

THE CRYSTAL STRUCTURE OF FRANZINITE, THE TEN-LAYER MINERAL OF THE CANCRINITE GROUP

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ABSTRACT

Franzinite, ideally $[(\text{Na},\text{K})_{30}\text{Ca}_{10}][\text{Si}_{30}\text{Al}_{30}\text{O}_{120}](\text{SO}_4)_{10}\cdot 2\text{H}_2\text{O}$, a feldspathoid belonging to the cancrinite group, has a ten-layer stacking sequence. Its structure was solved in the space group $P321$ to an R value of 5.96%. Cell parameters are a 12.916(1), c 26.543(3) Å. The framework, characterized by the stacking sequence $ABCABACABC$, contains regular alternations of two "cancrinite" and two "sodalite" cages along $[0\ 0\ z]$, and of two "sodalite" and one "losod" cages along $[\frac{2}{3}\ \frac{1}{3}\ z]$ and $[\frac{1}{3}\ \frac{2}{3}\ z]$. The Si:Al ratio is equal to 1, and the framework has a perfectly ordered Si,Al distribution, as was found in liottite and afghanite, the 6- and 8-layer cancrinite-like minerals, respectively. The cages host a complex distribution of extra-framework cations (Ca, Na, K), $(\text{SO}_4)^{2-}$ anions, and H_2O molecules. The "losod" cage contains two sulfate groups aligned along z and separated by a triplet of cations. Two additional triplets of cations are located around the two sulfate groups. The two bases of the cage are occupied by calcium atoms, which make favorable bond-distances with the apical atoms of oxygen of the SO_4 tetrahedra within the cage. The two "cancrinite" cages share a common base and contain a segment of the $\dots\text{Na}-\text{H}_2\text{O}\dots\text{Na}-\text{H}_2\text{O}\dots$ chain, which is a characteristic feature of the cancrinite, vishnevite, and pitiglianoite structures. The "sodalite" cages host one sulfate group that is always disordered and displaced from the three-fold axis. The (Ca, K, Na) cations are distributed among various split sites, each one with partial occupancy, showing marked similarity with the "sodalite" cages of the $(\text{SO}_4)^{2-}$ - and S^{2-} -bearing minerals with sodalite-type structure.

Keywords: franzinite, cancrinite group, sodalite group, feldspathoids, structure determination, IR spectroscopy, stacking sequences.

SOMMAIRE

La franzinite, un feldspathoïde dont la composition idéale est $[(\text{Na},\text{K})_{30}\text{Ca}_{10}][\text{Si}_{30}\text{Al}_{30}\text{O}_{120}](\text{SO}_4)_{10}\cdot 2\text{H}_2\text{O}$, fait partie du groupe de la cancrinite, et possède une période d'empilement de dix couches. Nous avons résolu sa structure dans le groupe spatial $P321$ jusqu'à un résidu R de 5.96%. Les paramètres réticulaires sont a 12.916(1) et c 26.543(3) Å. La trame, dont l'empilement répond à la séquence $ABCABACABC$, contient des alternances régulières de deux cages de type "cancrinite" et deux cages de type "sodalite" le long de $[0\ 0\ z]$, et de deux cages de type "sodalite" et d'une cage de type "losod" le long de $[\frac{2}{3}\ \frac{1}{3}\ z]$ et de $[\frac{1}{3}\ \frac{2}{3}\ z]$. Le rapport Si:Al est égal à 1, et la trame possède un degré d'ordre parfait, tout comme c'est le cas pour la liottite et l'afghanite, les membres du groupe de la cancrinite à six et à huit couches, respectivement. Les couches renferment un assemblage de cations ne formant pas partie de la trame (Ca, Na, K), des anions $(\text{SO}_4)^{2-}$ et des molécules de H_2O , tous distribués de façon très complexe. La cage "losod" contient deux groupes sulfate alignés le long de z et séparés par un triplet de cations. Deux triplets additionnels de cations sont situés autour des deux groupes de sulfate. Les deux bases de la cage sont le site d'atomes de calcium, disposés à une distance favorable des atomes apicaux d'oxygène des tétraèdres de SO_4 à l'intérieur de la cage. Les deux cages de type "cancrinite" partagent une base commune et contiennent un segment de la chaîne $\dots\text{Na}-\text{H}_2\text{O}\dots\text{Na}-\text{H}_2\text{O}\dots$, attribué des structures

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de la cancrinite, la vishnevitte et la pitiglianoïte. Les cages de type "sodalite" cages contiennent un groupe sulfate invariablement désordonné, et déplacé de l'axe de rotation trois. Les cations (Ca, K, Na) sont distribués sur plusieurs sous-sites, chacun étant incomplètement rempli, et ressemblant ainsi aux cages de type "sodalite" des minéraux sulfatés et sulfurés possédant la structure de la sodalite.

(Traduit par la Rédaction)

Mots-clés: franzinite, groupe de la cancrinite, groupe de la sodalite, feldspathoïde, détermination de la structure, spectroscopie infra-rouge, séquences d'empilement.

INTRODUCTION

A broad structural investigation of the minerals pertaining to the cancrinite group has recently been undertaken. These minerals are trigonal and hexagonal feldspathoids whose framework is built by the stacking along [001] of layers made of six-membered rings of TO_4 tetrahedra (T : Si, Al). The resulting topologies show cages and open channels that are filled by interstitial cations and anions. Approximately fourteen phases pertaining to this group are known, with two to twenty-eight layers per unit cell, corresponding to c parameters of 5 to 74 Å. For a detailed discussion of the features of the various members of the cancrinite family, see for example Ballirano *et al.* (1996a).

Franzinite was found in ejectum collected near Pitigliano, a well-known mineral-collecting locality of Tuscany, Italy (Merlino & Mellini 1976, Merlino & Orlandi 1977). The ejectum, mainly composed of diopside and vesuvianite, is the result of a metasomatic process involving a trachytic magma and the host carbonate-rich metasedimentary rocks. From Weissenberg and precession photographs, the Laue symmetry $3m1$ was determined, as well as the cell parameters $a \approx 12.9$, $c \approx 26.6$ Å. The c cell parameter pointed to a 10-layer stacking sequence. All attempts to solve the structure were unsuccessful, as R indices were never better than 25%. A stacking sequence $ABCABCACB$ was proposed by Merlino (1976); he suggested that the C layer in the sixth position was disordered and replaced periodically by an A layer with a one-third statistical probability. Rinaldi & Wenk (1979) confirmed the occurrence of stacking faults by means of both electron diffraction and transmission electron microscopy. They observed the presence of diffuse and weak satellite reflections, considered to be indicative of an incommensurate superstructure. This superstructure was ascribed to distortions of the framework due to ordering of extra-framework anions and cations or to long-range Si, Al order.

Samples of franzinite were subsequently reported from other localities of the Roman perpotassic province (Leoni *et al.* 1979, Franceschini & Orlandi 1989, Ballirano *et al.* 1996b). Important differences in chemical composition have been reported among the samples, especially with respect to the CO_2 and H_2O content, as discussed by Ballirano *et al.* (1996b).

The aims of this work are to establish the correct sequence of the layers in franzinite, and to check whether perfect Si,Al order occurs in franzinite as the other phases of this family (*e.g.*, davyne, liottite, afghanite).

