

CLINOMETABORITE, NATURAL β -METABORIC ACID, FROM LA FOSSA CRATER, VULCANO, AEOLIAN ISLANDS, ITALY

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ABSTRACT

Clinometaborite, β -HBO₂, the monoclinic modification of metaboric acid, which has long been known as a synthetic product only, was found as a sublimate in an active medium-temperature intracrater fumarole at La Fossa crater, Vulcano Island, Aeolian archipelago, Sicily, Italy. It forms aggregates of stout crystals up to 2 mm long in an altered pyroclastic breccia, associated with metaborite, sassolite and adranosite. The unaltered mineral is colorless and translucent, with a vitreous luster; it becomes chalky white after some months of exposure to open air. The crystal habit is prismatic; twinning was not observed. The strongest six lines in the X-ray powder-diffraction pattern [d_{obs} in Å(hkl)] are: 3.078(100)201, 4.193(20)111, 6.773(15)001, 2.550(10) $\bar{1}$ 22, 3.224(8) $\bar{2}$ 01, and 2.702(8)031. The unit cell is monoclinic, space group $P2_1/a$, with a 7.127(2), b 8.842(3), c 6.773(2) Å, β 93.21(1)°, V 426.1(2) Å³, Z = 12, values that virtually coincide with those of the synthetic compound. The calculated density is 2.049 g/cm³. The structure was refined to a final R of 0.043 for 977 observed reflections [$I > 2\sigma(I)$]. Both the mineral and the mineral name have been approved by the IMA Commission on New Minerals, Nomenclature and Classification (IMA 2010–022).

Keywords: clinometaborite, β -metaboric acid, new mineral species, monotropism, Vulcano Island, Aeolian Islands, Italy.

SOMMAIRE

Nous avons découvert la clinométaborite, β -HBO₂, modification monoclinique de l'acide métaborique connu depuis longtemps comme produit de synthèse, dans une fumarolle active intercratérique de température modérée au cratère La Fossa, île de Vulcano, archipel éolien, en Sicile, Italie. Les cristaux trappus, en agrégats d'une longueur de 2 mm dans une brèche pyroclastique altérée, sont associés à la métaborite, la sassolite et l'adranosite. Le minéral non altéré est incolore et translucide, avec un éclat vitreux; il devient blanc crayeux après quelques mois exposé à l'air. Les cristaux sont prismatiques; nous n'avons pas décelé de macle. Les six raies les plus intenses du spectre de diffraction X, méthode des poudres [d_{obs} en Å(hkl)] sont: 3.078(100)201, 4.193(20)111, 6.773(15)001, 2.550(10) $\bar{1}$ 22, 3.224(8) $\bar{2}$ 01, et 2.702(8)031. La maille élémentaire est monoclinique, groupe spatial $P2_1/a$, avec a 7.127(2), b 8.842(3), c 6.773(2) Å, β 93.21(1)°, V 426.1(2) Å³, Z = 12, valeurs qui coïncident pratiquement avec celles du composé synthétique. La densité calculée est 2.049 g/cm³. Nous en avons affiné la structure jusqu'à un résidu final R de 0.043 en utilisant 977 réflexions observées [$I > 2\sigma(I)$]. Le minéral et son nom ont reçu l'approbation de l'Association internationale de Minéralogie, Commission des nouveaux Minéraux, Nomenclature et Classification (IMA 2010–022).

(Traduit par la Rédaction)

Mots-clés: clinométaborite, acide métaborique β , nouvelle espèce minérale, monotropisme, île de Vulcano, archipel éolien, Italie.

INTRODUCTION

In the last few years, several new mineral species have been discovered at Vulcano, Aeolian Islands, Italy, as a result of a systematic investigation on the fumarolic products at the site (see Campostrini *et al.* 2011, Demartin *et al.* 2008a, 2008b, 2008c, 2009a, 2009b, 2009c, 2010a, 2010b, 2010c, 2010d, and references

therein). Here we report the discovery at La Fossa crater of an additional new mineral species: clinometaborite, β -HBO₂. It is the natural equivalent of β -metaboric acid or HBO₂ II, a substance already known as a synthetic product, and whose crystal structure has already been determined (Kracek *et al.* 1938, Zachariassen 1963, Freyhardt *et al.* 2000). The mineral is named after its composition and symmetry. Both the mineral and the

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mineral name have been approved by the IMA Commission on New Minerals, Nomenclature and Classification (IMA 2010–022). The holotype is deposited in the Reference Collection of Dipartimento di Chimica Strutturale e Stereochimica Inorganica of Università degli Studi di Milano (catalogue no. 2010–03).

