CAMERONITE, $Cu_{5-x}(Cu,Ag)_{3+x}Te_{10}$ (x = 0.43), FROM THE GOOD HOPE MINE, VULCAN, COLORADO: CRYSTAL STRUCTURE AND REVISION OF THE CHEMICAL FORMULA

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Abstract

The crystal structure of the mineral cameronite, a rare copper-silver telluride, was solved using intensity data collected from a crystal of the holotype material from the Good Hope mine, Vulcan, Colorado (USA). The study revealed that the structure is monoclinic, space group C2/c, with cell parameters: a 17.906(1), b 17.927(1), c 21.230(2) Å, β 98.081(8)°, and V 6747.2(8) Å³. The refinement of an anisotropic model led to an *R* index of 0.0216 for 9688 independent reflections. Nineteen Te sites and 14 metal (M) sites occur in the crystal structure of cameronite. Six M sites host Cu with variable amounts of Ag, whereas the remaining eight M sites host Cu only. All of the atoms are in tetrahedral coordination and complex MTe₄, M(Te₃M), M(Te₂M₂), TeM₄, Te(M₃Te), Te(M₂Te₂), Te(MTe₃), and TeTe₄ crystal-chemical environments are formed, as is typical of many intermetallic compounds. The short bond distances observed mainly among the Te atoms are discussed in relation to other copper and silver tellurides and pure metals. Electron microprobe analyses of the crystal used for the structural study led to the formula Cu_{7.08}Ag_{1.06}Fe_{0.01}S_{0.01}Te_{9.84}, on the basis of 18 atoms. On the basis of information gained from the structural and chemical characterization, the crystal-chemical formula was revised, yielding Cu_{5-x}(Cu,Ag)_{3+x}Te₁₀ (Z 14) with x = 0.43 instead of Cu₇AgTe₁₀ (Z 16) as previously reported.

Keywords: cameronite, crystal structure, tellurides, Good Hope, Colorado

INTRODUCTION

Cameronite was defined as a new mineral species with the chemical formula Cu₇AgTe₁₀ by Roberts *et al.* (1986) during a study of ores from the Good Hope mine, Vulcan, Colorado (USA). By means of X-ray photographic techniques (precession method), these authors studied four fragments of cameronite from a structural point of view and reported a tetragonal symmetry (with $P4_2/mmc$, $P4_2mc$, or $P\overline{42}c$ as possible space groups) with a 12.695(2), c 42.186(6) Å, V 6798.8(3.1) Å³ and Z = 16. Roberts *et al.* (1986) also noticed that the strongest X-ray reflections defined a subcell with a' = a/3 =4.232 and c' = c/7 = 6.027 Å, and that the precession single-crystal films were dominated by *hkl* reflections with *h*,*k* = 3n and *l* = 7n. Interestingly, a mineral with the same physical, optical, chemical, and diffraction characteristics had already been described by Cameron & Threadgold (1961) in their paper describing the mineral species vulcanite. These authors also carried out single-crystal studies which suggested a hexagonal symmetry. Although Roberts *et al.* (1986) concluded that the unknown phase studied by Cameron & Threadgold (1961) and cameronite were the same phase, they failed to locate any symmetry element in support of the hexagonal symmetry and the reason for such a difference (tetragonal *versus* hexagonal) remained unresolved at that time.

In the course of a research project dealing with the description and structural characterization of natural copper and silver tellurides (Bindi & Cipriani 2004a,

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2004b, 2004c, Bindi 2008, 2009, Bindi *et al.* 2004, 2005, 2009, 2013), we have examined a fragment from the holotype material (sample R-934) obtained in 1974 by one of us (WWP) on exchange from the Smithsonian Institution, Washington, D.C. (USA). The sample consists of large rickardite grains associated with vulcanite, cameronite, native tellurium, and arsenopyrite.

To help resolve the concerns relating to the structure of cameronite and those related to the different symmetry reported for cameronite in the literature, we present new crystal structure data for the mineral from its type locality, together with new chemical data.

