

# New arsenate minerals from the Arsenatnaya fumarole, Tolbachik volcano, Kamchatka, Russia. VI. Melanarsite, $K_3Cu_7Fe^{3+}O_4(AsO_4)_4$

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## ABSTRACT

The new mineral melanarsite,  $K_3Cu_7Fe^{3+}O_4(AsO_4)_4$ , was found in the sublimates of the Arsenatnaya fumarole at the Second scoria cone of the Northern Breakthrough of the Great Tolbachik Fissure Eruption, Tolbachik volcano, Kamchatka Peninsula, Russia. It is associated with dmiskolovite, shchurovskite, bradaczekite, hematite, tenorite, apthitalite, johillerite, arsmirandite, As-bearing orthoclase, hatertite, pharmazincite, etc. Melanarsite occurs as tabular to prismatic crystals up to 0.4 mm, separate or combined in clusters up to 1 mm across or in interrupted crusts up to  $0.02\text{ cm} \times 1\text{ cm} \times 1\text{ cm}$  covering basalt scoria. The mineral is opaque, black, with a vitreous lustre. Melanarsite is brittle. Mohs' hardness is  $\sim 4$  and the mean VHN =  $203\text{ kg mm}^{-2}$ . Cleavage was not observed and the fracture is uneven.  $D_{\text{calc}}$  is  $4.39\text{ g cm}^{-3}$ . In reflected light, melanarsite is dark grey. Birefractance is weak, anisotropism is very weak. Reflectance values [ $R_1$ – $R_2$ , % ( $\lambda$ , nm)] are 10.5–9.4 (470), 10.0–8.9 (546), 9.7–8.7 (589), 9.5–8.6 (650). The Raman spectrum is reported. Chemical composition (wt.%, electron microprobe) is  $K_2O$  10.70,  $CaO$  0.03,  $CuO$  45.11,  $ZnO$  0.24,  $Al_2O_3$  0.32,  $Fe_2O_3$  6.11,  $TiO_2$  0.12,  $P_2O_5$  0.07,  $As_2O_5$  36.86, total 99.56. The empirical formula, based on 20 O apfu, is  $(K_{2.81}Ca_{0.01})_{\Sigma 2.82}(Cu_{7.02}Fe_{0.95}^{3+}Al_{0.08}Zn_{0.04}Ti_{0.02})_{\Sigma 8.11}(As_{3.97}P_{0.01})_{\Sigma 3.98}O_{20}$ . Melanarsite is monoclinic,  $C2/c$ ,  $a = 11.4763(9)$ ,  $b = 16.620(2)$ ,  $c = 10.1322(8)\text{ \AA}$ ,  $\beta = 105.078(9)^\circ$ ,  $V = 1866.0(3)\text{ \AA}^3$  and  $Z = 4$ . The strongest reflections of the powder X-ray diffraction pattern [ $d$ ,  $\text{\AA}$  ( $hkl$ )] are 9.22(100)(110), 7.59(35)( $\bar{1}11$ ), 6.084(17)(111), 4.595(26)( $\bar{1}31$ , 220,  $\bar{2}21$ ), 3.124(22)( $\bar{3}31$ ,  $\bar{1}51$ ), 2.763(20)(400,  $\bar{1}52$ ), 2.570(23)(043) and 2.473(16)(260,  $\bar{2}61$ , 350). Melanarsite has a novel structure type. Its crystal structure, solved from single-crystal X-ray diffraction data ( $R = 0.091$ ), is based upon a heteropolyhedral pseudo-framework built by distorted  $Cu(1-3)O_6$  and  $(Fe,Cu)O_6$  octahedra and  $As(1-3)O_4$  tetrahedra. Two crystallographically independent  $K^+$  cations are located in the tunnels and voids of the pseudo-framework centring eight- and seven-fold polyhedra. The name reflects the mineral being an arsenate and its black colour (from the Greek μέλας, black).

**KEYWORDS:** melanarsite, new mineral, potassium copper iron arsenate, crystal structure, fumarole sublimate, Tolbachik volcano, Kamchatka.

## Introduction

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THIS paper continues a series of articles on new arsenates from the Arsenatnaya fumarole located at

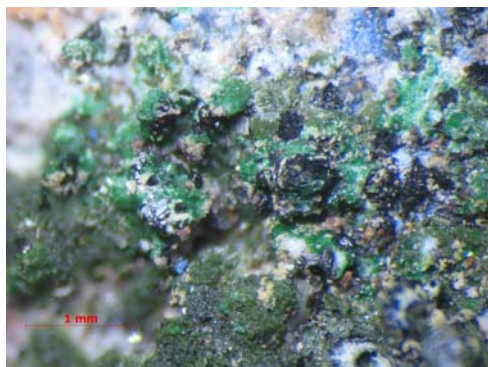


FIG. 1. Black crystals of melanarsite associated with bright green dmisokolovite, olive green shchurovskyite, blue bradaczekite and white As-bearing orthoclase. All these minerals form thin crusts on light-brown basalt scoria altered by fumarolic gas. Field of view: 4.4 mm. Photo: I.V. Pekov and A.V. Kasatkin.

the apical part of the Second scoria cone of the Northern Breakthrough of the Great Tolbachik Fissure Eruption, Tolbachik volcano, Kamchatka Peninsula, Far-Eastern Region, Russia (55°41'N 160°14'E, 1200 m asl). This active, hot fumarole, discovered by us in July 2012, was characterized in the first article of the series devoted to the new mineral yurmarinite  $\text{Na}_7(\text{Fe}^{3+}, \text{Mg}, \text{Cu})_4(\text{AsO}_4)_6$  (Pekov *et al.*, 2014b). In subsequent papers we described two polymorphous modifications of  $\text{Cu}_4\text{O}(\text{AsO}_4)_2$ , ericlxmanite and kozyrevskite (Pekov *et al.*, 2014c), popovite  $\text{Cu}_5\text{O}_2(\text{AsO}_4)_2$  (Pekov *et al.*, 2015b), structurally related shchurovskyite  $\text{K}_2\text{CaCu}_6\text{O}_2(\text{AsO}_4)_4$  and dmisokolovite



FIG. 2. Black, coarsely crystalline crust of melanarsite partially overgrown by bright green dmisokolovite. Field of view: 5.9 mm. Photo: I.V. Pekov and A.V. Kasatkin.

