

# Sbacchiite, $\text{Ca}_2\text{AlF}_7$ , a new fumarolic mineral from the Vesuvius volcano, Napoli, Italy

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**Abstract:** The new mineral sbacchiite (IMA 2017-097),  $\text{Ca}_2\text{AlF}_7$ , was found in a fossil fumarole (1944 eruption,  $T \approx 80^\circ\text{C}$ ) at the rim of the crater of the Vesuvius volcano, Napoli, Italy, associated with gearsksutite, usovite, creedite and opal. It forms elongated crystals up to about 60  $\mu\text{m}$  in length. On the basis of powder X-ray diffraction data and chemical analysis, the mineral was recognized to be identical to the corresponding synthetic phase. Crystals are transparent or translucent and colourless, with vitreous lustre and white streak. The tenacity is brittle. The measured density is 3.08(2)  $\text{g}/\text{cm}^3$ , the calculated density is 3.116  $\text{g}/\text{cm}^3$ . The empirical formula, (based on 10 atoms per formula unit, *apfu*) is  $\text{Ca}_{2.02}\text{Mg}_{0.03}\text{Al}_{0.99}\text{F}_{6.97}$ . Sbacchiite is orthorhombic, space group *Pnma*, with  $a = 7.665(2)$ ,  $b = 6.993(1)$ ,  $c = 9.566(2)$  Å,  $V = 512.2(2)$  Å<sup>3</sup> and  $Z = 4$ . The eight strongest X-ray powder diffraction lines are [ $d_{\text{obs}}$  Å( $I$ )( $hkl$ )]: 3.840(45)(200), 3.563(85)(201), 3.499(100)(020), 2.899(55)(013), 2.750(30)(212), 2.281(20)(104), 2.255(52)(302) and 2.173(36)(131). The structure was refined to  $R = 0.0479$  for 457 reflections with  $I > 2\sigma(I)$ . The asymmetric unit contains one  $\text{Al}^{3+}$  and two independent  $\text{Ca}^{2+}$  cations and five fluorine anions. Aluminium is octahedrally coordinated by six fluorine atoms; the arrangement of F around the 7-coordinated Ca(1) conforms to a distorted pentagonal bipyramid and that around Ca(2) to a very distorted polyhedron (in 7 + 1 coordination). All the fluorine atoms are threefold coordinated. The structure framework shows “isolated”  $[\text{AlF}_6]$  octahedra, whereas the coordination polyhedra around Ca are linked by common edges [sequence: Ca(1)–Ca(2)–Ca(1)...] along [010] and the same holds for the connection along [001]. Along [100], however, only the pentagonal bipyramids around Ca(1) are connected by bridging corners.

**Key-words:** sbacchiite; new mineral; calcium aluminium fluoride; crystal structure; fumaroles; Vesuvius volcano.

## 1. Introduction

The Somma-Vesuvius volcanic complex is one of the most studied volcanoes in the World. This area is potentially very dangerous for the more than 800 000 inhabitants living on the slopes of the Vesuvius and for this reason detailed studies in various fields of geology and geophysics have been carried out in this region, which is subject to continuous monitoring. However, little is known about the mineralogical phases that have formed or are forming in the fumaroles after the last eruption that occurred in 1944. The whole fumarolic area was studied by [Parascandola \(1951\)](#) between 1948 and 1960, when temperatures reached a maximum of about 800 °C in 1950, up to the 60s of the last century, when temperatures settled around 460 °C ([Parascandola, 1960, 1961](#)). No quantitative information about the composition of the fumarolic fluxes at that time is reported, although the presence of HCl, Cl<sub>2</sub> and H<sub>2</sub>S was indicated by Parascandola. On the basis of the mineral phases formed, the presence of SO<sub>2</sub> and HF was nevertheless conceivable. Currently, temperatures of the fumaroles at the eastern rim

of the crater are between 70 °C and 80 °C, and their chemistry reflects the composition of the air so that, in such conditions, there is no deposition of minerals. On the contrary, the chemical composition of the fumaroles of the crater bottom is of hydrothermal type, with the typical presence of H<sub>2</sub>S ([Chiodini et al., 2001](#)). We report here the description of the new mineral sbacchiite,  $\text{Ca}_2\text{AlF}_7$ , recently discovered in a fumarole at the rim of the crater. The mineral was recognized, on the basis of X-ray powder diffraction measurements and chemical analysis, to be identical to the corresponding  $\text{Ca}_2\text{AlF}_7$  synthetic phase ([Domsele & Hoppe, 1980](#)). The mineral and its name were approved by the IMA Commission on New Minerals, Nomenclature and Classification (IMA 2017-097). Sbacchiite is pronounced “sbaki-ite” and named after Dr. Massimo Sbacchi (b. 1958), biologist and mineral collector, for his long-time field collaboration on fumarolic minerals and for his continuous supply of interesting material for study. A specimen of the holotype material is deposited in the Reference Collection of the Dipartimento di Chimica, Università degli Studi di Milano, catalogue number 2017-01 and a co-type