CRYSTAL-STRUCTURE SOLUTION AND REFINEMENT

A small crystal fragment ($65 \times 85 \times 115 \mu m$) was selected for the single-crystal X-ray diffraction study. The intensity data collection (see Table 1 for details) was carried out by means of an Oxford Diffraction Xcalibur 3 single-crystal diffractometer (enhanced X-ray source, X-ray radiation MoK α , $\lambda = 0.71073$ Å) fitted with a Sapphire 2 CCD detector. A total of 1589 frames of data were collected at room temperature as eight sets of omega runs with an exposure time of 45 s per frame and a frame width of 1.00°. This afforded an overall data collection of 78245 reflections (9688 unique). The refined monoclinic unit-cell parameters are $a = 17.906(1), b = 17.927(1), c = 21.230(2) \text{ Å}, \beta = 98.081(8)^\circ$, and V 6747.2(8) $Å^3$, which are related, neglecting the monoclinic distortion, to those found by Roberts et al. (1986) [a 12.695(2), c 42.186(6) Å, V 6798.8(3.1) Å³] by the following transformation matrix $[\frac{1}{2}, \frac{1}{2}, 0/\frac{1}{2}, \frac{1}{2}, 0/0, 0, \overline{2}$]. The presence of a tetragonal substructure with $a' \approx 4.2$ and $c' \approx 6.03$ Å was confirmed, and found to be due to the typical disordered 'rickardite-type' basic structure (Forman & Peacock 1949, Schutte & de Boer 1993).

Data frames were processed using the CrysAlis software package (Oxford Diffraction 2006) running on the Xcalibur 3 control PC. The program ABSPACK (Oxford Diffraction 2006) was used for the absorption correction. The merging R for the data set decreased from 14.25% before the absorption correction to 3.66% after the correction. The observed reflection conditions (*hkl*: h+k = 2n; *h*0*l*: h,l = 2n; 0*kl*: k = 2n; *hk*0: h+k =2n; 0k0: k = 2n; h00: h = 2n; 00l: l = 2n), together with the statistical tests on the distribution of |E| values that strongly indicated the presence of an inversion center $(|E^2 - 1| = 0.976)$, suggested space group C2/c. The structure solution was then initiated in this space group. The positions of most of the atoms (all the M positions and most of the Te atoms) were determined by means of direct methods (Sheldrick 2008). A least-squares refinement on F^2 using these heavy-atom positions and isotropic temperature factors produced an R factor of 0.165. Three-dimensional difference Fourier synthesis yielded the position of the remaining Te atoms. The SHELXL program (Sheldrick 2008) was used for the refinement of the structure. The occupancy of all the sites was left free to vary (Ag *versus* vacancy for the M sites; Te *versus* vacancy). The occupancy of eight M sites (M7–M14; Table 2) was found to be consistent with complete occupation by Cu and then fixed to the resulting value. Neutral scattering curves for Ag, Cu, and Te were taken from the *International Tables for X-ray Crystallography* (Ibers & Hamilton 1974). In the final stage, with anisotropic atomic displacement parameters for all atoms and no constraints, the residual value settled at R = 0.0210 for 3839 observed reflections [$F_0 > 4\sigma(F_0)$] level] and 292 parameters, and at R = 0.0216 for all 9688 independent reflections. Inspection

TABLE 1. DATA AND EXPERIMENTAL DETAILS FOR THE SELECTED CAMERONITE CRYSTAL

Crystal data	
Formula Crystal size (mm) Form Color Crystal system Space group a (Å) b (Å) c (Å) β (°) V (Å ³) Z	$\begin{array}{l} Cu_{5-x}(Cu,Ag)_{3+x}Te_{10} \mbox{ (with } x=0.43)\\ 0.065\times0.085\times0.115\\ \mbox{black}\\ \mbox{black}\\ \mbox{monoclinic}\\ C2/c \mbox{ (\#15)}\\ 17.906(1)\\ 17.927(1)\\ 21.230(2)\\ 98.081(8)\\ 6747.2(8)\\ 14 \end{array}$
Data collection	
Lasta versat	Outrad Differentians Vanilihum 2
Instrument	Oxford Diffraction Acalibur 3 $M_{2}(k_{1}, \ell_{2}) = 0.34032$
	$MOK\alpha$ ($\lambda = 0.71073$)
Detector to sample distance (cm)	295(5) 2 6
Number of frames	1589
Measuring time (s)	45
Maximum covered 20 (°)	59.99
Absorption correction	multi-scan (<i>ABSPACK</i> ; Oxford Diffraction 2006)
Collected reflections	78245
Unique reflections	9688
Reflections with $F_0 > 4\sigma(F_0)$	3839
R _{int}	0.0366
R _σ	0.0582
Range of <i>h</i> , <i>k</i> , <i>l</i>	$-25 \le h \le 25, -25 \le k \le 25, -29 \le l \le 29$
Refinement	