EXPERIMENTAL

A crystal of colorless and transparent franzinite (approximately $0.2 \times 0.2 \times 0.2$ mm) was selected from a sample labeled as MMUR (Mineralogical Museum of the University of Roma) 24340, and used to collect the X-ray data. The hand specimen is an ejectum mainly composed of diopside and tuscanyite, which was collected near Sacrofano, in Latium, Italy.

The chemical data (Ballirano *et al.* 1996b), obtained on a CAMECA SX-50 electron microprobe, point to the empirical formula $(Na_{20.5}K_{6.7}Ca_{11.7})(Si_{29.8}Al_{30.2})O_{119.7}(SO_4)_{10.0}Cl_{0.3}F_{0.1}$.

The infrared (IR) spectrum of franzinite (Ballirano *et al.* 1996b) shows only relatively weak absorption bands in the 4000–3000 cm^{-1} range, typical of the O–H stretching vibrations, in contrast with the presence of large amounts of H_2O reported in the literature data. In order to investigate with more accuracy the 4000–3000 cm^{-1} range, we performed new IR analyses. Data were collected on a Perkin Elmer FT-IR 2000, averaging 48 scans with a nominal resolution of 4 cm^{-1} , using the conventional KBr pellet technique. A KBr pellet with a franzinite:KBr ratio of 5:200 (instead of the commonly used 2:200 ratio) was prepared, with the aim to improve the intensity of the relatively weak absorption bands associated with the O–H stretching vibrations. The IR spectra are reported in Figure 1. The analyses were repeated after heating the pellet in an oven at 80°C in order to remove the humidity adsorbed by both KBr and sample: the broad band, initially located between 3700 and 3000 cm^{-1} , became narrower, and a peak appears at around 3590 cm^{-1} as well as a small hump centered at 3450 cm^{-1} . These features seem to indicate the presence of a disordered distribution of H_2O molecules. In fact, the IR spectrum of carbonate-dominant cancrinite (Ballirano *et al.* 1995) shows two well-defined peaks located at 3607 and 3539 cm^{-1} , which correspond, according to the plot of Nakamoto *et al.* (1955), to (Si,Al)O...OH₂ distances of 3.1 and 3.0 Å, in excellent agreement with the known refinements of the structure

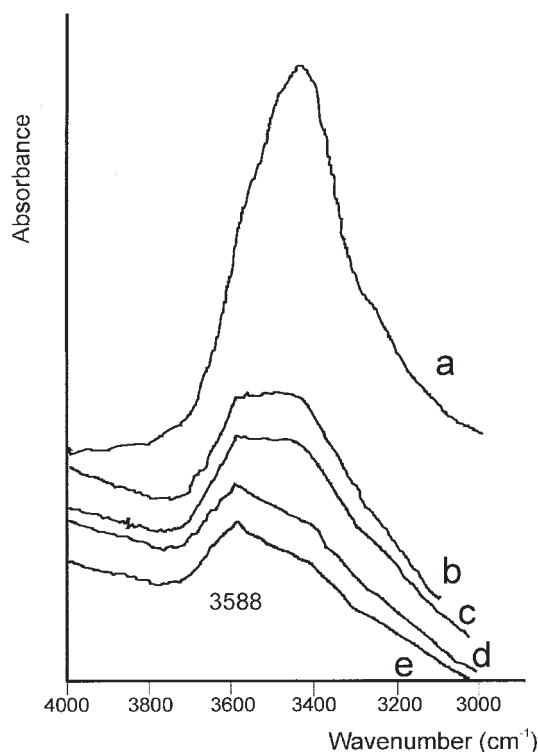


FIG. 1. IR spectra of franzinite MMUR 24340 (4000–3000 cm^{-1} range): a) immediately after the preparation of the pellet, b) after 1 h at 80°C, c) after 3 h at 80°C, d) after 4 days at 80°C, and e) after 10 days at 80°C. The pellet was stored in an oven in order to remove the adsorbed humidity.

of this mineral (Jarchow 1965, Smolin *et al.* 1981, Emiraliev & Yamzin 1982, Grundy & Hassan 1982).

Information on the X-ray single-crystal data collection are summarized in Table 1.

STRUCTURE DETERMINATION AND REFINEMENT

The stacking along [001] of layers made of six-membered rings of TO_4 tetrahedra may be discussed following the procedure first developed by Zhdanov (1945) for the close-packed stacking of layers of equal spheres. Sixteen different stacking sequences are possible for a ten-layer repeat (Patterson & Kasper 1959). Five sequences present $P\bar{3}m1$ symmetry: two of them, which may be denoted (8)(2) and (6)1(2)1 according to Zhdanov's notation (Patterson & Kasper 1959), present the center of inversion at the origin in the center of a hexagonal ring; the other three sequences, (7)1(1)1, (5)1(3)1 and (3)21(1)12, present the center of symmetry at the origin not in the center of a hexagonal ring, but in an "octahedral void" between successive layers (Patterson & Kasper 1959). Finally, two more sequences

TABLE 1. MISCELLANEOUS DATA FOR THE STRUCTURAL REFINEMENT OF FRANZINITE

Crystal dimensions:	0.2 x 0.2 x 0.2 mm
Instrument type:	Automatic 4-circle diffractometer Siemens P4 operating at 50 kV and 30 mA
Data collection:	ω scan
Radiation:	Graphite-monochromatized $\text{MoK}\alpha$ (0.71073 Å)
Cell parameters:	$a = 12.916(1)$ Å; $c = 26.543(3)$ Å based on 36 reflections ($15^\circ < 2\theta < 30^\circ$)
2θ max:	60°
Unique reflections:	4991
Reflections observed:	2841
$R_{\text{obs}} > 8\sigma(F_{\text{obs}})$	
Refined parameters:	482
Refinement program:	SHXLX1-93 (Sheldrick 1993).
Absorption correction:	Psi-scan empirical absorption correction
correction:	XEMP program (Siemens analytical X-ray Int. Inc.)
wR^2 (*)	0.1667
R (**)	0.0596

$$* wR^2 = \sqrt{\frac{\sum w(F_o^2 - F_c^2)^2}{\sum wF_o^4}}$$

$$(**) R = \frac{\sum |F_o| - |F_c|}{\sum F_o}$$

are possible, corresponding to $P\bar{3}m1$ symmetry, 4312 and 321211. We assumed, as indicated by the crystal structure of afghanite, the eight-layer member of the group, that the operator is in the center of a hexagonal ring, and we tested the two possible sequences, (8)(2) and (6)1(2)1: the comparison of the observed and calculated structure-factors clearly indicates that the correct sequence is (8)(2). The space-group symmetry $P\bar{3}m1$ does not allow an ordering of silicon and aluminum in the framework of tetrahedra. In keeping with the chemical composition of franzinite (Si:Al ratio of 1:1), and with the results of our recent refinements of the structures of liottite and afghanite (Ballirano *et al.* 1996a, 1997), we assumed an ordered distribution of the tetrahedrally coordinated cations in the space group $P321$, the maximum-order subgroup of $P\bar{3}m1$ allowing such distribution. The structure was completed by alternating difference Fourier maps and refinement cycles; the interpretation of the Fourier maps was made difficult by partial occupancies in the cationic sites, possible substitution involving Ca–Na–K, as well as the disordered distribution of sulfate groups. We have labeled as "A" the cationic sites in the center of the bases of the various cages, and "M" the cationic sites within the cages. During the first cycles of the refinement, strong maxima were identified in the difference-Fourier maps in positions expected for both S and Ca atoms, according to the model proposed by Ballirano *et al.* (1996b). However, a few of these Ca atoms are disordered over two or three sites (A3–A3a–A3b; A4–A4a; A6–A6a). Each set of split sites does not show a full occupancy, this fact being attributed to partial substitution of Ca with Na. The displacement parameters of the two sulfur atoms located inside the "losod" cage (S4 and S5) are