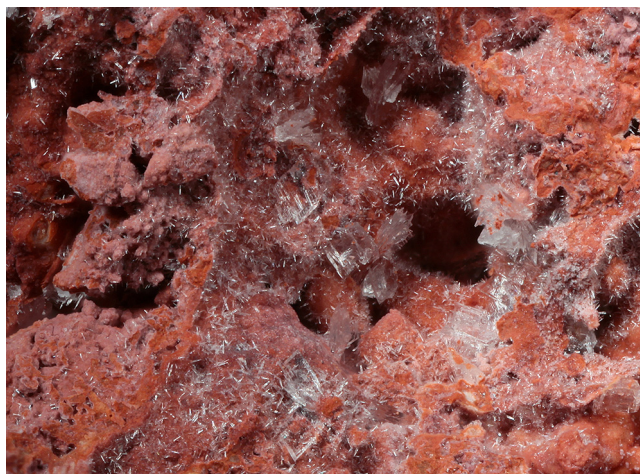


Fig. 1. Needle-shaped sbacchiite crystals with platy crystals of gearsksutite on red volcanic breccia (base width 3 mm).

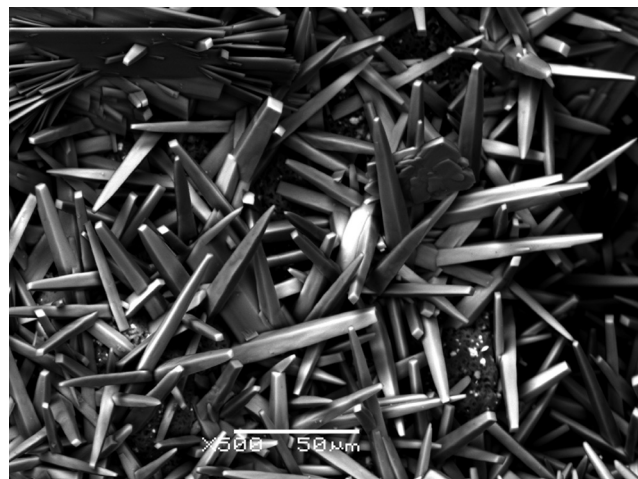


Fig. 2. Backscatter-electron image of sbacchiite crystals.

specimen in the collection of the Museum of Osservatorio Vesuviano (Ercolano, Napoli), catalogue number 2018-01.

## 2. Occurrence

The new mineral sbacchiite was found in a fossil fumarole at the eastern rim of the crater in the so-called “cotunnite pit” (latitude 40° 49′ 21.98″ N, longitude 14° 25′ 43.66″ E), where other rare minerals were also discovered such as artroite  $\text{Pb}[\text{AlF}_3(\text{OH})_2]$  (Campostrini & Gramaccioli, 2005), ammineite  $[\text{CuCl}_2(\text{NH}_3)_2]$  (Russo & Campostrini, 2011), fluorinatrocoulsellite  $\text{CaNa}_3\text{AlMg}_3\text{F}_{14}$  (Russo *et al.*, 2014), and parascandolaite  $\text{KMgF}_3$  (Demartin *et al.*, 2014). The formation of sbacchiite certainly took place between 1948 and 1960 or shortly thereafter (see below). It is not a real sublimate, but rather a high-temperature encrustation, as HF attacked the rock extracting aluminium and calcium. The mineral is extremely rare and was found in only one specimen of approximate diameter 7 cm, which was trimmed into five smaller samples. It forms small aggregates of crystals up to about 60  $\mu\text{m}$  in length, with a very steep bipyramidal habit, elongated along [100] and truncated by the (100) pinacoid (Figs. 1 and 2). Associated minerals are gearsksutite, usovite, creedite and opal.

## 3. Physical and optical properties

Crystals of sbacchiite are transparent or translucent and colourless, with vitreous lustre and white streak. The tenacity is brittle. No distinct cleavage is observed. The mineral does not fluoresce in long- or short-wave ultraviolet light. No twinning is apparent. The density measured by flotation in a diiodomethane/benzene solution is 3.08(2)  $\text{g}/\text{cm}^3$ , that calculated using the empirical formula and unit-cell data is 3.116  $\text{g}/\text{cm}^3$ . Due to the minute size of the crystals the Mohs hardness could not be determined. Optically, sbacchiite is biaxial (+), with  $\alpha = 1.379(4)$ ,  $\beta = 1.384(4)$ ,  $\gamma = 1.390(4)$  (measured in white light). The  $2V$  measured on a spindle

Table 1. Analytical data (in wt%) for sbacchiite (average of 10 analyses).