Refinement	Full-matrix least squares on F ²
Final $R_1 [F_0 > 4\sigma(F_0)]$	0.0210
Final R1 (all data)	0.0216
Number refined parameters	292
Δρ _{max} (e Å ⁻³)	1.47
Δρ _{min} (e Å ⁻³)	-2.01

Atom	s.o.f.	x/a	y/b	z/c	U _{eq}
M1	Cu _{0.78(1)} Ag _{0.22}	0.115992(9)	0.045060(7)	-0.036220(7)	0.01961(5)
M2	Cu _{0.69(1)} Ag _{0.31}	0.209464(9)	-0.042355(6)	0.030806(7)	0.01957(5)
M3	Cu _{0.92(1)} Ag _{0.08}	0.733704(8)	0.291402(6)	0.175078(7)	0.01447(5)
M4	Cu _{0.57(1)} Ag _{0.43}	0.54814(1)	-0.041555(7)	0.028293(9)	0.02767(5)
M5	Cu _{0.70(1)} Ag _{0.30}	-0.093670(9)	0.879470(9)	0.675700(8)	0.02702(6)
M6	Cu _{0.59(1)} Ag _{0.41}	-0.070723(8)	0.043392(7)	0.822239(6)	0.02390(5)
M7	Cu _{1.00}	0.38291(1)	0.788241(9)	0.032437(9)	0.02171(4)
M8	Cu _{1.00}	0.573341(9)	0.126046(8)	0.176352(8)	0.02044(5)
M9	Cu _{1.00}	0.76177(1)	0.457468(9)	0.323851(9)	0.02501(5)
M10	Cu _{1.00}	0.40567(1)	-0.042242(9)	0.176939(9)	0.02447(5)
M11	Cu _{1.00}	0.93115(1)	0.291467(8)	0.323874(8)	0.02066(5)
M12	Cu _{1.00}	0.59096(1)	0.293941(8)	0.320657(7)	0.02025(5)
M13	Cu _{1.00}	0.26180(1)	0.374311(9)	0.822078(8)	0.02389(5)
M14	Cu _{1.00}	0.71466(1)	0.124980(9)	0.035825(9)	0.02241(5)
Te1	Te _{1.00}	0.5	0.041272(5)	0.25	0.01766(3)
Te2	Te _{1.00}	0	0.210338(6)	0.25	0.02099(3)
Te3	Te _{1.00}	0	0.123149(6)	0.75	0.02071(3)
Te4	Te _{1.00}	0.835113(4)	0.376439(4)	0.251411(3)	0.02200(3)
Te5	Te _{1.00}	0.358290(4)	0.288967(4)	-0.108482(3)	0.02135(3)
Te6	Te _{1.00}	0.666249(5)	0.206467(4)	0.251036(4)	0.02069(3)
Te7	Te _{1.00}	0.311121(4)	0.041591(3)	0.109886(4)	0.02157(3)
Te8	Te _{1.00}	0.478428(5)	-0.122799(4)	0.110142(4)	0.02300(3)
Te9	Te _{1.00}	0.194542(6)	0.123686(5)	-0.108793(4)	0.02388(3)
Te10	Te _{1.00}	0.641929(6)	0.039075(5)	0.111922(6)	0.02365(3)
Te11	Te _{1.00}	0.476675(6)	0.204339(6)	0.110747(5)	0.02075(3)
Te12	Te _{1.00}	0.831844(6)	-0.041590(5)	0.747907(5)	0.02551(3)
Te13	Te _{1.00}	0.025378(7)	-0.041174(5)	-0.107167(6)	0.02215(4)
Te14	Te _{1.00}	0.646378(6)	0.376196(5)	0.106946(5)	0.02412(3)
Te15	Te _{1.00}	0.809574(5)	0.208891(4)	0.111224(4)	0.01983(3)
Te16	Te _{1.00}	0.452296(7)	0.376501(5)	-0.036558(5)	0.02051(3)
Te17	Te _{1.00}	0.546182(7)	0.293549(5)	0.036913(6)	0.02190(4)
Te18	Te _{1.00}	0.378792(8)	0.126708(6)	0.029071(7)	0.02185(3)
Te19	Te _{1.00}	0.284180(6)	0.210017(4)	-0.032196(5)	0.02119(3)