small and comparable; the sulfur atoms located inside the “sodalite” cages (S1, S2, and S3) have high and similar displacement-parameters. Following the procedure used in the structure refinement of liottite (Ballirano *et al.* 1996a) and afghanite (Ballirano *et al.* 1997), we displaced the S1, S2, and S3 sulfur atoms slightly from the three-fold axis. Further maxima were attributed to cations and to oxygen atoms pertaining to the sulfate groups. The occupancies of the oxygen sites were fixed as imposed by the correct geometry of the SO₄ tetrahedra, as explained below. The distribution of the cations is extremely complex, with multiple splitting of the cationic sites *M* (M1–M1a; M5–M5a–M5b–M5c; M6–M6a–M6b–M6c). The same anisotropic-displacement parameters have been refined for each set of split cation-sites. The corresponding occupancies were refined during the isotropic refinement and subsequently kept unchanged during the anisotropic refinement. Following the procedure described in the previous works on liottite and afghanite (Ballirano *et al.* 1996a, 1997), we refined the occupancies of the cation sites *M* and *A* in terms of the scattering power of Ca; successively, we calculated the actual cationic content as percentages of Na and (K,Ca) cations, assuming the full occupancy of each set of split sites.

The “cancrinite” cages show two relatively small maxima in electron density, *W* and *A2*, in positions expected for an anion and for a cation, respectively. According to the results of the chemical analysis and the IR data, it may be inferred that the *W* site is prevalently occupied by H₂O molecules, partially substituted by 0.3 atoms of Cl and 0.1 atoms of F. As regards the *A2* site, the very high displacement-parameter obtained using the scattering power of Ca indicates that most probably it is occupied by Na atoms, forming a fragment of the ...Na–H₂O...Na–H₂O... chain found in cancrinite, vishnevite and pitiglianoite.

The final *R* agreement index was 5.96% for 2841 *F*_{obs} > 8σ(*F*_{obs}) and 10.70 for all 4991 data; the *wR*² factor was 16.67% for 2841 *F*_{obs} > 8σ(*F*_{obs}) and 24.76% for all 4991 data. The number of refined parameters is 482 (Table 1). Positional parameters and equivalent displacement parameters are reported in Table 2; the bond distances between cations and oxygen atoms of the framework are listed in Table 3, whereas the bonds that involve the oxygen atoms of the sulfate groups within the different cages are reported under the corresponding figures. Table 4 shows the population of all the extra-framework sites, based on the refined electron-density and on crystal-chemical considerations. A table of structure factors is available at a nominal charge from the Depository of Unpublished Data, CISTI, National Research Council of Canada, Ottawa, Ontario K1A 0S2, Canada.

The franzzinite framework (Fig. 2) is composed of regular alternations of two “cancrinite” and two “sodalite” cages along [0 0 *z*], and two “sodalite” and one “losod” cages along [$\frac{1}{2}$ $\frac{1}{2}$ *z*] and [$\frac{1}{2}$ $\frac{1}{2}$ *z*] (Fig. 3).

TABLE 2. FRACTIONAL COORDINATES AND ISOTROPIC OR EQUIVALENT DISPLACEMENT PARAMETERS FOR FRANZZINITE

Atom	<i>x</i>	<i>y</i>	<i>z</i>	<i>U</i> _{iso} or <i>U</i> _{eq}
Si1	0.5834(3)	0.6684(3)	0.9018(2)	0.0158(9)
Al1	0.4190(3)	0.3332(3)	0.0978(2)	0.0135(9)
Al2	0.5866(4)	0.6658(3)	0.1963(2)	0.019(1)
Si2	0.4158(3)	0.3323(3)	0.8038(2)	0.0121(9)
Al3	0.5927(3)	0.6770(3)	0.6028(2)	0.0120(9)
Si3	0.4162(3)	0.3365(3)	0.3964(2)	0.0139(9)
Al4	0.7511(5)	0.7511(5)	0	0.018(2)
Si4	0.2472(5)	0.2472(5)	0	0.013(1)
Si5	0.7486(3)	0.7495(3)	0.2976(2)	0.0131(9)
Al5	0.2540(3)	0.2515(3)	0.7026(2)	0.015(1)
Si6	0.7485(4)	0.7485(4)	0.5	0.014(1)
Al6	0.2544(4)	0.2544(4)	0.5	0.012(1)
O1	0.4462(9)	0.218(1)	0.1051(3)	0.065(3)
O2	0.875(1)	0.7478(5)	0.2876(2)	0.028(1)
O3	0.2425(5)	0.1159(8)	0.5098(2)	0.025(1)
O4	0.322(1)	0.0005(9)	0.3468(5)	0.028(3)
O5	0.671(1)	0.002(1)	0.6482(5)	0.025(2)
O6	0.659(1)	0.6760(9)	0.5470(4)	0.022(2)
O7	0.353(1)	0.3302(9)	0.4504(4)	0.022(2)
O8	0.669(1)	0.650(1)	0.1495(6)	0.049(4)
O9	0.348(2)	0.351(1)	0.8495(7)	0.060(5)
O10	0.4501(5)	0.2340(8)	0.3901(2)	0.022(1)
O11	0.666(2)	0.686(2)	0.9481(6)	0.072(6)
O12	0.327(2)	0.307(2)	0.0468(7)	0.084(8)
O13	0.442(1)	0.549(1)	0.1995(2)	0.035(2)
O14	0.541(1)	0.0786(5)	0.3940(2)	0.026(1)
O15	0.656(1)	0.673(1)	0.2534(4)	0.027(2)
O16	0.344(1)	0.320(1)	0.7537(5)	0.053(4)
O17	0.888(1)	0.7821(5)	0.6994(2)	0.030(1)
O18	0.4275(7)	0.220(1)	0.8161(2)	0.037(2)
O19	0.9157(9)	0.454(1)	0.0863(3)	0.065(3)
O20	0.120(2)	0.233(1)	0.0106(3)	0.074(3)
S1	0.373(2)	0.676(4)	0.0546(3)	0.064(7)
S2	0.343(3)	0.706(1)	0.7415(3)	0.078(7)
S3	0.025(2)	0.023(2)	0.1561(4)	0.080(4)*
S4	1/3	2/3	0.3199(1)	0.0293(7)
S5	1/3	2/3	0.4884(1)	0.0241(6)
W	0.009(8)	0.039(4)	0.604(1)	0.11(1)*
A1	2/3	1/3	0.8189(1)	0.0370(7)
A2	0	0	0.490(3)	0.20(2)
A3	0	0	0.7218(5)	0.034(4)
A3a	0	0	0.7477(9)	0.034(4)
A3b	0	0	0.7009(9)	0.034(4)
A4	0	0	0.0318(8)	0.060(5)
A4a	0	0	0	0.060(5)
A5	2/3	1/3	0.3733(1)	0.0276(6)
A6	2/3	1/3	0.0635(4)	0.051(2)
A6a	2/3	1/3	0.1339(6)	0.051(2)
M1	0.503(1)	0.503(1)	0	0.047(2)
M1a	0.450(3)	0.544(3)	0.0150(5)	0.047(2)
M2	0.7815(3)	0.5649(2)	0.5960(1)	0.0289(4)
M3	0.4909(5)	0.5140(5)	0.3017(1)	0.0396(7)
M4	0.5015(4)	0.5015(4)	0.5	0.0284(6)
M5	0.849(1)	0.6967(9)	0.8973(3)	0.041(3)
M5a	0.806(2)	0.586(2)	0.9164(6)	0.041(3)
M5b	0.873(2)	0.760(2)	0.8870(6)	0.041(3)
M5c	0.829(4)	0.650(3)	0.903(1)	0.041(3)
M6	0.834(1)	0.6655(7)	0.2001(2)	0.034(2)
M6a	0.772(2)	0.556(2)	0.2169(5)	0.034(2)
M6b	0.873(2)	0.726(2)	0.1849(6)	0.034(2)
M6c	0.795(2)	0.605(2)	0.2139(7)	0.034(2)
OS1a	0.617(5)	0.32(2)	0.9054(8)	0.12(2)
OS1b	0.375(4)	0.603(4)	0.0207(6)	0.16(1)
OS1c	0.488(4)	0.736(7)	0.076(2)	0.16(3)
OS2a	0.401(2)	0.786(2)	0.7078(5)	0.10(7)
OS2b	0.730(5)	0.470(2)	0.225(1)	0.07(1)
OS2c	0.728(5)	0.399(5)	0.229(2)	0.12(3)
OS3a	0.046(5)	0.126(3)	0.131(1)	0.08(1)*
OS3b	0.113(3)	0.078(4)	0.184(1)	0.18(1)*
OS3c	0.045(5)	0.042(6)	0.113(1)	0.13(2)*
OS4a	1/3	2/3	0.2650(5)	0.093(6)
OS4b	0.396(2)	0.789(1)	0.3390(3)	0.053(2)
OS5a	1/3	2/3	0.5433(4)	0.058(4)
OS5b	0.397(1)	0.791(1)	0.4698(3)	0.042(2)