$\text{K}_3\text{Cu}_5\text{AlO}_2(\text{AsO}_4)_4$  (Pekov *et al.*, 2015c) and katiarsite  $\text{KTiO}(\text{AsO}_4)$  (Pekov *et al.*, 2016). In this paper the mineralogical description and crystal structure data for a new mineral melanarsite  $\text{K}_3\text{Cu}_7\text{Fe}^{3+}\text{O}_4(\text{AsO}_4)_4$  (Cyrillic: меланарсит) are given. The name reflects its being an arsenate and its black colour (from the Greek μέλαν, black), very rare for minerals of this chemical class. Both the new mineral and its name have been approved by the IMA Commission on New Minerals, Nomenclature and Classification (MA2014-048, Pekov *et al.*, 2014e). The type specimen is deposited in the systematic collection of the Fersman Mineralogical Museum of the Russian Academy of Sciences, Moscow, catalogue number 94992.

## Occurrence and general appearance

Specimens with melanarsite were collected by us in July 2013 in the southern part of the Arsenatnaya fumarole, at a depth ~0.5 m below the surface. The temperature measured using a chromel-alumel thermocouple immediately after uncovering the mineralized pockets was 360–380°C. Melanarsite was found in incrustations mainly consisting of arsenates, oxides and sulfates. We believe that the new mineral was deposited directly from the gas phase as a volcanic sublimate at temperatures not lower than 380°C (see Discussion).

Melanarsite occurs as well-shaped or, more commonly, coarse crystals up to 0.1 mm, rarely up to 0.4 mm in size, separate or combined in dense or open-work clusters up to 1 mm across, rarely forming interrupted incrustations up to 1 cm × 1 cm in area and up to 0.2 mm thick covering basalt scoria (Figs 1 to 3). Melanarsite crystals are tabular to prismatic, typically angled obliquely (Fig. 3). Usually they have a skeletal and/or blocky character. Their parallel or near-parallel intergrowths are sometimes combined in crusts (Fig. 3c,d). The most typical minerals associated intimately with melanarsite are dmisokolovite, shchurovskyite, bradaczekite  $\text{NaCu}_4(\text{AsO}_4)_3$ , hematite, As-bearing orthoclase (Figs 1, 2, 4 and 5), johillerite  $\text{Na}(\text{Mg}, \text{Cu})_3\text{Cu}(\text{AsO}_4)_3$ , arsmirandite  $\text{Na}_{18}\text{Cu}_{12}\text{Fe}^{3+}\text{O}_8(\text{AsO}_4)_8\text{Cl}_5$  (IMA2014-081, Pekov *et al.*, 2015a), apthitalite, tenorite, langbeinite, anhydrite, tilasite and fluorophlogopite. In some cases the new mineral is associated closely with hatertite  $\text{Na}_2(\text{Ca}, \text{Na})(\text{Fe}^{3+}, \text{Cu})_2(\text{AsO}_4)_3$  (Fig. 4) or pharmazincite  $\text{KZnAsO}_4$  (IMA2014-015, Pekov *et al.*, 2014a) (Fig. 5).

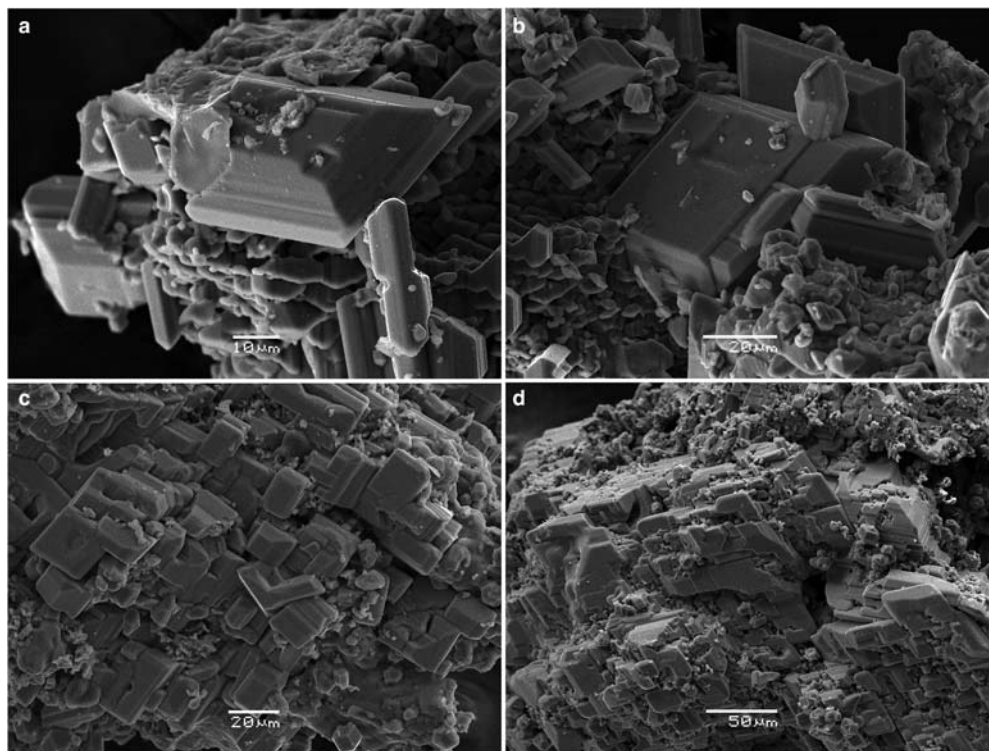


FIG. 3. Well-shaped crystals (*a, b*) and crystal crusts (*c, d*) of melanarsite. Scanning electron microscopy (SEM) secondary electron image.