TABLE 2. ATOMS, SITE OCCUPANCY FACTORS (s.o.f.), ATOMIC COORDINATES, AND EQUIVALENT ISOTROPIC DISPLACEMENT PARAMETERS (Å²) FOR THE SELECTED CAMERONITE CRYSTAL

of the difference Fourier map revealed that maximum positive and negative peaks were 1.47 and 2.01 e^{-/}Å³, respectively. Experimental details and *R* indices are given in Table 1. Fractional atomic coordinates and isotropic displacement parameters are reported in Table 2, whereas the anisotropic displacement parameters are given in Table 3. A list of the observed and calculated structure factors is available from the Depository of Unpublished Data, MAC website [document Cameronite CM52_423].

CHEMICAL COMPOSITION

The chemical composition was determined using wavelength dispersive analysis (WDS) conducted with a JEOL JXA-8200 electron microprobe from the same crystal used for the structural study. Concentrations of major and minor elements were determined at an accelerating voltage of 20 kV and a beam current of 40 nA, with 10 s as the counting time. For the WDS analyses the following lines were used: $SK\alpha$, $FeK\alpha$, $CuK\alpha$, $ZnK\alpha$, $AsL\alpha$, $SeL\alpha$, $AgL\alpha$, $SbL\beta$, $TeL\alpha$, Pb $M\alpha$, and Bi $M\beta$. The standards employed were: native elements for Cu and Ag, galena for Pb, pyrite for Fe and S, synthetic Sb₂Te₃ for Sb and Te, synthetic As₂S₃ for As, synthetic Bi₂S₃ for Bi, synthetic ZnS for Zn, and synthetic PtSe₂ for Se. The crystal fragment was found to be homogeneous within analytical uncertainty. The average chemical compositions (six analyses from different spots), together with ranges of wt.% of elements, are reported in Table 4. On the basis of 18 atoms, the chemical formula for cameronite can be written as $Cu_{7.08}Ag_{1.06}Fe_{0.01}S_{0.01}Te_{9.84}$ or, according to the revised chemical formula obtained on the basis of the structural results (see below), $Cu_{5-x}(Cu,Ag)_{3+x}Te_{10}$ with x = 0.43.

