# Brodtkorbite, Cu<sub>2</sub>HgSe<sub>2</sub>, from Příbram, Czech Republic: crystal structure and description

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Abstract: The crystal structure and second world occurrence of the rare selenide brodtkorbite, ideally Cu<sub>2</sub>HgSe<sub>2</sub>, are reported in this paper. Brodtkorbite occurs in the Příbram uranium and base-metal ore district, central Bohemia (Czech Republic) as anhedral aggregates up to 150 µm across, in intergrowths with tiemannite and eucairite in a calcite gangue. Other associated minerals include berzelianite, umangite and uraninite. In reflected light, brodtkorbite is cream-white somewhat darker than the associated tiemannite and moderately pleochroic with light pink or bluish tints. Between crossed polarisers, the anisotropy is strong, with violet-grey to yellow-grey rotation tints. Reflectance values in air (R<sub>min</sub>, R<sub>max</sub> in %, λ nm) are 29.4–36.9 (470), 27.8–34.4 (546), 27.2–33.7 (589), 26.4–33.4 (650). Its average composition (14 electron microprobe analyses) is Ag 0.49, Hg 43.08, Cu 23.86, Se 32.25, Te 0.04, S 0.05, total 99.78 wt%. The resulting empirical formula, on the basis 5 atoms per formula unit (apfu), is (Cu<sub>1.87</sub>Ag<sub>0.02</sub>)<sub>Σ1.89</sub>Hg<sub>1.07</sub>(Se<sub>2.03</sub>S<sub>0.01</sub>)<sub>Σ2.04</sub>. Brodtkorbite is monoclinic, a 7.5048(12), b 4.1851(9), c 8.0131(12) Å, β 124.430  $(16)^\circ$ , V 207.59(8) Å<sup>3</sup>, space group  $P_{21/c}$ , Z=2. The crystal structure of brodtkorbite was solved by the direct method and refined to R = 8.05% on the basis of 416 unique reflections  $[I > 3\sigma(I)]$  collected on a Rigaku (Oxford Diffraction) SuperNova diffractometer with an Atlas S2 CCD detector and using MoK $\alpha$  radiation. The crystal structure contains one special position of Hg and two general positions of Cu and Se. The crystal structure consists of layers of edge-sharing [CuSe<sub>4</sub>] distorted tetrahedra where Cu atoms show a 3+1 bond scheme. The [CuSe<sub>4</sub>] tetrahedra form mackinawite-like layers parallel to (100). The layers show an AA-type of stacking and are connected by linear Se-Hg-Se bonds. Each Cu atom shows three short contacts with three adjacent Cu atoms forming a pseudohexagonal net of metal-metal interactions within one layer. The brodtkorbite crystal structure shows similarities to that of imiterite (Ag<sub>2</sub>HgS<sub>2</sub>), but the two minerals are not isostructural.

Key-words: brodtkorbite; silver mercury selenide; selenide minerals; crystal structure; uranium deposit; Příbram; Czech Republic.

# 1. Introduction

Brodtkorbite is a very rare selenide of Cu and Hg, first described by Paar et al. (2002) from the rich selenide association of the Tumiñico selenium deposit, La Rioja, Argentina. The original crystallographic data indicated that brodtkorbite is monoclinic, space group P2/n, with a 7.492 Å, b 4.177 Å, c 7.239 Å, and β 114.2°, but its crystal structure was not determined. A new find of brodtkorbite at the Příbram uranium and base-metal ore district (Czech Republic) yielded material suitable for new descriptive data, which include results of the crystal structure analysis. Although the chemical composition of brodtkorbite is very simple, the synthetic phases Cu<sub>2</sub>HgSe<sub>2</sub> and its sulphide analogue Cu<sub>2</sub>HgS<sub>2</sub> have not yet been prepared.

#### 2. Occurrence

Selenide mineralizations are known from several ore (especially uranium) districts of the Bohemian Massif. The most important occurrences are Bukov (Kvaček, 1973, 1979, 1987), Předbořice (Johan et al., 1971a and b; Johan & Kvacek, 1971; Paar et al., 2005; Bindi et al., 2016), Oldřichov near Tachov (Čech & Vavřín, 1978, 1979) or Zálesí (Topa et al., 2010, Sejkora et al., 2011, 2012, 2014).

The Příbram ore area in central Bohemia, Czech Republic is well-known for the deposits of base-metal as well as U ores. There are two large ore districts in the Příbram ore area: the Březové Hory base-metal ore district, and the Příbram U and base-metal district. Both districts are located within the exocontact of the Permo-Carboniferous Central Bohemian

Fig. 1. Brodtkorbite (b) in association with tiemannite (t), eucairite (e), uraninite (u) and berzelianite (c). Backscatter electron (BSE) image. Horizontal field of view 1200  $\mu$ m, photo J. Sejkora, P. Škácha.

