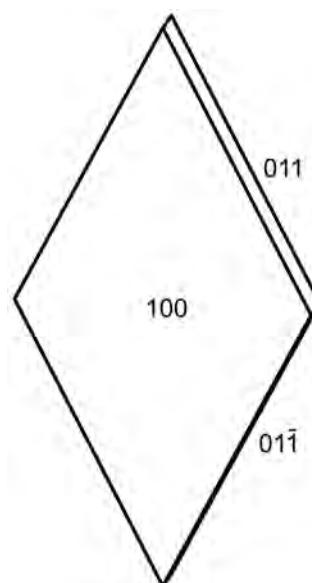


FIGURE 3. Crystal drawing of bairdite (clinographic projection).

calculations. Bairdite is biaxial (+), with indices of refraction $\alpha = 1.953$, $\beta = 1.966$, and $\gamma = 2.039$. These were calculated from the retardation, $\beta - \alpha = 0.013$, (measured with a Berek compensator), $2V_{\text{meas}} = 47(2)^\circ$ (measured directly on a spindle stage), and $n_{\text{av}} = 1.986$ (based upon the Gladstone-Dale relationship for the ideal composition; Mandarino 2007). The dispersion is strong, $r < v$. The optical orientation is: $Y = \mathbf{b}$, $Z \wedge \mathbf{a} = 34^\circ$ in obtuse angle β . Bairdite is strongly pleochroic: Z (pale green) << X (green)> Y (green).

INFRARED AND RAMAN SPECTROSCOPY

An infrared spectrum (Fig. 4) was obtained using a Nicolet Magna 860 FTIR with a KBr beamsplitter, DTGS detector, and a Nicolet Continuum infrared microscope operating with a 50 \times 50 μm aperture and without the use of a polarizer. The sample was a triangular half of a diamond-shaped platelet that varied between 9.1 and 9.6 μm thick. The Raman spectrum (Fig. 5) was obtained on the same crystal using a Renishaw M-1000 spectrometer with 20 mW argon ion laser operating at 514.5 nm. The spot was about 1 μm in diameter with about 5 mW at the sample when using a 100 \times objective lens at 100% laser power.

Prominent features visually apparent in the infrared spectrum are a set of broad bands at about 3356, 3117, 2638, 2351, 2021, 1723, and 1613 cm^{-1} . Sharper features occur at 1208 cm^{-1} (medium), strong overlapping features at about 1281 and 1060 cm^{-1} , weaker features at 973 and 896 cm^{-1} , a stronger band at 716 cm^{-1} , and probably features near 681 and 666 cm^{-1} , where noise begins to dominate the spectrum. The absorption features

in the 2400 to 3117 cm^{-1} region (and possibly specifically the 2638 cm^{-1} band) arise from OH stretching from either OH or H_2O . The broad band at 1613 cm^{-1} is attributable to the H_2O bending modes. The band at 716 cm^{-1} is assigned to TeO_6 and that at 1060 cm^{-1} is assigned to SO_4 .

The Raman spectrum is dominated by a feature at 721 cm^{-1} , with other significant features at 977, 634, 558, 518, 378, 336, 238, and 208 cm^{-1} . With the exceptions of the 977 cm^{-1} sulfate and 721 cm^{-1} tellurate features, there are comparatively few overlaps between the dominant IR and Raman features in the wavenumber region where the traces coincide. Other features in the Raman spectrum have not been definitely assigned, but it is likely that the 634 cm^{-1} feature is from sulfate and the features between 300 and 400 cm^{-1} are from tellurate.

CHEMICAL COMPOSITION

Quantitative chemical analyses (4) of bairdite were performed using a JEOL JXA-8200 electron microprobe at the Division of Geological and Planetary Sciences, California Institute of Technology. Analyses were conducted in WDS mode at 20 keV and 10 nA. A 10 μm beam diameter was used for the first analysis and a 1 μm beam diameter was used for three subsequent analyses. Crystals of bairdite proved very difficult to polish because of their fragile nature and perfect cleavage. The smaller beam diameter was used because flat areas on the sample were limited and generally very small. The sample was analyzed for As, Bi, Ca, Cu, Fe, P, Pb, S, Sb, Si, Te, V, and Zn, but only Ca, Cu, Pb, S, and Te were above the detection limits. The standards used were: anorthite (for Ca), cuprite (for Cu), galena (for Pb and S), and Sb_2Te_3 (for Te). Also, no other elements were detected in EDS analyses. Analytical results are given in Table 2. There was insufficient material for CHN analyses, so H_2O was calculated on the basis of 2 Te, charge balance and 17 total O atoms pfu, as determined by the crystal-structure analysis (see below). Infrared spectroscopy (see above) confirmed the presence of OH and H_2O and the absence of CO_3 . Note that bairdite is prone to electron beam damage, which contributes to the low analytical total. This is a common feature observed in most secondary tellurate species (e.g., Kampf et al. 2010a, 2010b, 2010c, 2010d, 2010e, 2010f, 2012, 2013a, 2013b; Mills et al. 2009, 2010).

The empirical formula (based on 17 O atoms pfu) is



The simplified formula is $\text{Pb}_2\text{Cu}_4^{2+}\text{Te}_2^{6+}\text{O}_{10}(\text{OH})_2(\text{SO}_4)(\text{H}_2\text{O})$, which requires PbO 36.24, CuO 25.83, TeO₃ 28.51, SO₃ 6.50, H₂O 2.92, total 100 wt%.