### Physical properties and optical data

Melanarsite is black, with a strong vitreous lustre. It is megascopically opaque; only very thin sections are translucent and dark green. Its streak is dark

green with an olive hue. The mineral is brittle. Cleavage or parting is not observed. The fracture is uneven. Mohs hardness is  $\sim 4$ . The mean micro-indentation hardness (VHN) is 203, range 178–225  $\text{kg mm}^{-2}$  (load 10 g). Density could not

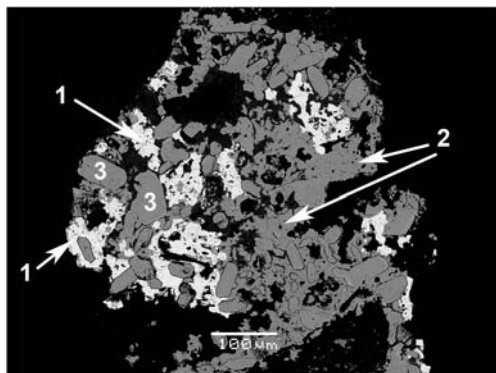


FIG. 4. Aggregates of melanarsite (1) and hatertite (2) overgrowing well-shaped crystals of hematite (3). Polished section. SEM back-scatter electron image.

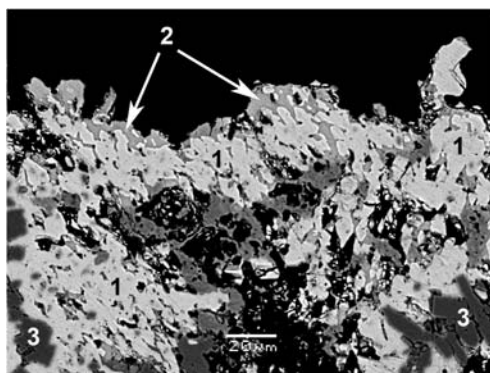


FIG. 5. Crystal crust composed by (1) melanarsite, (2) pharmazincite  $\text{KZnAsO}_4$  (IMA2014-015) and (3) hematite. Polished section. SEM back-scatter electron image.

TABLE 1. Reflectance data for melanarsite\*.

$\lambda$ , nm	$R_1$ , %	$R_2$ , %	$\lambda$ , nm	$R_1$ , %	$R_2$ , %
400	11.4	10.2	560	9.9	8.8
420	11.1	9.9	580	9.8	8.7
440	10.9	9.7	<b>589</b>	<b>9.7</b>	<b>8.7</b>
460	10.6	9.4	600	9.7	8.7
<b>470</b>	<b>10.5</b>	<b>9.4</b>	620	9.6	8.6
480	10.4	9.3	640	9.5	8.6
500	10.2	9.1	<b>650</b>	<b>9.5</b>	<b>8.6</b>
520	10.1	9.0	660	9.4	8.5
540	10.0	8.9	680	9.4	8.4
<b>546</b>	<b>10.0</b>	<b>8.9</b>	700	9.3	8.4

\*Values for wavenumbers ( $\lambda$ ) recommended by the Commission on Ore Mineralogy (COM) are given in bold.

be measured because monomineralic, massive (without caverns or inclusions) particles of melanarsite are too small for volumetric methods and because of a lack of heavy liquids of necessary density. Density calculated using its empirical formula is  $4.386 \text{ g cm}^{-3}$ .

Under the microscope in reflected light, melanarsite is dark grey and pleochroism was not observed. Bireflectance is weak,  $\Delta R = 1.0\%$  (589 nm). Anisotropism is very weak. The reflectance values measured by a MSF-21 microspectrophotometer (LOMO, Russia) using SiC standard (Zeiss, No. 545) are given in Table 1. Under the microscope in transmitted light, melanarsite

particles larger than  $5\text{--}7 \mu\text{m}$  are not transparent and optical data can be obtained only for tiny particles ( $3\text{--}5 \mu\text{m}$ ). Because of this and the high refractivity of the mineral, the data are of poor quality. Melanarsite is biaxial,  $\alpha = 1.80(1)$ ,  $\beta$  was not measured and  $\gamma = 1.91(1)$ . Optical sign and  $2V$  were not determined. Dispersion was not observed. Pleochroism, observed in the thinnest particles, is strong with  $Z$  (very dark greyish-green)  $> X$  (green to pale green).

### Raman spectroscopy

The Raman spectrum of melanarsite (Fig. 6) was recorded for a randomly orientated crystal using an EnSpectr R532 instrument (Department of Mineralogy, Moscow State University) with a green laser (532 nm) at room temperature. The output power of the laser beam was  $\sim 4 \text{ mW}$ . The spectrum was processed using the EnSpectr expert mode program in the range  $100$  to  $4000 \text{ cm}^{-1}$  with a holographic diffraction grating with  $1800 \text{ lines cm}^{-1}$  and a resolution of  $5\text{--}8 \text{ cm}^{-1}$ . The diameter of the focal spot on the sample was  $\sim 15 \mu\text{m}$ . The back-scattered Raman signal was collected with  $40\times$  objective; signal acquisition time for a single scan of the wavelength range was  $1500 \text{ ms}$  and the signal was averaged over 10 scans. Two strong bands in the range  $770$  to  $870 \text{ cm}^{-1}$  correspond to  $\text{As}^{5+}\text{--O}$  stretching vibrations of  $\text{AsO}_4^{3-}$  anions. Bands with frequencies lower than  $650 \text{ cm}^{-1}$  correspond to bending vibrations in  $\text{AsO}_4$  tetrahedra,  $\text{Cu}^{2+}\text{--O}$  stretching vibrations and lattice modes. In particular,

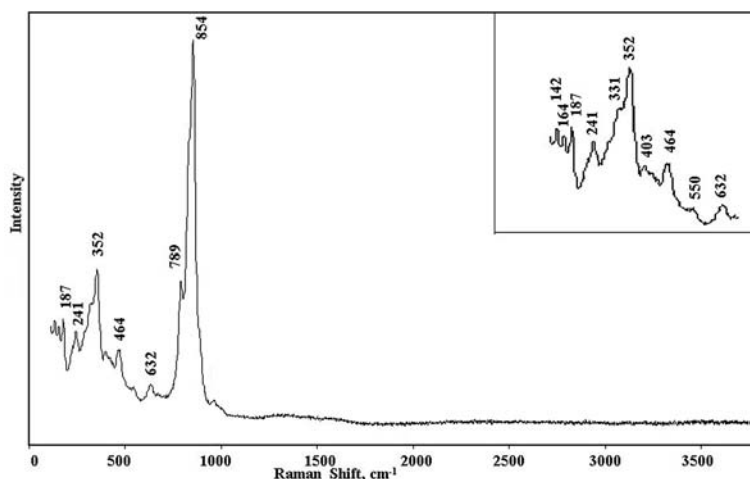


FIG. 6. Raman spectrum of melanarsite with an enlarged section showing the low-frequency region.