	U ₁₁	U ₂₂	U ₃₃	U ₁₂	U ₁₃	U ₂₃
M1	0.01992(8)	0.01966(7)	0.01927(8)	0.00002(4)	0.00285(5)	-0.00002(4)
M2	0.01992(8)	0.01955(7)	0.01927(7)	0.00001(3)	0.00281(5)	0.00002(3)
M3	0.01476(8)	0.01450(7)	0.01420(7)	-0.00003(3)	0.00217(4)	0.00003(3)
M4	0.02794(9)	0.02769(8)	0.02735(8)	0.00001(4)	0.00380(5)	-0.00002(4)
M5	0.02732(9)	0.02698(8)	0.02676(8)	-0.00001(4)	0.00383(6)	-0.00003(4)
M6	0.02407(9)	0.02397(8)	0.02364(8)	0.00005(4)	0.00329(5)	-0.00003(4)
M7	0.02196(9)	0.02168(8)	0.02149(8)	0.00001(5)	0.00310(6)	0.00013(5)
M8	0.02082(9)	0.02035(8)	0.02023(8)	-0.00002(4)	0.00320(6)	0.00000(4)
M9	0.0253(1)	0.02501(9)	0.02472(9)	-0.00011(5)	0.00364(7)	0.00009(5)
M10	0.0247(1)	0.02458(9)	0.02422(9)	-0.00005(5)	0.00364(7)	-0.00002(5)
M11	0.02105(9)	0.02067(8)	0.02028(8)	-0.00002(4)	0.00300(6)	-0.00002(4)
M12	0.02046(9)	0.02019(8)	0.02011(8)	-0.00004(4)	0.00294(6)	-0.00002(4)
M13	0.0241(1)	0.02392(9)	0.02366(9)	-0.00008(5)	0.00336(7)	0.00004(5)
M14	0.02272(9)	0.02240(8)	0.02210(8)	-0.00002(5)	0.00312(7)	0.00001(5)
Te1	0.02055(6)	0.02102(5)	0.02038(6)	0	0.00291(4)	0
Te2	0.02136(6)	0.02095(6)	0.02067(6)	0	0.00301(4)	0
Te3	0.02102(6)	0.02069(5)	0.02043(6)	0	0.00294(4)	0
Te4	0.02266(5)	0.02238(4)	0.02104(4)	0.00002(2)	0.00217(3)	0.00003(2)
Te5	0.02255(4)	0.02036(4)	0.02208(4)	0.00003(2)	0.00251(3)	-0.00002(2)
Te6	0.02100(5)	0.02072(5)	0.02036(5)	0.00005(2)	0.00237(3)	-0.00002(2)
Te7	0.02091(5)	0.02259(4)	0.02122(4)	0.00001(2)	0.00208(3)	0.00001(2)
Te8	0.02329(5)	0.02203(4)	0.02268(4)	-0.00002(2)	0.00220(3)	-0.00002(2)
Te9	0.02213(5)	0.02187(5)	0.02064(5)	0.00001(2)	0.00218(4)	0.00000(2)
Te10	0.02291(6)	0.02267(5)	0.02236(6)	-0.00003(3)	0.00215(4)	-0.00003(3)
Te11	0.02104(6)	0.02074(5)	0.02045(6)	-0.00006(3)	0.00227(4)	-0.00001(3)
Te12	0.02177(6)	0.02250(5)	0.02224(5)	-0.00001(3)	0.00256(4)	0.00004(3)
Te13	0.02240(7)	0.02215(6)	0.02190(6)	0.00004(3)	0.00256(5)	-0.00004(3)
Te14	0.02456(5)	0.02406(5)	0.02374(5)	0.00000(2)	0.00243(4)	-0.00004(2)
Te15	0.02010(5)	0.01986(5)	0.01954(5)	0.00004(2)	0.00283(3)	0.00001(2)
Te16	0.02076(6)	0.02152(5)	0.02028(5)	-0.00002(3)	0.00210(4)	0.00002(3)
Te17	0.02223(6)	0.02187(5)	0.02160(6)	0.00003(3)	0.00227(4)	-0.00001(3)
Te18	0.02209(6)	0.02215(6)	0.02203(6)	-0.00004(3)	0.00234(4)	0.00001(3)
Te19	0.02143(5)	0.02122(5)	0.02094(5)	-0.00004(2)	0.00202(3)	0.00001(2)

TABLE 3. ANISOTROPIC DISPLACEMENT PARAMETERS OF THE ATOMS FOR THE SELECTED CAMERONITE CRYSTAL

TABLE 4. ELECTRON MICROPROBE DATA AND ATOMIC RATIOS WITH THEIR STANDARD DEVIATIONS (σ) FOR THE SELECTED CAMERONITE CRYSTAL

	wt.%	Range	atomic ratios	σ
Ag	6.25	6.10-6.66	1.06	0.05
Cu	24.49	24.11-25.09	7.08	0.15
Bi	0.01	0.00-0.04	-	-
Pb	0.01	0.00-0.03	-	-
Zn	0.01	0.00-0.04	-	-
Fe	0.02	0.00-0.06	0.01	0.01
Sb	0.01	0.00-0.02	-	-
As	0.01	0.00-0.03	-	-
S	0.01	0.00-0.02	0.01	0.01
Se	0.02	0.00-0.05	-	-
Те	68.35	67.50-69.11	9.84	0.14
Total	99.19	98.90-100.55		

The means and ranges of electron microprobe data are in wt.% of elements. Atomic ratios are on the basis of 18 atoms.

DESCRIPTION OF THE STRUCTURE AND DISCUSSION

The crystal structure of cameronite (Figs. 1 and 2) represents a new structure type in the Cu–Te system (see Pashinkin & Fedorov 2003 for a review). It can be considered an intermetallic compound and, therefore, the classic polyhedral description is hardly applicable. There are 19 independent Te sites and 14 metal (M) sites in the crystal structure with Z = 14. The calculated density for cameronite is therefore 6.303 g/cm³ and not 7.144 g/cm³ as previously hypothesized on the basis of Z = 16 (Roberts *et al.* 1986).