Plutonic Complex with the Barrandian sedimentary unit (Ettler et al., 2010). In the Příbram U and base-metal district, there are four main mineralization stages (1) sideritesulphidic; (2) calcite; (3) calcite–uraninite; and (4) calcitesulphidic. The selenide mineralization occurs in close association with uraninite of the calcite-uraninite mineralization, but selenides are always younger than uraninite. It is not certain whether selenides in Příbram originated at the conclusion of the calcite-uraninite stage or at the beginning of the subsequent calcite-sulphidic stage, which is mentioned by some authors. Uranium ore is represented by uraninite, coffinite and U-bearing antraxolite (pyrobitumen). The age of the U mineralization, obtained by U-Pb radiometric age determination of uraninite, is middle Permian,  $275 \pm 4$  and  $278 \pm 4$  Ma (Anderson, 1987). The Se-U mineralization is of low-temperature hydrothermal origin and is associated with a calcite vein tens of centimetres to several metres thick. The main ore is represented by uraninite, whereas younger antraxolite predominates in deeper parts of the deposit.

The most important source of selenide mineralization in Příbram is the Bytíz deposit, which is the predominant U deposit in Příbram, with 52% of the overall U production from the whole district. Twenty-two primary selenides species were described here in the last few years (Litochleb *et al.*, 2004; Škácha & Sejkora, 2007; Škácha *et al.*, 2009, 2014, 2015, 2016a; Škácha, 2015), including the two new minerals příbramite (Škácha *et al.*, 2017a, this issue) and bytízite (Škácha *et al.*, 2016b).

The hand specimen containing brodtkorbite was found in a dump carrying material from shaft #16 (49 67'33.028" N, 14 05'89.919" E), located in the Háje deposit. Shaft #16 opened mainly the middle and deeper parts of the vein system of the Háje, Bytíz and Jerusalem deposits (from 500 m down to a depth of 1800 m below the surface). The majority of the selenide-bearing material is probably derived from the Bytíz deposit (Škácha *et al.*, 2016a).

#### 3. Physical properties

In Příbram, brodtkorbite was identified in the only fragment of carbonate gangue carrying abundant tiemannite, berzelianite, umangite and eucairite (Fig. 1). The only specimen containing brodtkorbite consists of up to 3 cm thick calcite vein, with hydrothermally altered and hematitized country rock (sample dimensions:  $8 \times 5 \times 4$ cm). Very thin green Cu secondary films (probably malachite) are present on the surface of calcite and its fractures in the proximity of selenides. Brodtkorbite forms anhedral aggregates up to 150 µm across, in intergrowths with tiemannite and eucairite, and it contains spheroidal uraninite inclusions (Fig. 2). Brodtkorbite is in part older than tiemannite, which replaces it along fractures (Fig. 3). Relatively abundant myrmekitic intergrowths of tiemannite and umangite, observed in some samples, are most probably brodtkorbite breakdown products (Fig. 4).

#### 4. Optical properties

Samples for microscopy and other methods were prepared as polished sections using standard diamond-polishing techniques. Qualitative optical properties of brodtkorbite in reflected light were studied using a Nikon Eclipse ME600 microscope (National Museum, Prague) with a Nikon DXM1200F digital camera. Brodtkorbite from Příbram has similar properties to the type material described by Paar *et al.* (2002). It is cream-white, somewhat darker than the associated tiemannite, and moderately pleochroic with light pink or bluish tints. Internal reflections were not observed. It is strongly anisotropic, with rotation tints varying from violet-grey to yellow-grey under crossed polarisers.

Quantitative reflectance values were obtained with a Leica microscope (objective  $50 \times$ ) equipped with a J & M Tidas MSP 400 spectrophotometer. After careful levelling of both specimen and standard (Zeiss, WTiC), measurements were taken at intervals of 5 nm from 400 to 700 nm in air. Table 1 summarizes the data collected from the most bireflectant grains of the mineral. Data of brodtkorbite from Příbram (Fig. 5) closely match the published data for Tumiñico brodtkorbite (Paar *et al.*, 2002).