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TABLE 2. Powder X-ray diffraction data for melanarsite.

$I_{\text{obs}}$	$d_{\text{obs}}$	$I_{\text{calc}}^*$	$d_{\text{calc}}^{**}$	$h\ k\ l$
100	9.22	100	9.220	110
35	7.59	42	7.578	$\bar{1}11$
17	6.084	20	6.085	111
4	4.958	4	4.955	130
6	4.905	4	4.892	002
7	4.779	8	4.770	$\bar{1}12$
26	4.595	4, 14, 9	4.644, 4.610, 4.569	$\bar{1}31, 220, \bar{2}21$
6	4.251	3	4.257	$\bar{2}02$
4	4.222	1, 3	4.227, 4.216	131, 022
14	3.979	6	3.979	112
12	3.826	9, 2	3.826, 3.824	221, 041
12	3.686	5	3.703	$\bar{1}32$
3	3.337	1	3.335	$\bar{3}12$
5	3.331	1	3.324	240
8	3.289	11, 3	3.295, 3.269	132, 202
22	3.124	1, 15, 8	3.135, 3.132, 3.096	311, $\bar{3}31, \bar{1}51$
8	3.078	10	3.073	330
14	3.038	23, 3	3.042, 3.036	222, 023
14	2.990	24	2.984	223
14	2.966	30	2.963	151
5	2.887	10	2.882	$\bar{1}33$
6	2.797	14	2.798	313
20	2.763	23, 1, 9	2.770, 2.766, 2.765	400, 331, $\bar{1}52$
11	2.739	18	2.735	402
14	2.599	36, 2, 9	2.604, 2.598, 2.589	312, $\bar{4}22, 133$
23	2.570	35	2.565	043
14	2.503	1, 22, 1	2.502, 2.501, 2.490	$\bar{1}14, \bar{3}51, \bar{2}04$
16	2.473	7, 14, 6	2.478, 2.471, 2.471	260, $\bar{2}61, 350$
2	2.433	3	2.426	223
2	2.408	2	2.410	062
7	2.317	1, 24	2.322, 2.316	$\bar{2}62, \bar{3}14$
6	2.176	12	2.180	402
9	2.130	12	2.129	$\bar{4}04$
5	2.113	1, 1, 10, 1	2.119, 2.111, 2.110, 2.098	531, 063, $\bar{4}43, 134$
9	2.077	14	2.078	080
2	2.055	3, 1	2.058, 2.054	530, 172
1	2.015	1	2.014	154
1	1.968	1	1.968	$\bar{5}33$
1	1.946	1, 1, 1	1.947, 1.945, 1.942	$\bar{3}72, 280, \bar{2}81$
1	1.897	2, 1	1.899, 1.897	$\bar{6}02, \bar{1}35$
2	1.868	2, 1, 1	1.873, 1.871, 1.864	154, 263, $\bar{5}52$
7	1.842	1, 1, 1	1.847, 1.844, 1.844	600, 173, 550
3	1.821	2, 2, 1	1.823, 1.822, 1.821	353, 190, 314
4	1.803	5, 3	1.805, 1.803	$\bar{1}91, 620$
3	1.783	3	1.781	$\bar{6}23$
3	1.779	1, 4	1.779, 1.777	553, 191
1	1.751	1	1.752	135
2	1.731	1, 2, 3, 3	1.734, 1.732, 1.731, 1.726	$\bar{6}41, \bar{1}74, \bar{1}92, \bar{1}55$
3	1.716	11	1.715	443
3	1.685	2, 8	1.684, 1.682	192, $\bar{4}45$
4	1.659	4, 1, 3	1.662, 1.662, 1.654	480, 010.0, $\bar{4}82$
6	1.632	3, 1, 16, 2, 2	1.635, 1.633, 1.631, 1.628, 1.628	404, $\bar{5}72, 006, \bar{7}12, \bar{5}35$

(continued)

TABLE 2. (contd.)

$I_{\text{obs}}$	$d_{\text{obs}}$	$I_{\text{calc}}^*$	$d_{\text{calc}}^{**}$	$h\ k\ l$
3	1.621	1, 1, 2	1.623, 1.623, 1.620	$\bar{7}11, \bar{3}92, 570$
4	1.599	2, 3, 1, 1, 2	1.601, 1.599, 1.598, 1.598, 1.598	641, 406, 391, 065, 481
3	1.592	2, 2, 1, 1	1.592, 1.590, 1.590, 1.589	2.10.0, $\bar{2}$ .10.1, $\bar{3}$ 36, 713
3	1.576	3, 1	1.576, 1.574	710, 0.10.2
2	1.561	2	1.565	$\bar{7}31$
3	1.513	4, 11, 2	1.515, 1.513, 1.509	$\bar{7}14, 392, 335$
6	1.488	2, 1, 6, 2, 2	1.492, 1.488, 1.487, 1.487, 1.485	$\bar{1}94, \bar{1}$ .11.1, $\bar{5}74, \bar{4}84, \bar{3}56$
1	1.476	3, 2	1.475, 1.474	$\bar{2}$ .10.3, 553

\*For the calculated pattern, only reflections with intensities  $\geq 1$  are given. \*\*For unit-cell parameters calculated from single-crystal data.

the band at  $632\text{ cm}^{-1}$  can be assigned as a Cu–O stretching mode and the strong band at  $352\text{ cm}^{-1}$  as an O–As–O bending mode. The absence of bands with frequencies higher than  $1000\text{ cm}^{-1}$  indicates the absence of groups with O–H, C–H, C–O, N–H, N–O and B–O bonds in the mineral.

### Chemical composition

The chemical composition of melanarsite was studied using a Jeol JSM-6480LV scanning electron microscope equipped with an INCA-Wave 500 wavelength-dispersive spectrometer (Laboratory of

TABLE 3. Crystal data, data collection information and structure refinement details for melanarsite.

