Debattistiite, $Ag_9Hg_{0.5}As_6S_{12}Te_2$, a new Te-bearing sulfosalt from Lengenbach quarry, Binn valley, Switzerland: description and crystal structure

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ABSTRACT

Debattistiite, ideally $Ag_9Hg_{0.5}As_6S_{12}Te_2$, is a new mineral (IMA-CNMNC 2011-098) from the Lengenbach quarry in the Binn Valley, Valais, Switzerland. It occurs as very rare tabular euhedral crystals up to 150 µm across in cavities in dolomitic marble, associated with realgar, rutile, trechmannite and hutchinsonite. Debattistiite is opaque with a metallic lustre and a grey streak. It is brittle; the Vickers hardness (VHN₂₅) is 80 kg mm⁻² (range: 65–94), corresponding to a Mohs hardness of $2-2\frac{1}{2}$. In reflected light debattistiite is dark grey, highly bireflectant and weakly pleochroic from dark grey to a slightly greenish grey. Between crossed polars it is highly anisotropic with brownish to blue rotation tints. Internal reflections are absent. Reflectance percentages for the four COM wavelengths (R_{min} and R_{max}) are 27.2, 34.5 (471.1 nm), 25.5, 31.0 (548.3 nm), 22.9, 28.4 (586.6 nm), and 20.1, 25.2 (652.3 nm), respectively.

Debattistiite is triclinic, space group $P\overline{1}$, with a = 7.832(5), b = 8.606(4), c = 10.755(5) Å, $\alpha = 95.563(9)$, $\beta = 95.880(5)$, $\gamma = 116.79(4)^{\circ}$, V = 635.3(6) Å³ and Z = 1. The crystal structure $[R_1 = 0.0826$ for 795 reflections with $I > 2\sigma(I)$] consists of corner-sharing AsS₃ pyramids forming three-membered distorted rings linked by Ag atoms in triangular or tetrahedral coordination.

The five strongest powder-diffraction lines [d in Å (I/I_0) (hkl)] are as follows: 10.56 (6) (001); 3.301 (5) ($\bar{2}12$); 2.991 (4) ($2\bar{1}2$); 2.742 ($\bar{2}\bar{1}1$) and 2.733 (10) ($\bar{2}30$). A mean of nine electron microprobe analyses gave: Ag 44.88, Hg 4.49, As 20.77, S 17.72, Te 11.82; total 99.68 wt.%, which corresponds to Ag_{9.02}Hg_{0.49}As_{6.012}S_{11.98}Te_{2.01}, on the basis of 29.5 atoms. The new mineral is named for Luca De Battisti, a systematic mineralogist and expert on the minerals of Lengenbach quarry.

KEYWORDS: debattistiite, new mineral, sulfosalts, tellurium, Lengenbach, Switzerland.

Introduction

LENGENBACH quarry in the Binn Valley, Valais, is one of the most famous mineral localities in Switerland and is well known for rare Pb-Cu-Ag-

* E-mail: fabrizio.nestola@unipd.it DOI: 10.1180/minmag.2012.076.3.21 As-Tl bearing sulfosalts. It is the type locality for the rare thallium-bearing sulfosalts dalnegroite, edenharterite, erniggliite, gabrielite, hatchite, hutchinsonite, imhofite, jentschite, rathite, sartorite, sicherite, stalderite and wallisite (Hofmann *et al.*, 1993; Graeser *et al.*, 2008; Nestola *et al.*, 2010; Bindi *et al.*, 2011). Tellurium-bearing minerals are very rare at the quarry. The presence of tellurium was noted by Graeser *et al.* (2008). Tellurium-rich canfieldite was subsequently reported by Bindi *et al.* (2012). Debattistiite is the second tellurium-bearing sulfosalt from Lengenbach.

Debattistiite has been identified in two specimens which were found in zone 1 of the quarry in 1996 (Greaser et al., 2008). It occurs with realgar, rutile, trechmannite and hutchinsonite in dolomitic marble. Debattistiite was approved as a new mineral by the IMA Commission (IMA 2011-098). The mineral name honours Luca De Battisti (b. 1958), a systematic mineralogist and expert on Lengenbach minerals. He discovered the new mineral dalnegroite (Nestola et al., 2010) in 2009 and also collected the specimens of debattistiite. He has published several descriptive articles on minerals and new finds from this famous locality (e.g. Guastoni and De Battisti, 2006). Holotype material is deposited in the mineralogical collections of the Museum of Mineralogy of the Department of Geosciences at the University of Padova, Italy, under catalogue number MMP M10680. To our knowledge, debattistiite has no synthetic analogues.