TABLE 2. Chemical analytical data for bairdite

Constituent	Average	Range	St.dev.	Normalized wt%
PbO	34.22	32.59–35.59	1.37	36.81
CaO	0.06	0.03–0.09	0.03	0.06
CuO	23.80	23.64–23.96	0.15	25.60
TeO ₃	26.34	25.88–26.69	0.35	28.33
SO ₃	5.74	5.38–5.98	0.27	6.17
H ₂ O*	2.81			3.02
Total	92.97			99.99†

* Based on the crystal structure (2 Te, charge balance and 17 O apfu).

† Rounding error.

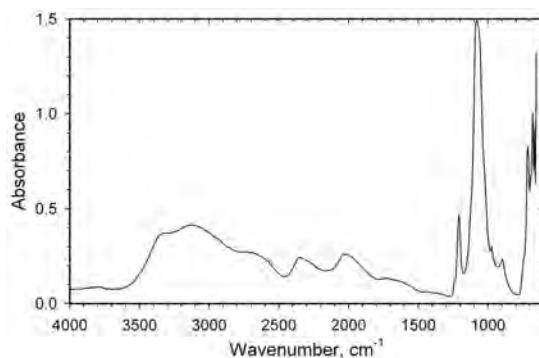


FIGURE 4. Transmission infrared spectrum through the (100) face of a crystal of bairdite.

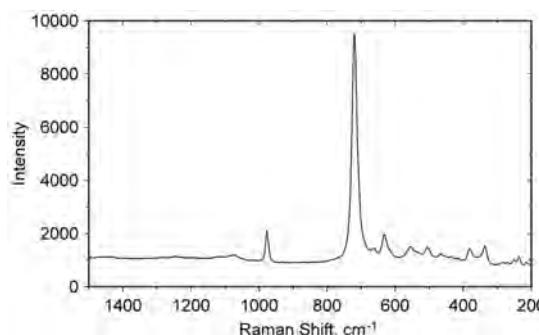


FIGURE 5. Raman spectrum obtained from the (100) face of a crystal of bairdite.

X-RAY CRYSTALLOGRAPHY AND STRUCTURE DETERMINATION

All powder and single-crystal X-ray diffraction data were obtained on a Rigaku R-Axis Rapid II curved imaging plate microdiffractometer utilizing monochromatized MoK α radiation. Observed powder d -values (with standard deviations) and intensities were derived by profile fitting using JADE 2010 software. Data (in angstroms) are given in Table 3. Unit-cell parameters refined from the powder data using JADE 2010 with whole-pattern fitting are: $a = 14.354(11)$, $b = 5.223(11)$, $c = 9.503(11)$ Å, $\beta = 107.04(2)^\circ$, and $V = 681.2(1.7)$ Å³. The observed powder data fit well with those calculated from the structure, also using JADE 2010. The relatively low precision of the cell refined from the powder data are attributable to the use of MoK α radiation.

The Rigaku CrystalClear software package was used for processing of the diffraction data, including the application of an empirical multi-scan absorption correction using ABSCOR (Higashi 2001). The structure was solved by direct methods using SHELXS-97 software and was refined using SHELXL-97 (Sheldrick 2008).

The S site refined to approximately half occupancy and three of the four O sites (O8, O9, and O10) associated with the SO₄ group also refined to roughly half occupancy, while the fourth (O7) refined to full occupancy. In the final refinement these sites were assigned half and full occupancies, accordingly, and isotropic displacement parameters were used for the O8, O9, and O10 sites. The occupancies of these sites are consistent

TABLE 4. Data collection and structure refinement details for bairdite

Diffractometer	Rigaku R-Axis Rapid II
X-ray radiation	MoK α ($\lambda = 0.71075$ Å)
Temperature	298(2) K
Ideal formula	Pb ₂ Cu ₄ ²⁺ Te ₂ ⁶⁺ O ₁₀ (OH) ₂ (SO ₄)(H ₂ O)
Space group	P2 ₁ /c
Unit-cell dimensions	$a = 14.3126(10)$ Å $b = 5.2267(3)$ Å $c = 9.4878(5)$ Å $\beta = 106.815(7)^\circ$
Z	2
Volume	679.41(7) Å ³
Density (for above formula)	6.021 g/cm ³
Absorption coefficient	35.303 mm ⁻¹
$F(000)$	1080
Crystal size	40 × 35 × 5 μm
θ range	4.1 to 27.48°
Index ranges	-18 ≤ h ≤ 18, -6 ≤ k ≤ 6, -12 ≤ l ≤ 12
Reflections collected/unique	15177/1554 [$R_{\text{int}} = 0.113$]
Reflections with $F_o > 4\sigma F$	1406
Completeness to $\theta = 25.01^\circ$	99.9%
Max. and min. transmission	0.8432 and 0.3325
Refinement method	Full-matrix least-squares on F^2
Parameters refined	138
GoF	1.173
Final R indices [$F_o > 4\sigma F$]	$R_1 = 0.0715$, $wR_2 = 0.1626$
R indices (all data)	$R_1 = 0.0774$, $wR_2 = 0.1656$
Extinction coefficient	0.0001(2)
Largest diff. peak/hole	+4.66/-2.38 e Å ⁻³

$$* R_{\text{int}} = \sum |F_o - F_c(\text{mean})| / \sum |F_o|. \quad \text{GoF} = S = \{\sum [w(F_o^2 - F_c^2)]^2 / (n - p)\}^{1/2}. \quad R_1 = \sum |F_o| - |F_c| / \sum |F_o|. \quad wR_2 = \{\sum [w(F_o^2 - F_c^2)] / \sum [w(F_o^2)]\}^{1/2}. \quad w = 1 / [\sigma^2(F_o^2) + (aP)^2 + bP] \text{ where } a = 0.0161, b = 120.54, \text{ and } P = [2F_c^2 + \text{Max}(F_o^2)]/3.$$

with one SO₄ group pfu, as is also indicated by the empirical formula. The full occupancy of the O7 site is consistent with it participating in the SO₄ group half of the time and being an H₂O

