Klaprothite, péligotite and ottohahnite, three new minerals with bidentate UO_7 -SO₄ linkages from the Blue Lizard mine, San Juan County, Utah, USA

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ABSTRACT

The new minerals klaprothite (IMA2015-087), Na₆(UO₂)(SO₄)₄(H₂O)₄, péligotite (IMA2015-088), $Na_6(UO_2)(SO_4)_4(H_2O)_4$ and ottohahnite (IMA2015-098), $Na_6(UO_2)_2(SO_4)_5(H_2O)_7 \cdot 1.5H_2O$, were found in the Blue Lizard mine, San Juan County, Utah, USA, where they occur together as secondary phases. All three minerals occur as yellowish-green to greenish-yellow crystals, are brittle with irregular fracture, have Mohs hardness of $\sim 2\frac{1}{2}$ and exhibit bright bluish-green fluorescence, and all are easily soluble in roomtemperature H₂O. Only klaprothite exhibits cleavage; perfect on {100} and {001}. Quantitative energydispersive spectroscopy analyses yielded the empirical formulas $Na_{6,01}(U_{1,03}O_2)(S_{0,093}O_4)_4(H_2O)_4$ $Na_{5.82}(U_{1.02}O_2)(S_{1.003}O_4)_4(H_2O)_4$ and $Na_{5.88}(U_{0.99}O_2)_2(S_{1.008}O_4)_5(H_2O)_{8.5}$ for klaprothite, péligotite and ottohahnite, respectively. Their Raman spectra exhibit similar features. Klaprothite is monoclinic, $P2_1/c$, a = 9.8271(4), b = 9.7452(3), c = 20.8725(15) Å, $\beta = 98.743(7)^{\circ}, V = 1975.66(17)$ Å³ and Z = 4. Péligotite is triclinic, $P\overline{1}$, a = 9.81511(18), b = 9.9575(2), c = 10.6289(8) Å, $a = 88.680(6)^{\circ}$, $\beta = 73.990(5)^{\circ}$, $\gamma = 89.205(6)^{\circ}$, V = 998.22(8) Å³ and Z = 2. Ottohahnite is triclinic, $P\overline{1}$, a = 9.97562(19), b = 11.6741(2), c = 14.2903(10)Å, $\alpha = 113.518(8)^{\circ}$, $\beta = 104.282(7)^{\circ}$, $\gamma = 91.400(6)^{\circ}$, V = 1464.59(14) Å³ and Z = 2. The structures of klaprothite ($R_1 = 2.22\%$) and péligotite ($R_1 = 2.28\%$) both contain [(UO₂)(SO₄)₄]⁶⁻ clusters in which one SO_4 group has a bidentate linkage with the UO₇ polyhedron; Na–O polyhedra link clusters into thick heteropolyhedral layers and link layers into frameworks; the structures differ in the configuration of Na-O polyhedra that link the layers. The structure of ottohahnite $(R_1 = 2.65\%)$ contains $[(UO_2)_4(SO_4)_{10}]^{12-1}$ clusters in which each UO_7 polyhedron has a bidentate linkage with one SO_4 group; Na–O polyhedra link clusters into a thin heteropolyhedral slice and also link the slices into a framework. The minerals are named for Martin Heinrich Klaproth (1743-1817), Eugène-Melchior Péligot (1811-1890) and Otto Hahn (1879-1968).

Keywords: klaprothite, péligotite, ottohahnite, new mineral, uranyl sulfate, crystal structure, Blue Lizard mine, Utah, USA.

Introduction

In the last few years, the Blue Lizard uranium mine in southeast Utah has proven to be a prolific source of new minerals, especially sodium uranyl sulfates (Table 1). To date, our investigations have yielded

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TABLE 1. New uranyl sulfate minerals from the Blue	Lizard mine, arranged by their charge deficiency	per anion (CDA), and including their molar proportions of
H_2O , S and U.		

Mineral	Structural formula	CDA (vu)	mol. H ₂ O (%)	mol. S (%)	mol. U (%)	Reference
Shumwayite*	$[(UO_2)(SO_4)(H_2O)_2]_2 \cdot H_2O$	0.10	0.56	0.22	0.22	Kampf <i>et al.</i> (2017 <i>b</i>)
Plášilite	$Na(H_2O)_2[(UO_2)(SO_4)(OH)]$	0.17	0.45	0.18	0.18	Kampf <i>et al.</i> (2015 <i>a</i>)
Alwilkinsite-(Y)	$Y(H_2O)_7[(UO_2)_3(SO_4)_2O(OH)_3](H_2O)_7$	0.20	0.72	0.09	0.14	Kampf <i>et al.</i> (2017 <i>a</i>)
Bobcookite	$Na(\tilde{H}_2O)_2Al(\tilde{H}_2O)_6[(UO_2)_2(SO_4)_4(\tilde{H}_2O)_2]\cdot 8H_2O$	0.22	0.69	0.15	0.08	Kampf <i>et al.</i> (2015 <i>b</i>)
Wetherillite	$Na_{2}(H_{2}O)_{6}(Mg,Zn)(H_{2}O)_{6}[(UO_{2})_{2}(SO_{4})_{4}(H_{2}O)_{2}]\cdot 4H_{2}O$	0.22	0.67	0.15	0.07	Kampf <i>et al.</i> (2015 <i>b</i>)
Oppenheimerite	$Na_{2}(H_{2}O)_{2}[(UO_{2})(SO_{4})_{2}(H_{2}O)]$	0.22	0.38	0.25	0.13	Kampf <i>et al.</i> $(2015c)$
Ottohahnite	$Na_{6}(H_{2}O)_{7}[(UO_{2})_{2}(SO_{4})_{5}](H_{2}O)_{2}$	0.25	0.41	0.23	0.09	This study
Fermiite	$Na_4(H_2O)_3[(UO_2)(SO_4)_3]$	0.29	0.27	0.27	0.09	Kampf et al. (2015c)
Meisserite	$Na_{5}(SO_{3}OH)(H_{2}O)[(UO_{2})(SO_{4})_{3}]$	0.29	0.13	0.35	0.09	Plášil et al. (2013)
Klaprothtite	$Na_{6}(H_{2}O)_{4}[(UO_{2}O)(SO_{4})_{4}]$	0.33	0.27	0.27	0.07	This study
Péligotite	$Na_{6}(H_{2}O)_{4}[(UO_{2})(SO_{4})_{4}]$	0.33	0.27	0.27	0.07	This study
Bluelizardite	$Na_7Cl(H_2O)_2[(UO_2)(SO_4)_4]$	0.33	0.13	0.27	0.07	Plášil et al. (2014)
Belakovskiite	$Na_7(SO_3OH)(H_2O)_3[(UO_2)(SO_4)_4(H_2O)]$	0.34	0.21	0.30	0.06	Kampf et al. (2014)

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*The type localities for shumwayite are the Giveaway-Simplot mine and the Green Lizard mine, but it has also been found at the Blue Lizard mine. vu – valence units.



FIG. 1. Klaprothite in incandescent light (top) and in 405 nm light (bottom); field of view = 1 mm across.

eight new uranyl sulfate minerals, belakovskiite, bluelizardite, fermiite, meisserite, oppenheimerite, plášilite, bobcookite and wetherillite. The ninth, tenth and eleventh new uranyl sulfates from this mine, klaprothite, péligotite and ottohahnite, are described herein. These three minerals, which occur in association with one another, are the first three minerals with structures containing bidentate linkages (shared edges) between UO₇ pentagonal pyramids and SO₄ tetrahedra. The twelfth new uranyl sulfate from the Blue Lizard mine, alwilkinsite-(Y) will be described in a subsequent paper and another new one occurring here, shumwayite, was recently described from the nearby Green Lizard and Giveaway-Simplot mines.

Klaprothite (/'klæp rout ait/) is named in honour of German chemist Martin Heinrich Klaproth (1743–1817), the discoverer of uranium (1789), zirconium (1789) and cerium (1803). It should be noted that the material that Klaproth identified as a new element and named uranium was actually an oxide of the element. There have been several previous attempts to name a mineral for Klaproth,

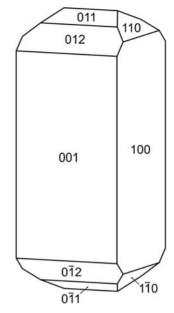


FIG. 2. Crystal drawing of klaprothite; clinographic projection in nonstandard orientation, **b** vertical.

each of which subsequently resulted in a discreditation. The most recent discreditation occurred in 1947, when klaprotholite was shown to be a mixture of wittichenite and emplectite. We think that sufficient time has passed to again propose the naming a new mineral in honour of this pioneer in analytical chemistry and mineralogy.

Péligotite (//pe li gou tait/) is named in honour of the French chemist Eugène-Melchior Péligot (1811–1890), who isolated the first sample of uranium metal in 1841 and proved that the material



FIG. 3. Péligotite with tamarugite (white); field of view = 1.1 mm across.

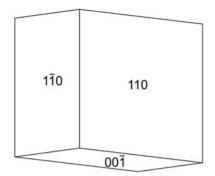


FIG. 4. Crystal drawing of péligotite; clinographic projection in standard orientation.

described by Klaproth as the element in 1789 was actually an oxide of uranium.

Ottohahnite (/au tOu 'ha:n ait/) is named in honour of German chemist Otto Hahn (1879–1968) who discovered nuclear fission (of uranium) in 1938, for which he received the Nobel Prize in Chemistry. The compound name is proposed to avoid the possible confusion of 'hahnite' with several other mineral names, i.e. hainite, gahnite and cahnite. In particular, in Russian, the pronunciation of hahnite would be essentially the same as that of gahnite.

The new minerals and their names were approved by the Commission on New Minerals, Nomenclature and Classification of the International Mineralogical Association: klaprothite (IMA2015-087), péligotite (IMA2015-088) and ottohahnite (IMA2015-098). The description of each mineral is based on five cotype specimens. Those deposited in the collections of the Natural History Museum of Los



FIG. 5. Ottohahnite with tamarugite (white); field of view = 0.7 mm across.

Angeles County, 900 Exposition Boulevard, Los Angeles, CA 90007, USA, have the catalogue numbers 65610 (cotype for klaprothite, péligotite and ottohahnite), 65611 (cotype for klaprothite), 65613 (cotype for klaprothite), 65613 (cotype for klaprothite), 65614 (cotype for péligotite and ottohahnite), 65616 (cotype for péligotite) and 65617 (cotype for ottohahnite). One cotype for each mineral is housed in the collections of the Fersman Mineralogical Museum of the Russian Academy of Sciences, Moscow, Russia, registration number 4778/1 (klaprothite), 4779/1 (péligotite) and 4782/1 (ottohahnite).

Occurrence

Klaprothite, péligotite and ottohahnite were found underground in the Blue Lizard mine, Red Canyon, White Canyon District, San Juan County, Utah, USA (37°33'26"N 110°17'44"W). The Blue Lizard mine is located ~72 km west of the town of Blanding, Utah and ~22 km southeast of Good Hope Bay on Lake Powell. It is on the north side of Red Canyon and close to the Markey mine. Information on the history and geology of the deposit is taken largely from Chenoweth (1993).