Physical and optical properties

Debattistiite occurs as very rare euhedral tabular crystals, up to 150 μ m in size, in cavities in dolomitic marble (Fig. 1). The crystals resemble wallisite and are closely associated with realgar, rutile, trechmannite and hutchinsonite. Debattistiite is black with a grey streak. It is opaque in transmitted light and has a metallic lustre. No cleavage was observed and the fracture is uneven. The calculated density using the empirical formula (see below) is 5.647 g cm⁻³.



FIG. 1. A backscattered-electron image of debattistiite crystals (bright) in dolomite matrix.

The density could not be measured because of the small grain size. Micro-indentation measurements carried out with a Vickers microhardness tester using a load of 25 g gave a mean value of 80 kg mm⁻² (range: 65–94) corresponding to a Mohs hardness of $2-2\frac{1}{2}$.

In plane-polarized incident light, debattistiite is dark grey, highly bireflectant and weakly pleochroic varying from dark grey to slightly greenish grey. Between crossed polars, debattistiite is highly anisotropic, with brownish to blue rotation tints. Internal reflections are absent and there is no optical evidence of growth zonation.

Reflectance measurements were performed in air using a Zeiss MPM-200 microphotometer equipped with an MSP-20 system processor on a Zeiss Axioplan ore microscope. The filament temperature was approximately 3350 K. An interference filter was adjusted to select four wavelengths for measurement (471.1, 548.3, 586.6 and 652.3 nm). Readings were taken from the specimen and an SiC calibration standard in the same conditions. The diameter of the measurement spot was 0.1 mm. Reflectance percentages, R_{min} and R_{max} , are 27.2, 34.5 (471.1 nm), 25.5, 31.0 (548.3 nm), 22.9, 28.4 (586.6 nm) and 20.1, 25.2 (652.3 nm), respectively.

X-ray crystallography and crystal structure determination

A crystal fragment $(70 \times 50 \times 40 \ \mu m)$ was selected for an X-ray single-crystal diffraction study which was carried out on a STOE-STADI IV CCD single-crystal diffractometer (Table 1). No systematic absences were observed in the collected data set, leading to the choice of space groups P1 and $P\overline{1}$. Statistical tests on the distribution of |E| values ($|E^2-1| = 0.965$) indicated the presence of an inversion centre, suggesting space group $P\overline{1}$. The positions of the heavy atoms were determined by the chargeflipping method (Oszlányi and Süto, 2008) using the JANA2006 software package (Petříček et al., 2006). A least-squares refinement on F^2 using these heavy-atom positions and isotropic temperature factors produced an R factor of 0.184. Three-dimensional difference-Fourier syntheses yielded the position of the remaining S atoms. The JANA2006 program (Petříček et al., 2006) was used to refine the structure.

Of the seven anion positions, one was found to be almost fully occupied by Te $[Te_{0.93(1)}S_{0.07}]$. It is important to note that the Ag4 and the Hg

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Crystal data	
Ideal formula	$Ag_{9}Hg_{0} As_{3}S_{12}Te_{2}$
Crystal system	triclinic
Space group	$P\overline{1}$
Unit-cell parameters (Å, °)	7.832(5) 8.606(4) 10.755(5) 95.563(9) 95.880(5) 116.79(4)
Unit-cell volume (Å ³) Z	635.3(6) 1
Crystal size (mm)	$0.070\times0.050\times0.040$
Data collection	
Diffractometer	STOE-STADI IV CCD
Temperature (K)	298(3)
Radiation, wavelength (Å)	ΜοΚα, 0.71073
2θ max for data collection (°)	86.18
Crystal-detector distance (mm)	50
h, k, l ranges	$-12 \rightarrow 12, -7 \rightarrow 16, -16 \rightarrow 20$
Axis, frames, width (°), time per frame (s)	ω-φ, 1380, 1.00, 25
Total reflections collected	8404
Unique reflections (R_{int})	5635 (0.125)
Unique reflections $I > 2\sigma(I)$	795
Data completeness to θ_{max} (%)	99.2
Absorption correction method	X-RED and X-SHAPE
Structure refinement	
Refinement method	Full-matrix least-squares on I
Data/restraints/parameters	795/0/133
$R_1 [I > 2\sigma(I)], wR_2 [I > 2\sigma(I)]$	0.0826, 0.1733
R_1 all, wR_2 all	0.1012, 0.2104
Goodness-of-fit on F^2	0.700
Largest diff. peak and hole $(e^{-} Å^{-3})$	3.05, -4.99

TABLE 1. Crystallographic data and refinement parameters for debattistiite.