Constituent	Mean	Range	Stand. dev.	Probe standard
F	54.67	54.06–55.22	0.51	Fluorite
Al	10.97	10.78–11.14	0.10	Spinel
Ca	33.41	32.98–34.57	0.32	Fluorite
Mg	0.26	0.17–0.30	0.05	Spinel
Total	99.31			

Notes: The empirical formula is  $(\text{Ca}_{2.02}\text{Mg}_{0.03})_{\Sigma 2.05}\text{Al}_{0.99}\text{F}_{6.97}$  (based on 10 *apfu*).

stage is 83(2)°; the calculated  $2V$  is 85.1°. The mean refractive index is 1.384, the one predicted using the empirical electronic polarizabilities of ions (Shannon & Fischer, 2016) is 1.382.

## 4. Chemical analysis

Quantitative chemical analyses (10) were carried out in energy-dispersive mode using a JEOL JSM 5500 LV scanning electron microscope equipped with an IXRF EDS 2000 microprobe (20 kV excitation voltage, 10 pA beam current, 2  $\mu\text{m}$  beam diameter). This analytical method was chosen because crystal intergrowths did not take a good polish and it was impossible to prepare a flat polished sample; moreover, the crystals were severely damaged by using the wavelength-dispersive spectrometry technique, even with a low voltage and current and a large diameter of the electron beam. In this case, as reported by Ruste (1979) and Acquafredda & Paglionico (2004), the EDS detector gives more accurate analyses of small volumes of investigated sample, also with a probe current lower than 1 nA. This method gives good results also when collecting X-rays emitted from a non perfectly flat surface of the specimen. X-ray intensities were converted to wt% by ZAF quantitative-analysis software. The standards employed were fluorite for Ca and F and spinel for Al and Mg.

Table 2. X-ray powder-diffraction data for sbacchiite and comparison with the synthetic analogue.

Sbacchiite*		Sbacchiite**		Synthetic Ca <sub>2</sub> AlF <sub>7</sub> (Domsele & Hoppe, 1980)			
<i>d</i> (Å) (calc.)	<i>I</i> / <i>I</i> <sub>0</sub> (calc.)	<i>d</i> (Å) (obs.)	<i>I</i> / <i>I</i> <sub>0</sub> (obs.)	<i>d</i> (Å) (obs.)	<i>d</i> (Å) (calc.)	<i>I</i> / <i>I</i> <sub>0</sub> (obs.)	<i>h k l</i>
5.982	5	5.986	6	5.9910	5.9871	7	1 0 1
4.546	3	4.550	2	4.5483	4.5493	3	1 1 1
3.833	38	3.840	45	3.8412	3.8426	40	2 0 0
3.558	72	3.563	85	3.5651	3.5648	90	2 0 1
3.497	100	3.499	100	3.4983	3.4989	100	0 2 0
3.361	3	3.366	5	3.3675	3.3682	3	2 1 0
3.171	10	3.174	15	3.1762	3.1764	13	2 1 1
3.018	1	3.020	2	3.0193	3.0209	3	1 2 1
2.991	8	2.990	10	2.9950	2.9935	13	2 0 2
2.944	4	2.941	3	2.9413	2.9408	3	1 0 3
2.901	46	2.899	55	2.8973	2.8974	53	0 1 3
2.823	1			2.8220	2.8223	3	0 2 2
2.750	22	2.750	30	2.7521	2.7523	33	2 1 2
2.713	8	2.710	10	2.7121	2.7111	13	1 1 3
2.649	2	2.647	5	2.6493	2.6493	7	1 2 2
2.494	2	2.495	4	2.4976	2.4971	7	2 2 1
2.468	2	2.470	5	2.4747	2.4742	7	3 0 1
2.451	6	2.449	8	2.4518	2.4513	10	2 0 3
2.391	1			2.3880	2.3873	3	0 0 4
2.328	1			2.3341	2.3327	3	3 1 1
2.313	4	2.314	5	2.3135	2.3135	7	2 1 3
2.283	17	2.281	20	2.2808	2.2798	30	1 0 4
2.263	5			2.2662	2.2660	3	0 3 1
2.253	61	2.255	52	2.2578	2.2573	50	3 0 2
2.172	23	2.173	36	2.1736	2.1735	43	1 3 1
2.029	8	2.028	10	2.0282	2.0278	13	2 0 4

\*Pattern calculated on the basis of the single-crystal data and structure refinement.