In the structure, six M sites (M1–M6; Table 2) host Cu with variable amounts of Ag (from 0.08 to 0.41 *apfu*), whereas the remaining eight M sites (M7–M14; Table 2) host Cu only. All the atoms are in tetrahedral coordination and complex MTe₄, M(Te₃M), M(Te₂M₂), TeM₄, Te(M₃Te), Te(M₂Te₂), Te(MTe₃), TeTe₄ crystal-chemical environments are formed (Table 5). The presence of Ag substituting for Cu at some of the M sites



FIG. 1. The crystal structure of cameronite down [001]. The horizontal direction is the *a* axis. The M1–M6 atoms are depicted as tetrahedra whereas the M7– M14 atoms are shown as small light grey spheres; Te: large dark grey spheres. The unit cell is outlined.



FIG. 2. The crystal structure of cameronite down [001]. The horizontal direction is the *a* axis. The M7–M14 atoms are depicted as tetrahedra whereas the M1– M6 atoms are shown as small white spheres; Te: large dark grey spheres. The unit-cell is outlined.

TABLE 5. BOND DISTANCES (IN Å) IN THE STRUCTURE OF CAMERONITE

M4		MO		Ted		T-42	
M1 T= 10	0 5000(0)		0.4007(0)	164	0.0045(0)	Ie12	0 5050(0)
T= 10	2.5260(3)	Te9	2.4007(3)	IVI9	2.0015(2)	IVI9	2.3639(3)
leis	2.5698(3)		2.5124(3)	MITU	2.6044(2)	IVI5	2.5897(3)
M2	2.5714(3)	le12	2.5859(3)	M11	2.6268(3)	M13	2.6231(2)
le9	2.6359(2)	le4	2.6015(2)	M3	2.7240(2)	M6	2.6610(2)
mean	2.5763	mean	2.5421	mean	2.6392	mean	2.6149
M2		M10		Te5		Te13	
Te14	2,5562(2)	Te8	2,5126(3)	M11	2,5239(2)	M5	2.4771(3)
M14	2,5654(3)	Te7	2,5444(3)	M13	2.6044(3)	M1	2,5698(3)
M1	2 5714(3)	Te1	2 6007(3)	Te16	2 6277(2)	Te16	2 5863(2)
Te7	2 7454(3)	Te4	2 6044(2)	Te19	2 6441(2)	M6	2 6039(3)
mean	2.6096	mean	2.5655	mean	2.6000	mean	2.5593
				T -0		T-44	
M3	a (aa (a)	M11	0.5445(0)	166	0 =000(0)	1014	a
le14	2.4924(2)	le8	2.5117(2)	M13	2.5639(3)	M3	2.4924(2)
Te15	2.5277(2)	Te5	2.5239(2)	M8	2.5726(3)	M2	2.5562(2)
Te6	2.6316(2)	Te2	2.5766(2)	M3	2.6316(2)	M6	2.5973(2)
Te4	2.7240(2)	Te4	2.6268(3)	M12	2.6498(2)	Te17	2.6220(2)
mean	2.5939	mean	2.5598	mean	2.6045	mean	2.5670
M4		M12		Te7		Te15	
Te18	2,4440(3)	Te3	2,5348(3)	M10	2,5444(3)	M3	2.5277(2)
M4	2,4603(4)	Te11	2,5814(2)	M13	2.5680(2)	M5	2,5919(2)
Te10	2 6876(3)	Te6	2 6498(2)	Te18	2 7064(2)	M14	2 6360(3)
Te8	2 7035(3)	Te9	2 6614(3)	M2	2 7454(3)	M7	2 6764(3)
mean	2.5739	mean	2.6069	mean	2.6411	mean	2.6080
M5		M13		Te8		Te16	
Te13	2.4771(3)	Te6	2.5639(3)	M11	2.5117(3)	M1	2.5280(3)
Te12	2.5897(2)	Te7	2.5680(2)	M10	2.5126(3)	Te13	2.5863(2)
Te15	2.5919(2)	Te5	2.6044(3)	M4	2.7035(3)	Te17	2.5949(2)
Te2	2.6739(3)	Te12	2.6231(3)	M7	2.7197(3)	Te5	2.6277(2)
mean	2.5832	mean	2.5899	mean	2.6119	mean	2.5842
M6		M14		To9		Te17	
To3	2 5585(2)	M2	2 5654(3)	MQ	2 4687(3)	M7	2 5407(3)
Te14	2.5505(2)	MZ	2.500+(3)	To10	2.7007(3)	To16	2.5+07(3)
Te 14	2.0973(2)		2.0194(3)	1019	2.0234(2)		2.5949(2)
Te13	2.0039(3)	Te10	2.0300(3)	IVI I M10	2.0339(2)	Te 14	2.0220(2)
leiz	2.0010(2)	leiu	2.0902(3)		2.0014(3)	lell	2.0009(2)
mean	2.0052	mean	2.0293	mean	2.5974	mean	2.0001
М7		le1		le10		le18	
Te17	2.5407(3)	M10 (×2)	2.6007(3)	M8	2.5065(3)	M4	2.4440(3)
M14	2.6194(3)	M8 (×2)	2.6564(2)	M9	2.5124(3)	Te19	2.4845(2)
Te15	2.6764(3)	mean	2.6286	M4	2.6876(3)	Te11	2.6760(3)
Te8	2.7197(3)			M14	2.6962(3)	Te7	2.7064(2)
mean	2.6391	Te2		mean	2.6007	mean	2.5777
		M11 (×2)	2.5766(2)				
M8		M5 (×2)	2.6739(3)	Te11		Te19	
Te11	2.4956(3)	mean	2.6253	M8	2.4956(3)	Te19	2.4275(2)
Te10	2.5065(3)			M12	2.5814(2)	Te18	2.4845(2)
Te6	2.5726(3)	Te3		Te17	2.6669(2)	Te9	2.6234(2)
Te1	2.6564(2)	M12 (×2)	2.5348(2)	Te18	2.6760(3)	Te5	2.6441(2)
mean	2.5578	M6 (×2)	2.5585(2)	mean	2.6050	mean	2.5449
		mean	2.5467				