TABLE 3. X-ray powder diffraction data for bairdite

l_{obs}	d_{obs}	d_{calc}	l_{calc}	$h \ k \ l$	l_{obs}	d_{obs}	d_{calc}	l_{calc}	$h \ k \ l$
18	13.9(4)	13.7007	19	1 0 0			2.2650	5	2 1 3
34	6.90(4)	6.8503	75	2 0 0			2.2078	11	5 1 1
50	4.77(9)	{ 4.8834 4.7393	9 20	1 1 0 1 0 2			2.1866	5	5 1 3
66	4.522(13)	{ 4.5411 4.5301 4.4989	12 10 57	0 0 2 0 1 1 1 1 1			2.1582	9	3 2 2
15	4.13(18)	4.1272	9	1 1 1			2.1423	12	1 0 4
21	3.98(10)	3.9803	18	1 0 2			2.0777	10	4 2 0
17	3.85(4)	3.8196	24	3 0 2			2.0665	6	5 0 4
		{ 3.5503 3.4768 3.4280 3.4252	7 22 5 24	2 1 1 3 1 1 0 1 2 4 0 0			2.0636	7	2 2 2
62	3.48(5)						2.0606	5	3 1 3
		{ 3.4250					2.0288	16	4 2 2
14	3.35(3)	3.3633	38	2 0 2			1.9849	6	6 1 3
39	3.241(15)	3.2185	50	4 0 2			1.9572	6	7 0 0
97	2.999(9)	{ 3.0066 2.9441	100 36	3 1 1 4 1 1			1.9522	7	4 2 1
		{ 2.8648 2.8359	5 8	4 1 0 3 0 2			1.8993	11	7 1 1
		{ 2.7198 2.7008 2.6810	20 27 67	5 0 2 1 1 3 2 1 3			1.8844	8	5 2 2
79	2.701(5)						1.8674	12	4 1 3
		{ 2.6197 2.6134	43 30	0 1 3 0 2 0			1.8378	5	6 0 2
100	2.614(5)	{ 2.5665 2.5602 2.5114	9 22 8	3 1 3 4 1 1 0 2 1			1.8327	11	3 0 4
		{ 2.4417 2.4360 2.4214 2.4188	6 6 6 7	2 2 0 1 2 1 2 2 1 4 0 2			1.7831	5	2 1 5
7	2.521(18)						1.7384	16	6 2 2
		{ 2.3282 2.3179	5 5	6 0 2 3 0 4			1.7357	7	4 1 5
13	2.436(10)	{ 2.282(18)	13 12	6 0 0 3 2 1			1.7195	6	6 2 0
		{ 2.2834 2.2790					1.7158	9	0 1 5
							1.7092	14	7 1 1
							1.7017	6	3 2 3
							1.6936	5	5 1 3
							1.6568	5	1 2 4
							1.6328	7	8 1 3
							1.6266	7	0 3 2
							1.5972	27	6 1 5
							1.5798	9	3 0 6
							1.5743	5	3 3 1
							1.5103	9	9 1 1
							1.5100	11	0 3 3
							1.5005	7	3 2 4

Note: Only calculated lines with intensities of 5 or greater are listed.

TABLE 5. Occupancies, fractional coordinates, and atomic displacement parameters for bairdite