\*\*Experimental pattern obtained using a Bruker D8 diffractometer (Cu K $\alpha$  radiation).

Element concentrations were measured using the K $\alpha$  lines. The mean analytical results are reported in Table 1. No amounts of other elements above 0.1 wt% were detected. The empirical formula is (Ca<sub>2.02</sub>Mg<sub>0.03</sub>) $\Sigma$ 2.05Al<sub>0.99</sub>F<sub>6.97</sub> (based on 10 apfu).

The simplified formula is Ca<sub>2</sub>AlF<sub>7</sub>, which requires 55.38 F, 11.24 Al and 33.38 Ca (wt%).

## 5. X-ray crystallography and crystal-structure determination

The X-ray powder-diffraction pattern (Table 2), obtained using a conventional Bruker D8 diffractometer, with graphite monochromatized Cu K $\alpha$  radiation, is in good agreement with that of the synthetic phase. The unit-cell parameters refined from powder data using the UNITCELL software (Holland & Redfern, 1997) are  $a = 7.674(1)$  Å,  $b = 6.996(1)$  Å,  $c = 9.553(1)$  Å,  $V = 512.9(1)$  Å<sup>3</sup>.

From a needle-shaped crystal fragment (approximately 0.05 × 0.01 × 0.01 mm), 3655 diffracted intensities, corresponding to a complete scan of the reciprocal lattice up to  $2\theta = 53.88^\circ$ , were collected at room temperature using a Bruker Apex II diffractometer, equipped with a 2K CCD detector and Mo K $\alpha$  radiation ( $\lambda = 0.71073$  Å). A 90s frame-time and a 0.5° frame width were used. The intensity data were reduced using the program SAINT (Bruker, 2001)

Table 3. Single-crystal diffraction data and refinement parameters for sbacchiite.

Crystal system	Orthorhombic
Space group	<i>Pnma</i> (no. 62)
<i>a</i> (Å)	7.665(2)
<i>b</i> (Å)	6.993(1)
<i>c</i> (Å)	9.566(2)
<i>V</i> (Å <sup>3</sup> )	512.2(2)
<i>Z</i>	4
Radiation	MoK $\alpha$
$\mu$ (mm <sup>-1</sup> )	2.47
<i>D</i> <sub>calc</sub> (g cm <sup>-3</sup> )	3.111
Reflections measured	3655
Independent reflections	590
Observed reflections [ $I > 2\sigma(I)$ ]	457
Parameters refined	55
Final <i>R</i> [ $I > 2\sigma(I)$ ], <i>wR</i> 2 (all data)	0.0479, 0.0863
Goof	1.157

Notes:  $R = \Sigma||F_o| - |F_c||/\Sigma|F_o|$ ;  $wR2 = \{\Sigma[w(F_o^2 - F_c^2)^2]/\Sigma[w(F_o^2)^2]\}^{1/2}$ ;  $w = 1/[\sigma^2(F_o^2) + (0.0328q)^2 + 0.8790q]$  where  $q = [\max(0, F_o^2) + 2F_c^2]/3$ ; Goof =  $\{\Sigma[w(F_o^2 - F_c^2)]/(n - p)\}^{1/2}$  where  $n$  is the number of reflections and  $p$  is the number of refined parameters.

and corrected for Lorentz, polarization and background. An absorption correction ( $\mu = 2.47$  mm<sup>-1</sup>,  $T_{\min} = 0.735$ ) was applied using the SADABS program (Sheldrick, 2000). Starting from the atomic positions reported by



Table 4. Atomic coordinates and displacement parameters [ $U_{eq}/U^{ij}$ , Å<sup>2</sup>] for sbacchiite.

Atom	Wyckoff	<i>x</i>	<i>y</i>	<i>z</i>	$U_{eq}$
Al	4c	0.3088(2)	1/4	0.05621(17)	0.0075(4)
Ca1	4c	0.11312(17)	1/4	0.74096(12)	0.0120(3)
Ca2	4c	0.80132(16)	1/4	0.10650(12)	0.0107(3)
F1	4c	0.3717(5)	1/4	0.6299(3)	0.0176(8)
F2	4c	0.0297(4)	1/4	0.5084(4)	0.0142(8)
F3	4c	0.0954(5)	1/4	0.1263(4)	0.0258(10)
F4	8d	0.2486(3)	0.07395(3)	0.92591(2)	0.0158(6)
F5	8d	0.3683(4)	0.0716(3)	0.1842(2)	0.0153(6)