(M1–M6) does not affect the bond distances (mean bond distances in the range 2.54–2.64 Å). Similarly, the Te-tetrahedra show mean bond distances in the range 2.54–2.64 Å, independent of the crystal-chemical environment adopted. Figures 3 and 4 show typical examples of "interpenetrated tetrahedra" that can be seen in the cameronite structure. Such a peculiar structural situation promotes the occurrence of very short bond distances between atoms. The shortest (Cu,Ag)–Te distance in cameronite of 2.4440(3) Å (*i.e.*, M4–Te18) is indeed much shorter than the 2.604(1) Å and 2.63(1) Å observed in vulcanite (CuTe, Pertlik 2001) and in the average structure of rickardite (Cu_{3–x}Te₂, Schutte & de Boer 1993), respectively. A similar behavior is



observable for the Cu–Cu and Te–Te contacts: the shortest Cu–Cu distance (*i.e.*, M4–M4) of 2.4603(4) Å is shorter than that observed in metallic copper (2.546 Å; Suh *et al.* 1988) and the shortest Te–Te distance (*i.e.*, Te19–Te19) of 2.4275(2) Å is shorter than that observed in native tellurium (Te–Te = 2.835 Å, Cherin & Unger 1967). The same features have been observed in the structure of weissite (Bindi *et al.* 2013), which shows even shorter Cu–Te [2.308(2) Å] and Cu–Cu [2.282(3) Å] distances and similar Te–Te [2.496(1) Å] contacts.