#### 5. Chemical composition

Chemical analyses were performed using a Cameca SX100 electron microprobe (National Museum, Prague) operating in wavelength-dispersive mode (25 kV, 20 nA and 2  $\mu$ m wide beam). The following standards and X-ray lines were used to minimize line overlaps: Ag (AgL $\alpha$ ), Au (AuM $\alpha$ ), Bi (BiM $\beta$ ), CdTe (CdL $\alpha$ ), Co (CoK $\alpha$ ), chalcopyrite (CuK $\alpha$ ), FeS<sub>2</sub> (FeK $\alpha$ , SK $\alpha$ ), HgTe (HgM $\alpha$ ), NiAs (NiK $\alpha$ , AsL $\alpha$ ), PbS (PbM $\alpha$ ), PbSe (SeL $\alpha$ ), PbTe (TeL $\alpha$ ), Sb<sub>2</sub>S<sub>3</sub> (SbL $\alpha$ ), Tl (BrI) (TlL $\alpha$ ) and ZnS (ZnK $\alpha$ ). Peak counting times were 20 s for all elements, and one half of the peak time for each background. Other elements, such as Au, Bi, Cd, Co, Fe, Ni, As, Pb, Sb, Tl and Zn, were found to be below the detection





Fig. 2. (a) Brodtkorbite aggregate with bluish tint with tiemannite (yellowish) and dark spheroidal uraninite; reflected-light photomicrograph (one polariser); (b) the same brodtkorbite aggregate under crossed polarisers. Horizontal field of view 210 µm, photos P. Škácha.



Fig. 3. Brodtkorbite (dark grey) partly replaced by later tiemannite (white) in association with spheroidal uraninite. BSE image. Horizontal field of view  $300 \,\mu$ m, photo J. Sejkora, P. Škácha.

limits (0.02–0.05 wt%). Raw intensities were converted to the concentrations of elements using the "*PAP*" (Pouchou & Pichoir, 1985) matrix-correction software.

The chemical composition of brodtkorbite from Příbram (Table 2) corresponds to the ideal formula Cu<sub>2</sub>HgSe<sub>2</sub>. Only minor contents of Ag ( $\leq 0.06 apfu$ ), S ( $\leq 0.01 apfu$ ) and Te ( $\leq 0.003 apfu$ ) were observed. Although a minor Fe content was reported for the type specimen (Paar *et al.*, 2002), no Fe was found in the Příbram material. Its average empirical formula (14 analyses) based on 5 apfu is (Cu<sub>1.87</sub>Ag<sub>0.02</sub>)<sub> $\Sigma 1.89$ </sub>Hg<sub>1.07</sub> (Se<sub>2.03</sub>S<sub>0.01</sub>)<sub> $\Sigma 2.04$ </sub>.

## 6. X-ray powder diffraction data

An X-ray powder diffraction pattern of brodtkorbite was obtained from hand-picked grains using a Bruker D8 Advance diffractometer equipped with solid-state 1D



Fig. 4. Myrmekitic aggregates of tiemannite (white) and umangite (violet), probably a product of brodtkorbite breakdown. Horizontal field of view  $210 \,\mu$ m, reflected-light photomicrograph (one polariser) J. Sejkora.

LynxEye detector using CuK $\alpha$  radiation and operating at 40 kV and 40 mA (National Museum, Prague). In order to minimize the background, the powdered sample was placed on the surface of a flat silicon wafer from an acetone suspension. The powder pattern was collected using Bragg–Brentano geometry in the range 4–70° 2 $\theta$ , in 0.01° steps with a counting time of 30 s/step (total duration of experiments was *ca.* 3 days). The positions and intensities of diffractions were found and refined using the Pearson VII profile-shape function of the ZDS program package (Ondruš, 1993) and the unit-cell parameters were refined by the least-squares program of Burnham (1962).

The X-ray powder diffraction pattern for brodtkorbite from Příbram (Table 3) is similar to that of the type material from Tumiñico (Paar *et al.*, 2002), obtained with a 114.6-mm

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 $\lambda$  (nm)  $\lambda$  (nm)  $R_{\rm min}/R_{\rm max}$  $R_{\rm min}/R_{\rm max}$ 29.4/37.4 400 27.7/34.1 560 29.9/37.9 420 27.4/33.8 580 30.0/37.9 440 27.2/33.7 589 29.7/37.5 460 27.1/33.6 600 29.4/36.9 470 26.7/33.5 620 29.1/36.4 480 26.5/33.4 640 500 28.6/35.6 26.4/33.4 650 28.2/34.9 520 26.4/33.3 660 27.8/34.4 540 26.5/33.1 680 546 700 27.8/34.4 26.5/32.7

Table 1. Reflectance data of brodtkorbite from Příbram.

Debye–Scherrer camera; better sensitivity and resolution of data from the diffractometer led to additional diffraction peaks, yielding refined unit-cell parameters for Příbram brodtkorbite of *a* 7.5137(5), *b* 4.1840(2), *c* 8.0183(6) Å,  $\beta$  124.394(3)°, *V*208.01(4) Å<sup>3</sup>, in excellent agreement with the single-crystal data.