	Occ.	<i>x/a</i>	<i>y/b</i>	<i>z/c</i>	<i>U</i> _{eq}	<i>U</i> ₁₁	<i>U</i> ₂₂	<i>U</i> ₃₃	<i>U</i> ₂₃	<i>U</i> ₁₃	<i>U</i> ₁₂
Pb	0.873(9)	0.34148(7)	0.0795(3)	0.39773(15)	0.0218(5)	0.0184(6)	0.0217(9)	0.0261(7)	0.0012(6)	0.0077(4)	-0.0008(4)
PbA	0.049(9)	0.3421(13)	0.965(7)	0.437(3)	0.016(8)						
PbB	0.076(7)	0.3415(11)	0.850(5)	0.4112(18)	0.033(6)						
Te	1.0	0.16195(10)	0.4933(3)	0.49077(15)	0.0156(5)	0.0155(8)	0.0209(8)	0.0114(7)	-0.0005(5)	0.0053(5)	-0.0003(5)
Cu1	1.0	0.7926(2)	0.4758(5)	0.8472(3)	0.0167(7)	0.0181(14)	0.0214(15)	0.0116(13)	-0.0001(10)	0.0059(10)	0.0001(10)
Cu2	1.0	0.9079(2)	0.5282(5)	0.1916(3)	0.0184(7)	0.0206(14)	0.0195(15)	0.0166(14)	0.0012(11)	0.0079(11)	0.0008(11)
S	0.5	0.5334(8)	0.087(3)	0.1993(13)	0.024(3)	0.013(5)	0.035(7)	0.023(6)	0.005(5)	0.003(4)	0.003(5)
O1	1.0	0.8577(12)	0.144(3)	0.8362(15)	0.019(3)	0.036(9)	0.019(8)	0.007(7)	-0.001(6)	0.012(6)	0.008(7)
O2	1.0	0.0766(12)	0.209(3)	0.4836(18)	0.023(4)	0.020(8)	0.026(9)	0.018(8)	-0.009(7)	-0.001(7)	-0.004(7)
O3	1.0	0.2700(12)	0.291(3)	0.6011(18)	0.020(3)	0.023(8)	0.016(8)	0.023(8)	0.001(7)	0.009(7)	-0.002(6)
O4	1.0	0.0453(11)	0.696(3)	0.3687(16)	0.019(3)	0.017(8)	0.024(9)	0.011(7)	-0.005(6)	0.000(6)	-0.002(6)
O5	1.0	0.1757(13)	0.156(3)	0.808(2)	0.026(4)	0.025(9)	0.026(9)	0.033(10)	0.003(8)	0.017(8)	0.003(7)
O6	1.0	0.7648(12)	0.291(3)	0.012(2)	0.024(4)	0.021(8)	0.011(8)	0.040(10)	0.004(7)	0.008(7)	0.007(6)
O7	1.0	0.6267(18)	0.091(5)	0.183(3)	0.055(7)	0.052(14)	0.077(18)	0.048(13)	0.033(13)	0.032(12)	0.023(13)
O8	0.5	0.530(3)	0.032(7)	0.354(4)	0.030(8)						
O9	0.5	0.493(3)	0.862(9)	0.115(5)	0.046(11)						
O10	0.5	0.480(3)	0.329(8)	0.142(5)	0.038(9)						

half of the time, providing one H₂O pfu. It is also significant that O7 forms a long bond (2.46 Å) to Cu1, while O8, O9, and O10 bond to Pb and do not participate in either Cu or Te octahedra. The bond-valence sums (BVS) for O7 as half sulfate oxygen and half water oxygen support these assignments, although the BVS when it participates in the SO₄ group is rather high (2.36 v.u.). Because the complexities of the partial occupancies make a comprehensive assignment of hydrogen bonds from O7 acting as an H₂O problematic, we did not include those hydrogen bond contributions in our bond-valence analysis. The BVS for O4 (1.40 v.u.) is indicative of it being an OH group, although this value is somewhat high for an OH. The BVS for O2 (1.60 v.u.) is quite low for an O atom; however, the short distance between O4 and O2 (2.58 Å) is indicative of a strong hydrogen bond, which serves to balance the BVS for O4 and O2.

In the difference Fourier map, significant residual electron density (17.65 e/Å³) was noted, centered 1.04 Å from the Pb site. This was best modeled as two separate satellite Pb peaks

TABLE 6. Selected bond lengths (Å) in bairdite

Pb-O6	2.463(16)	Cu1-O5	1.913(18)	Te-O6	1.880(16)
Pb-O8 (x½)	2.60(4)	Cu1-O6	1.979(18)	Te-O2	1.912(17)
Pb-O10 (x½)	2.62(3)	Cu1-O1	1.984(16)	Te-O1	1.915(15)
Pb-O5	2.665(18)	Cu1-O3	2.002(16)	Te-O3	1.916(16)
Pb-O3	2.671(16)	Cu1-O2	2.415(17)	Te-O5	1.958(17)
Pb-O7	2.74(3)	Cu1-O7	2.46(2)	Te-O4	2.034(16)
Pb-O3	2.785(16)	<Cu-O>	2.126	<Te-O>	1.936
Pb-O9 (x½)	2.82(4)				
Pb-O7	2.85(3)	Cu2-O1	1.945(14)	S-O7	1.39(2)
Pb-O8 (x½)	2.86(4)	Cu2-O2	1.980(16)	S-O9	1.45(5)
Pb-O10 (x½)	2.99(4)	Cu2-O4	2.002(17)	S-O10	1.50(4)
Pb-O1	3.281(17)	Cu2-O5	2.040(18)	S-O8	1.51(4)
Pb-O9 (x½)	3.42(5)	Cu2-O4	2.355(15)	<S-O>	1.46
<Pb-O>	2.811*	Cu2-O6	2.572(18)		
		<Cu-O>	2.149	Hydrogen bond	
				O4-O2	2.58(2)

* Based upon 10-coordination with bond lengths to half-occupied O atoms given 0.5 weight.