$R_{\rm int} = (n/n - 1)^{1/2} [F_{\rm o}^2 - F_{\rm o} ({\rm mean})^2] / \Sigma F_{\rm o}^2$
$R_1 = \Sigma F_0 - F_c / \Sigma F_0 ; \ wR_2 = \{ \Sigma [w(F_0^2 - F_c^2)^2] / \Sigma [w(F_0^2)^2] \}^{1/2};$
* $w = 1/[\sigma^2(F_o^2) + (aP)^2 + bP]$, where $P = (max(F_o^2, 0) + 2F_c^2)/3$;
GooF = { $\Sigma[w(F_o^2 - F_c^2)^2]/(n - p)$ } ^{1/2}

1/2 2 2 2

where n is the number of reflections and p is the number of refined parameters.

positions, which are very close (separation 1.137 Å), are only partially (50%) occupied (Table 2). The occupancy of these sites was allowed to vary and they were found to be 50% occupied (the analytical uncertainty of the occupancy values is 0.02). After this had been determined, the occupancy of these sites was fixed at 50% in the structure refinement. There is, therefore, a statistical distribution between them (see below). Neutral scattering curves for Ag, Hg, As, S and Te were taken from the *International Tables for X-ray Crystallography* (Ibers and Hamilton, 1974). At the last stage of the refinement, with anisotropic atomic displacement

parameters for all atoms and no constraints, the refinement converged at $R_1 = 0.0826$ for 795 observed reflections $(2\sigma(I) \text{ level})$ and 133 parameters, and at $R_1 = 0.1012$ for all 5623 independent reflections. The relatively small size of the crystal and its poor quality accounts for the significant difference between the total number of unique reflections measured and those having $I > 2\sigma(I)$.

Inspection of the difference-Fourier map revealed maximum positive and negative peaks of 3.05 and -4.99 e⁻ Å⁻³, respectively. Experimental details and *R* indices are listed in Table 1. Fractional atom coordinates and isotropic

Atom	Wyckoff	Site occupancy	x/a	y/b	z/c	U_{eq}
Ag1	2 <i>i</i>	Ag _{1.00}	-0.0372(4)	0.5057(3)	0.3695(3)	0.0453(9)
Ag2	2i	Ag _{1.00}	0.5761(4)	0.9562(4)	0.3442(3)	0.060(1)
Ag3	2i	$Ag_{1.00}$	0.2110(4)	0.6290(3)	0.1574(4)	0.067(1)
Ag4	2i	Ag _{0.50}	0.5814(6)	0.5688(5)	0.6241(4)	0.019(1)
Hg4	1h	Hg _{0.50}	1/2	1/2	1/2	0.024(2)
Ag5	2i	Ag _{1.00}	0.5248(4)	0.2238(4)	0.1972(4)	0.079(1)
As1	2i	As _{1.00}	0.1403(4)	0.9740(4)	0.3703(3)	0.0205(8)
As2	2i	As _{1.00}	-0.2694(4)	0.6768(4)	0.1302(3)	0.0237(8)
As3	2i	As _{1.00}	-0.0383(4)	0.1421(4)	0.1233(3)	0.0190(7)
S1	2i	$S_{1.00}$	-0.165(1)	0.7423(9)	0.3469(8)	0.028(2)
S2	2i	S _{1.00}	0.052(1)	0.1946(9)	0.3412(9)	0.032(2)
S3	2i	$S_{1.00}$	-0.1406(9)	0.3513(8)	0.1274(7)	0.016(2)
S4	2i	$S_{1.00}$	-0.344(1)	0.908(1)	0.1133(8)	0.030(2)
S5	2i	S _{1.00}	-0.570(1)	0.4765(9)	0.1487(9)	0.028(2)
S6	2i	S _{1.00}	0.219(1)	0.929(1)	0.1780(8)	0.024(2)
Te1	2i	$Te_{1.86(2)}S_{0.14}$	0.3521(3)	0.7157(3)	0.4993(2)	0.0258(8)