#### 7. Single-crystal X-ray diffraction

A  $0.015 \times 0.011 \times 0.011$  mm large crystal fragment of brodtkorbite extracted from the polished section was used for intensity-data collection on a Rigaku (Oxford Diffraction) SuperNova single-crystal diffractometer equipped with an Atlas S2 CCD detector utilising Mo $K\alpha$ radiation provided by the micro-focus X-ray tube and monochromatised by primary mirror optics. According to the single-crystal X-ray diffraction experiment, the studied brodtkorbite is monoclinic with a = 7.271(3), b = 4.2029(13), c = 7.522(4) Å,  $\beta = 114.19(6)^{\circ}$ , which is very similar to the cell in the setting proposed by Paar et al. (2002). This cell was later transformed to the conventional cell settings in a monoclinic space group  $P2_1/c$  using the transformation matrix (101/010/001) resulting in:  $a = 7.505(1), b = 4.1851(9), c = 8.013(1) \text{ Å}, \beta = 124.43$ (2)° with V=207.59(8)Å<sup>3</sup>, Z=2. The  $\omega$  rotational scans (frame width 1.0°, counting time 300 s) were adopted for the acquisition of the three-dimensional intensity data. From the total of 2777 reflections, 460 were independent and 416 classified as unique and observed with  $I > 3\sigma(I)$ . Corrections for background, Lorentz effects and polarisation were applied during the data reduction using the Crysalis package supplied by Rigaku. The empirical correction for absorption (spherical-harmonics) was applied leading to the dataset with  $R_{int}$  of 0.068.

The structure of brodtkorbite was solved by the chargeflipping algorithm implemented in the SHELXT program (Sheldrick, 2015). The solution provided a complete structure, which was subsequently refined by full-matrix least-squares using Jana2006 (Petříček *et al.*, 2014) based on  $F^2$ . All atoms were refined with anisotropic atomic displacement parameters. The refinement for 27 parameters converged smoothly to the final R = 0.0811 and wR = 0.1729 for 416 unique observed reflections. Details of data collection, crystallographic data and refinement



are given in Table 4. Atom coordinates, displacement parameters and tentative bond-valence sums are listed in Table 5; selected interatomic distances are provided in Table 6.

# 8. Crystal structure of brodtkorbite and related structures

The crystal structure of brodtkorbite contains one special position of Hg and two general positions of Cu and Se. The site-scattering experiment of using Ag entering the Cu site was tested but did not return physically realistic values, probably due to the limited quality of the diffraction data. As is indicated in Fig. 6, the Cu site displays a quasi-tetrahedral coordination by Se atoms with a (3+1) bonding scheme. The Cu site has three close contacts with Se atoms at distances ranging from 2.415 to 2.531 Å forming a trigonal base of the deformed tetrahedron. The coordination is completed by the fourth Se atom at a distance of 2.891 Å. The distortion of the [CuSe<sub>4</sub>] tetrahedron is also reflected by a variation of Se–Cu–Se angles ranging from 79.2° to 119.0°, instead of being 109.47° for a regular tetrahedron. In this particular tetrahedron, the Cu atom is in a very off-centred position, with the distance between the Cu atom to the centre of the tetrahedron being 0.53 Å. The Hg atoms display typical linear coordination by two Se atoms at distances of 2.479 Å. Their coordination is further completed by four long Hg-Se distances at 3.397 Å and 3.420 Å, so the complete coordination sphere of Hg can be described as a strongly flattened octahedron. There are no short Cu-Hg interactions (<3.3 Å) in the structure of brodtkorbite.

The crystal structure of brodtkorbite is composed of layers of edge-sharing [CuSe<sub>4</sub>] distorted tetrahedra running parallel to (100) (Fig. 7a). It resembles the layers of [FeS<sub>4</sub>] tetrahedra observed in the mackinawite crystal structure (Makovicky, 2006). However, while the [FeS<sub>4</sub>] tetrahedra in the mackinawite structure are almost perfectly regular (Lennie *et al.*, 1995), the [CuSe<sub>4</sub>] tetrahedra in brodtkorbite are quite distorted. The CuSe layers show an AA-type stacking and are connected by linear Se–Hg–Se bonds. Each Cu atom also shows three short metal–metal contacts with