(PbA and PbB), which together with the main Pb peak have a total refined occupancy of almost exactly 1 Pb. The significance of these sites is not clear as their distances from the nearest O

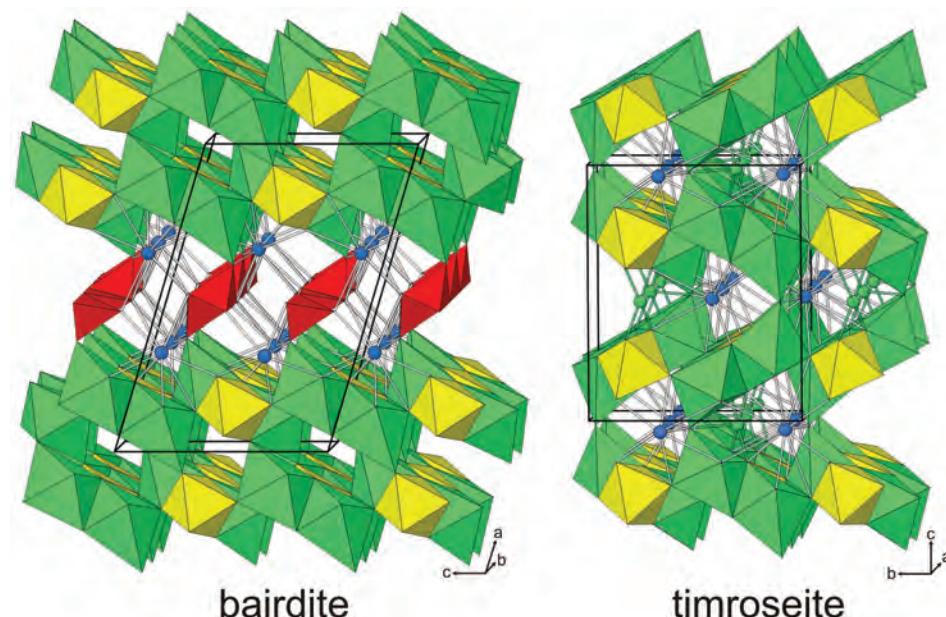


FIGURE 6. The structures of bairdite and timroseite. Pb atoms are dark gray (blue online), SO₄ tetrahedra are very dark gray (red online), TeO₆ octahedra are light gray (yellow online), CuO₆ octahedra are gray (green online) and 5-coordinate Cu atoms and corresponding bonds (for timroseite) are gray (green online).

TABLE 7. Bond valence sums for bairdite (values are expressed in valence units)

	O1	O2	O3	O4	O5	O6	O7 _S	O7 _W	O8	O9	O10	Σ
Pb	0.07		0.24		0.24	0.36	0.20 ^{x½} →	0.20 ^{x½} →	0.27 ^{x½} →	0.17 ^{x½} →	0.26 ^{x½} →	1.98
Cu1	0.44	0.14	0.42		0.53	0.44	0.16 ^{x½} →	0.16 ^{x½} →	0.16 ^{x½} →	0.05 ^{x½} →	0.12 ^{x½} →	2.09
Cu2	0.49	0.44		0.42	0.16	0.38	0.09					1.98
Te	1.01	1.02	1.01	0.82	0.94	1.08		1.88		1.36	1.60	5.76
S											1.40	6.24
H		0.26		0.74					1.36			1.00
Σ	2.01	1.86	1.86	2.14	2.09	1.97	2.36	0.48	1.79	1.82	1.78	

Notes: O7 is a sulfate oxygen half of the time (O7_S) and an H₂O half of the time (O7_W). Pb²⁺-O bond strengths are from Krivovichev and Brown (2001); Te⁶⁺-O bond strengths are from Mills and Christy (2013); Cu²⁺-O and S⁶⁺-O bond strengths are from Brown and Altermatt (1985) as is the hydrogen bond strength for O2...O4.

atoms are too short for Pb-O bonds and they do not appear to represent the 6s² lone electron pair of the Pb²⁺. Note that, because of the widely dispersed remaining residual electron density, it was not possible to locate the H atoms associated with the OH and H₂O groups.

Details concerning data collection and structure refinement are provided in Table 4. Fractional coordinates and atom displacement parameters are provided in Table 5, selected interatomic distances in Table 6 and bond valences in Table 7. Other data and a CIF are available on deposit¹.

DESCRIPTION OF THE STRUCTURE

In the structure of bairdite (Fig. 6), individual TeO₆ octahedra and pairs of edge-sharing Jahn-Teller distorted CuO₆ octahedra link by edge-sharing into chains along **b** (Fig. 7). The chains are linked to one another by corner-sharing to form stair-step-like layers parallel to {100}. The same corner-sharing joins two such layers into a thick double layer. The region between the double layers contains half-occupied SO₄ tetrahedra and Pb²⁺. The Pb²⁺ bonds to seven fully occupied O sites and six approximately half occupied O sites, yielding an effective coordination of 10 (Fig. 8). The Pb²⁺-O bonds cover a fairly broad range (2.46 to 3.42 Å); however, there is not a pronounced lopsided distribution of bond lengths typical of Pb²⁺ with stereoactive 6s² lone-pair electrons.

The same types of chains forming stair-step-like layers are found in the structures of timroseite and paratimroseite (Kampf et al. 2010b). The layer in bairdite is parallel to {100}, while those in timroseite and paratimroseite are parallel to {001}. As a consequence, the three minerals have two similar cell dimensions, those corresponding to the dimensions in these planes: for bairdite *b* = 5.2257 and *c* = 9.4848, for timroseite *a* = 5.2000 and *b* = 9.6225 Å, and for paratimroseite *a* = 5.1943 and *b* = 9.6198 Å. In the structures of timroseite and paratimroseite, the stair-step-like layers are assembled into frameworks by corner-sharing with each successive layer reversed in orientation. This arrangement for timroseite can be seen in Figure 6. It should be noted that one of the two Cu²⁺ polyhedra participating in the chain in the paratimroseite structure is a Cu²⁺O₅ square pyramid, but the overall chain and layer topology is otherwise the same as in timroseite. Another difference between the structures of timroseite and paratimroseite is that in timroseite the stair-step