The deposit was first recognized in the summer of 1898 by John Wetherill, while leading an archaeological expedition into Red Canyon. He noted yellow stains around a petrified tree. At that spot, he built a rock monument, in which he placed a piece of paper to claim the minerals. Although he never officially recorded his claim, 45 years later, in 1943, he described the spot to Preston V. Redd of Blanding, Utah, who went to the site, found Wetherill's monument and claimed the area as the Blue Lizzard claim (note alternate spelling). Underground workings to mine uranium were not developed until the 1950s.

Mineralized channels are in the Shinarump member of the Chinle Formation. The Shinarump member consists of medium- to coarse-grained sandstone, conglomeratic sandstone beds and thick siltstone lenses. Ore minerals were deposited as replacements of wood and other organic material and as disseminations in the enclosing sandstone. Since the mine closed in 1978, hydration-oxidation weathering of primary uranium minerals, mainly uraninite, by acidic solutions derived from the decomposition of associated sulfides such as pyrite, marcasite and chalcopyrite in the humid underground environment has produced a variety of

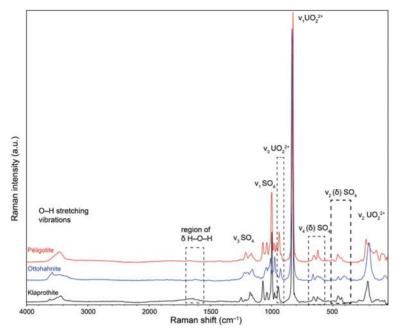


FIG. 6. The Raman spectra of klaprothite, péligotite and ottohahnite measured with a 532 nm laser.

secondary minerals, mainly sulfates, as efflorescent crusts on the surfaces of mine walls.

The new minerals are rare in the secondary uranyl sulfate mineral assemblage. Other secondary minerals found in direct association with klaprothite, péligotite and ottohahnite include blödite, bluelizardite, bobcookite, epsomite, gypsum, hexahydrite, konyaite, plášilite and tamarugite. The bulk of the matrix is comprised of subhedral to euhedral, equant quartz crystals that are recrystallized counterparts of the original grains of the sandstone. Other minerals remaining from the original sandstone include feldspar and rare almandine and zircon. Minerals related to the ore deposition include baryte, bornite, chalcopyrite, covellite, pyrite and uraninite. Other secondary minerals in the general assemblage include aluminocoquimbite, alwilkinsite-(Y), atacamite, belakovskiite, boyleite, brochantite, calcite, chalcanthite, cobaltoblödite, copiapite, coquimbite, cvanotrichite, d'ansite-(Mn), dickite, fermiite, ferrinatrite, dietrichite. gerhardtite, gordaite, halite, johannite, kaolinite, kieserite, kröhnkite, lishizhenite, manganoblödite, meisserite, metavoltine. natrozippeite, oppenheimerite, pseudojohannite, pickeringite. rhomboclase. römerite, rozenite, sideronatrite, shumwayite, thérèsemagnanite, wetherillite and other potentially new minerals currently under investigation.

ture by evaporative processes at the surface of a rock with high relative porosity. The environment is relatively oxidizing (high Eh) and solutions are generally acidic (low pH). The relative acidity prevalent during the formation of any given phase can be interpreted from its charge deficiency per anion (CDA) calculated using the bond-valence approach (cf. Hawthorne and Schindler, 2008; Hawthorne, 2012); higher CDA correlates with higher pH. The molar proportions of S and U in the formulas are indicative of the relative concentrations of these elements in solution during formation and the molar proportion of H2O can be expected to increase with the relative humidity (RH) during formation. Table 1 provides the CDA and molar proportions of H₂O, S and U for each of the new uranyl sulfates found at the Blue Lizard mine. Klaprothite and péligotite are likely to have formed from less acidic solutions than did ottohahnite, and most of the other uranyl sulfates at the Blue Lizard mine. Klaprothite and péligotite are presumed to have formed under relatively low RH, while ottohahnite is likely to have formed at moderate RH. The molar proportions of S and U for these

A detailed discussion of the origin and condi-

tions of formation of the uranyl sulfate assemblages at the Blue Lizard mine was provided by Plášil *et al.*

(2014). The minerals formed at ambient tempera-

		Klaprothite			Péligotite		Ottohahnite				
	Mean	Range	S.D.	Mean	Range	S.D.	Mean	Range	S.D.		
Na ₂ O	21.06	19.67-22.85	1.08	20.55	18.98-22.56	1.46	13.73	12.92-15.66	0.84		
UO ₃	33.14	31.07-36.92	1.82	33.27	29.61-38.14	3.53	42.68	41.01-44.39	0.99		
SO ₃	35.93	33.15-38.40	1.38	36.60	33.88-39.20	1.88	30.44	28.86-31.45	0.97		
H ₂ O	8.15*			8.22*			11.55§				
Total	98.28			98.64			98.40				

TABLE 2. Chemical analyses (in wt.%) for klaprothite, péligotite and ottohahnite.

S.D. - standard deviation.

*Calculated by stoichiometry on the basis of 22 O apfu. [§]Calculated by stoichiometry on the basis of 32.5 O apfu.

phases, with mol. S >> mol. U, suggest that they formed relatively far from the source of U (uraninite).

Physical and optical properties

Klaprothite

Klaprothite crystals are equant to prismatic (Figs 1 and 2), up to ~ 1 mm in maximum dimension, but usually much smaller. Crystal faces often are skeletal with rounded edges, suggestive of incomplete growth or later dissolution/deliquescence. Crystals typically occur in parallel intergrowths with surfaces made up of many stepped faces. Prisms are elongated on [010] and crystals exhibit the forms {100}, {001}, {110}, {011} and {012}. No twinning was observed.

Crystals are yellow green and transparent with a vitreous lustre. The streak is pale yellow green. The mineral fluoresces bright bluish green under both longwave and shortwave ultraviolet light. The Mohs hardness is ~2½. Crystals are brittle (slightly sectile) with perfect cleavages on {100} and {001} and have irregular fracture. The mineral is slightly deliquescent and is easily soluble in room-temperature H₂O. The density measured by flotation in a mixture of methylene iodide and acetone is 2.90(2) g cm⁻³. The calculated densities are 2.923 g cm⁻³ based on the empirical formula and 2.905 g cm⁻³ based on the ideal formula.

Optically, klaprothite is biaxial (-) with $\alpha = 1.497(1)$, $\beta = 1.517(1)$, $\gamma = 1.519(1)$ (measured in white light). The 2V measured directly on a spindle stage is $34(1)^{\circ}$; the calculated 2V is 34.7° . Dispersion is r > v, distinct. Pleochroism is X = colourless, Y = light yellowish green, Z = light yellowish green; $X < Y \approx Z$. The optical orientation is $Y = \mathbf{b}$, $X \wedge \mathbf{c} = 10^{\circ}$ in obtuse angle β .

Péligotite

Péligotite crystals are generally equant relatively simple rhombs (Figs 3 and 4), up to ~0.5 mm in maximum dimension, but usually much smaller. Crystal faces often are concave with rounded edges, suggestive of incomplete growth or later dissolution/ deliquescence. Crystals typically occur in subparallel aggregates and drusy intergrowths. Crystals exhibit the forms {001}, {110} and {110}; more complex forms are probably present, but are difficult to measure because crystals occur in tight intergrowths. No twinning was observed.

Crystals are yellow green and transparent with a vitreous lustre. The streak is pale yellow green. The mineral fluoresces bright bluish green under both longwave and shortwave ultraviolet light. The Mohs hardness is ~2½. Crystals are brittle (slightly sectile) with no cleavage and have irregular fracture. The mineral is slightly deliquescent and is easily soluble in room-temperature H₂O. The density measured by flotation in a mixture of methylene iodide and acetone is 2.88(2) g cm⁻³. The calculated densities are 2.878 g cm⁻³ based on the empirical formula and 2.875 g cm⁻³ based on the ideal formula.

Optically, péligotite is biaxial (-) with 1.493(1), $\beta = 1.511(1)$, $\gamma = 1.515(1)$ (measured in white light). The 2V measured directly on a spindle stage is 50(1)°; the calculated 2V is 50.0°. Dispersion is r > v, distinct. Pleochroism is X =colourless, Y = light yellowish green, Z = light yellowish green; $X < Y \approx Z$. The optical orientation is $X \wedge \mathbf{c} = 3^\circ$, $Y \wedge \mathbf{b} = 43^\circ$, $Z \wedge \mathbf{a} = 40^\circ$.

Ottohahnite

Ottohahnite crystals are generally equant (Fig. 5), up to ~ 0.1 mm in maximum dimension, but

TABLE 3.	Powder X-ray	data for k	laprothite,	péligotite and	ottohahnite.