	Mean	1	2	3	4	5	6	7	8	9	10	11	12	13	14
Ag	0.49	0.20	0.27	0.18	0.68	0.17	0.20	0.41	1.26	0.19	0.94	0.33	1.35	0.36	0.37
Hg	43.08	42.89	44.38	43.71	46.32	43.69	43.83	44.58	44.60	42.41	42.72	40.61	42.71	40.75	39.92
Cu	23.86	22.72	23.01	22.74	22.85	23.11	23.10	22.40	23.42	23.90	24.62	25.23	24.64	25.74	26.51
Se	32.25	32.82	32.52	32.11	31.65	32.57	32.42	30.81	32.02	32.49	32.02	32.75	31.78	32.90	32.71
Те	0.04	0.11	0.07	0.07	0.00	0.07	0.00	0.00	0.00	0.06	0.07	0.07	0.00	0.00	0.00
S	0.05	0.00	0.07	0.00	0.07	0.05	0.05	0.08	0.08	0.05	0.07	0.08	0.05	0.06	0.06
Total	99.78	98.74	100.33	98.82	101.57	99.66	99.60	98.28	101.38	99.10	100.45	99.06	100.52	99.80	99.57
Ag	0.023	0.010	0.013	0.008	0.031	0.008	0.009	0.020	0.058	0.009	0.043	0.015	0.061	0.016	0.017
Hg	1.068	1.081	1.106	1.107	1.155	1.092	1.097	1.144	1.100	1.054	1.047	0.992	1.046	0.986	0.961
Cu	1.867	1.807	1.810	1.818	1.799	1.824	1.825	1.815	1.824	1.875	1.904	1.946	1.906	1.966	2.014
Se	2.032	2.100	2.058	2.066	2.005	2.069	2.061	2.009	2.006	2.052	1.993	2.032	1.978	2.023	1.999
Те	0.001	0.003	0.003	0.000	0.000	0.000	0.000	0.000	0.000	0.003	0.003	0.003	0.000	0.000	0.000
S	0.009	0.000	0.012	0.000	0.010	0.008	0.008	0.013	0.012	0.007	0.011	0.012	0.008	0.009	0.010

Table 2. Chemical composition of brodtkorbite from Příbram (wt%).

Coefficients of empirical formula were calculated on the basis of 5 apfu.

Table 3. X-ray powder diffraction pattern of brodtkorbite from Příbram.

h	k	l	d <sub>obs.</sub>	I <sub>obs.</sub>	d <sub>calc.</sub>	$d^*_{obs.}$	$I_{obs.}^{*}$	h	k	l	d <sub>obs.</sub>	Iobs.	d <sub>calc.</sub>
1	0	0	6.201	4.6	6.200	6.259	10	2	1	2	1.6605	1.4	1.6604
1	0	2	4.007	100.0	4.006	3.991	70	0	0	4	1.6541	4.3	1.6541
1	1	1	3.571	57.8	3.571	3.576	50	2	2	3	1.6374	3.6	1.6376
0	1	1	3.537	37.8	3.536	3.534	50	1	2	3	1.6276	2.1	1.6276
1	1	0	3.468	39.7	3.468	3.476	10	4	1	1	1.6095	3.6	1.6095
2	0	2	3.425	81.3	3.425	3.414	50	3	2	2	1.6038	4.3	1.6039
2	0	0	3.100	8.0	3.100	3.095	20	3	1	1	1.5871	1.8	1.5877
1	1	2	2.894	3.3	2.894	2.887	10	1	2	2	1.5801	2.6	1.5793
2	1	1	2.785	21.8	2.785	2.784	20	4	0	0	1.5500	4.4	1.5500
1	1	1	2.737	82.5	2.737	2.730	100	0	1	4	1.5379	0.3	1.5383
2	1	2	2.650	6.9	2.650	2.644	10	3	1	5	1.4920	7.8	1.4920
3	0	2	2.4979	20.4	2.4980	2.495	30	2	1	5	1.4794	2.1	1.4795
2	1	0	2.4920	20.1	2.4909			3	2	0	1.4704	2.1	1.4702
1	0	2	2.4079	57.5	2.4080	2.401	30	5	0	4	1.4665	1.9	1.4664
2	1	3	2.2280	64.7	2.2276	2.223	70	5	0	2	1.4644	2.0	1.4644
1	1	3	2.2029	22.9	2.2027	2.199	30	2	2	4	1.4468	1.1	1.4468
3	1	2	2.1448	9.1	2.1448	2.140	40	4	1	5	1.4234	4.3	1.4233
0	2	0	2.0921	15.1	2.0920			3	2	4	1.4173	3.8	1.4176
1	1	2	2.0878	34.7	2.0870			5	1	3	1.4143	1.1	1.4143
3	1	1	2.0786	27.2	2.0785	2.072	50	1	0	4	1.4118	1.9	1.4119
3	0	0	2.0668	3.8	2.0667			1	2	4	1.4004	1.5	1.4005
1	2	1	2.0013	4.5	2.0008			4	2	$\overline{2}$	1.3923	2.0	1.3924
0	2	1	1.9946	16.6	1.9947	1.989	20	4	2	3	1.3869	1.6	1.3870
1	2	0	1.9821	6.8	1.9822			2	1	3	1.3709	5.1	1.3710
0	1	3	1.9514	17.7	1.9510	1.947	10	1	3	1	1.3665	4.0	1.3666
3	0	4	1.9278	12.7	1.9276	1.922	5						
1	0	4	1.8855	24.8	1.8853	1.883	20						
1	2	2	1.8542	13.2	1.8544	1.851	15						
2	0	2	1.8089	15.0	1.8090	1.803	20						
2	2	$\overline{2}$	1.7849	1.4	1.7853	1.782	3						
0	2	2	1.7680	7.3	1.7682	1.765	10						
3	1	4	1.7507	6.8	1.7508	1.746	3						
2	2	0	1.7342	9.7	1.7341	1.731	10						
1	1	4	1.7193	0.8	1.7189								
4	0	$\overline{4}$	1.7124	7.5	1.7123	1.711	5						
4	1	3	1.6940	7.3	1.6941	1.689	20						