* Thermal parameters which were not refined anisotropically.

TABLE 3. SELECTED BOND-DISTANCES (Å) FOR THE OXYGEN ATOMS OF THE FRAMEWORK OF FRANZINITE

Si1	-O1	1.57(2) Å	Si2	-O9	1.58(1) Å	Si3	-O5	1.62(1) Å
	-O8	1.61(1) Å		-O13	1.59(2) Å		-O7	1.63(1) Å
	-O11	1.57(1) Å		-O16	1.58(1) Å		-O10	1.60(1) Å
	-O19	1.61(2) Å		-O18	1.56(1) Å		-O14	1.58(1) Å
Al1	-O1	1.71(2) Å	Al2	-O8	1.72(1) Å	Al3	-O4	1.73(1) Å
	-O9	1.71(1) Å		-O13	1.72(2) Å		-O6	1.72(1) Å
	-O12	1.72(1) Å		-O15	1.74(1) Å		-O10	1.74(1) Å
	-O19	1.68(2) Å		-O18	1.77(1) Å		-O14	1.76(1) Å
Si4	-O12	1.55(1) Å	Si5	-O2	1.67(1) Å	Si6	-O3	1.58(1) Å
	-O12	1.55(1) Å		-O4	1.59(1) Å		-O3	1.58(1) Å
	-O20	1.58(2) Å		-O15	1.61(1) Å		-O6	1.64(1) Å
	-O20	1.58(2) Å		-O17	1.63(1) Å		-O6	1.64(1) Å
Al4	-O11	1.70(1) Å	Al5	-O2	1.66(1) Å	Al6	-O3	1.74(1) Å
	-O11	1.70(1) Å		-O5	1.75(1) Å		-O3	1.74(1) Å
	-O20	1.69(2) Å		-O16	1.72(1) Å		-O7	1.75(1) Å
	-O20	1.69(2) Å		-O17	1.69(1) Å		-O7	1.75(1) Å
A1	-O13	2.57(1) Å x 3	A2	-O3	2.71(1) Å x 3			
	-O18	2.68(1) Å x 3		-O3	2.77(1) Å x 3			
A3	-O2	2.83(1) Å x 3	A3a	-O2	2.97(1) Å x 3	A3b	-O2	2.84(1) Å x 3
	-O17	2.51(1) Å x 3		-O17	2.75(1) Å x 3		-O17	2.44(1) Å x 3
A4	-O20	2.66(1) Å x 3	A4a	-O20	2.62(1) Å x 6	A5	-O10	2.47(1) Å x 3
	-O20	2.84(1) Å x 3					-O14	2.90(1) Å x 3
A6	-O1	2.70(1) Å x 3	A6a	-O1	2.58(1) Å x 3			
	-O19	2.85(1) Å x 3		-O19	3.06(1) Å x 3			
M1	-O11	2.64(2) Å x 2	M1a	-O11	2.82(4) Å			
	-O12	2.71(3) Å x 2		-O12	2.78(4) Å			
	-O19	2.48(1) Å x 2		-O12	2.96(4) Å			
				-O19	2.69(1) Å			
				-O19	2.75(2) Å			
M2	-O4	2.87(1) Å	M3	-O4	2.58(1) Å	M4	-O6	2.49(1) Å x 2
	-O5	2.79(1) Å		-O5	2.61(1) Å		-O7	2.46(1) Å x 2
	-O6	2.91(1) Å		-O13	2.88(1) Å		-O14	2.95(1) Å x 2
	-O7	2.97(1) Å		-O14	2.70(1) Å			
				-O15	2.46(1) Å			
				-O16	2.65(2) Å			
M5	-O8	2.71(2) Å	M5a	-O9	2.86(3) Å	M5b	-O8	2.69(3) Å
	-O9	2.65(2) Å		-O11	2.82(3) Å		-O9	2.84(3) Å
	-O11	2.66(2) Å		-O12	2.59(3) Å		-O11	2.86(3) Å
	-O12	2.69(3) Å		-O18	2.67(2) Å		-O18	2.83(2) Å
	-O18	2.57(1) Å		-O20	2.81(2) Å		-O20	2.72(2) Å
	-O20	2.57(1) Å						
M5c	-O8	2.81(5) Å	M6	-O1	2.82(1) Å	M6a	-O1	2.97(2) Å
	-O9	2.65(4) Å		-O2	2.50(1) Å		-O2	2.85(2) Å
	-O11	2.66(4) Å		-O8	2.44(2) Å		-O8	2.84(3) Å
	-O12	2.62(5) Å		-O9	2.59(2) Å		-O15	2.78(3) Å
	-O18	2.47(3) Å		-O15	2.74(2) Å		-O16	2.93(3) Å
	-O20	2.64(3) Å		-O16	2.68(2) Å			
M6b	-O1	2.87(2) Å	M6c	-O1	2.95(2) Å			
	-O2	2.74(2) Å		-O2	2.53(2) Å			
	-O8	2.49(2) Å		-O8	2.62(3) Å			
	-O9	2.44(3) Å		-O9	2.91(3) Å			
	-O16	2.83(2) Å		-O15	2.58(3) Å			
				-O16	2.75(3) Å			

The refined T -O bond distances confirm an ordered Si,Al distribution. In fact, the average Si-O bond distance is 1.60(3) Å, whereas the average Al-O bond dis-

tance is 1.72(3) Å with a $d_{\text{Si-O}}/d_{\text{Al-O}}$ ratio of 0.93 (the values in brackets represent the actual dispersion of the T -O bond distances from their averaged values).