OCCURRENCE AND PROPERTIES

Clinometaborite occurs in a medium-temperature ($\sim 250^\circ\text{C}$) intracrater fumarole developed on altered pyroclastic breccia, as stout crystals up to 2 mm long. It is accompanied by sassolite, adranosite, $(\text{NH}_4)_4\text{NaAl}_2(\text{SO}_4)_4\text{Cl}(\text{OH})_2$ (Demartin *et al.* 2010c), and metaborite; the last species is usually known to form good transparent rhombododecahedra in saline deposits (Lobanova & Avrova 1964), but its presence has never been noted in volcanic fumaroles.

The crystal habit is prismatic, but the faces are too rough for an assignment of indices (Fig. 1). Microscopic observations reveal that these “crystals” consist of aggregates of fibrous individuals and are most likely pseudomorphous. The unaltered material is transparent and colorless, the streak is white, and the luster is vitreous; it becomes chalk-white upon a few months’ exposure to open air and is transformed into orthoboric acid. The mineral is not fluorescent either under short-wave or long-wave ultraviolet radiation. It is brittle, and a perfect cleavage can be observed. The calculated density is 2.049 g/cm^3 . The optical properties of the synthetic compound reported by Kracek *et al.* (1938) are: biaxial (–) with α 1.434, β 1.570, γ 1.588 (589 nm).

The infrared spectrum, recorded on a Jasco IRT–3000 spectrometer, shows absorption bands at 1194 and 1457 cm^{-1} typical of the E’ vibration modes of the BO_2^- ion (Hisatsune & Haddock Suarez 1963).

Analyses with an electron microprobe (EDS mode) indicate only the presence of oxygen, a trace amount of Na due to contaminant material, and no elements with atomic number greater than 11. No attempt to analyze the mineral with WDS was made as its composition was unequivocally established by crystal-structure refinement.

X-RAY DATA AND REFINEMENT OF THE STRUCTURE

X-ray powder-diffraction data have been obtained at room temperature using a Philips PW1830 diffractometer, with $\text{CuK}\alpha$ radiation (Table 1). A least-squares fit of these data provided the following unit-cell parameters: monoclinic system, space group $P2_1/a$, with a 7.1243(6), b 8.8468(10), c 6.7699(7) Å, β 93.23(1)°, V 426.00(6) Å³, Z = 12. These data are close to those of the corresponding synthetic compound (β -metaboric acid or HBO_2 II) reported by other authors (Table 2).

The crystal structure of the synthetic counterpart was determined and refined from single-crystal data by Zachariassen (1963) and Freyhardt *et al.* (2000).

For our structure refinement of the natural material, single-crystal diffraction data were collected at room temperature by means of ω scans on a crystal fragment of about $0.20 \times 0.03 \times 0.02$ mm, with $\text{MoK}\alpha$ radiation ($\lambda = 0.71073$ Å), using a BRUKER Apex II diffractometer equipped with a 2K CCD detector. A

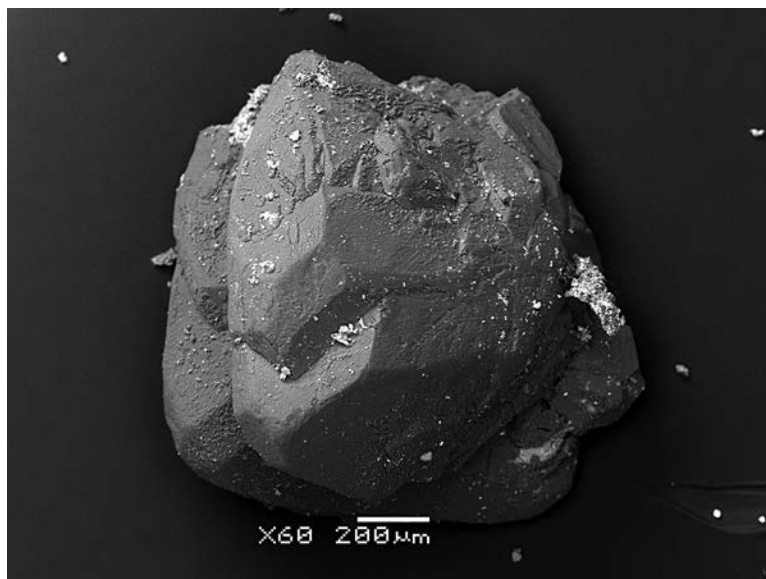


FIG. 1. Scanning electron microscope (SEM) image (BSE) of clinometaborite.