TABLE 2. Atoms, Wyckoff symbol, site occupancy, fractional atom coordinates (Å) and isotropic atomic displacement parameters (Å²) for debattistiite.

displacement parameters are reported in Table 2 (anisotropic atomic displacement parameters are listed in the accompanying CIF). Bond distances are given in Table 3. Calculated X-ray powder diffraction data, computed using the atom coordinates and occupancies reported in Table 2, are listed in Table 4. Structure factors and a crystallographic information file have been deposited with the Principal Editors of *Mineralogical Magazine* and are available at http://www.minersoc.org/pages/e_journals/dep_mat.html.

Chemical composition

A preliminary chemical analysis by energydispersive spectrometry was performed using a JEOL-5610 LV scanning electron microscope at the Museum of Natural History of Milan. The chemical composition was determined by wavelength-dispersive spectrometry (WDS) using a CAMECA-CAMEBAX electron microprobe. Major and minor elements were determined at 15 kV accelerating voltage and 15 nA beam

As1-S6 -S1	2.265(9) 2.289(8)	As2-S5 -S1	2.249(8) 2.314(9)	As3-S3 -S4	2.275(7) 2.317(9)
-S2	2.329(8)	-S4	2.327(8)	-S2	2.317(9)
<as1-s></as1-s>	2.294	<as2-s></as2-s>	2.297	<as1-s></as1-s>	2.303
Ag1-S1	2.664(8)	Ag2-S4	2.662(9)	Ag3-S6	2.539(8)
-S3	2.670(8)	-Te1	2.853(4)	-85	2.589(7)
-Te1	2.860(4)	-Te1	2.928(4)	-S3	2.671(7)
-Te1	2.950(4)				
<ag1-s></ag1-s>	2.786	< <u>Ag2</u> - <u>S</u> >	2.814	< <u>Ag3</u> – <u>S</u> >	2.600
Ag4-S5	2.516(9)	Hg4-Te1	2.594(2)	Ag5-S6	2.560(8)
-S2	2.623(9)	-Tel	2.594(3)	-\$3	2.571(7)
-Te1	2.916(5)	<hg-te></hg-te>	2.594	-S5	2.675(8)
-Te1	2.947(5)				
<ag4–s></ag4–s>	2.751	<hg-te></hg-te>	2.594	<ag5-s></ag5-s>	2.6025

TABLE 3. Selected bond distances (Å) for debattistiite.

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I _{rel}	$d_{\rm calc}$	h	k	l	I _{rel}	$d_{\rm calc}$	h	k	l
62	10.5585	0	0	1	25	2.4800	1	1	3
10	7.5818	0	1	0	14	2.4576	ī	ī	4
15	6.9630	ī	1	0	11	2.3941	0	2	3
12	6.8961	1	0	0	14	2.3512	0	2	4
6	6.2648	ī	0	1	15	2.2978	2	1	4
6	5.3826	1	0	1	9	2.2421	3	2	2
8	4.2634	ī	2	0	7	2.1159	ī	3	1
5	4.1707	1	1	2	6	2.1141	2	3	3
14	3.5195	0	0	3	8	2.0926	2	ī	4
8	3.4481	2	0	0	7	2.0521	1	3	4
14	3.3980	0	2	1	20	1.9514	4	2	1
14	3.3453	0	2	2	6	1.9217	3	3	3
47	3.3010	2	1	2	26	1.9202	0	4	1
7	3.2888	2	2	1	5	1.9099	1	4	3
15	3.1293	2	0	1	8	1.8981	4	1	1
10	3.0263	1	1	2	6	1.8817	4	3	1
11	3.0123	0	1	3	7	1.8583	2	3	4
13	2.9972	ī	1	3	8	1.8379	2	1	5
40	2.9906	2	1	2	10	1.7598	0	0	6
10	2.9449	1	0	3	6	1.7588	3	3	4
11	2.8680	0	2	2	5	1.7058	ī	4	3
18	2.8155	0	2	3	5	1.6988	1	1	6
95	2.7419	2	1	1	6	1.6505	4	2	4
100	2.7329	2	3	0	5	1.6444	4	4	2
31	2.6396	0	0	4	5	1.6044	2	5	3
6	2.6265	1	3	2	8	1.5101	2	3	6
10	2.6180	2	3	1	8	0.7723	5	6	5
6	2.6140	ī	2	3	9	0.7710	6	1	12
21	2.6129	ī	0	4	22	0.7710	3	3	13
6	2.6016	1	2	1					

TABLE 4. Calculated X-ray powder diffraction data for debattistiite.