Klap I _{obs}	d_{obs}		$d_{\rm calc}$	$I_{\rm calc}$	hkl	$I_{\rm obs}$	$d_{\rm obs}$	$d_{\rm cal}$	$I_{\rm calc}$	h k l	$I_{\rm obs}$	$d_{\rm obs}$	$d_{\rm cale}$	$I_{\rm calc}$	hkl	$I_{\rm obs}$	$d_{\rm obs}$	d_{cale}	$I_{\rm calc}$	h k
30	10.27		10.3150	28	002	31	2.716	∫ 2.729		224	16	2.0613	f 2.0630	13	0 0 10	11	1.6489	∫ 1.6547	3	511
8	9.72		9.7129	76	100			2.706		131	10	2.0244	2.0365	4	243			1.6525	3	43
			8.8115	4	011			2.702		134			2.0328	3	414			1.6455	2	42
			7.6777	7	102			2.696	5 6	320			2.0261	2	138			1.6391	2	51
7	7.09		7.0838	100	012			r 2.667) 2	232			2.0103	2	336			(1.6374	2	13
			6.8795	4	110	22	2.646	2.665	6	313			1.9957	2	423	23	1.6233	{ 1.6288	2	33
			6.5891	6	102			2.654	5 6	315			(1.9825	2	245			1.6272	6	11
			6.3250	2	111			2.648	5 5	231	17	1.9504	1.9584	6	1010			1.6242	3	41
36	6.04		6.0309	40	112			2.624	4 4	206			1.9549	4	341			1.6232	2	25
			5.6186	3	013			(2.613		126			(1.9404	3	051			1.6210	2	53
31	5.467		5.4585	41	112	35	2.570	2.590		134	48	1.9136	1.9254	13	512			1.6188	2	60
77	5.158		5.1575	82	004			2.578		008			1.9217	3	341			(1.6152	2	61
28	4.835		4.8722	25	104			2.574		Ī26			1.9196		343	12	1.5931	1.6020	4	31
	Titlererer		4.8565	3	200			2.559		232			1.9186		335			1.5943	2	35
			4.7421	11	021			2.552		035			1.9165	2	504			1.5919	3	35
14	4.557		4.5585	13	014			2.534		216			1.9156		1 2 10			(1.5810	3	43
	4.201	12	4.3750	2	211			(2.527		Ĩ35			1.9147	3	431	12	1.5640	{ 1.5689	3	16
58	4.330	ſ	4.3553	32	120	17	2.500	{ 2.524		324			1.9118	2	424	12	1,5040	1.5601	4	34
10	4.550	1	4.3213	9	121	17	2.500	2.521		027			1.9109	9	150	16	1.5361	{ 1.5404	3	33
			4.2929	23	104			2.492		018			1.9080	7	151	10	1.5501	1.5392	2	61
16	1.104	ſ			212															
40	4.194	1	4.2160	33				2.479		234	27	1.0751	1.8976	5	151			1.5350	2	33
			4.2038	2	121			2.475		316	27	1.8751	1.8833	7	418		1 6100	(1.5341	2	62
			4.1571	14	202			(2.436		040			1.8778	5	147	11	1.5129	1.5291	2	21
		ſ	4.1411	2	211	8	2.415	2.428		400			1.8752	5	427			1.5265	3	45
13	3.960	í	3.9757	5	023			2.419		041			1.8737	4	3010	12/22/	1000000	(1.5091	2	23
		ł	3.9286	7	114			2.403		108			1.8717	2	511	10	1.4898	{ 1.4980	2	41
28	3.811	3	3.8237	17	212			(2.374		412			(1.8702	2	432			1.4899	2	15
		(3.7983	11	123	34	2.357	{ 2.367		127	17	1.8492	1.8597	8	2210			1.4879	2	26
			3.7194	3	115			2.363		218			1.8555	2	153			(1.4848	3	05
13	3.572		3.5710	14	123			2.361		036			1.8483	9	318	12	1.4687	{ 1.4736	2	0.0
		1	3.5419	3	024			2.357	5 3	141	27	1.8145	1.8298	5	343			1.4724	2	45
00	3.434	3	3.4537	8	221			2.352	2 6	235			1.8252	6	239			1.4715	4	$\overline{2}4$
		U	3.4453	15	124			2.338	1 7	404			1.8171	2	1 2 10			(1.4645	2	52
			3.4397	68	220			(2.333.	5 3	118			1.8133	2	338	13	1.4456	{ 1.4508	3	62
			3.4083	5	106	19	2.303	\$ 2.310		226			1.8114	9	247			1.4473	2	54
		1	3.3739	15	222			2.306		331			1.8092	3	523			1.4467	2	15
37	3.227	Ł	3.2424	9	016			2.303) 2	411			1.8045	2	520			1.4404	2	36
		l	3,2376	14	300			2.296		043			r 1.8027	3	2010			(1.4336	2	51
			3.2211	6	124			2.291	4 3	332	16	1.7881	1.7980	5	139	17	1.4213	{ 1.4292	2	63
			3.2172	9	Ī16			(2.274		324			1.7931	3	251			1.4257	3	12
			3.2089	4	031	10	2.260	2.265		326			1.7908	3	238			1.4225	2	16
			3.1625	4	222			2.260	5	143			1.7835	4	524			1.4204	3	44
			3.1210	5	214			(2.252		331			c 1.7709	2	048			(1.4158	2	31
55	3.082	J	3.1049	15	311	11	2.215	2.226		412	16	1.7528	1.7623	5	055	10	1.4010	1.4106	4	22
~	0.002	l	3.0966	10	106			2.216		236			1.7541	4	155			1.4013	3	36
			3.0807	6	130			2.209		143			1.7508		434			(1.3929		16
				5											3210	10	1.3795	{ 1.3929		71
			3.0725		$\frac{310}{312}$			2.196		306			1.7488			15	1.3795			
	2.012	ſ	3.0682	38				2.192		421	20	1 71 50	1.733		$\frac{426}{441}$			1.3842		45
51	3.012	1	3.0257	10	131	10		(2.189		332	26	1.7159	{ 1.7294					1.3830		
		6	3.0154	41	224	19	2.179	{ 2.184		334			1.7199		440	0.5	10.000	(1.3770		
		ſ	2.9917	12	132			2.181		241			1.7140		155	9	1.3531	{ 1.3544		40
2	2.930	1	2.9518	10	223			2.176		235			1.7124		443			1.3474		71
		C	2.9136	7	132			2.173		420			1.7063		514	15	1.3361	{ 1.3416		44
		ſ	2.8951	2	216			2.158		423			1.7053		4010			1.3406		14
6	2.816	1	2.8349	14	312			2.142		406			(1.697)		347			1.3391	4	27
		L	2.8248	3	314			(2.128		327	25	1.6852	{ 1.6930		0112			1.3356		71
			2.8210	2	017	18	2.100	{ 2.118		243			1.6922	3	255			(1.3262		45
			2.8078	7	215			2.108		424			1.6907		2 2 10	9	1.3166	{ 1.3189		15
			2.7929	7	126			2.102	3 3	335			1.6820	3	435			1.3157	2	17
			2.7622	2	133			2.096		218			1.6808		532	11	1.2955	1.3046		22
								2.092		416			1.679		531			1.2960		71
								2.083		T 0 10			1.678		2112			1,2928		11
																		1.2889		

usually much smaller. Crystal faces and edges are often rounded, suggestive of incomplete growth or later dissolution/deliquescence. Crystals are typically intergrown with tamarugite. No crystals forms could be measured. No twinning was observed.

Crystals are green yellow and transparent with a vitreous lustre. The streak is pale green yellow. The mineral fluoresces bright bluish green under both

longwave and shortwave ultraviolet light. The Mohs hardness is ~2½. Crystals are brittle (slightly sectile) with no cleavage and have irregular fracture. The mineral is slightly deliquescent and is easily soluble in room-temperature H₂O. The small size and difficulty of separating crystals from their intergrowths with tamarugite made the measurement of the density impractical. The calculated densities are 2.858 g cm⁻³ based on

TABLE 3. (contd.)

lobs	gotite d_{obs}		denle	Lat	h k l	I _{obs}	$d_{\rm obs}$	d_{calc}	Law	h k l	I _{obs}	dobs	d_{cale}	Lat	hkl	I _{obs}	d_{obs}	<i>d</i>	Icale	h k
005	Pobs		cale	cate		-00s	obs	calc	caic		-00s	CODS	cale	caje		-obs	- 00s	cale	-cate	
39	10.19		10.2146	42	001	18	2.816	2.8199	3	ĪĪ3	7	2.1394	2.1448	3	242			(1.7115	4	35
			9.9546	11	010			2.8161	14	132			2.1378	6	420	13	1.6908	{ 1.7052	2 2	440
48	9.51		9.4341	70	100			2.8083	5	032			(2.1221	7	411			1.7039	2	05:
32	8.15	1	8.1378	49	101			2.7966	3	301	3	2.0958	{ 2.0842	3	115			1.7024	2	000
00	7.11	3	7.2014	44	011			2.7552	7	231			2.0823	2	124			1.6952	4	53:
		ų	7.0589	56	011			(2.7511	2	321			2.0622	3	215			1.6893	6	311
			6.8744	8	110	24	2.717	2.7295	9	223	11	2.0459	2.0429	16	005			1.6756	2	25:
			6.8209	12	T10			2.7240	5	230			(2.0242	2	204			(1.6725	2	01
			6.3716	10	111			2.7146	4	321	9	1.9919	2.0037	3	234	8	1.6605	1.6659	2	52
32	6.23		6.2316	34	111			2.7107	5	231			2.0022	5	414			1.6636	2	32.
	Crimer,		6.1384	3	101			2.6934	7	311			1.9909	5	050			1.6555	3	13:
			5.2411	11	111			2.6913	6	311			1.9875	2	421			1.6524	2	443
			5.2088	10	ÎÎI			2.6700	8	223			1.9837	2	431			1.6347	2	530
			5.1209	3	102			(2.6492	2	320			1.9656	5	432			1.6334	2	44
63	5.14		5.1073	71	002	21	2.639	2.6450	7	322			1.9590	2	340			(1.6316	2	13
8						21	2.039	2.6324		313				2	340	15	1 6325		2	50
	4.79	ſ	4.8175	8	201				10	222			1.9557			15	1.6235	1.6276		
43	4.54	1	4.5994	15	112			2.6206	10	$\frac{2}{2}$ $\frac{2}{2}$ $\frac{2}{2}$ $\frac{2}{2}$			1.9535	2	502			1.6247	2	22:
			4.5815	22	012			2.6044	4	222			(1.9516	2	315			1.6230	5	351
			4.5100	23	021	12521	12/12/22	(2.5605	2	204	26	1.9434	{ 1.9449	20	150			1.6208	4	020
		ſ	4.4396	2	011	17	2.552	2.5566	6	123			1.9409	4	332			1.6185	2	161
53	4.307	1	4.3880	22	120			2.5537	5	004			1.9357	2	125			1.6110	2	154
		(4.3663	2	211			2.5508	3	132			1.9335	6	315			1.6084	2	51
			4.3071	2	211			2.5434	3	232			(1.9238	2	151			1.6056	2	43:
			4.2896	13	121			2.5269	4	231	30	1.9131	{ 1.9229	4	334			r 1.6044	2	16
			4.2757	18	210			(2.4886	2	040			1.9163	7	225	6	1.5889	1.5988	2	611
		1	4.2498	11	210	9	2.471	2.4855	6	014			1.9153	2	430			1.5905	2	254
21	4.078	Į	4.0689	12	202			2.4653	3	$2\bar{1}4$			1.9132	16	511			1.5841	3	110
28		ļ	4.0485	18	102			2.4617	4	014			1.9101	8	412			(1.5752	2	34:
13	3.852	Ţ	3.8792	4	121			2.4509	2	401			1.8917	2	105	7	1.5602	1.5690	3	260
10	01004	l	3.8530	9	121			2.4396	3	321			(1.8893	2	143	- C.	110000	1.5645	2	531
			3.7967	3	212			2.4291	4	041	12	1.8716	1.8879	3	323			1.5638	2	045
16	3.754	ſ	3.7667	7	112			2.4110	2	140	14	1.0710	1.8868	2	500			1.5604	3	33.
10	3.754	1																		
	3 600	ì	3.7369	15	212		0.000	2.4088	2	402			1.8853	2	503			1.5483	3	33.
19	3.598	1	3.6274	14	211	18	2.392	2.4069	4	041			1.8684	3	424	- 11	1.5317	1.5424	3	130
			3.6133	3	122			2.4005	5	033			1.8670	6	224			1.5358	4	43:
			3.6007	7	022			2.3951	2	141			(1.8549	3	343			1.5346	2	40.
20	3.526		3.5282	27	103			2.3921	4	213	10	1.8430	{ 1.8514	3	251			1.5331	2	330
			3.4921	9	221			2.3870	5	411			1.8460	3	513			1.5243	2	262
		٢	3.4372	3	220			2.3806	3	2ī3			1.8425	2	052			1.5171	2	207
73	3.418	3	3.4319	35	221			(2.3663	5	124			1.8374	2	144			1.5151	3	414
		ų	3.4105	28	220	20	2.340	{ 2.3585	4	400			1.8288	2	251			(1.4844	2	111
			3.4049	29	003			2.3530	3	033			1.8264	2	521	4	1.4599	\$ 1.4693	2	424
		1	3.3182	6	030			2.3460	9	331			r 1.8190	2	431			1.4624	3	510
11	3.249	1	3.2695	8	301			2.3319	2	412	15	1.8065	{ 1.8134	5	343	3	1.4307	{ 1.4356	2	15:
		t	3.2021	6	013			2.3303	2	312			1.8088	2	522			1.4299	2	363
		1	3.1601	6	122			2.3197	2	304			1.8058	6	135			1.4286	2	62
74	3.121	Į	3.1379	31	130			(2.3119	3	141			1.8045	2	252			1.4100	2	220
		l	3.1198	40	311	11	2.279	{ 2.3008	5	141			1.7955	4	152			(1.4078	2	170
			3.1158	20	222			2.2910	5	410			1.7783	3	125	8	1.3829	1.3922	3	040
			3.0972	10	131			2.2816	3	332			1.7742	2	523		110000	1.3908	2	712
			3.0840	6	302			2.2600	3	142			1.7689	2	520			1.3831	5	15
					213				2	1142				4	044					46
24	2.059	ſ	3.0793	6	213			2.2579	2				1.7647			e	1 2670	£ 1.3822	2	
34	3.058	1	3.0684	6	$\frac{\overline{2}}{\overline{2}}\frac{2}{\overline{2}}\frac{1}{1}$			2.2547		$2 \overline{2} 4$ $\overline{2} \overline{3} 2$	10	1 750 4	∫ 1.7624	2	106	3	1.3578	1.3548		53:
			3.0650	30		10	0.010	(2.2454	2		10	1.7584	1.7593		304		1.226-	1.3471	3	71.
	0.000	ſ	3.0485	6	131	16	2.212	2.2371	8	332			1.7576	6	441	8	1.3356	{ 1.3401	2	51
11	2.983	1	3.0054	13	310			2.2198	5	$0\frac{4}{4}2$			1.7522	4	434			1.3395	2	72
		C	2.9919	2	310			2.2172	3	142			(1.7473	2	514			1.3387	2	514
			2.9625	2	312			2.2082	4	240	16	1.7263	{ 1.7351	9	441			1.3369	2	710
			2.9533	2	103			2.2052	3	413			1.7338	2	053			1.3361	2	25
			2.9296	3	312			2.1940	3	240			1.7296	3	314			1.3350	3	44
			2.9110	2	123			2.1876	4	421			1.7286	3	403			1.3342	2	72:
			2.8468	4	123			2.1857	4	413			1.7242	2	342					
			2.8429	2	113			2.1639	2	322			1.7186	2	234					
			2.8368	3	023			2.1535	2	422			1.7166	4	153					
			210200		0.00				~				1.1100		100					