one-minute exposure and 0.5° frame width were used. A total of 3218 reflections were measured, corresponding to a complete scan of the reciprocal lattice up to $2\theta_{\max} = 63.54^\circ$. Of these, 1372 were found to be unique ($R_{\text{int}} = 0.0256$). The unit-cell dimensions obtained from the single crystal were refined by least-squares from

TABLE 1. X-RAY POWDER-DIFFRACTION DATA FOR CLINOMETABORITE AND A COMPARISON WITH ITS SYNTHETIC COUNTERPART

clinometaborite				synthetic**		h	k	l
d_{obs} (Å)	d_{calc} (Å)*	<i>l</i> / <i>o</i>	<i>F</i> / <i>W</i> <i>M</i>	d_{calc} (Å)	<i>l</i> / <i>o</i>			
6.773	6.759	15	0.16	6.753	33	0	0	1
				5.542	1	1	1	0
5.378	5.371	1	0.16	5.368	7	0	1	1
4.371	4.382	2	0.14	4.380	22	1	1	1
4.193	4.197	20	0.17	4.195	100	1	1	1
3.557	3.556	5	0.20	3.556	5	2	0	0
3.318	3.255	4	0.20	3.324	5	1	2	1
3.296	3.300	3	0.20	3.297	9	2	1	0
3.224	3.223	8	0.14	3.222	10	2	0	1
				3.157	3	0	1	2
3.078	3.077	100	0.18	3.074	100	2	0	1
				3.027	2	1	1	1
2.947	2.944	3	0.20	2.944	1	1	1	2
2.836	2.831	4	0.20	2.831	2	1	1	2
				2.723	7	1	3	0
2.702	2.703	8	0.15	2.701	10	0	3	1
2.685	2.685	3	0.15	2.685	10	0	2	2
2.550	2.546	10	0.14	2.550	11	1	2	2
2.518	2.542	7	0.15	2.545	9	1	3	1
				2.524	12	2	2	1
2.422	2.425	3	0.13	2.425	6	2	1	2
2.385	2.384	1	0.13	2.383	1	2	0	2
				2.301	1	2	1	2
				2.290	2	3	1	0
				2.211	10	0	4	0
				2.206	8	1	1	1
2.189	2.191	3	0.17	2.190	17	2	2	2
				2.144	3	1	3	2
				2.128	3	2	3	1
2.112	2.111	2	0.16	2.111	2	1	4	0
2.025	2.026	2	0.20	2.025	2	1	4	1
2.007	2.008	4	0.20	2.008	15	0	2	3
1.958	1.958	3	0.18	1.958	6	1	2	3
				1.916	1	2	3	2
1.907	1.907	2	0.20	1.907	3	1	2	3
1.849	1.850	5	0.20	1.849	10	3	1	2
1.819	1.818	1	0.16	1.818	1	3	2	2
				1.802	4	3	3	1
				1.795	4	2	4	1
				1.791	3	0	3	3
1.778	1.778	2	0.18	1.778	2	4	0	0
				1.762	1	3	3	1
1.719	1.718	2	0.20	1.717	2	1	3	3
				1.711	3	0	5	1
1.697	1.696	3	0.25	1.696	2	4	0	1
				1.665	1	4	1	1
1.659	1.659	2	0.16	1.658	3	1	5	1
1.652	1.651	2	0.20	1.652	3	1	3	1
				1.622	1	4	2	1
				1.617	1	3	4	0
1.612	1.612	1	0.20	1.612	1	4	0	2
1.596	1.597	1	0.20	1.597	1	1	1	4
1.583	1.584	1	0.20	1.583	1	4	2	1

* Calculated from the unit cell obtained from least-squares refinement of the data in the first column using the program UNITCELL (Holland & Redfern 1997). E.s.d. on $2\theta = 0.030^\circ$, average $\Delta 2\theta = 0.025^\circ$.