The calculated X-ray powder pattern was computed on the basis of a = 7.832(5), b = 8.606(4), c = 10.755(5) Å, $\alpha = 95.563(9)$, $\beta = 95.880(5)$, $\gamma = 116.79(4)^{\circ}$, and with the atomic coordinates and occupancies reported in Table 2. The five strongest lines are indicated in bold face.

current with a 1 µm beam diameter using 15 s counting times for both peaks and background. The X-ray lines used for the WDS analyses were as follows: SKa, FeKa, CuKa, ZnKa, AsLa, SeLa, TeLa, AgLa, SbL β , PbMa, BiM β , and HgLa. The elements Fe, Cu, Zn, Se, Sb, Pb and Bi were at or below the limit of detection of the electron microprobe (0.01 wt.%). The estimated analytical precision for the remaining elements is ± 0.13 for Ag, ± 0.09 for As, ± 0.08 for S, ± 0.06 for Te and ± 0.04 for Hg. The standards employed were Ag metal for Ag, elemental Te for Te, synthetic As₂S₃ for As, cinnabar for Hg and pyrite for S. The crystal fragment was found to be homogeneous to within analytical error. The average chemical composition (nine analyses on different spots), and the wt.% ranges for the elements is reported in Table 5. On the basis of 29.5 atoms, the empirical formula for debattistiite is $Ag_{9.01}Hg_{0.49}As_{6.01}S_{11.98}Te_{2.01}$. The simplified ideal formula is $Ag_9Hg_{0.5}As_6S_{12}Te_2$, which requires: Ag 44.93, Hg 4.64, As 20.81, S 17.81, Te 11.81; total 100.00 wt.%.

Results and discussion

Description of the crystal structure

The crystal structure of debattistiite (Fig. 2) consists of three corner-sharing AsS_3 pyramids forming distorted rings that are linked by Ag polyhedra. A similar feature is present in the structure of trechmannite, $AgAsS_2$ (Matsumoto and Nowacki, 1969). In detail, Ag3 and Ag5 link three S atoms with mean bond distances of 2.600

Element	Mean (wt.%)	Range	SD	Standard
Ag	44.88	44.43-45.26	0.13	Ag metal
Hg	4.49	4.11-4.68	0.04	cinnabar
As	20.77	20.21-21.30	0.09	synthetic As ₂ S ₃
S	17.72	17.19-18.01	0.08	pyrite
Те	11.82	11.44-12.09	0.06	elemental Te
Total	99.68	98.96-100.68		

TABLE 5. Electron-microprobe data (means and ranges in wt.% of elements on 9 analyses) with standard deviations (1σ) and analytical standards for debattistiite.

and 2.602 Å (Table 3). These values are in excellent agreement with those observed for silver atoms in triangular coordination in the structure of minerals of the pearceite-polybasite group (Bindi et al., 2006, 2007; Evain et al., 2006). The Ag2 links to one S and two Te/S atoms (i.e. the Te1 position) in a distorted triangular arrangement. The Ag1 and Ag4 sites have an interesting crystal-chemical environment, with a distorted tetrahedral coordination with 2 S and 2 Te/S atoms. These tetrahedra are strongly distorted with mean bond distances of 2.786 and 2.751 Å (Table 3). It is noteworthy that the two distances to S atoms are in keeping with those reported for tetrahedrally coordinated silver atoms in the structure of minerals of the pearceitepolybasite group (Bindi *et al.*, 2006, 2007; Evain *et al.*, 2006) and the two longest Te distances are similar to those reported in the hessite structure (2.8415(7)-3.034(1) Å; Van der Lee and de Boer, 1993).

The Hg atom is in a nearly perfect linear coordination with two Te/S atoms. The mean distance of 2.594 Å is greater than that in fettelite ($\langle Hg-S \rangle = 2.395$ Å; Bindi *et al.*, 2009) and values reported in synthetic [Hg{S(CH₂)₂NH₃}₂] (C1)₂ ($\langle Hg-S \rangle : 2.333$ and 2.338 Å; Kim *et al.*, 2002). This reflects the larger radius of Te in comparison to S. It is important to note that the Ag4 and the Hg positions (separation 1.137 Å) are partially occupied (50%). Thus there is a statistical distribution (50:50) between



FIG. 2. The crystal structure of debattistiite projected down [100]. The Ag, Hg, As, S and Te atoms are indicated as white, red, black, yellow and orange spheres, respectively. The long As–S distances (lone pairs) are depicted using dashed lines. The unit cell is outlined.