the empirical formula and 2.973 $\,{\rm g~cm^{-3}}$ based on the ideal formula.

Optically, ottohahnite is biaxial (-) with $\alpha = 1.511(1)$, $\beta = 1.514(1)$, $\gamma = 1.537(1)$ (measured in white light). The 2V measured directly on a spindle stage is 42(2)°; the calculated 2V is 40.2°. Dispersion is r > v, strong. Pleochroism is X = light yellowish green, Y = colourless, Z = yellowish

green; Y < X < Z. The optical orientation is $X \land \mathbf{a} = 43^{\circ}$, $Y \land \mathbf{b} = 11^{\circ}$, $Z \land \mathbf{c} = 29^{\circ}$.

Raman spectroscopy

The Raman spectra of klaprothite, péligotite and ottohahnite were recorded on a Thermo-Scientific DXR Raman microscope interfaced to an Olympus

TABLE	3.	(contd.))
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Otto	hahnite														
I _{obs}	d_{obs}		d _{cale}	Icalc	hkl	I _{obs}	d_{obs}	$d_{\rm calc}$	Icale	hkl	I _{obs}	dobs	d_{calc}	Icale	hkl
12	12.67		12.5801	10				3.3808							$4\overline{2}1$
3	10.64		10.6016	2				3.2077	2	1 2 2	6	2 109	$ \left\{ \begin{array}{c} 2.2322 \\ 2.2136 \\ 2.1935 \\ 2.1900 \\ 2.1405 \end{array} \right. $	2 3	$\frac{421}{154}$
3	10.04	1	10.5726	2	$010 \\ 011$	25	3.156	5.2077	22	1 3 3	0	2.196	2.2130	2	$\frac{1}{3}$ $\frac{3}{4}$ 0
10	9.54					55	5.150	{ 3.1607 3.1410	25	222			2.1955	2 2	$\frac{340}{142}$
10			9.5744	10	$\frac{1}{1}$ 0 0	21	2 000	3.1410	4	311	20	0.100	$\left\{\begin{array}{c} 2.1900\\ 2.1405\\ 2.1312\end{array}\right.$	2	$\frac{1}{3}\frac{4}{4}\frac{2}{4}$
7	8.92		8.9150	11	$\overline{1}01$	51	3.088	3.1023	22 2	$\frac{1}{2}\frac{1}{2}\frac{1}{3}\frac{1}{1}$	29	2.122	{ 2.1405	2	344
32	7.64		7.6513	27	$\overline{1}10$			3.0984	2	231				6	$\underline{2}\overline{1}5$
	(1) (1) (1)	ſ	7.4672	6	$\bar{1} \bar{1} 1$			3.0831	2	031			2.1269	2	$\overline{2}$ $\underline{4}$ 2
41	6.81	4	6.8221	20	0 1 1			3.0651		$\bar{1} \bar{3} 4$			2.1248	3	$\bar{3}\bar{1}6$
		C	6.7990	11	$0\bar{1}2$			$\left\{\begin{array}{c} 3.0489\\ 2.9948\\ 2.9923\end{array}\right.$	3	$\bar{2}04$			2.1234	2	124
			6.7766	3	$1\bar{1}1$	63	2.977	{ 2.9948	4	313			2.1203	3	050
			6.7616	4	101			2.9923	2	$1\bar{1}4$			2.1058	4	411
			6.6621	5	110			2.9861	17	113			2.0931	2	$\bar{2}\bar{5}3$
100	6.21	C	6.2047	100	112			2.9798	2	$\bar{2} \bar{3} 3$	10	2.071	(2.0682	6	336
27	5.28	ł	5.3008	12	020			2.9763	15	311	9	2.0106	$\Big\{\begin{array}{c} 2.0682\\ 2.0140\\ 2.0095 \end{array}$	2	$\bar{2} 3 4$
		l	5.2863	15	$0\bar{2}2$			2.9670	2	232			2.0095	2	4 32
			5.0567	2	$1\overline{1}2$			2.9456	4	310			1 0000	22	$\bar{2}\bar{3}7$
19	4.992		5.0333	11	121			$\Big\{ \begin{array}{c} 2.9431 \\ 2.9176 \\ 2.9078 \end{array} \Big.$	15	$2\overline{2}3$	10	1.9642	$\left\{\begin{array}{c} 1.9882\\ 1.9774\\ 1.9754\end{array}\right.$	3	117
	000038077		4.9591	2	$\bar{2}01$	42	2.913	2.9176	14	0 4 2	1000		1.9754	2	025
				8	$\overline{1}$ $\overline{2}$ 1			2 9078	15	221			1.9477		513
30	4.650	J	4.6894	7	102			2 8579	4	$\begin{array}{c}2&2&1\\\bar{3}\bar{2}&2\end{array}$			$\begin{cases} 1.9477 \\ 1.9377 \\ 1.9146 \\ 1.9135 \end{cases}$	2 2 2	413
22.	1.050	1	4.6826	11	112			$\left\{\begin{array}{c} 2.8399\\ 2.8180\\ 2.7910\end{array}\right.$	2	$\overline{1}$ $\overline{2}$ $\overline{5}$ $\overline{1}$ $\overline{1}$ $\overline{5}$	35	1 9076	1 9146	2	163
			4.6268	6	012	16	2 705	28180	2	115	55	1.5070	1 0135	8	441
			4.6148	7	$0 \overline{1} \overline{3}$	10	4.195	2.0100	2	203			1.9103	2	513
			4.6086	4	$\frac{0}{2}$ 1 0			2.7910	3	131			1.9088	9	043
			4.5296	7	$\frac{2}{2}$ $\frac{1}{1}$ $\frac{1}{1}$			2.7850	4				1.9088		$043 \\ 047$
24	1 165				211			2.7727	4	025				4	
24	4.465		4.4575	14	202	-	2	$\left\{\begin{array}{c} 2.7644\\ 2.6558\\ 2.6447\end{array}\right.$	2	$0 \overline{2} 5 \\ 3 1 1$	22	1.0004	1.8815	2	353
			4.3875	2	$\frac{1}{2} \frac{2}{1} \frac{0}{2}$	5	2.660	2.6558	2	311	23	1.8604	1.8752	3	3 5 3
			4.3796	2	$\overline{\underline{2}} \overline{1} 2$			2.6447	4 2 2 2 2 2 2 2	032			1.8687	3	<u>4</u> <u>4</u> 3
15	4.301		4.3096	10	$\overline{1}03$			(2.5385	2	$3\bar{3}1$			1.8624	2	162
		ſ	4.2539	2	121	10	2.513	{ 2.5284	2	$2\overline{2}4$			1.8593	4	151
18	4.153	1	4.1934	3	003			2.5160	4	005			1.8553	4	$\overline{4} \ \underline{0} \ 6$
		C	4.1529	8	210	1000		$ \left\{\begin{array}{c} 2.6447 \\ 2.5385 \\ 2.5284 \\ 2.5160 \\ 2.4824 \\ 2.4824 \\ 2.4670 \\ \end{array}\right. $	3	401			(1.8399	2	127
			4.1084	6	201	10	2.456	$\left\{\begin{array}{c} 2.4824\\ 2.4679\\ 2.4316\end{array}\right.$	4	$\frac{\overline{2}}{\overline{1}} \frac{4}{4} \frac{0}{1}$	12	1.7895	$\Big\{ \begin{array}{c} 1.8399 \\ 1.8074 \\ 1.7914 \end{array} \Big.$	2	$\overline{5}\ \overline{2}\ 1$
			3.8466	2	$0\overline{3}1$			2.4316	2	141				3	$\overline{1}$ 3 5
			3.8301	2	$\overline{2} \overline{1} 3$			$\Big\{ \begin{array}{c} 2.4301 \\ 2.4257 \\ 2.4029 \end{array} \Big.$	2 2	314			1.7851	3	531
12	3.802		3.8048	6	1 1 3	11	2.402	{ 2.4257	2	135			1.7693	2	422
		٢	3.6405	2	121			2.4029	2	4 12			1.7491	2	206
12	3.558	1	3.5916	3	132			2.3992	4	124		1.7343		5	260
		L	3.5339	2	030			2.3813	3	331	13	1.7021		3	352
		r	3.6405 3.5916 3.5339 3.5242	4	033			$\left\{\begin{array}{c} 2.3587\\ 2.3017\\ 2.3017\\ 2.2871\end{array}\right.$	3 2 3 2	$\frac{\overline{2}}{\overline{3}}\frac{\overline{4}}{\overline{3}}\frac{1}{\overline{2}}$			1.7113	2	535
52	3.462	ł	3.4917	3	$\overline{1}$ $\overline{3}$ 1	11	2.294	2.3017	3	332			1.7083	4	266
		l	3.4750	32	130			2.2871	2	$4\bar{1}1$			1.6962	3	265
			3.4440	3	$\bar{1} \bar{2} 4$			(2.2775	23	233			1.6904	2	404
			3.4316	3	$0\bar{1}4$	13	2.255	$\Big\{ \begin{array}{c} 2.2775 \\ 2.2586 \\ 2.2539 \end{array} \Big.$	3	41 4				1	
			21.12.10	100			2.200	2 2520	5	303					

Only calculated lines with I > 2 are shown.