** PDF entry # 00-022-1109 (calculated pattern).

1425 reflections with $I > 5\sigma(I)$; results are given in Table 3, together with other details concerning the data collection and refinement. The intensity data were reduced using the Bruker program SAINT, and corrected for Lorentz, polarization, and background. No violations of the systematic absences of the space group $P2_1/a$ were observed. An absorption correction was applied (maximum and minimum transmission factors 1.000 and 0.691, respectively) using the SADABS program (Sheldrick 2000). The structure was refined starting from the atom coordinates obtained by Freyhardt *et al.* (2000) and transformed into the $P2_1/a$ setting using the SHELXL97 program (Sheldrick 2008) implemented in the WINGX suite (Farrugia 1999). The final coordinates and displacement parameters of the atoms are reported in Table 4. The final R is 0.043 for 977 unique observed reflections [$I > 2\sigma(I)$]. A list of structure factors and a cif file are available from the Depository of Unpublished Data, on the Mineralogical Association of Canada website [document Chinometaborite CM49_1273].

RESULTS AND DISCUSSION

Metaboric acid exists in three crystalline phases, which are prepared by dehydration of orthoboric acid and subsequent heat treatment (Wells 1984, Cotton &

TABLE 2. CLINOMETABORITE: UNIT-CELL DIMENSIONS

	<i>a</i> (Å)	<i>b</i> (Å)	<i>c</i> (Å)	β (°)	<i>V</i> (Å ³)
Our data (single crystal)	7.127(2)	8.842(3)	6.773(2)	93.21(1)	426.1(2)
Our data (powder diffraction)	7.1243(6)	8.8468(10)	6.7699(7)	93.23(1)	426.00(6)
Zachariassen (1963)	7.122(2)	8.842(2)	6.771(2)	93.26(1)	425.70(2)
Freyhardt <i>et al.</i> (2000)	7.075(2)	8.844(1)	6.758(1)	93.50(1)	422.07(2)

TABLE 3. CLINOMETABORITE: SINGLE-CRYSTAL DATA AND REFINEMENT PARAMETERS

Crystal system	monoclinic	Space group	$P2_1/a$ (no.14)
<i>a</i> (Å)	7.127(2)	β (°)	93.21(1)
<i>b</i> (Å)	8.842(3)	<i>V</i> (Å ³)	426.1(2)
<i>c</i> (Å)	6.773(2)	<i>Z</i>	12
Radiation	MoK α	μ (mm ⁻¹)	0.203
D_{calc} (g/cm ³)	2.049		
Measured reflections		3218	
<i>hkl</i> range		-10, 10/ -12, 12/ -8, 9	
Independent reflections		1372	
Observed reflections [$I > 2\sigma(I)$]		977	
Parameters refined		94	
Final R and $wR2$		0.043, 0.144	
<i>S</i>		1.098	

Notes: $R = \sum |F_o| - |F_c| / \sum |F_o|$; $wR2 = \{\sum [w(F_o^2 - F_c^2)^2] / \sum [w(F_o^2)^2]\}^{1/2}$; $w = 1/[\sigma^2(F_o^2) + (0.0885q)^2]$, where $q = [\max(0, F_o^2) + 2F_c^2]/3$; $S = \{\sum [w(F_o^2 - F_c^2)] / (n - p)\}^{1/2}$, where n is the number of reflections and p is the number of refined parameters.

Wilkinson 1980). The present X-ray-diffraction data and structure refinement confirm the new mineral species to be the monoclinic polymorph of metaboric acid, β -HBO₂ or HBO₂ II.

In the cubic γ -HBO₂ phase (or HBO₂ I), corresponding to the mineral metaborite, all the boron atoms are equivalent by symmetry and display a tetrahedral coordination. The BO₂(OH)₂ tetrahedra are linked by sharing their corners to form a unique framework, analogous to that of cristobalite, so that the mineral can be classified as a mega-tectoborate, according to Strunz & Nickel (2001). A trigonal-planar coordination for all the boron atoms is observed instead in the orthorhombic α -HBO₂ phase (or HBO₂ III), known as a synthetic product only (Peters & Milberg 1964). Here the triangular building-blocks share their vertices to form [B₃O₆]³⁻ rings that are linked by H⁺ ions into pseudohexagonal sheets parallel to (001), and van der Waals interactions occur between sheets. Therefore the presence of isolated triborate anions allows one to classify this phase as neso-triborate.