Atom	$U^{11}$	$U^{22}$	$U^{33}$	$U^{23}$	$U^{13}$	$U^{12}$
Al	0.0092(9)	0.0059(8)	0.0074(9)	0	0.0004(7)	0
Ca1	0.0128(7)	0.0109(6)	0.0122(6)	0	0.0014(5)	0
Ca2	0.0132(7)	0.0073(6)	0.0116(6)	0	-0.0033(5)	0
F1	0.017(2)	0.019(2)	0.016(2)	0	-0.004(2)	0
F2	0.008(2)	0.018(2)	0.018(2)	0	-0.005(2)	0
F3	0.011(2)	0.030(2)	0.036(2)	0	0.011(2)	0
F4	0.022(2)	0.008(1)	0.017(1)	-0.001(1)	-0.010(1)	-0.001(1)
F5	0.026(2)	0.008(1)	0.012(1)	0.002(1)	-0.002(1)	0.001(1)

Notes: The anisotropic displacement factor exponent takes the form:  $-2\pi^2(U^{11}h^2(a^*)^2 + \dots + 2U^{12}hka^*b^* + \dots)$ ;  $U_{eq}$  according to Fischer & Tillmanns (1988).

Domsele & Hoppe (1980), the structure was refined anisotropically with the *SHELXL-2017* program (Sheldrick, 2015), implemented in the WinGX program (Farrugia, 1999). The refinement converged to a final  $R = 0.0479$  for 457 observed reflections [ $I > 2\sigma(I)$ ]. Details about the data collection and refinement are summarized in Table 3. Final atom coordinates and anisotropic displacement parameters are reported in Table 4. Selected interatomic distances are reported in Table 5.

## 6. Description of the crystal structure and discussion

Perspective views of the crystal structure of sbacchiite are reported in Figs. 3 and 4. The asymmetric unit contains one Al<sup>3+</sup> and two independent Ca<sup>2+</sup> cations and five fluorine anions. Aluminium is octahedrally coordinated by six fluorine atoms with Al–F distances in the range 1.768(4)–1.812(2) Å. The arrangement of F around Ca(1) (Fig. 5) conforms to a distorted pentagonal bipyramid (coordination number C.N. 7; Ca–F distances in the range 2.224(4)–2.392(2) Å), similar to that observed in jakobssonite CaAlF<sub>5</sub> (Balić-Zunić *et al.*, 2012) and that around Ca(2) to a very distorted polyhedron (C.N. 7 + 1, Ca–F distances in the range 2.262(4)–3.004(4) Å). All the fluorine atoms have C.N. 3; F(1), F(2) and F(3) are surrounded by three cations in a plane, whereas in the case of F(4) and even more so for F(5) the anions are slightly out of plane. The bond-valence analysis (Table 6) confirms the crystal-chemical soundness of the crystal structure and the expected values for the valence of the Ca<sup>2+</sup>, Al<sup>3+</sup> and F<sup>-</sup> ions. The whole structure framework is made up of linkages of “isolated” [AlF<sub>6</sub>] octahedra, [CaF<sub>7</sub>] and [CaF<sub>8</sub>] polyhedra. “Isolated” [AlF<sub>6</sub>] octahedra have also been observed in carlhintzeite Ca<sub>2</sub>AlF<sub>7</sub>·H<sub>2</sub>O (Kampf *et al.*, 2010), where two independent

Table 5. Selected interatomic distances (Å) in sbacchiite.

Al–F2	1.802(4)	Al–F4 (×2)	1.812(2)
Al–F3	1.768(4)	Al–F5 (×2)	1.806(2)
<Al–F>	1.801		
Ca1–F1	2.224(4)	Ca2–F1	2.325(3)
Ca1–F1	2.249(4)	Ca2–F2	2.354(4)
Ca1–F2	2.315(4)	Ca2–F3	2.262(4)
Ca1–F4 (×2)	2.392(2)	Ca2–F3	3.004(4)
Ca1–F5 (×2)	2.318(2)	Ca2–F4 (×2)	2.318(2)
<Ca1–F>	2.315	Ca2–F5 (×2)	2.414(2)
		<Ca2–F>	2.426

Al<sup>3+</sup> cations are present but with a different environment. In carlhintzeite, one Al octahedron is sharing two opposite faces with two Ca polyhedra and the other one is edge-sharing two opposite F–F edges. In sbacchiite, only one face is shared with the adjacent Ca(2) polyhedron and, on the opposite side, an edge and a corner of the same face are shared with two adjacent Ca(1) polyhedra (Fig. 6). The structure of jakobssonite, CaAlF<sub>5</sub> contains instead vertex-sharing chains of [AlF<sub>6</sub>] octahedra, interconnected by chains of [CaF<sub>7</sub>] pentagonal bipyramids. In the other minerals with similar chemical constituents, whose structure is known *i.e.* gearsutite CaAlF<sub>4</sub>(OH)·H<sub>2</sub>O (Marchetti & Perchiazzi, 2000) and prosopite CaAl<sub>2</sub>F<sub>4</sub>(OH)<sub>4</sub> (Giacovazzo & Menchetti, 1969; Pudovkina *et al.*, 1973), the Al octahedra are not isolated but arranged as edge-sharing dimers for the former or as edge-sharing chains for the latter.