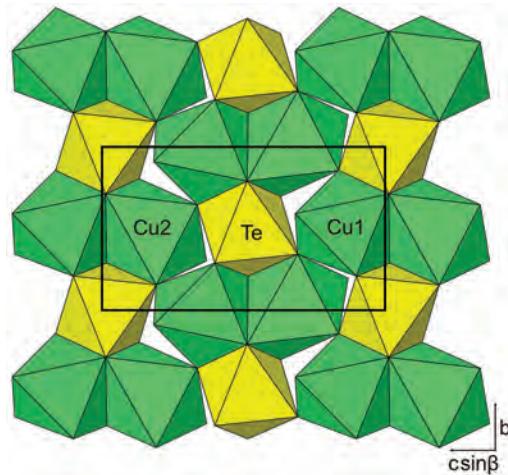


FIGURE 7. Stair-step-like layer of edge-sharing TeO₆ and CuO₆ octahedra, linked via shared corners in the structure of bairdite. Note that there are three stair-steps, increasing in elevation from left to right. (Color online.)

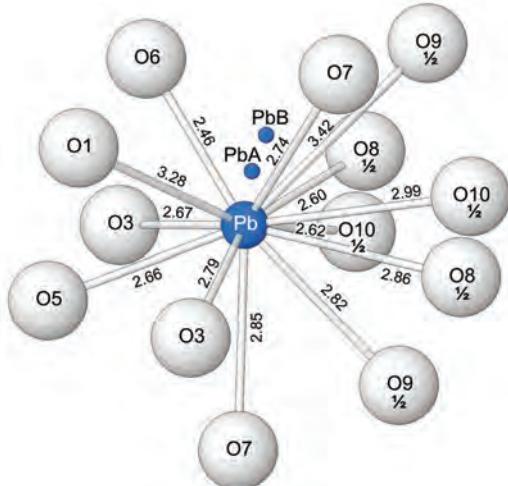


FIGURE 8. Pb coordination in bairdite showing Pb-O bond lengths in angstroms, Å. (Color online.)

layers are further linked to one another by an additional Cu²⁺O₅ square pyramid.

An interesting feature of the stair-step-like layers in the structures of bairdite, timroseite, and paratimroseite is that they are based upon hexagonal close packing (HCP), not only in terms

¹ Deposit item AM-13-707, CIF and data. Deposit items are available two ways: For a paper copy contact the Business Office of the Mineralogical Society of America (see inside front cover of recent issue) for price information. For an electronic copy visit the MSA web site at <http://www.minsocam.org>, go to the American Mineralogist Contents, find the table of contents for the specific volume/issue wanted, and then click on the deposit link there.

of the individual steps (or chains), but even with respect to the continuous assembly of steps. In the timroseite and paratimroseite structures, the HCP nature is flipped in successive layers as described above, so it does not extend over the entire framework. In the structure of bairdite, the entire thick double layer exhibits HCP and successive layers are in the same orientation, so the structure can be described as a stacking of stepped HCP layers interrupted by thick interlayer regions containing PbO_{10} polyhedra and SO_4 groups.

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Observed and calculated structure factors for bairdite

Page 1

h	k	l	10Fo	10Fc	10s	h	k	l	10Fo	10Fc	10s	h	k	l	10Fo	10Fc	10s	h	k	l	10Fo	10Fc	10s
3	0	0	938	677	20	8	5	0	412	426	35	-12	3	1	1076	1092	30	-1	6	1	121	30	39
4	0	0	2706	2646	50	9	5	0	235	214	40	-11	3	1	174	218	44	0	6	1	829	604	20
5	0	0	397	364	19	10	5	0	0	49	1	-10	3	1	1105	1158	26	1	6	1	663	473	25
6	0	0	3196	3044	71	11	5	0	805	755	31	-9	3	1	1082	1151	22	2	6	1	52	50	51
7	0	0	2469	2354	52	12	5	0	274	315	60	-8	3	1	319	316	20	3	6	1	638	427	23
8	0	0	966	913	27	0	6	0	373	191	44	-7	3	1	589	633	17	4	6	1	604	489	23
9	0	0	1038	973	26	1	6	0	1080	965	29	-6	3	1	626	649	16	5	6	1	185	108	37
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8	2	0	751	743	20	9	1	1	2834	2741	51	-2	4	1	1548	1500	35	13	0	2	1368	1225	43
9	2	0	513	518	19	10	1	1	613	504	20	-1	4	1	527	489	33	14	0	2	450	408	44
10	2	0	384	446	15	11	1	1	1289	1225	26	0	4	1	1565	1530	31	15	0	2	285	213	54
11	2	0	574	626	20	12	1	1	629	585	22	1	4	1	991	1012	17	16	0	2	747	604	43
12	2	0	1039	1108	25	13	1	1	741	674	24	2	4	1	858	843	29	-18	1	2	169	207	59
13	2	0	128	126	62	14	1	1	327	256	30	3	4	1	218	192	17	-17	1	2	121	42	66
14	2	0	179	178	40	15	1	1	1650	1556	36	4	4	1	1525	1546	37	-16	1	2	235	218	39
15	2	0	274	150	41	16	1	1	326	305	30	5	4	1	500	494	17	-15	1	2	459	422	25
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17	2	0	205	225	205	-17	2	1	346	362	45	7	4	1	1002	1075	23	-13	1	2	492	472	19
1	3	0	1324	1307	15	-16	2	1	398	428	37	8	4	1	678	738	24	-12	1	2	597	604	17
2	3	0	1083	1079	13	-15	2	1	789	779	35	9	4	1	412	381	28	-11	1	2	176	120	26
3	3	0	610	572	17	-14	2	1	549	575	28	10	4	1	490	493	29	-10	1	2	425	368	16
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5	3	0	1583	1630	25	-12	2	1	661	657	18	12	4	1	0	91	1	-8	1	2	371	303	12
6	3	0	372	314	18	-11	2	1	536	499	17	13	4	1	508	584	39	-7	1	2	416	397	11
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  Refinement of F^2^ against ALL reflections. The weighted R-factor wR  