The basic structural units observed in clinometaborite are six-membered B₃O₃ rings containing two three-coordinated boron atoms [B(2) and B(3)] and an additional boron atom [B(1)] in a distorted tetrahedral coordination, because of interaction with a H₂O molecule [O(5)]. The tetrahedral B(1) atom is displaced toward O(5) from the otherwise nearly planar

B₃O₃ ring by about 0.40 Å. In agreement with the fibrous aspect of the crystals, the six-membered rings are linked through the O(4) atoms to form polymeric chains extending along [010] in a zig-zag fashion (Fig. 2) and arranged to form layers parallel to (201). These chains are linked *via* hydrogen bonds between exocyclic O(6)–H(3) hydroxyl groups and endocyclic O(3) atoms within the same layer and *via* interlayer hydrogen bonds between the donor H₂O molecule O(5) and the acceptor O(1) and O(6) atoms of different layers (see Table 5, Fig. 3). Clinometaborite is therefore an ino-triborate. Three additional weaker contacts, reported in Table 5, may be considered as parts of a polyfurcated hydrogen-bonding system. Triborate rings like those found in clinometaborite, consisting of two triangular and one tetrahedral boron atoms, are rare, the only other known example being ameghinite, Na[B₃O₃(OH)₄] (Dal Negro *et al.* 1975), which is, however, a neso-triborate.

The discovery of the β -form of metaboric acid (or HBO₂ II) at Vulcano together with metaborite emphasizes the remarkable mineralogical and physicochemical interest of this locality. The overall complex picture points out the coexistence of different phases of boric acid, with the possibility of various combinations of planar BO₃ and tetrahedral BO₄ groups, depending on the conditions of formation. However, according to Wells (1962) and to Kracek *et al.* (1938), the polymorphism of HBO₂ provides a very good illustration

TABLE 4. COORDINATES AND DISPLACEMENT PARAMETERS [$U_{eq}/U_{(i,j)}$, Å²] OF ATOMS IN CLINOMETABORITE

Atom	X/a	Y/b	Z/c	$U_{eq}/U_{(i,j)}$		
B(1)	0.23416(8)	-0.04313(6)	0.83071(7)	0.0149(1)		
B(2)	0.26124(7)	0.21091(6)	0.97679(8)	0.0142(1)		
B(3)	0.40077(7)	0.16241(6)	0.67183(8)	0.0146(1)		
O(1)	0.21653(5)	0.06075(3)	0.99463(5)	0.0148(1)		
O(2)	0.35331(5)	0.26191(4)	0.81577(5)	0.0193(1)		
O(3)	0.35674(5)	0.01313(4)	0.68317(5)	0.0174(1)		
O(4)	0.28810(5)	-0.19372(4)	0.87931(5)	0.0178(1)		
O(5)	0.03355(5)	-0.04735(4)	0.72342(5)	0.0184(1)		
O(6)	0.49777(5)	0.21595(4)	0.51815(5)	0.0202(1)		
H(1)	0.012(1)	-0.130(1)	0.650(1)	0.0540(2)		
H(2)	-0.068(1)	-0.048(1)	0.809(1)	0.0680(3)		
H(3)	0.534(1)	0.136(1)	0.443(1)	0.0350(2)		

Atom	U_{11}	U_{22}	U_{33}	U_{23}	U_{13}	U_{12}
B(1)	0.0221(2)	0.0106(2)	0.0128(2)	0.0012(2)	0.0078(2)	0.0004(2)
B(2)	0.0190(2)	0.0122(2)	0.0117(2)	0.0000(2)	0.0040(2)	0.0000(2)
B(3)	0.0208(2)	0.0116(2)	0.0119(2)	0.0011(2)	0.0046(2)	0.0000(2)
O(1)	0.0242(1)	0.0084(1)	0.0126(1)	-0.0018(1)	0.0068(1)	-0.0021(1)
O(2)	0.0322(2)	0.0110(1)	0.0159(1)	-0.0021(1)	0.0118(1)	-0.0040(1)
O(3)	0.0279(1)	0.0100(1)	0.0155(1)	0.0006(1)	0.0120(1)	-0.0009(1)
O(4)	0.0299(2)	0.0091(1)	0.0151(1)	0.0015(1)	0.0091(1)	0.0025(1)
O(5)	0.0229(1)	0.0157(2)	0.0169(1)	-0.0037(1)	0.0048(1)	-0.0010(1)
O(6)	0.0331(2)	0.0115(1)	0.0172(1)	0.0024(1)	0.0131(1)	0.0005(1)

The anisotropic displacement factor exponent takes the form: $-2\pi^2[U_{11}h^2(a^*)^2 + \dots + 2U_{12}hka^*b^* + \dots]$; $U_{eq} = 1/3(U_{11} + U_{22} + U_{33})$.