Formula	$\text{K}_3\text{Cu}_7\text{Fe}(\text{AsO}_4)_4\text{O}_4$
Formula weight	1237.61
Temperature, K	293(2)
Radiation and wavelength, Å	MoK $\alpha$ ; 0.71073
Crystal system, space group, $Z$	Monoclinic, $C2/c$ ; 4
Unit-cell dimensions, Å, °	$a = 11.4763(9)$ $b = 16.620(2)$ $\beta = 105.078(9)$ $c = 10.1322(8)$
$V$ , Å <sup>3</sup>	1866.0(3)
Absorption coefficient $\mu$ , mm <sup>-1</sup>	16.384
$F_{000}$	2312
Crystal size, mm	$0.03 \times 0.05 \times 0.06$
Diffractometer	Xcalibur S CCD
$\theta$ range for data collection, °	2.69–26.37
Index ranges	$-14 \leq h \leq 14$ , $-20 \leq k \leq 20$ , $-12 \leq l \leq 12$
Reflections collected	14,009
Independent reflections	1918 ( $R_{\text{int}} = 0.1256$ )
Independent reflections with $I > 2\sigma(I)$	1652
Structure solution	direct methods
Refinement method	full-matrix least-squares on $F^2$
Number of refined parameters	160
Final $R$ indices [ $I > 2\sigma(I)$ ]	$R1 = 0.0914$ , $wR2 = 0.1746$
$R$ indices (all data)	$R1 = 0.1095$ , $wR2 = 0.1823$
Goof	1.278
Largest diff. peak and hole, $e\text{ Å}^{-3}$	2.355 and $-1.569$

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TABLE 4. Atom coordinates, equivalent isotropic displacement parameters ( $\text{\AA}^2$ ) and site multiplicities ( $Q$ ) for melanarsite.

Atom	$x/a$	$y/b$	$z/c$	$U_{\text{eq}}$	$Q$
As(1)	0.5	0.65808(14)	0.75	0.0106(5)	4
As(2)	0.5	0.77774(14)	0.25	0.0121(6)	4
As(3)	0.79713(15)	0.56128(10)	0.24785(17)	0.0096(4)	8
(Fe <sub>0.5</sub> Cu <sub>0.5</sub> )*	0.9638(2)	0.66941(14)	0.0859(2)	0.0143(6)	8
Cu(1)	0.45787(19)	0.95338(12)	0.0939(2)	0.0111(5)	8
Cu(2)	0.76820(18)	0.42242(12)	0.4141(2)	0.0105(5)	8
Cu(3)	0.69130(17)	0.70371(12)	0.0866(2)	0.0091(5)	8
K(1)	0.5	0.4451(4)	0.75	0.0347(17)	4
K(2)	0.2957(4)	0.3422(4)	0.4385(6)	0.0405(13)	8
O(1)	0.8811(11)	0.9024(7)	0.2321(13)	0.016(3)	8
O(2)	0.6378(12)	0.0269(7)	0.7750(13)	0.018(3)	8
O(3)	0.7452(11)	0.5435(7)	0.3863(13)	0.018(3)	8
O(4)	0.3845(13)	0.2805(8)	0.7594(15)	0.028(3)	8
O(5)	0.8846(12)	0.6440(8)	0.2740(17)	0.031(4)	8
O(6)	0.9493(12)	0.6662(8)	0.6054(14)	0.023(3)	8
O(7)	0.9369(10)	0.5613(7)	0.0250(11)	0.011(2)	8
O(8)	0.8151(10)	0.9149(6)	0.8902(12)	0.011(2)	8
O(9)	-0.0094(11)	0.7844(7)	0.1091(12)	0.013(3)	8
O(10)	0.6970(10)	0.8140(7)	0.0296(11)	0.010(2)	8

\*The site content was refined and fixed in the last stages of the refinement.

local methods of matter investigation, Faculty of Geology, Moscow State University), with an acceleration voltage of 20 kV, a beam current of 20 nA and a 3  $\mu\text{m}$  beam diameter. The following standards were used: orthoclase (K),  $\text{CaWO}_4$  (Ca),  $\text{CuFeS}_2$  (Fe), ZnS (Zn),  $\text{Al}_2\text{O}_3$  (Al), ilmenite (Ti), GaP (P) and FeAsS (As). The average (seven spot analyses) chemical composition of melanarsite (wt.%, ranges are in parentheses) is  $\text{K}_2\text{O}$  10.70 (10.40–10.91),  $\text{CaO}$  0.03 (0.00–0.10),  $\text{CuO}$  45.11 (44.51–46.25),  $\text{ZnO}$  0.24 (0.00–0.55),  $\text{Al}_2\text{O}_3$  0.32 (0.00–1.36),  $\text{Fe}_2\text{O}_3$  6.11 (5.49–7.07),  $\text{TiO}_2$  0.12 (0.00–0.47),  $\text{P}_2\text{O}_5$  0.07 (0.00–0.25),  $\text{As}_2\text{O}_5$  36.86 (35.94–37.77), total 99.56. Contents of other elements with atomic numbers higher than carbon are below detection limits. Fe is considered as  $\text{Fe}^{3+}$  based on the structure data (see below) and taking into account the extremely oxidizing conditions of mineral deposition in the Arsenatnaya fumarole; all other known iron minerals from the locality contain only  $\text{Fe}^{3+}$  (Pekov *et al.*, 2014b, 2015c). The empirical formula of melanarsite, calculated based on 20 O apfu, is  $(\text{K}_{2.81}\text{Ca}_{0.01})_{\Sigma 2.82}(\text{Cu}_{7.02}\text{Fe}_{0.95}\text{Al}_{0.08}\text{Zn}_{0.04}\text{Ti}_{0.02})_{\Sigma 8.11}(\text{As}_{3.97}\text{P}_{0.01})_{\Sigma 3.98}\text{O}_{20}$ . The simplified formula is  $\text{K}_3\text{Cu}_7\text{Fe}^{3+}\text{O}_4(\text{AsO}_4)_4$ ,

which requires  $\text{K}_2\text{O}$  11.42,  $\text{CuO}$  44.99,  $\text{Fe}_2\text{O}_3$  6.45,  $\text{As}_2\text{O}_5$  37.14, total 100.00 wt.%.