and  

  goodness of fit S are based on F^2^, conventional R-factors R are based  

on F, with F set to zero for negative F^2^. The threshold expression of  

F^2^ > 2sigma(F^2^) is used only for calculating R-factors(gt) etc. and  

is  

  not relevant to the choice of reflections for refinement. R-factors  

based  

  on F^2^ are statistically about twice as large as those based on F, and  

R-  

  factors based on ALL data will be even larger.  

;  

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P=(Fo^2^+2Fc^2^)/3'
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O6 O 0.7648(12) 0.291(3) 0.012(2) 0.024(4) Uani 1 1 d . .
O7 O 0.6267(18) 0.091(5) 0.183(3) 0.055(7) Uani 1 1 d . .
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All esds (except the esd in the dihedral angle between two l.s. planes)
are estimated using the full covariance matrix. The cell esds are taken
into account individually in the estimation of esds in distances, angles
and torsion angles; correlations between esds in cell parameters are
only

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used when they are defined by crystal symmetry. An approximate
(isotropic)
treatment of cell esds is used for estimating esds involving l.s.
planes.
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Pb O10 2.62(4) 4_566 ?
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O3 Pb O9 139.1(8) 4_565 4_566 ?
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O8 PbA Cu1 85.5(12) 3_666 2_656 ?
O7 PbA Cu1 123.7(15) 2_655 2_656 ?
O10 PbA Cu1 88.8(12) 4_576 2_656 ?
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O8 PbA Cu2 142.9(13) 3_666 2_655 ?
O7 PbA Cu2 103.3(10) 2_655 2_655 ?
O10 PbA Cu2 135.8(14) 4_576 2_655 ?
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O10 PbA Cu1 161.9(14) 4_576 3_666 ?
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Cu2 PbA Cu1 56.8(4) 2_655 3_676 ?
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O8 PbB S 94.1(10) 3_666 2_655 ?
O9 PbB S 65.3(11) 4_576 2_655 ?
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O6 PbB O10 163.4(14) 2_655 2_655 ?
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O9 PbB O10 60.9(13) 4_576 2_655 ?
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Pb PbB Cu1 118.3(9) 1_565 3_666 ?
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O8 PbB Cu1 161.3(12) 3_666 3_666 ?
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O3 Te Cu2 87.1(5) . 2_645 ?
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O1 Cu2 O5 92.3(7) 4_565 3_666 ?
O2 Cu2 O5 79.7(7) 2_655 3_666 ?
O4 Cu2 O5 161.2(7) 2_645 3_666 ?
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O2 Cu2 O4 98.5(6) 2_655 1_655 ?