TABLE 5. CLINOMETABORITE: INTERATOMIC DISTANCES (Å) AND ANGLES (°)

B(1)–O(1)	1.4518(7)	B(2)–O(1)	1.3723(7)	B(3)–O(2)	1.3696(7)
B(1)–O(3)	1.4516(7)	B(2)–O(2)	1.3797(7)	B(3)–O(3)	1.3599(8)
B(1)–O(4)	1.4195(8)	B(2)–O(4) ^a	1.3502(7)	B(3)–O(6)	1.3660(7)
B(1)–O(5)	1.5676(7)				
<B(1)–O>	1.4726	<B(2)–O>	1.3674	<B(3)–O>	1.3652
O(1)–B(1)–O(3)	113.01(4)	O(3)–B(1)–O(5)	104.92(4)	O(2)–B(2)–O(4) ^a	121.40(5)
O(1)–B(1)–O(4)	116.75(4)	O(4)–B(1)–O(5)	108.26(4)	O(2)–B(3)–O(3)	121.09(5)
O(1)–B(1)–O(5)	104.70(4)	O(1)–B(2)–O(2)	120.59(4)	O(2)–B(3)–O(6)	118.39(5)
O(3)–B(1)–O(4)	108.39(4)	O(1)–B(2)–O(4) ^a	118.01(5)	O(3)–B(3)–O(6)	120.50(4)
B(1)–O(1)–B(2)	120.95(4)	B(2)–O(2)–B(3)	120.00(4)	B(1)–O(3)–B(3)	121.41(4)
B(1)–O(4)–B(2)	132.47(4)				

Hydrogen – bonding geometry

D – H...A	D – H	H...A	D...A	D – H...A
O(5) – H(1)...O(6) ^b	0.893(8)	1.774(8)	2.658(1)	170.5(7)
O(5) – H(2)...O(1) ^c	0.949(9)	1.752(9)	2.687(1)	167.5(8)
O(6) – H(3)...O(3) ^d	0.917(6)	1.774(6)	2.682(1)	170.4(6)
O(5) – H(2)...O(4) ^a	0.949(9)	2.559(9)	3.102(1)	116.6(7)
O(5) – H(2)...O(2) ^f	0.949(9)	2.594(9)	2.916(1)	100.2(6)
O(5) – H(1)...O(4) ^a	0.893(8)	2.771(9)	3.102(1)	103.4(5)
O(5) – H(1)...O(5) ^g	0.893(8)	2.976(9)	3.151(1)	92.9(6)
O(6) – H(3)...O(4) ^d	0.917(6)	2.633(7)	3.174(1)	118.5(5)

Symmetry codes: a = 1/2 – x, y + 1/2, 2 – z; b = 1/2 – x, y – 1/2, 1 – z; c = –x, –y, 2 – z; d = 1 – x, –y, 1 – z; e = x – 1/2, –1/2 – y, z; f = x – 1/2, 1/2 – y, z; g = –x, –y, 1 – z.

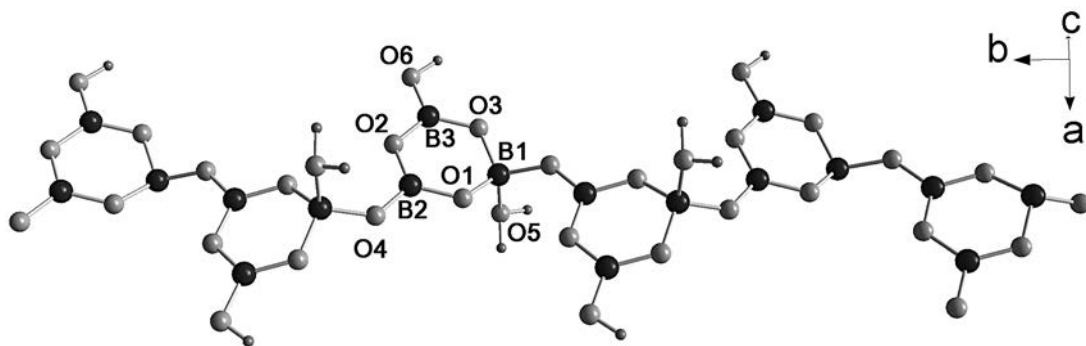


FIG. 2. A polymeric chain in the structure of clinometaborite.