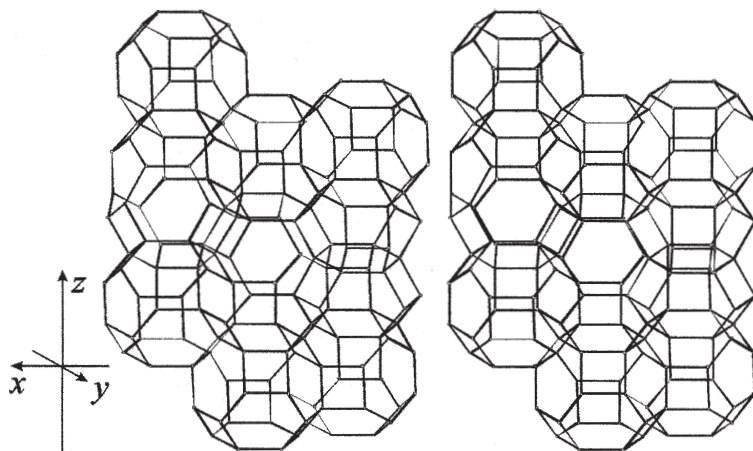


FIG. 2. Stereoplot of the framework of franzinite.

Because of the great complexity of the structure, we will describe separately the individual features of the various cages that constitute the franzinite structure.

TABLE 4. POPULATION OF THE EXTRAFRAMEWORK SITES (A AND M) IN FRANZINITE

Site	Refined occupanc y	% Na (*)	% K,Ca (*)	atoms Na	atoms (K,Ca)	atoms Ca (**)
A1	1	0	100	-	-	2.00
A2	0.50	100	0	1.00	-	-
A3	0.36	53	47	1.06	-	0.94
A3a	0.18					
A3b	0.22					
A4	0.27	89	11	0.89	-	0.11
A4a	0.07					
A5	1	0	100	-	-	2.00
A6	0.37	83	17	1.66	0.34	-
A6a	0.25					
M1	0.42	66	34	1.98	1.02	-
M1a	0.14					
M2	0.91	19	81	1.14	4.86	-
M3	0.66	76	24	4.56	-	1.44
M4	0.81	42	58	1.26	-	1.74
M5	0.29	72	28	4.32	-	1.68
M5a	0.12					
M5b	0.15					
M5c	0.11					
M6	0.33	60	40	3.60	2.40	-
M6a	0.15					
M6b	0.13					
M6c	0.12					
Σ_{atoms}				21.47	8.62	9.91

(*) These values were calculated assuming that, actually, each group of sites is fully occupied. For example, the total occupancy of the sites A3, A3a, A3b, refined on the basis of the scattering power of Ca, is 0.76. A full occupancy would require the presence of only 15.2 electrons, which corresponds to 53% Na and 47% Ca (we include also K, which has almost the same scattering power of Ca).

(**) In some cases it was possible to discriminate between Ca and (K, Ca) cations, on the basis of the observed distances with the oxygen atoms.

"LOSOD" CAGES

The two symmetry-equivalent "losod" (17-hedra) cages are located along $[\frac{2}{3} \frac{1}{3} z]$ and $[\frac{1}{3} \frac{2}{3} z]$, respectively (Fig. 4). Each cage hosts two sulfate groups (S4 and S5), which are separated by a triplet of symmetry-equivalent *M2* cations. The *M2* cations mostly consist of K,Ca (81% K,Ca; 19% Na). Two further triplets of cations *M3* and *M4* are located, respectively, around the S4 and S5 sulfate groups. *M3* is dominantly occupied by Na atoms (76% Na; 24% K,Ca), whereas the *M4* site has a 58% K,Ca and a 42% Na occupancy. On the basis of the cation-oxygen distances, we may assume a distribution of cations similar to that of davynite (Bonaccorsi *et al.* 1990), with the "internal" site *M2* mainly occupied by K and the "external" sites *M3* and *M4* mainly occupied by Na and Ca atoms (Table 4).

The A1 and A5 cations occupy the bases of the cage and, according to the refined distribution of electron density and the cation-oxygen distances, they are calcium atoms. The two sulfur atoms are located on the triad axis and are bonded to two triplets of oxygen atoms, OS4b and OS5b, respectively, and to two apical atoms of oxygen, OS4a and OS5a, respectively. The bond distances are very similar and range from 1.46(1) to 1.47(1) Å. This type of sulfate coordination is allowed by the position of the calcium cations A1 and A5, which occur at favorable distances with the corresponding apical atoms of oxygen of the two sulfate groups [A1-OS4a = 2.23(1) Å; A5-OS5a = 2.21(1) Å]. There are some differences with respect to the distribution of cations and anions of the "losod" cage of liottite. In fact, the sulfate groups of that cage are disordered in liottite, because one of the calcium atoms located near the bases of the "losod" cage is distributed between two neighboring sites. Moreover, there are differences also with respect to the distribution of cations.

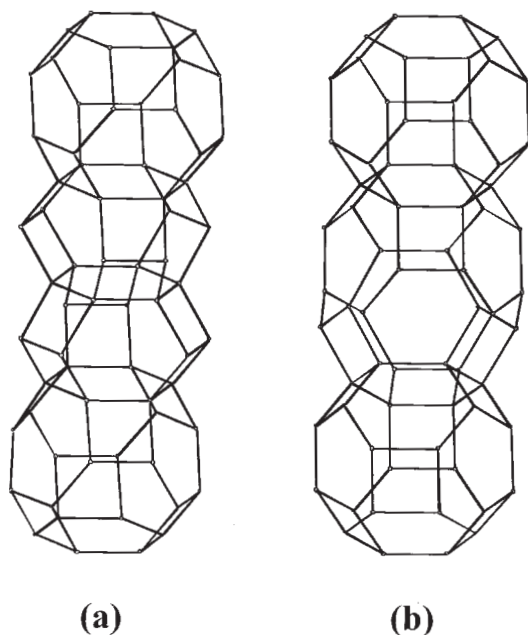


FIG. 3. (a) Superposition of the two "cancrinite" and two "sodalite" cages along $[0\ 0\ z]$; (b) superposition of two "sodalite" and one "losod" cages along $[\frac{2}{3}\ \frac{1}{3}\ z]$ and $[\frac{1}{3}\ \frac{2}{3}\ z]$.