d<sup>\*</sup><sub>obs.</sub>, I<sup>\*</sup><sub>obs.</sub> – X-ray powder diffraction data for brodtkorbite from Tumiñico (Paar et al., 2002).

Formula	Cu <sub>2</sub> HgSe <sub>2</sub>
Crystal system	Monoclinic
Space group	$P2_1/c$
Unit-cell parameters:	7.5048(12), 4.1851(9),
a, b, c [Å]	8.0131(12)
β [°]	124.430(16)
Unit-cell volume [Å <sup>3</sup> ]	207.59(8)
Ζ	2
Calculated density [g/cm <sup>3</sup> ]	7.769 (for above formula)
Crystal size [mm]	0.015  imes 0.01  imes 0.01
Diffractometer	Rigaku SuperNova with Atlas S2 CCD
Temperature [K]	297
Radiation, wavelength [Å]	MoK $\alpha$ , 0.71073 (50 kV, 30 mA)
$\theta$ range for data collection [°]	5 09-28 65
Limiting Miller indices	$h = -10 \rightarrow 9$ $k = -5 \rightarrow 5$
	$l = -10 \rightarrow 9$
Axis, frame width (°), time per frame (s)	ω, 1.0, 150
Total reflections collected	2777
Unique reflections	490
Unique observed reflections.	416. $[I > 3\sigma(I)]$
criterion	
Absorption coefficient $[mm^{-1}]$ .	64.34: empirical
type	· ···· · · · · · · · · · · · · · · · ·
$T_{\rm min}/T_{\rm max}$	0.00107/0.00825
Data completeness to $\theta_{max}$ (%).	99.28, 0.076
R <sub>int</sub>	,
Structure refinement	Full-matrix least-squares on $F^2$
No of param restraints	27 0 0
constraints	27, 8, 8
R = wR (obs)	0.0805 0.1782
$R \ wR$ (all)	0.0888 0.1811
GOF obs/all	2.78, 2.59
Weighting scheme, weights	$\sigma, w = 1/(\sigma^2(I) + 0.00016I^2)$
Largest diffraction peak and	3.952.34
hole $(e^{-}/Å^3)$	2

Table 4. Summary of data collection conditions and refinement parameters for brodtkorbite.

three adjacent Cu atoms at distances of 2.535 and 2.670 Å, respectively, forming a pseudohexagonal net of metal–metal interactions (Fig. 7b) within one layer.

The crystal structure of brodtkorbite appears to be unique; there are no exact structural analogues between similar chalcogenides. The mackinawite-type layer is a characteristic structural feature for the thalcusite-rohaite series (Makovicky et al., 1980). Tetrahedral coordination of Cu has also been observed in crystal structures of ternary phases from the Tl–Cu–Se system; e.g., in crookesite (TlCu<sub>7</sub>Se<sub>4</sub>; Eriksson *et al.*, 1991) and the synthetic phases  $TlCu_4Se_3$ (Berger, 1987), TlCu<sub>5</sub>Se<sub>3</sub> (Berger et al., 1990) and TlCu<sub>3</sub>Se<sub>2</sub> (Berger & Eriksson, 1990). However, in these compounds, the Tl atoms are coordinated by eight Se atoms in roughly cubic arrangement, whereas Hg atoms in brodtkorbite display a linear coordination. Imiterite, Ag<sub>2</sub>HgS<sub>2</sub> (Guillou et al., 1985), shows a stoichiometry that is analogous to that of brodtkorbite (Cu<sub>2</sub>HgSe<sub>2</sub>); both minerals exhibit the same symmetry  $(P2_1/c)$  and an identical sequence of occupied Wyckoff positions (i.e., they are isopontial structures).



Fig. 6. *Quasi*-tetrahedral coordination of Cu in the brodtkorbite crystal structure showing Cu–Se bond distances. The Cu–Cu interactions are indicated by dashed lines.