## X-ray crystallography

Powder X-ray diffraction data of melanarsite (Table 2) were collected with a Rigaku R-Axis Rapid II single-crystal diffractometer equipped with a cylindrical image plate detector using Debye-Scherrer geometry ( $d = 127.4$  mm;  $\text{CoK}\alpha$  radiation). Parameters of the monoclinic unit cell calculated from the powder data are  $a = 11.460(8)$ ,  $b = 16.630(6)$ ,  $c = 10.146(8)$   $\text{\AA}$ ,  $\beta = 105.21(5)^\circ$  and  $V = 1866(3)$   $\text{\AA}^3$ .

Single-crystal X-ray studies of melanarsite were carried out using an Xcalibur S diffractometer equipped with a CCD detector. The data were corrected for Lorentz and polarization effects. Data reduction was performed using *CrysAlis Pro*, Version 1.171.35.21 (Agilent Technologies, 2012), including analytical numeric absorption correction using a multifaceted crystal model based on expressions derived by Clark and Reid (1995). The crystal structure of the mineral was

TABLE 5. Selected interatomic distances (Å) in the structure of melanarsite.

As(1)	–O(1)	1.666(12) × 2	Cu(2)	–O(10)	1.901(12)
As(1)	–O(9)	1.697(12) × 2		–O(1)	1.981(13)
<As(1)–O>		1.68		–O(7)	1.988(11)
				–O(3)	2.040(12)
As(2)	–O(4)	1.664(13) × 2		–O(8)	2.418(12)
	–O(6)	1.706(13) × 2		–O(2)	2.574(12)
<As(2)–O>		1.68	<Cu(2)–O>		2.15
As(3)	–O(5)	1.682(13)	Cu(3)	–O(10)	1.928(11)
	–O(8)	1.683(11)		–O(10)	1.975(11)
	–O(3)	1.687(13)		–O(8)	1.989(11)
	–O(2)	1.688(12)		–O(4)	1.991(13)
<As(3)–O>		1.68		–O(9)	2.485(12)
(Fe,Cu)	–O(7)	1.900(12)		–O(5)	2.706(14)
	–O(10)	1.931(12)	<Cu(3)–O>		2.18
	–O(9)	1.941(11)	K(1)	–O(3)	2.800(13) × 2
	–O(5)	1.983(14)		–O(1)	2.862(14) × 2
	–O(5)	2.354(16)		–O(8)	2.888(11) × 2
	–O(6)	2.400(14)		–O(4)	3.052(15) × 2
<(Fe,Cu)–O>		2.08	<K(1)–O>		2.90
Cu(1)	–O(7)	1.916(12)	K(2)	–O(1)	2.720(13)
	–O(7)	1.929(11)		–O(3)	2.722(13)
	–O(2)	1.958(12)		–O(6)	2.734(14)
	–O(6)	1.995(13)		–O(9)	2.790(12)
	–O(2)	2.401(14)		–O(2)	3.033(14)
	–O(3)	2.777(13)		–O(4)	3.210(17)
<Cu(1)–O>		2.16		–O(4)	3.308(15)
			<K(2)–O>		2.93

TABLE 6. Bond-valence calculations for melanarsite\*.

	As(1)	As(2)	As(3)	(Fe,Cu)	Cu(1)	Cu(2)	Cu(3)	K(1)	K(2)	Σ
O(1)	1.31 <sup>x2↓</sup>					0.44		0.14 <sup>x2↓</sup>	0.20	2.09
O(2)			1.24		0.47	0.09			0.09	2.03
					0.14					
O(3)			1.24		0.05	0.38		0.16 <sup>x2↓</sup>	0.20	2.03
O(4)		1.32 <sup>x2↓</sup>					0.43	0.08 <sup>x2↓</sup>	0.05	1.92
									0.04	
O(5)			1.26	0.49			0.06			1.99
				0.18						
O(6)		1.18 <sup>x2↓</sup>		0.16	0.43				0.19	1.96
O(7)				0.62	0.53	0.43				2.09
					0.51					
O(8)			1.25			0.14	0.43	0.13 <sup>x2↓</sup>		1.95
O(9)	1.21 <sup>x2↓</sup>			0.56			0.11		0.17	2.05
O(10)				0.56		0.55	0.51			2.07
							0.45			
Σ	5.04	5.00	4.99	2.57	2.13	2.03	1.99	1.02	0.94	

\*Bond-valence parameters were taken from Brese and O'Keeffe (1991).



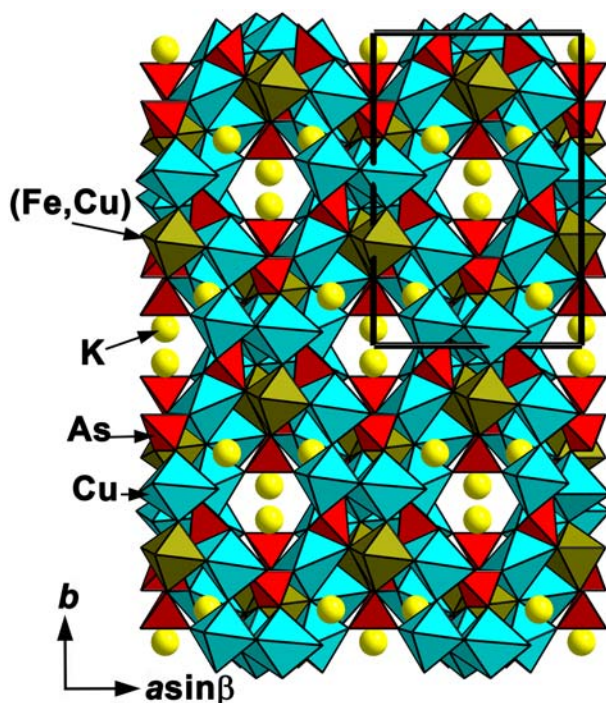


FIG. 7. The crystal structure of melanarsite projected along the  $c$  axis. The unit cell is outlined.

solved by direct methods and refined with the use of SHELX-97 software package (Sheldrick, 2008) to  $R = 0.091$ . The unit-cell parameters and the experimental details are presented in Table 3, atom coordinates and displacement parameters in Table 4, selected interatomic distances in Table 5 and bond-valence calculations in Table 6. Even the best of all tested single crystals of melanarsite showed, unfortunately, not very high quality because of a slightly divergent character responsible for the relatively high final value of  $R = 9.1\%$ . However reasonable values of displacement parameters (Table 3), interatomic distances (Table 4) and bond-valence parameters (Table 5), as well as good agreement between the measured and calculated powder X-ray diffraction data (Table 2) demonstrate that the determined structure is correct.