"SODALITE" CAGES

The "sodalite" (cubo-octahedra) cages contain one sulfate group, which is strongly disordered; according to the extremely high U_{11} and U_{22} displacement parameters, the sulfur atoms seem to be displaced from the three-fold axis. Generally speaking, the cations are distributed among four different horizontal planes (Fig. 5): (a) and (b) planes are located in correspondence of the two bases of the cage; (c) and (d) are planes that bisect the six-membered rings of the lateral walls of the cage. Furthermore, each one of these planes presents multiple splitting of sites. Because of the complexity of the distribution of anions and cations, each one of the three independent "sodalite" cages will be described in turn. The identification number of the various cages refers to the label of the corresponding sulfur atom contained in the "sodalite" cage (example: "sodalite" cage 1 refers to the cage that contains S1).

"Sodalite" cage 1

The "sodalite" cage 1 shares its bases with a "losod" cage and with the "sodalite" cage 2. Three partially occupied triplets of oxygen atoms (OS1a, OS1b, and OS1c, respectively) set up the coordination of the sulfate group (in Fig. 6a, only one of the three possible orientations of the sulfate tetrahedron is shown). The

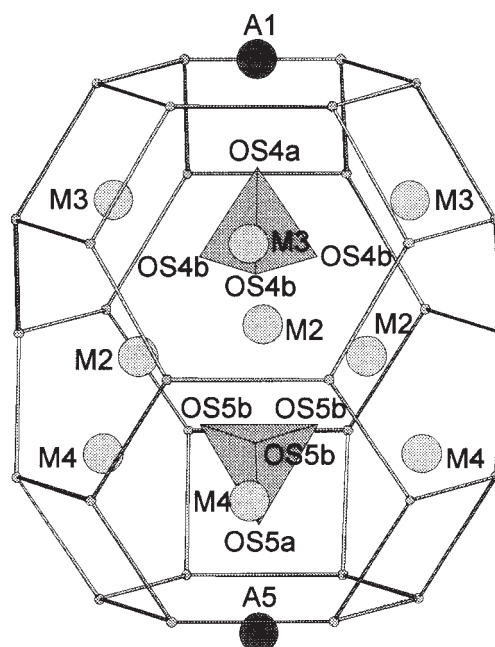


FIG. 4. The "losod" cage and its content. The relevant bond-distances among the extra-framework cations and anions also are reported.

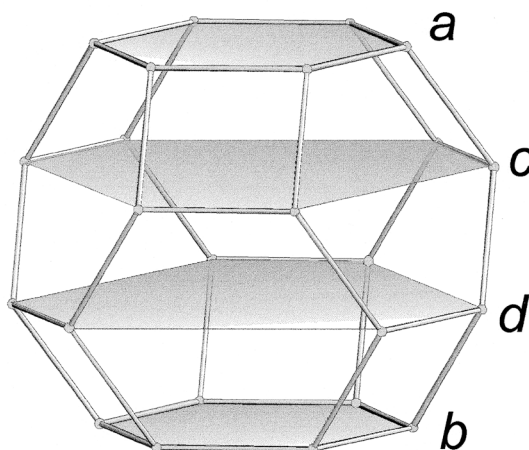


FIG. 5. Within the "sodalite" cages, the cationic sites are located on four different levels, indicated with the letters a, b, c, and d.

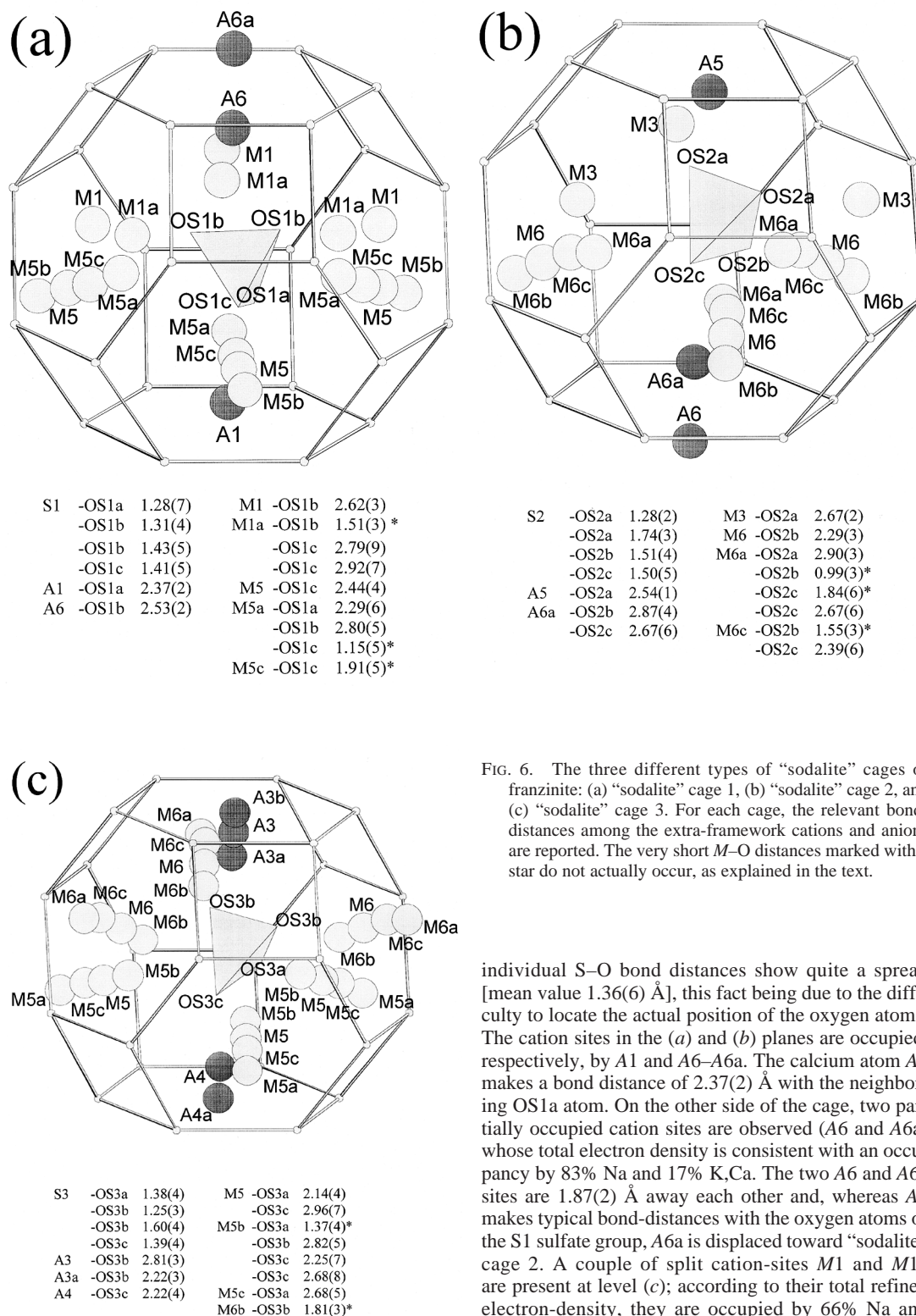


FIG. 6. The three different types of “sodalite” cages of franzinite: (a) “sodalite” cage 1, (b) “sodalite” cage 2, and (c) “sodalite” cage 3. For each cage, the relevant bond-distances among the extra-framework cations and anions are reported. The very short $M-O$ distances marked with a star do not actually occur, as explained in the text.