In sbacchiite, the coordination polyhedra around Ca are linked by common edges in the sequence Ca(1)–Ca(2)–Ca(1) ... along [010] and the same holds for the connection along [001]. Along [100] only the pentagonal bipyramids around Ca(1) are connected by bridging corners. In carlhintzeite all the Ca cations are instead 8-coordinated, and their linkage can be envisioned as a framework in which

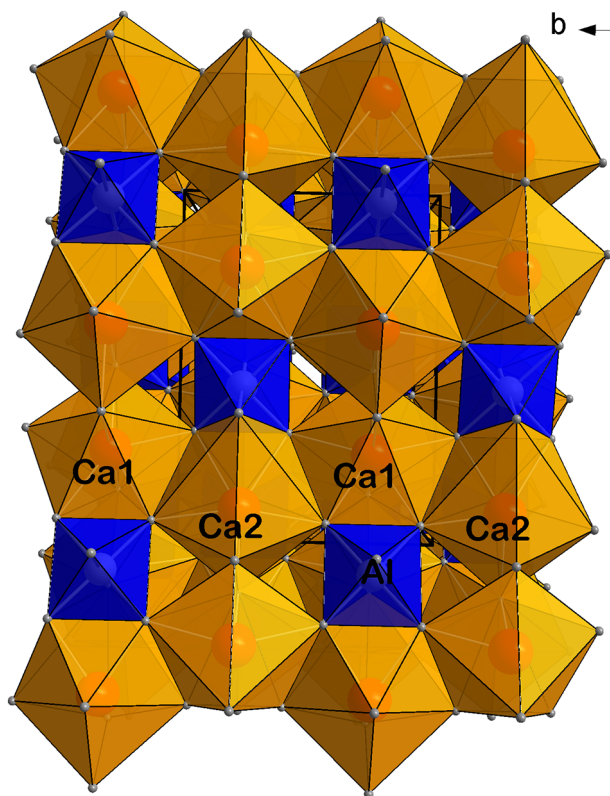


Fig. 3. View of the crystal structure of sbacchiite along [100].

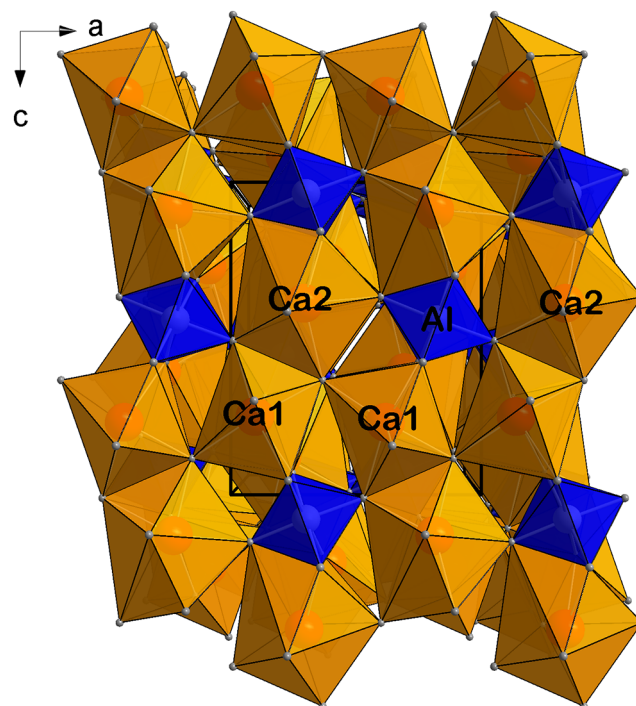


Fig. 4. View of the crystal structure of sbacchiite along [010].

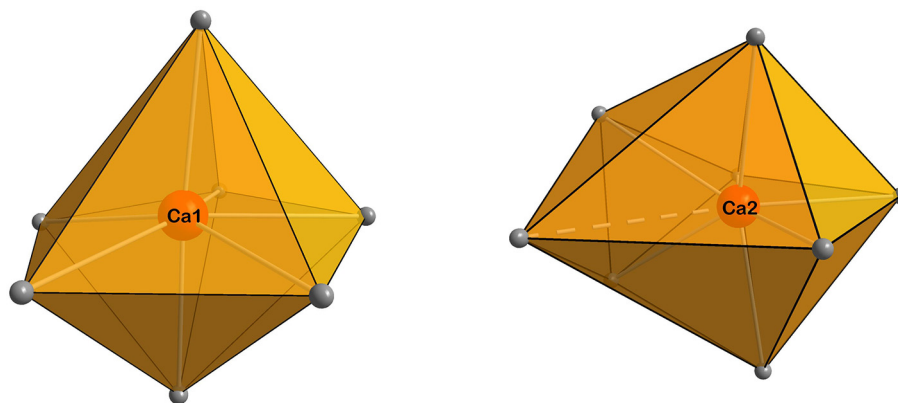
Fig. 5. The coordination polyhedron of the  $\text{Ca}^{2+}$  cations.