The data in these three tables are also deposited with the Principal Editor of *Mineralogical Magazine* and are available from http://www.minersoc.org/pages/e_journals/dep_mat_mm.html.

microscope (objective $50\times$). The spectrometer was calibrated by a software-controlled calibration procedure (within *Omnic 8* software) using multiple neon emission lines (wavelength calibration), multiple polystyrene Raman bands (laser frequency calibration), and standardized white light sources (intensity calibration). Spectral manipulation such as background correction and band-component

analysis was done with *Omnic* δ software. Spectra of all three minerals are displayed in Fig. 6.

OH stretching and H₂O bending vibrations

The bands attributable to the O-H stretching vibrations were observed for all three minerals, in

	Klaprothite	Péligotite	Ottohahnite
Structural formula	Na ₆ (UO ₂)(SO ₄) ₄ (H ₂ O) ₄	Na ₆ (UO ₂)(SO ₄) ₄ (H ₂ O) ₄	Na ₆ (UO ₂) ₂ (SO ₄) ₅ (H ₂ O) ₇ ·1.5H ₂ O
Space group	$P2_1/c$	$P\overline{1}$	$P\overline{1}$
Unit-cell dimensions	a = 9.8271(4) Å	a = 9.81511(18) Å	a = 9.97562(19) Å
	b = 9.7452(3) Å	b = 9.9575(2) Å	b = 11.6741(2) Å
	c = 20.8725(15) Å	c = 10.6289(8) Å	c = 14.2903(10) Å
	$\beta = 98.743(7)^{\circ}$	$\alpha = 88.680(6)^{\circ}$	$\alpha = 113.518(8)^{\circ}$
		$\beta = 73.990(5)^{\circ}$	$\beta = 104.282(7)^{\circ}$
		$\gamma = 89.205(6)^{\circ}$	$\gamma = 91.400(6)^{\circ}$
V	1975.66(17) Å ³	998.22(8) Å ³	1464.59(14) Å ³
Ζ	4	2	2
Density (for above formula)	$2.906 \text{ g} \cdot \text{cm}^{-3}$	$2.875 \text{ g} \cdot \text{cm}^{-3}$	$2.978 \text{ g} \cdot \text{cm}^{-3}$
Absorption coefficient	8.877 mm^{-1}	8.784 mm^{-1}	11.611 mm^{-1}
<i>F</i> (000)	1624	812	1216
Crystal size (µm)	$80 \times 60 \times 30$	$60 \times 45 \times 25$	$80 \times 60 \times 50$
θ range	3.02 to 27.49°	3.20 to 27.44°	3.00 to 27.44°
Index ranges	$-12 \le h \le 12$	$-12 \le h \le 11$	$-11 \le h \le 12$
	$-12 \le k \le 12$	$-12 \le k \le 12$	$-15 \le k \le 15$
	$-27 \le l \le 27$	$-13 \le l \le 13$	$-18 \le l \le 18$
Reflections collected; unique	$23,627/4507; R_{\rm int} = 0.035$	$17,794/4512; R_{\rm int} = 0.039$	$26,283/6658; R_{int} = 0.041$
Reflections with $F_{o} > 4\sigma(F)$	4034	4146	5721
Completeness to θ_{max}	99.5%	99.1%	99.6%
Restraints/parameters	12/325	12/325	0/425
Goof	1.098	1.089	1.040
Final <i>R</i> indices $[F_0 > 4\sigma(F)]$	$R_1 = 0.0222, wR_2 = 0.0455$	$R_1 = 0.0228, wR_2 = 0.0490$	$R_1 = 0.0265, wR_2 = 0.0581$
R indices (all data)	$R_1 = 0.0271, wR_2 = 0.0470$	$R_1 = 0.0273, wR_2 = 0.0506$	$R_1 = 0.0341, wR_2 = 0.0610$
Largest diff. peak/hole	$+1.09/-0.86 \ e \ A^{-3}$	$+1.14/-0.67 \ e \ A^{-3}$	$+1.35/-1.38 \ e \ A^{-3}$

TABLE 4. Data collection and structure refinement details for klaprothite, péligotite and ottohahnite.

 $*R_{int} = \Sigma |F_o^2 - F_o^2(mean)| / \Sigma [F_o^2]. \text{ Goof} = S = \{\Sigma [w(F_o^2 - F_c^2)^2] / (n-p)\}^{1/2}. R_1 = \Sigma ||F_o| - |F_c|| / \Sigma |F_o|. wR_2 = \{\Sigma [w(F_o^2 - F_c^2)^2] / \Sigma [w(F_o^2)^2]\}^{1/2}; w = 1 / [\sigma^2 (F_o^2) + (aP)^2 + bP] \text{ where } P \text{ is } [2F_c^2 + Max(F_o^2, 0)] / 3; \text{ for klaprothite } a \text{ is } 0.0153 \text{ and } b \text{ is } 5.0943; \text{ for péligotite } a \text{ is } 0.2090 \text{ and } b \text{ is } 1.0734; \text{ for ottohahnite } a \text{ is } 0.0288 \text{ and } b \text{ is } 2.1310.$

TABLE 5. Atom coordinates and displacement parameters (Å²) for klaprothite, péligotite and ottohahnite.

Klaprothite						22	22	22	12	12
	x/a	y/b	z/c	$U_{\rm eq}$	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
Na1	0.90437(17)	0.30192(16)	0.89756(8)	0.0239(3)	0.0252(8)	0.0193(8)	0.0268(8)	0.0013(7)	0.0025(6)	-0.0041(7)
Na2	0.26877(17)	0.23763(17)	0.75503(8)	0.0271(4)	0.0240(9)	0.0303(9)	0.0256(8)	-0.0052(7)	-0.0008(7)	0.0031(7)
Na3	0.30983(19)	0.22928(17)	0.99039(8)	0.0303(4)	0.0376(10)	0.0256(9)	0.0247(8)	0.0045(7)	-0.0048(7)	-0.0117(7)
Na4	0.40612(17)	0.81668(16)	0.90064(8)	0.0254(4)	0.0240(8)	0.0267(9)	0.0242(8)	-0.0004(7)	-0.0003(6)	-0.0067(7)
Na5	0.3506(2)	0.62409(19)	0.74617(9)	0.0410(5)	0.0575(13)	0.0289(10)	0.0411(11)	-0.0083(8)	0.0219(9)	-0.0097(9)
Na6	0.0	0.5	0.5	0.0325(6)	0.0346(14)	0.0409(14)	0.0241(12)	0.0123(11)	0.0115(10)	0.0166(12)
Na7	0.5	0.5	0	0.0988(19)	0.075(3)	0.163(4)	0.0460(19)	0.059(2)	-0.0328(18)	-0.092(3)
U	0.99601(2)	0.20230(2)	0.60250(2)	0.01284(4)	0.01114(7)	0.01244(7)	0.01492(7)	0.00186(5)	0.00187(4)	-0.00010(5)
S1	0.66470(9)	0.53766(9)	0.87958(4)	0.01520(18)	0.0154(4)	0.0136(4)	0.0170(4)	0.0001(3)	0.0036(3)	-0.0022(3)
52	0.16743(10)	0.04653(9)	0.88370(4)	0.01691(18)	0.0154(4)	0.0165(4)	0.0190(4)	0.0007(4)	0.0034(3)	-0.0018(4)
\$3	0.26752(9)	0.45861(9)	0.60626(4)	0.01456(18)	0.0114(4)	0.0134(4)	0.0190(4)	-0.0005(3)	0.0028(3)	-0.0004(3)
54	0.22619(9)	0.49007(9)	0.87535(4)	0.01472(18)	0.0137(4)	0.0136(4)	0.0170(4)	-0.0004(3)	0.0026(3)	0.0013(3)
D1	0.3707(3)	0.0712(3)	0.68903(13)	0.0309(7)	0.045(2)	0.0300(17)	0.0176(14)	-0.0040(12)	0.0032(13)	0.0023(14)
02	0.5873(3)	0.6234(3)	0.91893(13)	0.0247(6)	0.0220(15)	0.0252(15)	0.0283(15)	-0.0042(12)	0.0087(12)	0.0022(12)
03	0.6440(3)	0.3917(3)	0.89178(14)	0.0272(6)	0.0321(17)	0.0141(13)	0.0345(16)	0.0026(12)	0.0024(13)	-0.0063(12)
D4	0.8155(3)	0.5630(3)	0.90098(15)	0.0300(7)	0.0157(15)	0.0303(16)	0.0445(18)	0.0038(14)	0.0061(13)	-0.0052(13)
05	0.1284(4)	0.0692(4)	0.81476(15)	0.0445(9)	0.040(2)	0.071(3)	0.0219(16)	0.0061(16)	0.0023(14)	0.0007(18)
D6	0.8981(3)	0.6436(3)	0.57752(14)	0.0253(6)	0.0229(15)	0.0180(14)	0.0353(16)	0.0055(12)	0.0052(12)	-0.0015(12)
D7	0.3169(3)	0.0514(3)	0.90312(14)	0.0274(7)	0.0146(14)	0.0322(16)	0.0361(17)	-0.0025(13)	0.0068(12)	-0.0046(12)
28	0.8771(3)	0.4064(3)	0.59786(17)	0.0354(8)	0.0303(18)	0.0149(14)	0.060(2)	-0.0011(14)	0.0026(15)	0.0069(13)
D9	0.2088(3)	0.5943(3)	0.59533(18)	0.0373(8)	0.0297(17)	0.0169(15)	0.068(2)	0.0035(15)	0.0151(16)	0.0094(13)
D10	0.3031(3)	0.4250(3)	0.67504(13)	0.0282(7)	0.0286(16)	0.0387(18)	0.0172(14)	-0.0008(12)	0.0030(12)	-0.0108(14)
D11	0.3860(3)	0.4425(3)	0.57219(13)	0.0211(6)	0.0150(13)	0.0248(14)	0.0244(14)	0.0000(12)	0.0061(11)	0.0023(11)
D12	0.1596(3)	0.3577(3)	0.57522(13)	0.0236(6)	0.0216(15)	0.0271(15)	0.0214(14)	-0.0001(12)	0.0007(11)	-0.0119(12)
D13	0.2590(3)	0.4444(3)	0.81359(14)	0.0298(7)	0.0390(18)	0.0257(15)	0.0275(15)	-0.0070(13)	0.0139(13)	0.0030(14)
D14	0.3122(3)	0.4266(3)	0.92971(14)	0.0263(6)	0.0196(15)	0.0293(16)	0.0276(15)	0.0103(12)	-0.0043(12)	-0.0002(12)
D15	0.0770(3)	0.4683(3)	0.87985(13)	0.0211(6)	0.0139(13)	0.0159(13)	0.0334(15)	-0.0014(11)	0.0035(11)	-0.0026(11)
D16	0.2383(3)	0.6437(3)	0.87958(13)	0.0198(6)	0.0149(13)	0.0117(12)	0.0339(15)	-0.0009(11)	0.0072(11)	-0.0002(10)
D17	0.0502(3)	0.2256(3)	0.68669(13)	0.0218(6)	0.0190(14)	0.0271(15)	0.0197(13)	-0.0016(11)	0.0042(11)	0.0008(12)
D18	0.9439(3)	0.1845(3)	0.51757(13)	0.0255(6)	0.0268(15)	0.0305(16)	0.0182(13)	0.0017(12)	0.0007(11)	-0.0029(13)
OW1	0.8704(4)	0.2477(6)	0.78456(18)	0.0576(12)	0.037(2)	0.106(4)	0.032(2)	-0.010(2)	0.0101(17)	-0.001(2)
H1a	0.937(5)	0.203(5)	0.783(3)	0.069						
H1b	0.880(6)	0.317(4)	0.761(3)	0.069						
OW2	0.4328(3)	0.7935(3)	0.01370(14)	0.0270(6)	0.0331(17)	0.0228(15)	0.0272(15)	0.0002(13)	0.0115(13)	0.0009(14)

TABLE 5. (contd.)