## Discussion

No mineral or synthetic compound close to melanarsite in terms of crystal structure is known. Two recently discovered anhydrous oxyarsenates with species-defining Cu and K, shchurovskyite  $K_2CaCu_6O_2(AsO_4)_4$  and dmisokolovite

$K_3Cu_5AlO_2(AsO_4)_4$  (Pekov *et al.*, 2015c), are related to one another but differ from melanarsite in both stoichiometry and structure. All three of these K-Cu arsenates are closely associated and can be easily distinguished from each other due to their quite different colours (Fig. 1).

The crystal structure of melanarsite (Fig. 7) is based upon the heteropolyhedral pseudo-framework built by distorted  $Cu(1-3)O_6$  and  $(Fe, Cu)O_6$  octahedra and  $As(1-3)O_4$  tetrahedra. All Cu-centred octahedra are characterized by strong Jahn-Teller distortion with four short Cu–O distances in the ranges 1.916(12)–1.995(13) Å [Cu(1)], 1.901(12)–2.040(12) Å [Cu(2)] and 1.928(11)–1.991(13) Å [Cu(3)] and two elongated Cu–O bonds, 2.401(14) and 2.777(13) Å for Cu(1), 2.418(12) and 2.574(12) Å for Cu(2), and 2.485(12) and 2.706(14) Å for Cu(3). The  $(Fe, Cu)O_6$  octahedron, occupied by  $Fe^{3+}$  and  $Cu^{2+}$  cations as  $(Fe_{0.5}Cu_{0.5})$ , is also characterized by Jahn-Teller distortion but not as strong; four short  $(Fe, Cu)$ –O distances are in the range 1.900(12)–1.983(14) Å whereas two elongated bonds are 2.354(16) and 2.400(14) Å.  $Cu(1)O_6$  octahedra are connected *via* common edges to form zig-zag chains running

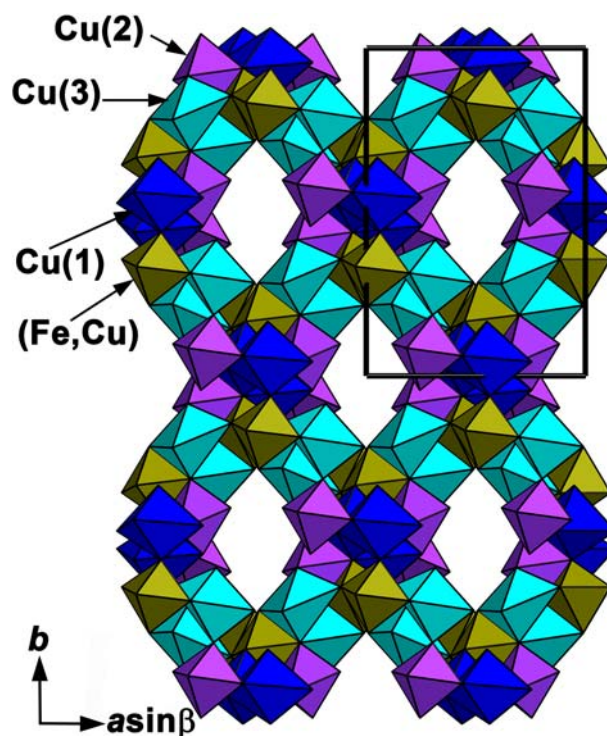


FIG. 8. The motif formed by Cu- and (Fe,Cu)-centred polyhedra in the crystal structure of melanarsite. The unit cell is outlined.

along the  $c$  axis. The Cu(3)-centred octahedra share common edges forming dimers, as well as the (Fe, Cu)-centred octahedra. The  $[\text{Cu}(3)_2\text{O}_{10}]$  and  $[(\text{Fe}, \text{Cu})_2\text{O}_{10}]$  dimers share common edges building complicated chains connected on two sides by isolated Cu(2) $\text{O}_6$  octahedra. These fragments are linked by the chains of Cu(1)-centred octahedra (Fig. 8). For better clarity, octahedra occupied by three crystallographically independent Cu cations are shown in different colours in Fig. 8. Three crystallographically independent As atoms occupy  $\text{AsO}_4$  tetrahedra connected with Cu- and (Fe,Cu)-centred octahedra via common O vertices to form a heteropolyhedral pseudo-framework with tunnels running along the  $c$  axis. Two crystallographically independent  $\text{K}^+$  cations, K(1) and K(2), are located in the tunnels and voids of the pseudo-framework centring eight- and seven-fold polyhedra, respectively.

Oxygen atoms O(7) and O(10) in melanarsite are bound only to Cu and (Fe,Cu) and, therefore, the crystal structure of the mineral could be described

in terms of anion-centred tetrahedra. The O-centred tetrahedra  $\text{OCu}_3(\text{Fe,Cu})$  are connected via common Cu(1)–Cu(1), Cu(3)–Cu(3) and Cu(2)–(Fe,Cu) edges to form infinite, isolated chains of composition  $[\text{O}_4\text{Cu}_6(\text{Fe,Cu})_2]^\infty$  (Fig. 9). Similar chains were reported in piypite  $\text{K}_8\text{Cu}_9\text{O}_4(\text{SO}_4)_8\text{Cl}_2$  (Effenberger and Zemann, 1984; described as caratiite), its synthetic sodium analogue  $\text{Na}_4\text{Cu}_4\text{O}_2(\text{SO}_4)_4 \cdot \text{MeCl}$  ( $M = \text{Na}, \text{Cu}, \square$ ) (Kahlenberg *et al.*, 2000), coparsite  $\text{Cu}_4\text{O}_2[(\text{As}, \text{V})\text{O}_4]\text{Cl}$  (Starova *et al.*, 1998) and recently discovered wulfite  $\text{K}_3\text{NaCu}_4\text{O}_2(\text{SO}_4)_4$  and parawulfite  $\text{K}_5\text{Na}_3\text{Cu}_8\text{O}_4(\text{SO}_4)_8$  (Pekov *et al.*, 2014d).