Table 6. Bond-valence analysis for sbacchiite.

	F1	F2	F3	F4	F5	$\Sigma c$
Al		0.499	0.547	$0.485 \times 2 \rightarrow 0.485 \downarrow$	$0.491 \times 2 \rightarrow 0.491 \downarrow$	2.998
Ca1	$0.355 \rightarrow 0.333 \downarrow$	0.279		$0.226 \times 2 \rightarrow 0.226 \downarrow$	$0.276 \times 2 \rightarrow 0.276 \downarrow$	1.971
Ca2	0.272	0.247	$0.323 \rightarrow 0.043 \downarrow$	$0.278 \times 2 \rightarrow 0.278 \downarrow$	$0.215 \times 2 \rightarrow 0.215 \downarrow$	1.871
$\Sigma a$	0.960	1.025	0.913	0.989	0.982	

Notes: Bond-valence parameters from [Brown & Altermatt \(1985\)](#), [Brese & O'Keeffe \(1991\)](#). Values are expressed in valence units (vu).

edge-sharing chains along [010] are cross-linked by edge sharing.

As already observed in carlhintzeite and gearsutite, the distortion from the idealized geometry in the Ca polyhedra is clearly related to the way in which these polyhedra are

linked. The lengths of their F··F edges can be correlated with their linkages to the  $[\text{AlF}_6]$  octahedra and to the other Ca polyhedra, the shortest F··F edges corresponding to those shared with the  $[\text{AlF}_6]$  octahedra. It is interesting to point out that in sbacchiite, prosopite, gearsutite and

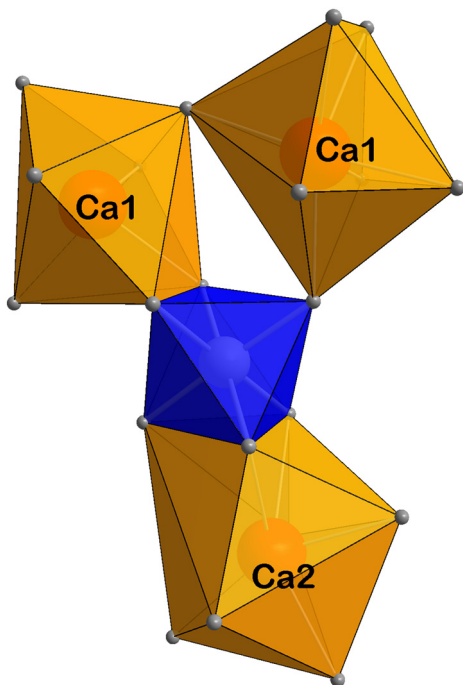


Fig. 6. The environment of the Al site.

carlhintzeite the direction of the chains made by the Ca polyhedra determines a comparable unit-cell parameter. The chains in prosopite and gearsksutite are parallel to the *c* axis and correspond to *c* cell parameters of 7.32 and 6.978(1) Å, respectively. In carlhintzeite the chain direction is parallel to the *b* unit-cell edge of 6.9670(5) Å, and for sbacchiite the corresponding value is 6.993(1) Å.

The discovery of sbacchiite in the fumaroles related to the 1944 eruption adds a new piece to the knowledge of the products formed when temperatures were higher than today, and confirms that fluorides deriving from reaction of HF with the surrounding rocks were the most common phases deposited. Formation of opal, which normally encrusts most of the occurring minerals and derives from leaching of the rock-forming silica and silicates, seems to be the last stage of a solid phase deposition. At present only water vapour is the main constituent of the fumarolic gases and formation of other mineral species is not observed.

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## References

- Acquafredda, P. & Paglionico, A. (2004): SEM-EDS microanalyses of micro-phenocrysts of Mediterranean obsidians: a preliminary approach to source discrimination. *Eur. J. Mineral.*, **16**, 419–429.
- Balić-Zunić, T., Garavelli, A., Mitolo, D., Acquafredda, P., Leonardsen, E. (2012): Jakobssonite, CaAlF<sub>5</sub>, a new mineral from fumaroles at the Eldfell and Hekla volcanoes, Iceland. *Mineral. Mag.*, **76**, 751–760.