Klaprothite	(contd.)									
	x/a	y/b	z/c	U_{eq}	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
H2a	0.414(5)	0.869(3)	0.023(2)	0.032						
H2b	0.393(5)	0.744(4)	0.035(2)	0.032						
OW3	0.4801(3)	0.1886(3)	0.82167(13)	0.0251(6)	0.0274(16)	0.0260(16)	0.0212(14)	-0.0022(12)	0.0019(12)	-0.0053(13)
H3a	0.526(4)	0.252(4)	0.840(2)	0.030						
H3b	0.454(5)	0.143(4)	0.8493(18)	0.030						
OW4	0.4137(4)	0.8479(3)	0.78808(15)	0.0333(7)	0.0398(19)	0.0300(17)	0.0296(17)	0.0014(14)	0.0036(15)	-0.0016(15)
H4a	0.376(4)	0.903(4)	0.763(2)	0.040						
H4b	0.499(3)	0.868(5)	0.791(2)	0.040						
Péligotite										
	x/a	y/b	z/c	U_{eq}	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
Na1	0.0173(2)	0.21679(18)	0.78354(16)	0.0321(4)	0.0449(11)	0.0279(10)	0.0257(9)	-0.0057(7)	-0.0139(8)	0.0180(8)
Na2	0.13381(19)	0.60608(18)	0.49903(16)	0.0312(4)	0.0367(10)	0.0336(10)	0.0223(9)	-0.0064(7)	-0.0056(7)	-0.0115(8)
Na3	0.32017(19)	0.28083(18)	0.98490(17)	0.0302(4)	0.0367(10)	0.0285(10)	0.0291(9)	-0.0105(7)	-0.0152(8)	0.0145(8)
Na4	0.48329(18)	0.30651(18)	0.21999(16)	0.0290(4)	0.0303(10)	0.0307(10)	0.0298(9)	-0.0052(7)	-0.0146(7)	0.0090(8)
Na5	0.6224(2)	0.1241(2)	0.52127(18)	0.0405(5)	0.0411(11)	0.0445(12)	0.0295(10)	-0.0090(8)	0.0028(8)	-0.0229(9)
Na6	0	0.5	0	0.0255(5)	0.0246(13)	0.0317(13)	0.0201(11)	0.0046(9)	-0.0062(9)	0.0020(10)
Na7	0.5	0	0	0.0826(15)	0.084(3)	0.137(4)	0.0279(16)	-0.0204(18)	-0.0191(16)	0.089(3)
U	0.09294(2)	0.80723(2)	0.80855(2)	0.01259(5)	0.01248(8)	0.01214(8)	0.01266(7)	-0.00023(5)	-0.00267(5)	-0.00012(5)
S1	0.23359(10)	0.02517(10)	0.23419(9)	0.01805(19)	0.0154(5)	0.0173(5)	0.0218(5)	0.0013(4)	-0.0059(4)	0.0024(4)
S2	0.27155(10)	0.47726(9)	0.76028(9)	0.01702(19)	0.0156(5)	0.0146(5)	0.0201(5)	-0.0015(3)	-0.0038(4)	0.0032(4)
S3	0.15332(10)	0.45960(9)	0.21644(9)	0.01466(18)	0.0136(5)	0.0137(5)	0.0164(4)	-0.0018(3)	-0.0035(3)	-0.0008(4)
S4	0.33207(10)	0.01303(9)	0.76845(9)	0.01533(18)	0.0143(5)	0.0147(5)	0.0167(4)	-0.0023(3)	-0.0036(3)	-0.0015(4)
01	0.2244(4)	0.0221(4)	0.3731(3)	0.0397(8)	0.049(2)	0.048(2)	0.0190(16)	0.0021(14)	-0.0037(14)	0.0052(17)
O2	0.3376(3)	0.1233(3)	0.1628(3)	0.0256(6)	0.0203(15)	0.0266(16)	0.0273(15)	0.0020(12)	-0.0022(12)	-0.0029(12)
O3	0.2676(4)	0.8920(3)	0.1791(3)	0.0364(8)	0.047(2)	0.0201(17)	0.045(2)	-0.0065(14)	-0.0165(16)	0.0092(15)
04	0.0921(3)	0.0632(3)	0.2179(4)	0.0401(9)	0.0176(17)	0.0341(19)	0.072(3)	0.0093(17)	-0.0188(16)	0.0027(14)
05	0.2572(4)	0.4774(3)	0.6279(3)	0.0386(8)	0.059(2)	0.038(2)	0.0192(15)	-0.0025(13)	-0.0108(15)	0.0101(17)
O6	0.1923(3)	0.3676(3)	0.8403(3)	0.0281(7)	0.0319(17)	0.0185(15)	0.0330(17)	0.0052(12)	-0.0078(13)	-0.0001(13)
O7	0.4194(3)	0.4725(3)	0.7621(4)	0.0408(9)	0.0172(17)	0.036(2)	0.071(2)	-0.0167(17)	-0.0140(16)	0.0066(14)
08	0.2080(3)	0.6058(3)	0.8259(3)	0.0266(7)	0.0309(17)	0.0192(15)	0.0303(16)	-0.0019(12)	-0.0099(13)	0.0081(13)
O9	0.1136(3)	0.5881(3)	0.1683(3)	0.0287(7)	0.0282(17)	0.0177(15)	0.0409(18)	0.0014(13)	-0.0111(14)	0.0068(12)
O10	0.1011(3)	0.4443(3)	0.3582(3)	0.0302(7)	0.0358(18)	0.0379(19)	0.0155(14)	-0.0028(12)	-0.0042(12)	-0.0103(14)
O11	0.3065(3)	0.4391(3)	0.1685(3)	0.0242(6)	0.0130(14)	0.0239(16)	0.0367(17)	-0.0038(12)	-0.0083(12)	0.0028(12)
O12	0.0866(3)	0.3520(3)	0.1543(3)	0.0223(6)	0.0247(16)	0.0211(15)	0.0223(14)	-0.0027(11)	-0.0082(12)	-0.0068(12)

(continued)

TABLE 2	5. (conta	!.)
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Péligo	tite (contd.)	17	,		T T 1 1	U^{22}	U^{33}	± 73	U^{13}	U^{12}
	x/a	y/b	z/c	$U_{\rm eq}$	U^{11}	U^{22}	U^{ss}	U^{23}	U^{13}	U^{12}
013	0.4120(3)	0.0350(3)	0.6332(3)	0.0308(7)	0.0260(17)	0.0421(19)	0.0198(15)	-0.0039(13)	0.0018(12)	-0.0100(14
O14	0.3831(3)	0.0928(3)	0.8573(3)	0.0254(6)	0.0262(16)	0.0253(16)	0.0291(16)	-0.0104(12)	-0.0144(13)	0.0019(13
015	0.1781(3)	0.0374(3)	0.7821(3)	0.0200(6)	0.0141(14)	0.0168(14)	0.0305(15)	0.0014(11)	-0.0085(11)	0.0001(11
O16	0.3355(3)	0.8660(3)	0.8039(3)	0.0235(6)	0.0175(15)	0.0137(14)	0.0404(17)	0.0007(12)	-0.0099(12)	0.0003(11
017	0.1527(3)	0.7912(3)	0.6350(3)	0.0218(6)	0.0268(16)	0.0217(15)	0.0159(13)	-0.0011(11)	-0.0040(11)	-0.0041(12
O18	0.9593(3)	0.1796(3)	0.0190(3)	0.0241(6)	0.0292(16)	0.0253(16)	0.0164(13)	-0.0030(11)	-0.0038(11)	-0.0017(13
OW1	0.0212(6)	0.2052(6)	0.5629(4)	0.0723(15)	0.109(4)	0.078(3)	0.034(2)	-0.008(2)	-0.028(3)	0.044(3)
H1a	0.030(7)	0.129(3)	0.539(7)	0.087	()					()
H1b	0.106(4)	0.233(7)	0.536(7)	0.087						
OW2	0.5681(3)	0.3171(3)	0.9800(3)	0.0285(7)	0.0239(17)	0.0277(18)	0.0301(17)	0.0024(13)	-0.0013(13)	0.0008(14
H2a	0.588(5)	0.392(3)	0.944(5)	0.034		()	()	()	()	× ×
H2b	0.642(4)	0.277(4)	0.961(5)	0.034						
OW3	0.3404(3)	0.6990(3)	0.3532(3)	0.0289(7)	0.0315(18)	0.0283(18)	0.0245(16)	-0.0021(13)	-0.0037(13)	0.0029(14
H3a	0.321(5)	0.742(4)	0.294(4)	0.035		× ,	× /	()	()	`
H3b	0.403(4)	0.647(4)	0.317(4)	0.035						
OW4	0.4291(7)	0.2766(5)	0.4487(4)	0.0841(18)	0.126(5)	0.081(4)	0.034(2)	0.004(2)	-0.005(3)	0.052(3)
H4a	0.359(5)	0.303(7)	0.510(6)	0.101						
H4b	0.486(7)	0.342(7)	0.454(8)	0.101						
Ottoha	ahnite									
	x/a	y/b	z/c	U_{eq}	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
U1	0.76518(2)	0.92310(2)	0.59278(2)	0.01693(6)	0.01595(10)	0.01423(9)	0.01643(9)	0.00435(7)	0.00092(7)	0.00139(7
U2	0.43880(2)	0.47704(2)	0.20323(2)	0.01589(5)	0.01530(10)	0.01535(9)	0.01348(8)	0.00398(7)	0.00148(7)	0.00216(7
S1	0.64717(12)	0.22386(10)	0.18552(9)	0.0213(3)	0.0182(6)	0.0170(5)	0.0233(6)	0.0042(5)	0.0036(5)	0.0043(5)
S2	0.74046(12)	0.58141(10)	0.43301(8)	0.0173(2)	0.0161(6)	0.0153(5)	0.0150(5)	0.0029(4)	0.0011(4)	0.0015(4)
S3	0.42597(12)	0.80033(10)	0.39688(9)	0.0180(2)	0.0164(6)	0.0160(5)	0.0164(5)	0.0043(4)	-0.0003(4)	0.0041(4)
S4	0.01287(13)	0.99886(11)	0.79367(9)	0.0228(3)	0.0211(6)	0.0225(6)	0.0176(5)	0.0052(5)	-0.0016(5)	0.0036(5)
S5	0.86118(12)	0.63587(11)	0.93924(9)	0.0214(2)	0.0182(6)	0.0210(6)	0.0189(5)	0.0064(4)	-0.0012(5)	-0.0023(5)
Na1	0.8619(2)	0.28948(17)	0.81894(15)	0.0293(5)	0.0302(12)	0.0235(10)	0.0255(10)	0.0055(8)	0.0013(9)	0.0041(9)
Na2	0.5133(2)	0.33816(19)	0.41652(16)	0.0312(5)	0.0337(12)	0.0311(11)	0.0323(11)	0.0166(9)	0.0093(9)	0.0046(9)
Na3	0.0619(2)	0.6873(2)	0.61423(17)	0.0344(5)	0.0444(14)	0.0287(11)	0.0371(12)	0.0165(9)	0.0174(10)	0.0182(10
Na4	0.8728(2)	0.1325(2)	0.02411(18)	0.0390(5)	0.0364(13)	0.0345(12)	0.0441(13)	0.0218(10)	-0.0008(10)	-0.0028(10
Na5	0.6421(3)	0.8889(2)	0.02774(19)	0.0401(6)	0.0458(14)	0.0238(11)	0.0443(13)	0.0058(10)	0.0160(11)	0.0064(10
										(continued

(continued)

TABLE 5	. (cont	<i>d</i> .)