Fig. 7. (a) Polyhedral representation of the crystal structure of brodtkorbite emphasising the mackinawite-type layers of edge-sharing [CuSe<sub>4</sub>] tetrahedra. The linear Se–Hg–Se bonds are indicated. (b) A (100) mackinawite-type layer of [CuSe<sub>4</sub>] tetrahedra showing the pseudohexagonal net of Cu–Cu interactions.

However, both crystal structures differ to a certain extent. On one hand, a linear coordination of Hg occurs in both structures and Cu (Ag) atoms display a 3 + 1 bonding scheme by Se (S) atoms, respectively. On the other hand, the imiterite structure shows a different topology of metal–chalcogen bonds and lacks the system of metal–metal interactions found in brodtkorbite. Within the experimental study of the Hg–Cu–S system by Ollitrault-Fichet *et al.* (1984), a partial solid solution along the HgS–Cu<sub>2</sub>S line with a limiting composition of Cu<sub>2.15</sub>Hg<sub>0.92</sub>S<sub>2</sub> was described. This phase is chemically very close to the brodtkorbite. However, this solid solution shows cubic symmetry and is only stable at high temperatures (>416 °C). It is not quenchable.

## 9. Raman spectroscopy

Raman spectra of brodtkorbite from Příbram were collected in the range 25–1800 cm<sup>-1</sup> using a DXR dispersive Raman Spectrometer (Thermo Scientific) mounted on a confocal Olympus microscope (National Museum, Prague). The Raman signal was excited by a NIR 780 nm diode-pumped solid-state laser and detected by a CCD detector. The experimental parameters were:  $50 \times$  objective, 10 s exposure, 100 exposures, 830 lines/mm grating, 50 µm slit spectrograph aperture, and 0.2 mW laser power level. The

Atom	x		v	Ζ	$U_{ m eq}$	BV	
Hg	0	0.5		0.5	0.0595(8)	1.951(8)	
Se	-0.2715(3)	0.8416(6)		0.2152(3)	0.0448(11)	1.697(8)	
Cu	-0.4707(6)	0.9080(10)		0.3666(5)	0.0675(19)	0.721(6)	
	$U^{11}$	$U^{22}$	$U^{33}$	$U^{12}$	$U^{13}$	$U^{23}$	
Hg	0.0537(9)	0.0656(11)	0.0497(9)	0.0100(6)	0.0234(7)	0.0093(6)	
Se	0.0445(12)	0.0489(14)	0.0418(12)	0.0002(9)	0.0247(9)	0.0004(9)	
Cu	0.080(2)	0.077(2)	0.060(2)	0.0206(19)	0.0486(18)	0.0113(17)	

Table 5. Atom positions, displacement parameters  $(Å^2)$  and bond-valence (BV) sums for brodtkorbite.

Table 6. Interatomic distances (Å) in brodtkorbite and imiterite (Guillou *et al.*, 1985).

	Brodtkorbite		Imiterite
Hg–Se	2.479(2) (2×)	Hg–S	2.3759 (2×)
Cu–Se	2.415(6) 2.456(5) 2.891(5) 2.531(5)	Ag–S	2.495 2.5996 2.6539
Cu–Cu	2.671(6) (2×) 2.534(7)		

instrument was set up by a software-controlled calibration procedure using multiple Ne emission lines (wavelength calibration), multiple polystyrene Raman bands (laserfrequency calibration) and standardized white-light sources (intensity calibration). Spectral manipulations were performed using the Omnic 9 software (Thermo Scientific).

The dominant feature of the brodtkorbite spectrum (Fig. 8, Table 7) is the presence of intense vibration bands in the  $220-170 \text{ cm}^{-1}$  region. According to the literature (Nakamoto, 2009; Kharbish et al., 2007, 2009; Škácha et al., 2014; Števko et al., 2015), these bands can be assigned to stretching vibrations of the CuSe<sub>4</sub> tetrahedra and Hg-Se bonds. As the CuSe<sub>4</sub> tetrahedra are distorted with different internal bond-lengths and form complex CuSe layers connected by linear Se-Hg-Se bonds, we were not able to determine and assign the precise character of the observed vibration bands. The most intense band at  $189 \text{ cm}^{-1}$  can be connected with vibrations of Hg–Se bonds, similar to that observed in tiemannite, HgSe  $(191 \text{ cm}^{-1}, \text{ our study}; 200-220 \text{ cm}^{-1}, \text{ RRUFF R050550},$ R060463). The vibrations connected with Hg–S bonds are observed at  $275-261 \text{ cm}^{-1}$  (grumiplucite, HgBi<sub>2</sub>S<sub>4</sub>, Števko *et al.*, 2015), 283 cm<sup>-1</sup> (imiterite, Ag<sub>2</sub>HgS<sub>2</sub>, our study, Fig. 9) or 255-290 cm<sup>-1</sup> (cinnabar, HgS, Frost et al., 2002). A substantial decrease of frequencies from X–S to X–Se bonds (about  $100 \text{ cm}^{-1}$  and more), according to Hook's rule, is described, e.g., for Nb(S/Se)<sub>4</sub> and Mo( $\overline{S}$ / Se)<sub>4</sub> tetrahedra (Nakamoto, 2009) or for famatinite Cu<sub>3</sub>SbS<sub>4</sub> and permingeatite Cu<sub>3</sub>SbSe<sub>4</sub> (Škácha et al., 2014).