- Breese, N.E. & O’Keeffe, M. (1991): Bond-valence parameters for solids. *Acta Crystallogr.*, **B47**, 192–197.
- Brown, I.D. & Altermatt, D. (1985): Bond-valence parameters from a systematic analysis of the inorganic crystal structure database. *Acta Crystallogr.*, **B41**, 244–247.
- Bruker (2001): SAINT. Bruker AXS Inc., Madison, Wisconsin.
- Campostrini, I. & Gramaccioli, C.M. (2005): Artroeite del Monte Somma-Vesuvio, secondo ritrovamento mondiale. *Rivista Mineralogica Italiana*, **29**, 1, 50–52.
- Chiodini, G., Marini, L., Russo, M. (2001): Geochemical evidence for the existence of high-temperature hydrothermal brines at Vesuvio volcano, Italy. *Geochim. Cosmochim. Acta*, **65**, 2129–2147.
- Demartin, F., Campostrini, I., Castellano, C., Russo, M. (2014): Parascandolaite, KMgF<sub>3</sub>, a new perovskite-type fluoride from Vesuvius. *Phys. Chem. Minerals*, **41**, 403–407.
- Domsele, R. & Hoppe, R. (1980): The crystal structure of Ca<sub>2</sub>AlF<sub>7</sub>. *Z. Kristallogr.*, **153**, 317–328.
- Farrugia, L.J. (1999): WinGX suite for small-molecule single-crystal crystallography. *J. Appl. Crystallogr.*, **32**, 837–838.
- Fischer, R.X. & Tillmanns, E. (1988): The equivalent isotropic displacement factor. *Acta Crystallogr.*, **C44**, 775–776.
- Giacovazzo, C. & Menchetti, S. (1969): The crystal structure of prosopite. *Naturwissenschaften*, **56**, 86–87.
- Holland, T.J.B. & Redfern, S.A.T. (1997): Unit cell refinement from powder diffraction data: the use of regression diagnostics. *Mineral. Mag.*, **61**, 65–77.
- Kampf, A.R., Colombo, F., Gonzalez del Tanago, J. (2010): Carlhintzeite, Ca<sub>2</sub>AlF<sub>7</sub>·H<sub>2</sub>O from the Gigante granitic pegmatite, Cordoba Province, Argentina: description and crystal structure. *Mineral. Mag.*, **74**, 623–632.
- Marchetti, F. & Perchiazzi, N. (2000): The crystal structure of gearsksutite CaAlF<sub>4</sub>(OH)·H<sub>2</sub>O. *Am. Mineral.*, **85**, 231–235.
- Parascandola, A. (1951): I minerali del Vesuvio nella eruzione del marzo 1944 e quelli formati durante l’attuale periodo di riposo. *Bollettino della Società Geologica Italiana*, **70**, 523–526.
- (1960): Notizie vesuviane. Il Vesuvio dal marzo 1948 al dicembre 1958. *Boll. Soc. Naturalisti in Napoli*, **68**, 184.
- (1961): Notizie vesuviane. Il Vesuvio dal gennaio 1959 al dicembre 1960. *Boll. Soc. Naturalisti in Napoli*, **69**, 263–298.
- Pudovkina, Z.V., Chernitsova, N.M., Pyatenko, Y.A. (1973): Refinement of the crystalline structure of prosopite CaAl<sub>2</sub>F<sub>4</sub>(OH)<sub>4</sub>. *J. Struct. Chem.*, **14**, 345–347.
- Russo, M. & Campostrini, I. (2011): Ammineite, matlockite and post 1944 eruption fumarolic minerals at Vesuvius. *Plinius*, **37**, 312–312.
- Russo, M., Campostrini, I., Demartin, F. (2014): Fumarolic minerals after the 1944 Vesuvius eruption. *Plinius*, **40**, 306–306.
- Ruste, J. (1979): X-ray spectrometry. in “Microanalysis and Scanning Electron Microscopy”, Maurice, F., Meny, L., Tixier, R., eds. Summer School St-Martin-d’Hères, France, September 11–16, 1978. Les Editions de Physique, Orsay, 215–267.
- Shannon, R.D. & Fischer, R.X. (2016): Empirical electronic polarizabilities of ions for the prediction and interpretation of refractive indices: oxides and oxysalts. *Am. Mineral.*, **101**, 2288–2300.
- Sheldrick, G.M. (2000): *SADABS Area-Detector Absorption Correction Program*. Bruker AXS Inc., Madison, WI, USA.
- (2015): Crystal structure refinement with SHELXL. *Acta Crystallogr.*, **C71**, 3–8.

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