individual S–O bond distances show quite a spread [mean value 1.36(6) Å], this fact being due to the difficulty to locate the actual position of the oxygen atoms. The cation sites in the (a) and (b) planes are occupied, respectively, by A1 and A6–A6a. The calcium atom A1 makes a bond distance of 2.37(2) Å with the neighboring OS1a atom. On the other side of the cage, two partially occupied cation sites are observed (A6 and A6a) whose total electron density is consistent with an occupancy by 83% Na and 17% K,Ca. The two A6 and A6a sites are 1.87(2) Å away each other and, whereas A6 makes typical bond-distances with the oxygen atoms of the S1 sulfate group, A6a is displaced toward “sodalite” cage 2. A couple of split cation-sites M1 and M1a are present at level (c); according to their total refined electron-density, they are occupied by 66% Na and

34% K,Ca. Four further sites *M5*, *M5a*, *M5b*, and *M5c* are located on the plane (*d*). Sites *M5*, *M5a*, and *M5c* are within the cage 1, whereas *M5b* is displaced toward the neighboring "sodalite" cage 3. The refined electron-density and the observed distances from the oxygen atoms are compatible with an occupancy of 72% Na and 28% Ca for these sites (Table 4).

As a general remark, some of the bond distances between cation sites and oxygen atoms of the sulfate groups are unacceptably short, which means that a simultaneous occupancy of the cation and oxygen sites does not actually occur (Figs. 6a, b, c).

"Sodalite" cage 2

This cage shares its bases with a "losod" cage and with the "sodalite" cage 1. Three triplets of oxygen atoms (OS2a, OS2b, and OS2c, respectively) occur within the cage, each one with partial occupancy. One of the three resulting symmetry-related SO₄ tetrahedra is shown in Figure 6b. As in "sodalite" cage 1, the S–O bond distances show a broad range. On the (*a*) and (*b*) planes there are, respectively, the A5 calcium cation and the A6–A6a split sites already described in descriptions of "sodalite" cage 1.

The triplet of *M3* cations is located exactly at the center of three hexagonal rings of TO₄ tetrahedra (shared with the "losod" cage) in correspondence of the (*c*) plane. As previously indicated, the chemical content of the *M3* site, on the basis of the refined electron-density and of the *M3*–O distances, is 76% Na and 24% Ca.

Four further split sites are present on plane (*d*), *M6*, *M6a*, *M6b*, and *M6c*, which show a 60% Na and a 40% K,Ca occupancy. *M6* atoms are located at the center of the hexagonal rings, whereas *M6a* and *M6c* are displaced toward the center of "sodalite" cage 2, and *M6b*, toward the neighboring "sodalite" cage 3.

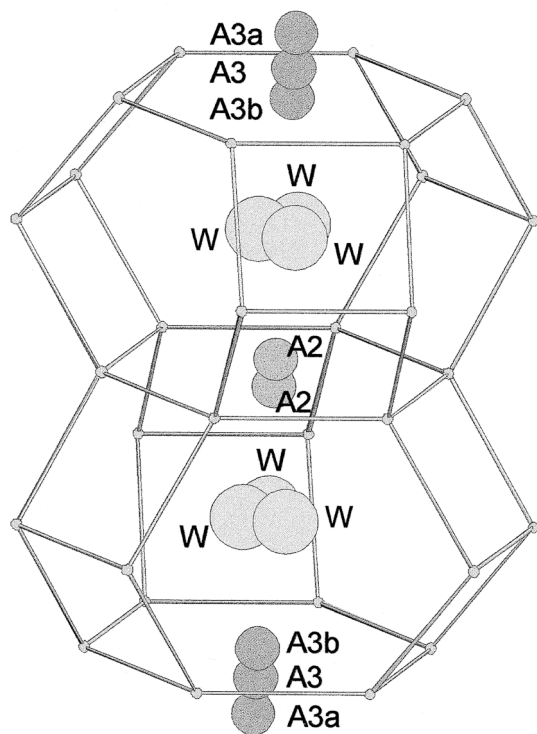
"Sodalite" cage 3

The two symmetry-related "sodalite" cages 3 share one base, the other base being shared with a superimposed "cancrinite" cage. The SO₄ polyhedron is distorted, and is placed off-axis (Fig. 6c). On the two bases of the cage we find, respectively, a triplet (A3, A3a, and A3b) and a pair (A4 and A4a) of electron-density maxima. The sites A3, A3a, and A3b contain 53% Na and 47% K,Ca, whereas the A4 and A4a sites are mostly occupied by Na (89% Na and 11% K,Ca). In both sites, the occurrence of short cation–oxygen distances [for example, A3a–OS3b at 2.22(3) Å and A4–OS3c at 2.22(4) Å] suggests that the sites cannot be occupied by potassium cations, and that their actual chemical content is 53%Na, 47%Ca and 89%Na, 11%Ca, respectively. The four *M5* sites (shared with the "sodalite" cage 1) and the four *M6* sites (shared with the "sodalite" cage 2) occupy the planes (*c*) and (*d*).

During the isotropic refinement, we made an attempt to refine the occupancy of the oxygen sites of the sulfate groups in the various "sodalite" cages. The occupancies of the oxygen sites of sulfate groups S1 and S2 are consistent with the required geometry of the polyhedra, whereas sulfate group S3 did not show an adequate value of the electron density compatible with the presence of four oxygen atoms around the corresponding sulfur atom (only 3/4 of the oxygen atoms seem to be present). On the other side, the sulfur site is fully occupied (actually S3 is slightly displaced from the special position at 0, 0, *z*, and statistically occupies one of the three symmetry-related positions). Three possible interpretations may be proposed: a) an error in the calculated electron-density may be due to the correlation existing between the thermal parameter and the corresponding occupancy, b) a partial replacement of (SO₄)^{2−} by Cl[−] occurs, as in sodalite; this hypothesis seems to be in contradiction with the chemical data, which do not reveal any deficiency in sulfur content, or c) a partial replacement of (SO₄)^{2−} by S^{2−} occurs, as in lazurite, a member of the related sodalite-group minerals (Hassan *et al.* 1985).

No strong support of any one of these hypotheses has been found; therefore, we preferred to continue the refinement considering the "sodalite" cage 3 as fully occupied by one (SO₄)^{2−} group, despite the real or apparent deficiency in oxygen.