Table 7. Band-components of the Raman spectrum of brodtkorbite.

Position/rel. intensity/FWHM (cm <sup>-1</sup> /%/cm <sup>-1</sup> )	Tentative assignment
378/2/16	Overtones
197/13/6 189/100/3	Cu–Se and Hg–Se stretching vibrations
147/19/8 103/7/5	Cu–Se and Hg–Se bending vibrations
91/3/6 83/2/6 75/2/6 67/3/8 55/9/13 42/20/8 37/22/5	External modes (transl- ations and librations)

The less intense bands of brodtkorbite in the region  $160-100 \text{ cm}^{-1}$  can be assigned to bending vibrations of Hg–Se and Cu–Se bonds and those below  $100 \text{ cm}^{-1}$  should correspond to external modes (translations and librations). The broad band at  $378 \text{ cm}^{-1}$  probably belongs to overtones.

#### **10.** Conditions of formation

Formation temperatures of the selenide mineralization of the Příbram area are similar to those at other selenide occurrences in the Czech Republic at around 100 °C. Berzelianite forms at temperatures exceeding 112 °C, and it alters to umangite at lower temperatures (Simon & Essene, 1996). Both minerals are relatively abundant in the studied selenide association. Locally, athabascaite occurs and indicates temperatures below 100 °C (Harris *et al.*, 1970). Results of a fluid-inclusion study in carbonates of the uraninite–antraxolite stage indicate crystallization temperatures in the range 140–80 °C (Žák & Dobeš, 1991). Similar temperatures (130–60 °C), based on a study of selenide mineralization and fluid inclusions, were obtained from the Tumiñico Se deposit, Argentina (Paar *et al.*, 2002).

Parasyuk *et al.* (1999) summarized the experimental study of the Cu<sub>2</sub>Se–HgSe *quasi*-binary system. The system is of the eutectic type, and no ternary phases were detected in







Fig. 9. Raman spectrum of imiterite, mine Lill, Příbram, Czech Republic;  $Ag_{2.04}Hg_{1.01}As_{0.01}S_{1.94}$ , monoclinic, *a* 4.038(1), *b* 8.006(1), *c* 6.582(1)Å,  $\beta$  107.09(2)° (Plášil *et al.*, 2009).

the performed experiments. It is worth mentioning that the solubility region of Cu<sub>2</sub>Se ( $\alpha$ -solid solution) extends to 54 mol% HgSe at the eutectic temperature (715 °C). The solubility diminishes as the temperature decreases. The non-existence of the ordered ternary phase at temperatures close to the eutectic point also supports the formation of brodtkorbite at low temperatures.

The abundance of calcite in the gangue indicates a neutral to weakly alkaline environment of these hydrothermal fluids (Dymkov, 1985; Kvaček, 1987). The selenide mineralization is formed from hydrothermal fluids at conditions of  $fO_2$  above the hematite–magnetite buffer, probably in the range of 5.8 log units (hematite–ferroselite univariant reaction) to 7 log units (anglesite–galena buffer) above the hematite–magnetite buffer (Simon *et al.*, 1997). Such high  $fO_2$  values result in geochemical separation of Se from S in the hydrothermal fluids, a high Se<sub>2</sub>/S<sub>2</sub> fugacity ratio and in the deposition of various selenide minerals. The quantity and variability of the selenide association which formed under oxidising hydrothermal conditions are a function of  $fSe_2$  and reflects the concentration of available Se in solutions (Simon *et al.*, 1997).

The selenide mineralization in the Příbram U district shows variations in  $fSe_2$  and  $fS_2$  (Se, -12 to -18; S, -17to -24), which are similar to those published in detail (Simon *et al.*, 1997) for "selenide-bearing uncorformityrelated uranium deposits". Considering the absence of krutaite in the association we can expect an origin of the Se mineralisation at conditions with lower  $fSe_2$  than those of the krutaite stability field, which according to the stability diagram at 100 °C (Simon *et al.*, 1997), corresponds to log $fSe_2 < -12$  (Škácha *et al.*, 2017b).

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