The crystal chemical formula of melanarsite is  $\text{K}_{(1)}^{(2)}\text{K}_{(2)}^{(2)}\text{Cu}_{(1)}^{(2)}\text{Cu}_{(2)}^{(2)}\text{Cu}_{(2)}^{(3)}\text{Cu}_{(2)}^{(\text{Fe,Cu})}(\text{Fe}_{0.5}^{3+}\text{Cu}_{0.5})_2\text{O}_4(\text{AsO}_4)_4$  (Table 3) and corresponds to the idealized formula  $\text{K}_3\text{Cu}_7\text{Fe}^{3+}\text{O}_4(\text{AsO}_4)_4$ . The presence of iron in the (Fe,Cu) site as  $\text{Fe}^{3+}$  is confirmed by bond-valence calculations (Table 5) and the distinctly smaller size of this octahedron in comparison with Cu(1–3)-centred octahedra; average interatomic distances are 2.16, 2.15 and 2.18 Å for the latter

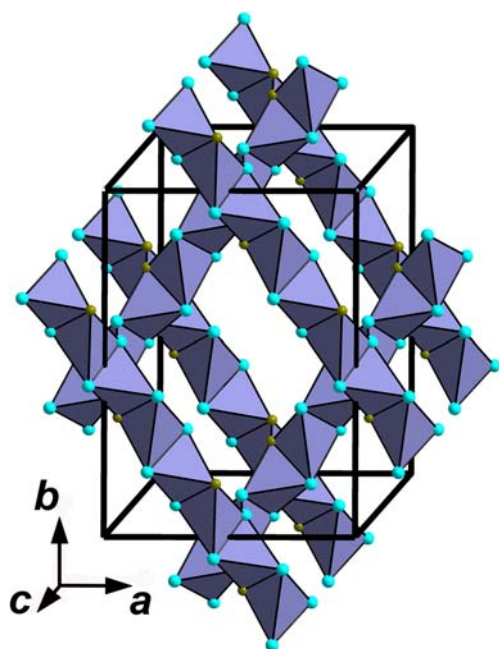


FIG. 9. The chain motif formed by oxocentred tetrahedra  $\text{OCu}_3(\text{Fe,Cu})$  in melanarsite. Blue circles are Cu(1–3) sites, olive green circles are (Fe,Cu) sites. The unit cell is outlined.

and 2.08 Å for the former (Table 4). The Fe:Cu ratio in this site is close to 1.00, ( $\text{Fe}_{0.5}^{3+}\text{Cu}_{0.5}^{2+}$ ). A very similar situation was reported for the Cu-rich variety of lyonsite,  $\text{Cu}_{3+x}(\text{Fe}_{4-2x}^{3+}\text{Cu}_{2x})(\text{VO}_4)_6$ , found in the Yadovitaya fumarole at the same Second scoria cone of the Northern Breakthrough of the Great Tolbachik Fissure Eruption; here the (Fe,Cu) site content is ( $\text{Fe}_{0.55}^{3+}\text{Cu}_{0.45}^{2+}$ ) (Pekov *et al.*, 2013). Significant substitution of  $\text{Fe}^{3+}$  by  $\text{Cu}^{2+}$  was reported for another sublimate mineral from the Tolbachik fumaroles, hatertite  $\text{Na}_2(\text{Ca,Na})(\text{Fe}^{3+},\text{Cu})_2(\text{AsO}_4)_3$ , in which the  $M(2)$  site contains ( $\text{Fe}_{0.45}^{3+}\text{Cu}_{0.20}^{2+}\text{Al}_{0.19}\text{Zn}_{0.16}$ ) (Krivovichev *et al.*, 2013).

The presence of only  $\text{Fe}^{3+}$  or its strong preference over  $\text{Fe}^{2+}$  is a typical feature of the sublimate minerals formed in the fumaroles at the Second scoria cone of the Northern Breakthrough of the Great Tolbachik Fissure Eruption, characterized by strongly oxidizing conditions (Vergasova and Filatov, 2012; Pekov *et al.*, 2014b). This was obviously caused by the mixing of volcanic gases with the atmospheric air here (Meniaylov *et al.*, 1980; Zelenski *et al.*, 2012). High oxygen fugacity in gases is clearly confirmed by the presence of S, As and Mo only in the forms of sulfate, arsenate

and molybdate minerals, respectively, in incrustations of fumaroles at the Second scoria cone of the Northern Breakthrough of the Great Tolbachik Fissure Eruption.

Under reducing conditions, As in volcanic fumaroles is incorporated mainly in different sulfides or in As-bearing amorphous sulfur formed at relatively low temperatures (below 300°C), as both empirical data [e.g. for sulfide-bearing fumaroles at such volcanoes as Vulcano (Aeolian Archipelago, Italy), Satsuma-Iwojima (Kyushu, Japan) and Mutnovsky (Kamchatka, Russia)] and thermodynamic modelling show (Symonds *et al.*, 1987; Symonds and Reed, 1993; Cheynet *et al.*, 2000; Africano *et al.*, 2002; Zelenski and Bortnikova, 2005). However, an increase of oxygen fugacity (as a result of the mixing of volcanic gases with the air) may reduce the volatility of some elements, including As (Africano *et al.*, 2002), and this could increase the temperature of crystallization of arsenic (at the first time, of  $\text{As}^{5+}$ ) phases in a fumarolic system. This assumption is in a good agreement with our temperature measurements in pockets with arsenate minerals in the Arsenatnaya fumarole, 360–380°C for the assemblage with melanarsite and up to 420°C for assemblages with mainly alluaudite-group arsenates and tilasite (measured immediately after the uncovering of pockets). In general, in fumaroles belonging to different types in terms of  $f_{\text{O}_2}$ , the temperature of formation of arsenic mineralization is different. Arsenates crystallize at significantly higher temperatures than arsenic sulfides and As-containing native sulfur.

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