Weissite and cameronite can be considered as complex superstructures of the basic tetragonal P4/nmm substructure with $a' \approx 4.2$ and $c' \approx 6.03$ Å (the disordered 'rickardite-type' basic structure; Forman & Peacock 1949, Schutte & de Boer 1993) where the arrangement of the Cu (and Ag) ions depends on the composition of the phase and thermal history. The formation of such superstructures, as others have observed in the Cu-Te system (Blachnik et al. 1983, Pashinkin & Fedorov 2003), is generally ascribed to the mobility of the copper ions in different valence states, Cu+ or Cu2+, in a rigid Te framework. The presence of Ag in cameronite seems fundamental to stabilize the new structure type never before recognized in the deeply studied Cu-Te system. However, Ag does not order at a specific site in the crystal structure of cameronite, as would be expected by its unitary stoichiometric content in the chemical formula when normalized on the basis of 18 atoms. This implies that the structure determined here could still be an average structure (although neither commensurate nor incommensurate satellites have been observed) and the real structure might be an even larger superstructure or, more likely, a complex incommensurately modulated structure like that observed in synthetic rickardite (Schutte & de Boer 1993) with a displacive modulation of the atoms associated with a Cu^+/Cu^{2+} valence fluctuation. Such a phenomenon has also been observed in calaverite, AuTe₂, with a fluctuation between Au⁺ and Au³⁺ in the structure (Bindi *et al.* 2009).

Table 6 shows the X-ray powder pattern calculated using the structural data obtained in this study. It appears very similar to that reported by Roberts *et al.* (1986) for cameronite. The presence of two minor peaks (at 2.531 and 2.344 Å) reported in the original study (Roberts *et al.* 1986) and not present in the calculated pattern is noted. Possible reasons could be linked to the fact that cameronite exhibits complex intergrowths with rickardite and vulcanite.

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	1		2	
	I		2	
hkl	d _{calc} (Å)	I _{calc}	d _{obs} (Å)	I _{rel}
-334 333	3.4544	100	3.456	100
-335	3.1325	3	-	-
061	2.9581	5	2.958	20
- 008	2.6274	2	2.635 2.531	10 10
-	-	-	2.344	5
067 608 606	2.1180 2.1167	25 25	2.118	100
-661 -33 <u>11</u>	2.1115 1.8089	26 15	1.815	30
33 <u>10</u> –394 393	1.8024	15	1.804	60
-935 932	1.8008	15		
-395 394	1.7516	2	1.755	3
-936 933	1.7502	2		

TABLE 6. X-RAY POWDER DIFFRACTION DATA FOR CAMERONITE

-669	1.6459	2	1.650	5
667				
00 <u>14</u>	1.5014	4	1.507	10
0 <u>12</u> 0	1.4939	4	1.496	20
– <u>12</u> 02	1.4922	5		
-93 <u>12</u>	1.3734	6	1.377	40
939				
-995	1.3705	8		
992				
–66 <u>15</u>	1.2236	7	1.227	20
66 <u>13</u>				
-6 <u>12</u> 8	1.2205	8		
6 <u>12</u> 6				
– <u>12</u> 69	1.2198	8	1.222	40
– <u>12</u> 65				
–33 <u>18</u>	1.1553	3	1.158	5
33 <u>17</u>				
-99 <u>12</u>	1.1515	3	1.151	30
999				
–3 <u>15</u> 4	1.1502	3		
3 <u>15</u> 3				
– <u>15</u> 36	1.1490	3		
<u>15</u> 31				
0 <u>1214</u>	1.0590	2	1.062	10
– <u>1212</u> 2	1.0557	2	1.058	5
–39 <u>18</u>	1.0136	2	1.017	10
39 <u>17</u>				
-93 <u>19</u>	1.0134	2	1.014	5
93 <u>16</u>				
–3 <u>1511</u>	1.0113	2	1.012	20
3 <u>1510</u>				
– <u>15</u> 3 <u>13</u>	1.0105	3		
<u>15</u> 38				
-9 <u>15</u> 5	1.0099	2		
9 <u>15</u> 2				
– <u>15</u> 96	1.0094	2		
<u>15</u> 91				
06 <u>21</u>	0.9491	1	0.9521	10
0 <u>18</u> 7	0.9453	1	0.9465	30
-6 <u>18</u> 1	0.9447	1		
- <u>18</u> 63	0.9438	1		
-99 <u>19</u>	0.9137	2	0.9164	10
99 <u>16</u>			0.9140	30

Notes: 1 = Calculated powder pattern and indexing for cameronite on the basis of *a* = 17.906(1), *b* = 17.927(1), *c* = 21.230(2) Å, β = 98.081(8)° and with the atom coordinates reported in Table 2. Intensities were calculated using XPOW software, version 2.0 (Downs *et al.* 1993). 2 = observed powder pattern originally reported by Roberts *et al.* (1986).

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