The "sodalite" cages of lazurite (Hassan *et al.* 1985) and nosean (Hassan & Grundy 1989) show features similar to those just described for the "sodalite" cage of franzinite. The cations are distributed among the four (*a*), (*b*), (*c*), and (*d*) planes, and they show multiple splitting of sites; however, the occupancy of one of the split sites is incompatible with the presence inside the cage of one (SO₄)^{2−} group. This fact has been explained as due to either S^{2−} or H₂O *versus* (SO₄)^{2−} partial substitutions. The corresponding frameworks are distorted because the different content of anions is reflected in the relative dimensions of the cages. In the case of nosean (Hassan & Grundy 1989), the mineral has an average structure that results from the presence of two distinct structures, respectively with a cubic cell edge of 8.9 Å [nosean containing H₂O only], and a cubic cell edge of 9.2 Å [nosean containing only (SO₄)^{2−}]. The same situation is found in lazurite (Hassan *et al.* 1985), owing to the occurrence of S^{2−}-bearing and (SO₄)^{2−}-bearing cages. As a matter of fact, "sodalite" cage 3 has an extremely disordered distribution (A3–A3a–A3b; A4–A4a; *M5*–*M5a*–*M5b*–*M5c*; *M6*–*M6a*–*M6b*–*M6c*), which is coupled with an apparent deficiency of oxygen atoms of the sulfate group. On the basis of the preceding discussion, this situation seems to be compatible with a partial S^{2−} ⇌ (SO₄)^{2−} substitution.



W	-A2	2.53(7)	W	-A2	3.07(7)
	-A3	3.15(3)		-A3b	2.61(3)

FIG. 7. The two superimposed "cancrinite" cages and their content. Bond distances between the H₂O molecules W and the neighboring cations also are reported.

CANCRINITE CAGES

The two symmetry-equivalent "cancrinite" cages share a common base, the other base being shared with the "sodalite" cage 3. Each "cancrinite" cage contains a molecule of H₂O that is statistically disordered over three symmetry-related positions (Fig. 7). The Na atom is not exactly located at the center of the common basis, but is slightly displaced (A2). This displacement from the special position 0, 0, ½ leads to a disordered distribution of the Na atom into two equivalent half-occupied positions at $z \approx 0.49$ and $z \approx 0.51$. Owing to the very short distance between the two positions (~ 0.5 Å), only one of them is actually occupied. The H₂O molecule makes two different bond-distances with A2 [2.53(7) and 3.07(7) Å], depending on which one of the two positions is occupied.

Three different electron-density maxima were observed near the center of the base shared with the "sodalite" cage 3 (A3, A3a, and A3b), each one with partial occupancy. As previously pointed out, the sites

A3, A3a, and A3b contain 53% Na and 47% K,Ca. The H₂O molecule makes two different bond-distances with A3b [2.61(4) Å] and A3 [3.15(4) Å], whereas A3a is displaced toward the neighboring "sodalite" cage 3.

The structure refinements of carbonate-dominant cancrinite (Jarchow 1965, Smolin *et al.* 1981, Emiraliev & Yamzin 1982, Grundy & Hassan 1982), Ge-substituted cancrinite (Belokoneva *et al.* 1986), basic (*i.e.*, hydroxyl-dominant) cancrinite (Bresciani Pahor *et al.* 1982, Hassan & Grundy 1991), vishnevite (Hassan & Grundy 1984, Pushcharovskii *et al.* 1989), pitiglianoite (Merlino *et al.* 1991), bystrite (Pobedinskaya *et al.* 1991a), and afghanite (Pobedinskaya *et al.* 1991b) have shown that the H₂O molecule of the "cancrinite" cage has one short distance (2.4 Å) and one long distance (2.9 Å) to the cations. Accordingly, we propose that the short W-A2 bond distance [2.53(7) Å] is coupled with the occupancy of the A3 site, leading to a long [3.15(4) Å] W-A3 bond distance. Conversely, the long W-A2 bond distance [3.07(7) Å] may be coupled with the occupancy of the A3b site [W-A3b bond distance of 2.61(4) Å]. The explanation of the occurrence of the A3a site is more difficult. If we consider the partial $S^{2-} \leftrightarrow (SO_4)^{2-}$ substitution into the "sodalite" cage 3 as a possibility, the occurrence of the A3a site could be related with the presence of S^{2-} . In any case, A3a makes an acceptable bond-distance with the oxygen OS3b of the sulfate group of the "sodalite" cage 3.

DISCUSSION

The structure of franzinite has been refined to an R value of 5.96% in the space group $P321$. The stacking sequence is ABCABACABC. The space group $P321$ allows an ordered distribution of Si and Al inside the framework, a pattern of order that is confirmed by the refined Si-O [1.60(3) Å] and Al-O [1.72(3) Å] bond distances. The result is in keeping with the Si:Al ratio, equal to 1, as derived from the electron-microprobe data. According to the recent refinement of the structure of liottite (Ballirano *et al.* 1996a) and afghanite (Ballirano *et al.* 1997), which correspond to the 6- and 8-layer members, respectively, of the family of cancrinite-like minerals, an ordered Si,Al distribution may be expected for the phases showing a 1:1 Si,Al compositional ratio. So far, only cancrisilite (Khomyakov *et al.* 1991) shows a significant deviation from the ideal 1:1 ratio. Three different types of cages are found in the franzinite structure: "cancrinite" cages, "sodalite" cages, and "losod" cages. Two "cancrinite" cages and two "sodalite" cages are stacked along [0 0 z], whereas a "losod" cage and two "sodalite" cages repeat along [$\frac{2}{3}$ $\frac{1}{3}$ z] and [$\frac{1}{3}$ $\frac{2}{3}$ z]. The "cancrinite" cage hosts a molecule of H₂O that is displaced from the three-fold axis and disordered over three symmetry-related positions, forming a Na-H₂O...Na-H₂O sequence, similar to that found in many cancrinite-like minerals. Three distinct "sodalite" cages have been observed, differing in the extremely

complex pattern of distribution of cations. The sulfate groups hosted by the cages are displaced off-axis. In the case of the "sodalite" cage 3, the geometry of the sulfate group does not seem complete because of an apparent lack of oxygen atoms. This fact has been tentatively attributed to a partial $S^{2-} \leftrightarrow (SO_4)^{2-}$ substitution, as found in the related mineral lazurite. The "losod" cage contains two sulfate groups that are perfectly superimposed along z ; it features a relatively simple distribution of cations, similar to that observed in the "losod" cage of liottite. From the structural refinement, a formula of $[Na_{21.5}(Ca,K)_{8.6}Ca_{9.9}[Si_{30}Al_{30}O_{120}](SO_4)_{10} \cdot 2H_2O$ was obtained (Table 4), which is in agreement with $[Na_{20.5}K_{6.7}Ca_{11.7}[Si_{29.8}Al_{30.2}O_{119.7}](SO_4)_{10.0}Cl_{0.3}F_{0.1} \cdot nH_2O$ derived from the results of the electron microprobe-analysis and IR spectroscopy. The ideal formula may be written as $[(Na,K)_{30}Ca_{10}][Si_{30}Al_{30}O_{120}](SO_4)_{10} \cdot 2H_2O$. In order to obey electroneutrality, 50 positive charges have to be supplied by the cations, and this value may only be obtained with the presence of ten calcium atoms; every calcium in excess may be balanced by a non-stoichiometric proportion of the cations or a partial substitution of the H_2O molecules with Cl^- and F^- anions. The sample we studied contains 11.7 calcium atoms, and a nonstoichiometric proportion of the cations has been detected by means of the electron-microprobe analysis, leading to a total of 39 cations instead of the maximum allowable number of 40; moreover, small amounts of Cl^- and F^- anions have been documented by electron-microprobe analysis.

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