 $Hg(Te,S)_2$ and $AgS_2(Te,S)_2$ polyhedra in the structure (Fig. 3).

The Ag1–Ag1 and Ag4–Ag4 distances of 2.83 and 2.76 Å, respectively, are similar to those reported in face centred cubic silver ($R_{Ag-Ag} = 2.89$ Å; Suh *et al.*, 1988) and hexagonal close packed silver ($R_{Ag-Ag} = 2.93$ Å; Petruk *et al.*, 1970).

Possible origin and mineral association

Debattistiite is the second Te-bearing mineral to be found at Lengenbach after Te-bearing canfieldite (Bindi et al., 2012). It occurs with trechmannite and hutchinsonite in dolomitic marble. In hydrothermal quartz veins, tellurium is found in tellurides such as muthmannite, krennerite, petzite and calaverite. The formation of sulfotellurides in hydrothermal systems indicates a relatively low fTe2, insufficient for the formation of most Sfree tellurides (Novoselov et al., 2006). Estimates of the crystallization temperatures for sulfotellurides such as cervelleite, Ag4TeS, have been proposed as >270°C by Karup-Møller (1976) and 160-260°C by Helmy (1999). Metastable sulfotellurides in association with native tellurium are thought to have formed at temperatures between 100 and 160°C (Maslennikov, 1999). On the basis of their textural relationships and mineral associations it is thought that cervelleite and a sulfotelluride with the formula Ag₂Cu₂TeS



FIG. 3. The crystal-chemical environment of the Ag4 site and Hg atoms. The Ag, Hg, S and Te atoms are indicated as white, red, yellow and orange spheres, respectively. The Ag4 site and Hg are very close (separation 1.137 Å), so there is a statistical distribution (50:50) between Hg(Te,S)₂ and AgS₂(Te,S)₂ polyhedra in the structure. The Hg-(Te,S) distances are indicated using solid lines whereas the Ag-S and the Ag-(Te,S) distances are indicated using dashed lines.

formed at temperatures close to 400°C in a skarn deposit (Cook and Ciobanu, 2003). These data show that estimating the temperature of formation is difficult for sulfotellurides due to the rarity or absence of thermodynamic data and the wide variety of geological environments in which they occur (Novoselov et al., 2006). Recent work by Voudouris et al. (2011) indicates that sulfotellurides and cervelleite-like phases could have exsolved from galena during cooling below 200°C. These authors suggest that the initial temperatures for the formation of sulfotellurides could be close to 300°C. Oberthür and Weiser (2008) studied an assemblage of Au-Bi-Te-S minerals (joseite-A, joseite-B and hessite) and proposed that such an assemblage can form in reducing conditions at relatively low fS_2 and fTe_2 at temperatures of ~340°C.

The Lengenbach mineralization is in a Triassic dolostone marble which formed during greenschist to lower amphibolite facies Alpine metamorphism in the temperature range 400–500°C. Metamorphic processes remobilized the primary mineralization, which was originally a stratiform carbonate-hosted orebody (Vokes, 1971; Hofmann and Knill, 1996). Cooling of sulfide melts led to the formation of massive to wellcrystallized sulfides and sulfosalts from aqueous hydrothermal fluids at estimated temperatures >330°C (Hofmann, 1994). It is therefore suggested that debattistiite formed from hydrothermal solutions at similar temperatures to those reported for the other Lengenbach sulfosalts.

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Refinement of $F^{2^{a}}$ against ALL reflections. The weighted R-factor wR and goodness of fit S are based on $F^{2^{a}}$, conventional R-factors R are based on F, with F set to zero for negative $F^{2^{a}}$. The threshold expression of $F^{2^{a}} > 2 \text{sigma}(F^{2^{a}})$ is used only for calculating R-factors(gt) etc. and is not relevant to the choice of reflections for refinement. R-factors based on $F^{2^{a}}$ are statistically about twice as large as those based on F, and R-factors based on ALL data will be even larger.

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