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05 Cu2 O4 93.8(7) 3_666 1_655 ?
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02 Cu2 O6 86.4(6) 2_655 . ?
04 Cu2 O6 70.2(6) 2_645 . ?
05 Cu2 O6 92.0(6) 3_666 . ?
04 Cu2 O6 173.0(6) 1_655 . ?
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05 Cu2 Te 40.7(5) 3_666 2_655 ?
04 Cu2 Te 100.4(4) 1_655 2_655 ?
06 Cu2 Te 86.6(4) . 2_655 ?
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02 Cu2 Cu1 48.5(5) 2_655 1_554 ?
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05 Cu2 Cu1 86.6(5) 3_666 1_554 ?
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09 S O10 114(2) 1_545 . ?
07 S O8 114.5(19) . . ?
09 S O8 105(2) 1_545 . ?
010 S O8 111(2) . . ?
07 S O10 112.5(16) . 2_645 ?
09 S O10 77(2) 1_545 2_645 ?
010 S O10 132.1(12) . 2_645 ?
08 S O10 29.9(18) . 2_645 ?
07 S PbB 35.0(11) . 2_645 ?
09 S PbB 113.7(19) 1_545 2_645 ?
010 S PbB 76.0(17) . 2_645 ?
08 S PbB 133.5(16) . 2_645 ?
010 S PbB 145.5(13) 2_645 2_645 ?
07 S Pb 50.3(11) . 2_655 ?
09 S Pb 128.4(19) 1_545 2_655 ?
010 S Pb 61.3(16) . 2_655 ?
08 S Pb 125.1(16) . 2_655 ?
010 S Pb 148.5(13) 2_645 2_655 ?
PbB S Pb 18.5(4) 2_645 2_655 ?
07 S Pb 52.8(13) . 2_645 ?
09 S Pb 52.1(18) 1_545 2_645 ?
010 S Pb 142.9(17) . 2_645 ?
08 S Pb 106.0(15) . 2_645 ?
010 S Pb 81.8(12) 2_645 2_645 ?
PbB S Pb 80.5(5) 2_645 2_645 ?
Pb S Pb 98.9(3) 2_655 2_645 ?
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09 S Cu1 110.2(19) 1_545 4_565 ?
010 S Cu1 125.9(17) . 4_565 ?
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O10 S Pb 79.5(16) . . ?
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Pb S Pb 131.4(4) 2_655 . ?
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Pb S Pb 76.8(3) 2_655 4_565 ?
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Cu1 S Pb 141.7(4) 4_565 4_565 ?
Pb S Pb 84.7(2) . 4_565 ?
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O10 S Pb 126.9(17) . 3_656 ?
O8 S Pb 28.7(14) . 3_656 ?
O10 S Pb 40.4(12) 2_645 3_656 ?
PbB S Pb 109.5(5) 2_645 3_656 ?
Pb S Pb 108.1(3) 2_655 3_656 ?
Pb S Pb 87.9(3) 2_645 3_656 ?
Cu1 S Pb 56.97(18) 4_565 3_656 ?
Pb S Pb 72.2(2) . 3_656 ?
Pb S Pb 153.1(4) 4_565 3_656 ?
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Te O1 Cu1 97.3(6) 2_646 . ?
Cu2 O1 Cu1 135.8(8) 4_566 . ?
Te O1 Pb 96.8(6) 2_646 3_656 ?
Cu2 O1 Pb 77.5(5) 4_566 3_656 ?
Cu1 O1 Pb 91.0(6) . 3_656 ?
Te O1 Cu1 58.3(4) 2_646 1_545 ?
Cu2 O1 Cu1 74.0(5) 4_566 1_545 ?
Cu1 O1 Cu1 135.3(7) . 1_545 ?
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Te O1 Cu2 59.7(4) 2_646 1_556 ?
Cu2 O1 Cu2 148.8(7) 4_566 1_556 ?
Cu1 O1 Cu2 57.7(4) . 1_556 ?
Pb O1 Cu2 133.7(5) 3_656 1_556 ?
Cu1 O1 Cu2 117.8(4) 1_545 1_556 ?
Te O1 Cu2 126.2(7) 2_646 3_766 ?
Cu2 O1 Cu2 69.5(5) 4_566 3_766 ?
Cu1 O1 Cu2 92.8(6) . 3_766 ?
Pb O1 Cu2 135.9(4) 3_656 3_766 ?
Cu1 O1 Cu2 131.9(4) 1_545 3_766 ?
Cu2 O1 Cu2 83.2(3) 1_556 3_766 ?
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Te O2 Cu1 90.2(6) . 2_646 ?
Cu2 O2 Cu1 93.6(7) 2_645 2_646 ?
Te O2 Cu2 77.0(5) . 3_666 ?
Cu2 O2 Cu2 169.3(8) 2_645 3_666 ?
Cu1 O2 Cu2 76.2(4) 2_646 3_666 ?
Te O2 Cu2 86.9(6) . 1_455 ?
Cu2 O2 Cu2 77.5(5) 2_645 1_455 ?
Cu1 O2 Cu2 170.0(7) 2_646 1_455 ?
Cu2 O2 Cu2 112.4(5) 3_666 1_455 ?
Te O2 Cu2 137.8(7) . 4_466 ?
Cu2 O2 Cu2 122.0(7) 2_645 4_466 ?
Cu1 O2 Cu2 88.6(5) 2_646 4_466 ?
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Pb O3 Cu2 151.2(6) 4_566 2_645 ?
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Cu2 O3 Cu2 93.8(4) 2_645 3_666 ?

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Cu2 O4 Te 109.1(7) 2_655 . ?
Cu2 O4 Cu2 113.5(7) 2_655 1_455 ?
Te O4 Cu2 126.1(8) . 1_455 ?
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Te O4 Cu1 60.3(4) . 2_656 ?
Cu2 O4 Cu1 168.5(6) 1_455 2_656 ?
Cu2 O4 Cu1 74.4(5) 2_655 3_666 ?
Te O4 Cu1 68.8(4) . 3_666 ?
Cu2 O4 Cu1 92.5(5) 1_455 3_666 ?
Cu1 O4 Cu1 98.9(4) 2_656 3_666 ?
Cu2 O4 Cu2 134.1(6) 2_655 2_645 ?
Te O4 Cu2 55.2(4) . 2_645 ?
Cu2 O4 Cu2 71.5(4) 1_455 2_645 ?
Cu1 O4 Cu2 115.5(4) 2_656 2_645 ?
Cu1 O4 Cu2 59.7(3) 3_666 2_645 ?
Cu2 O4 Cu1 89.2(5) 2_655 4_475 ?
Te O4 Cu1 147.5(6) . 4_475 ?
Cu2 O4 Cu1 64.1(4) 1_455 4_475 ?
Cu1 O4 Cu1 105.3(4) 2_656 4_475 ?
Cu1 O4 Cu1 143.5(4) 3_666 4_475 ?
Cu2 O4 Cu1 128.0(4) 2_645 4_475 ?
Cu1 O5 Te 126.4(10) 2_646 4_566 ?
Cu1 O5 Cu2 131.9(10) 2_646 3_666 ?
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Te O5 Pb 103.5(8) 4_566 4_566 ?
Cu2 O5 Pb 92.7(7) 3_666 4_566 ?
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PbA O6 Cu1 108.6(11) 2_645 1_554 ?
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Pb O10 Cu1 59.4(8) 4_565 3_666 ?
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