of monotropism, *i.e.*, of the actual existence of various forms, some of which (*e.g.*, clinometaborite) are not thermodynamically stable at any condition. For this reason, in view of absence of chemical equilibrium, the conditions of formation of this mineral cannot be deduced with a firm foundation. A suggestion from Wells (1962), that this form “is readily prepared by the dehydration of orthoboric acid in an open vessel at 130–140°C; heating it in a sealed tube for a long period at higher temperature gives the most stable [cubic] form” could well apply to our occurrence.

Such conditions should reasonably occur in some part of the fumaroles, where sassolite (orthoboric acid) is a common product.

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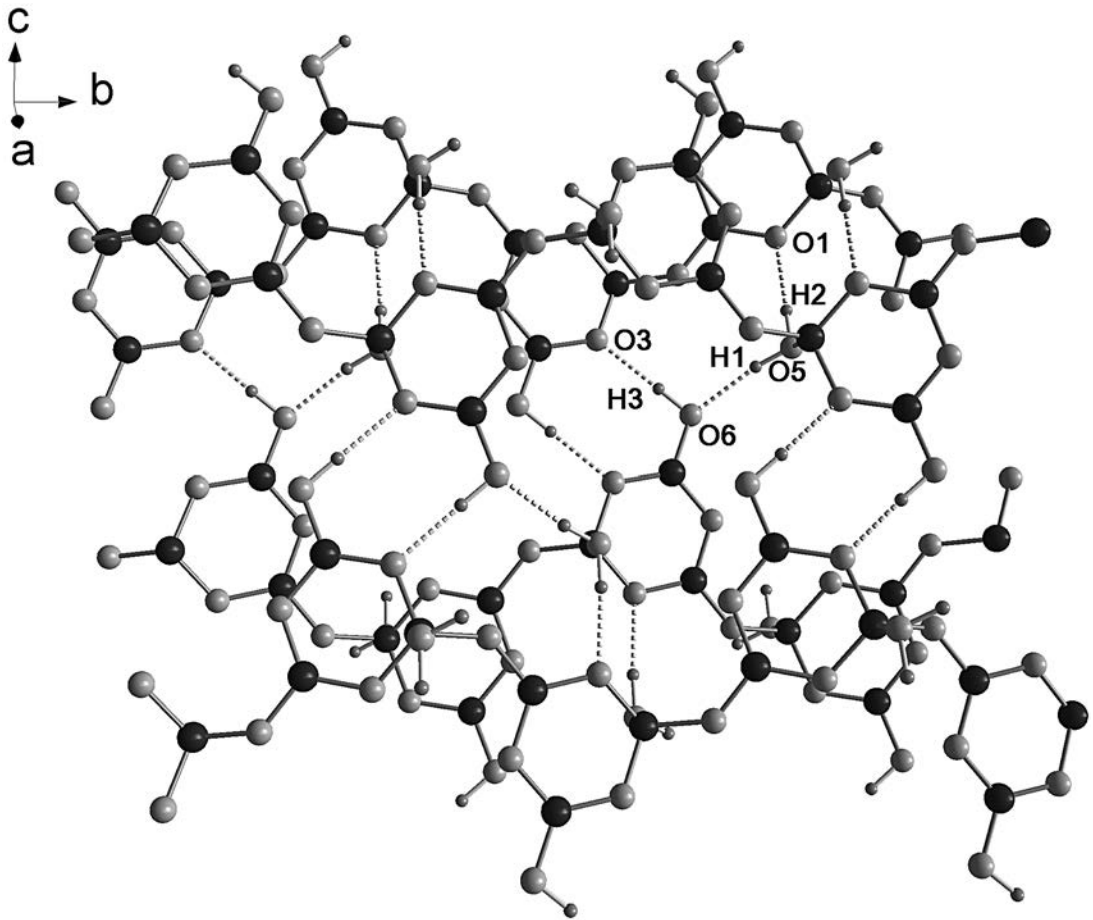


FIG. 3. The pattern of hydrogen bonds in clinometaborite, with O...O interactions less than 2.7 Å.

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