Ottohahnite (contd.)					- 11	- 22	22	22	12	- 12
	x/a	y/b	z/c	$U_{\rm eq}$	U^{11}	U^{22}	U^{33}	U^{23}	U^{13}	U^{12}
03	0.6328(5)	0.2402(4)	0.2889(3)	0.0503(12)	0.071(3)	0.059(3)	0.045(2)	0.033(2)	0.033(2)	0.028(2)
04	0.5596(4)	0.3102(3)	0.1485(3)	0.0324(9)	0.033(2)	0.0294(19)	0.0307(19)	0.0100(16)	0.0048(16)	0.0193(16)
05	0.8601(4)	0.5142(3)	0.4276(3)	0.0337(9)	0.029(2)	0.033(2)	0.040(2)	0.0154(17)	0.0095(17)	0.0177(17)
O6	0.6499(4)	0.5452(3)	0.4854(3)	0.0314(8)	0.032(2)	0.036(2)	0.0237(18)	0.0104(16)	0.0064(16)	-0.0033(17)
07	0.6615(4)	0.5619(3)	0.3236(2)	0.0258(8)	0.0211(18)	0.0326(19)	0.0181(16)	0.0086(14)	0.0003(14)	-0.0040(15)
08	0.7937(4)	0.7186(3)	0.4916(3)	0.0312(8)	0.035(2)	0.0154(16)	0.0300(19)	-0.0014(14)	0.0043(16)	0.0025(15)
09	0.3898(4)	0.7619(4)	0.4731(3)	0.0409(10)	0.031(2)	0.061(3)	0.037(2)	0.031(2)	0.0046(17)	0.000(2)
O10	0.4102(4)	0.6908(3)	0.2950(3)	0.0333(9)	0.042(2)	0.0164(17)	0.0274(18)	0.0001(14)	0.0009(17)	0.0086(16)
011	0.3319(4)	0.8863(3)	0.3723(3)	0.0364(9)	0.037(2)	0.0284(19)	0.038(2)	0.0119(17)	0.0023(17)	0.0173(17
012	0.5713(4)	0.8642(3)	0.4383(3)	0.0295(8)	0.0222(19)	0.0328(19)	0.0256(18)	0.0115(15)	-0.0047(15)	-0.0036(15
013	0.1514(4)	0.0527(4)	0.8109(3)	0.0450(11)	0.021(2)	0.042(2)	0.062(3)	0.019(2)	-0.0029(19)	-0.0038(18)
014	0.9971(5)	0.9828(4)	0.8860(3)	0.0468(11)	0.062(3)	0.056(3)	0.032(2)	0.027(2)	0.015(2)	0.023(2)
015	0.9695(4)	0.8768(3)	0.6967(3)	0.0306(9)	0.031(2)	0.0190(17)	0.0269(18)	0.0028(14)	-0.0058(15)	0.0099(15
O16	0.9052(3)	0.0764(3)	0.7644(3)	0.0250(8)	0.0218(18)	0.0180(16)	0.0242(17)	0.0028(14)	-0.0022(14)	0.0049(14
O17	0.9285(4)	0.6428(3)	0.0433(3)	0.0287(8)	0.0247(19)	0.035(2)	0.0229(17)	0.0145(15)	-0.0020(15)	-0.0024(16
O18	0.0473(4)	0.3070(4)	0.0998(3)	0.0418(10)	0.039(2)	0.051(2)	0.032(2)	0.0154(19)	0.0073(18)	-0.0171(19)
019	0.7297(4)	0.6946(3)	0.9403(3)	0.0326(9)	0.026(2)	0.0208(17)	0.0309(19)	-0.0012(15)	-0.0059(16)	0.0053(15
O20	0.8068(3)	0.5023(3)	0.8604(3)	0.0235(7)	0.0138(16)	0.0215(16)	0.0266(17)	0.0049(14)	0.0001(14)	0.0025(13
O21	0.8678(4)	0.9717(3)	0.5275(3)	0.0296(8)	0.026(2)	0.0302(19)	0.0319(19)	0.0116(16)	0.0104(16)	0.0011(16)
O22	0.6598(4)	0.8690(3)	0.6539(3)	0.0302(8)	0.0245(19)	0.037(2)	0.0274(18)	0.0136(16)	0.0039(15)	0.0010(16
O23	0.4849(4)	0.5240(3)	0.1106(3)	0.0286(8)	0.028(2)	0.037(2)	0.0225(17)	0.0142(15)	0.0058(15)	0.0061(16
O24	0.3923(3)	0.4298(3)	0.2963(3)	0.0265(8)	0.0201(18)	0.0322(19)	0.0268(18)	0.0150(15)	0.0020(14)	-0.0026(15
OW1	0.7096(5)	0.9473(4)	0.9026(3)	0.0545(12)	0.050(3)	0.062(3)	0.041(2)	0.015(2)	0.010(2)	-0.015(2)
OW2	0.9016(4)	0.5982(4)	0.6761(3)	0.0448(11)	0.036(2)	0.064(3)	0.035(2)	0.022(2)	0.0098(18)	0.003(2)
OW3	0.7482(4)	0.2485(4)	0.9332(3)	0.0381(9)	0.038(2)	0.044(2)	0.035(2)	0.0197(18)	0.0096(18)	0.0092(19
OW4	0.9939(5)	0.3217(4)	0.7067(3)	0.0489(11)	0.059(3)	0.046(2)	0.040(2)	0.017(2)	0.014(2)	-0.003(2)
OW5	0.1271(4)	0.7924(4)	0.5167(3)	0.0446(10)	0.038(2)	0.054(3)	0.055(3)	0.033(2)	0.018(2)	0.007(2)
OW6	0.2298(4)	0.5100(4)	0.9070(3)	0.0369(9)	0.035(2)	0.043(2)	0.036(2)	0.0196(18)	0.0093(18)	0.0060(18
OW7a*	0.5335(8)	0.802(2)	0.1177(13)	0.076(5)	0.037(4)	0.113(14)	0.091(9)	0.068(10)	0.001(4)	-0.013(5)
OW7b*	0.5358(19)	0.906(5)	0.180(3)	0.090(16)	0.034(10)	0.15(4)	0.11(3)	0.09(3)	0.004(11)	-0.018(13)
OW8	0.7656(7)	0.7849(5)	0.2689(4)	0.0800(17)	0.095(5)	0.084(4)	0.078(4)	0.045(3)	0.033(3)	0.018(3)
OW9*	0.6236(12)	0.0097(9)	0.3142(8)	0.063(3)	0.087(8)	0.050(6)	0.070(7)	0.028(5)	0.048(6)	0.019(5)

*Occupancies: OW7a: 0.69(4); OW7b: 0.31(4); OW9: 0.5

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Klaprothite									
Na1–O9	2.325(3)	Na4–OW2	2.345(3)		S101	1.458(3)	U017	1.772(3)
Na1–OW1	2.390(4	4)	Na4-016	2.353(3)		S1O2	1.464(3)	U018	1.777(3)
Na1-015	2.415(3)	Na4-011	2.376(3)		S1-O3	1.464(3)	U-012	2.301(3)
Na1–O6	2.472	3)	Na4–OW4	2.381(4)		S1-04	1.502(3)	U-08	2.306(3)
Na1-O18	2.479(3)	Na4-07	2.453(3)		<s1-0></s1-0>	1.472	U–O4	2.341(3)
Na1–O3	2.690(3)	Na402	2.580(3)				U-015	2.435(3)
Na1–O4	2.694(3)	<na4–o></na4–o>	2.415		S2-O5	1.449(3)	U016	2.456(3)
<na1-o></na1-o>	2.495					S206	1.457(3)	$< U - O_{Ur} >$	1.775
			Na5–OW4	2.396(4)		S2-O7	1.464(3)	<u-o<sub>eq></u-o<sub>	2.368
Na2–OW3	2.366(3)	Na5–OW3	2.425(4)		S2-O8	1.502(3)	сų	
Na2013	2.366(.	3)	Na5010	2.445(3)		<s2–o></s2–o>	1.468		
Na2017	2.394(3)	Na5–OW1	2.481(5)					
Na2–O1	2.440(3)	Na5013	2.498(4)		S3–O9	1.447(3)		
Na2O10	2.531(3)	Na501	2.910(4)		S3-O10	1.462(3)		
Na2–O5	2.583(4	4)	<na5–o></na5–o>	2.526		S3011	1.462(3)		
<na2–o></na2–o>	2.447					S3O12	1.518(3)		
			Na6-O6(×2)	2.464(3)		<s3–o></s3–o>	1.472		
Na3014	2.305(3)	Na6–O12(×2)	2.470(3)					
Na3011	2.427(3)	Na6–O8(×2)	2.686(4)		S4-013	1.446(3)		
Na306	2.450(3)	Na6–O9(×2)	2.788(4)		S4-014	1.447(3)		
Na3–O2	2.467(3)	<na6–o></na6–o>	2.602		S4-015	1.498(3)		
Na3–O7	2.523(3)				S4-016	1.503(3)		
Na3–OW2	2.553(4	4)	Na7-O14(×2)	2.290(3)		<s4–o></s4–o>	1.474		
Na3012	2.613(3)	Na7–O2(×2)	2.342(3)					
<na3-o></na3-o>	2.477		Na7–OW2(×2)) 2.959(3)					
Hydrogen bo	onds								
$D-H\cdots A$		D–H	$H \cdots A$	$D \cdots A$	<dha< td=""><td>1</td><td></td><td></td><td></td></dha<>	1			
OW1−H1a…	·O5	0.79(3)	2.30(3)	3.060(6)	161(6))			
OW1-H1b··	·O17	0.84(3)		2.905(6)	102(6)				
OW2−H2a··	·O11	0.80(3)	2.13(3)	2.914(4)	166(5				
OW2−H2b··	·O3	0.79(3)	2.10(3)	2.859(4)	161(5				
OW3−H3a…	·O3	0.83(3)	2.00(3)	2.817(4)	174(4)				
OW3-H3b··	·O7	0.80(3)		2.842(4)	159(5)				
OW4–H4a…	·O1	0.80(3)		2.987(4)	154(4))			
OW4–H4b⋯	·O10	0.86(3)		2.873(5)	163(5))			

TABLE 6. Selected bond distances (Å) and angles	(°)) for kla	prothite,	péligotite and ottohahnite.

the approximate range $3650-3300 \text{ cm}^{-1}$. The bands are very broad and of low intensity. Approximate O–H···O hydrogen bond lengths inferred from the corresponding stretching wavenumbers using Libowitzky's empirical relation (Libowitzky, 1999) vary in the range >3.2–2.77 Å.

Very weak bands at ~1650 cm⁻¹, attributed to the v₂ (δ) H–O–H bending mode of the H₂O molecules, were observed in the spectra of all three minerals. The low intensities of bands is typical of Raman spectra, in contrast to the higher intensity of v₂ (δ) mode bands in infrared spectra.

Vibrations of SO₄ groups

The vibrations of the sulfate tetrahedra are well resolved in the spectra of all three minerals. The activated antisymmetric vibration v_3 (SO₄) and the splitting of the degenerate modes suggest departure from the ideal T_d symmetry. The distortion of the tetrahedra due to the bidentate linkage between the SO₄ and UO₇ bipyramids is considerable (Plášil *et al.*, 2015). Multiple bands for v_1 (SO₄) in all three spectra are indicative of symmetrically non-equivalent SO₄ groups.

TABLE 6. (contd.)

Péligotite								
Na1-09	2.302(3)		Na4011	2.342(3)	S101	1.454(3)	U018	1.769(3)
Na1-OW1	2.340(5)		Na4–OW4	2.353(5)	S1-O3	1.459(3)	U-017	1.785(3)
Na1-O15	2.365(3)		Na4016	2.422(3)	S1-O2	1.463(3)	U-04	2.290(3)
Na1-O18	2.429(3)		Na4-07	2.443(4)	S1-04	1.489(3)	U-08	2.316(3)
Na1–O6	2.501(4)		Na4–OW2	2.455(4)	<s1–o></s1–o>	1.466	U-012	2.332(3)
Na1–O3	2.932(4)		Na402	2.519(3)			U-015	2.435(3)
Na1–O4	3.004(4)		<na4–o></na4–o>	2.422	S2-O5	1.452(3)	U-016	2.447(3)
<na1-o></na1-o>	2.553				S2-O7	1.457(3)	$< U - O_{Ur} >$	1.777
			Na5013	2.262(3)	S2-O6	1.462(3)	<u-o<sub>eq></u-o<sub>	2.364
Na2-O10	2.309(3)		Na5–OW3	2.325(4)	S2-O8	1.514(3)	eq	
Na2–OW3	2.371(4)		Na5–O13	2.401(3)	<s2–o></s2–o>	1.471		
Na2–O5	2.403(4)		Na5017	2.511(3)				
Na2017	2.408(3)		Na501	2.537(4)	S3–O9	1.455(3)		
Na2-O10	2.437(4)		Na5–OW4	2.679(6)	S3-O10	1.456(3)		
Na2-OW1	2.588(5)		<na5-o></na5-o>	2.453	S3011	1.461(3)		
<na2–o></na2–o>	2.419				S3O12	1.518(3)		
			Na6-O12(×2)	2.492(3)	<s3–o></s3–o>	1.473		
Na3014	2.313(3)		Na6–O9(×2)	2.532(3)				
Na306	2.378(3)		Na6-O6(×2)	2.540(3)	S4-014	1.442(3)		
Na3–OW2	2.452(4)		Na6–O8(×2)	2.567(4)	S4013	1.448(3)		
Na3–O2	2.470(3)		<na6–o></na6–o>	2.533	S4-015	1.494(3)		
Na3011	2.510(3)				S4016	1.505(3)		
Na3012	2.596(3)		Na7–O14(×2)	2.311(3)	<s4–o></s4–o>	1.472		
Na3–O7	2.961(4)		Na7–O2(×2)	2.359(3)				
<na3–o></na3–o>	2.526		Na7–O3(×2)	2.750(3)				
			<na7–o></na7–o>	2.473				
Hydrogen b	onds							
D–H··· A		D–H	$H \cdots A$	$D \cdots A$	<dha< td=""><td></td><td></td><td></td></dha<>			
OW1-H1a·	··O1	0.80(3)	2.46(3)	3.037(6)	130(6)			
OW1-H1b.	··none	0.85(3)						
OW2−H2a·	··O11	0.83(3)	2.15(3)	2.955(4)	165(4)			
OW2-H2b·	··O3	0.80(3)	2.27(4)	2.892(5)	135(5)			
OW3−H3a·	··O3	0.82(3)		2.856(4)	165(5)			
OW3-H3b·	··O7	0.82(3)	2.08(3)	2.893(4)	172(5)			
OW4–H4a∙	··O5	0.85(3)		2.963(6)	146(7)			
OW4–H4b∙	··none	0.87(3)						

The bands located at ~1250 cm⁻¹ are attributable to split triply degenerate v_3 (SO₄) antisymmetric stretching vibrations. Raman bands located in the range ~1080 to ~960 cm⁻¹ are attributable to v_1 (SO₄) symmetric stretching vibrations. The intensities of the bands related to the symmetric stretching mode are considerably higher than those attributed to the activated antisymmetric stretching vibrations.

The medium-weak bands located at ~650 cm⁻¹ are attributed to the v_4 (δ) triply degenerated antisymmetric stretching vibrations of SO₄ tetrahedra and the bands at ~450 cm⁻¹ are assigned to the

split $v_2(\delta)$ doubly degenerate bending vibrations of the SO₄ groups. The number of split $v_2(\delta)$ vibrations is in accordance with the presence of several non-equivalent SO₄ tetrahedra in the structures of all three minerals.

Vibrations of UO₂²⁺

The bands related to the activated v_3 antisymmetric stretching vibration of the uranyl ion, $(UO_2)^{2+}$, were observed in the spectra of all three minerals. The reason for activation of this mode in the Raman spectra is the lowering of the factor group

TABLE 6. (contd.)

Ottohahnite	e*						
Na1-017	2.349(4)	Na4018	2.343(4)	S101	1.449(4)	Hydrogen bor	nds
Na1016	2.377(4)	Na4–OW3	2.399(4)	S1-O2	1.452(4)	OW1014	2.968(7)
Na1–OW3	2.386(4)	Na4–OW1	2.404(5)	S1-O3	1.455(4)	OW1-022	3.198(5)
Na1–O20	2.425(4)	Na4014	2.406(4)	S1-04	1.516(4)	OW2-06	3.053(5)
Na1–OW4	2.439(5)	Na4–O2	2.634(5)	<s1–o></s1–o>	1.470	OW2018	2.842(6)
Na1–O11	2.853(4)	Na4-014	2.655(5)			OW3-OW6	2.793(6)
Na1–O10	2.867(4)	<na4–o></na4–o>	2.474	S2-O5	1.446(4)	OW3–OW7a	2.710(9)
<na1-0></na1-0>	2.528			S2-06	1.457(4)	OW3-OW7b	3.01(2)
		Na5–O1	2.339(4)	S2-07	1.489(3)	OW4–OW5	2.838(6)
Na2–O3	2.352(5)	Na5–OW7a	2.349(9)	S2-08	1.489(3)	OW4–OW8	2.739(7)
Na209	2.357(4)	Na5–OW1	2.380(5)	<s2–o></s2–o>	1.468	OW5-09	2.839(6)
Na206	2.439(4)	Na5019	2.417(4)			OW5-021	3.064(5)
Na2024	2.462(4)	Na5013	2.522(5)	S309	1.448(4)	OW6-017	2.757(5)
Na206	2.486(4)	Na5–OW7b	2.59(3)	S3010	1.473(3)	OW6-023	2.941(5)
Na2-O22	2.615(4)	Na5-O1	2.711(5)	S3-011	1.476(4)	OW7a–OW8	2.824(10)
<na2–o></na2–o>	2.452	Na5–O4	2.897(4)	S3012	1.477(4)	OW7a-OW9	2.790(17)
		<na5-o></na5-o>	2.527*	<\$3–O>	1.469	OW7b-O1	2.93(5)
Na3–OW5	2.375(5)					OW7b-OW8	3.00(3)
Na3–O5	2.385(4)	Na6–OW6	2.306(4)	S4-013	1.426(4)	OW8-013	2.773(7)
Na3015	2.386(4)	Na6–OW2	2.357(5)	S4-014	1.449(4)	OW8-OW9	2.931(11)
Na3–OW2	2.403(5)	Na6–OW4	2.384(5)	S4-015	1.495(3)	OW9-03	2.867(10)
Na3–O2	2.576(5)	Na6–O2	2.402(5)	S4-016	1.509(3)	OW9-09	3.180(10)
Na3–O5	2.846(4)	Na6–O5	2.638(5)	<s4–o></s4–o>	1.470		
Na308	2.934(4)	Na6018	2.885(5)				
Na3–O3	2.961(6)	Na6–O7	2.961(5)	S5-017	1.441(3)		
<na3–o></na3–o>	2.608	<na6-o></na6-o>	2.562	S5018	1.452(4)		
				S5019	1.496(4)		
U1021	1.762(4)	U2-023	1.770(3)	S5-O20	1.500(3)		
U1-022	1.768(4)	U2024	1.779(3)	<\$5–O>			
U1-08	2.313(3)	U204	2.285(3)				
U1-011	2.365(3)	U2-07	2.350(3)				
U1012	2.388(3)	U2010	2.374(3)				
U1-015	2.419(3)	U2019	2.430(3)				
U1016	2.420(3)	U2020	2.447(3)				
$< U1 - O_{Ur} >$	1.765	$< U2 - O_{Ur} >$	1.775				
$< U1 - O_{eq}^{U1} >$	2.381	<U2 $-$ O _{eq} $>$	2.377				
eq		cq					

*The contributions of Na–OW7a and Na–OW7b to <Na–O> are weighted according to the refined occupancies of the OW7a and OW7b sites.

(symmetry) due to the relatively low symmetry (monoclinic or triclinic) of these minerals. The bands are of low or very low intensity and are observed in the range ~960 to ~920 cm⁻¹. The very strong bands observed at ~830 cm⁻¹ are attributed to the v₁ symmetric stretching vibration of the uranyl ion.

The approximate U–O bond-lengths inferred from the observed energies of the $(UO_2)^{2^+}$ vibrations using the empirical relation of Bartlett and Cooney (1989) is 1.79 Å, consistent with the bond-

lengths obtained in our X-ray studies (klaprothite \sim 1.78 Å, péligotite \sim 1.79 Å, ottohahnite \sim 1.77 Å) and identical to the most common bond-length given by Burns *et al.* (1997) for the uranyl pentagonal bipyramid, 1.79 Å.

The very weak bands at ~550 and ~500 cm⁻¹ may be attributable to the v (U–O_{*ligand*}) vibrations. Bands of medium intensity at low wavenumbers, ~250 cm⁻¹, can be assigned to the v₂ (δ) doubly degenerate bending vibrations of (UO₂)²⁺ (e.g. Bullock and Parret, 1970; Ohwada, 1976; Brittain

Klaprothite														
	Na1	Na2	Na3	Na4	Na5	Na6	Na7	U	S1	S2	S3	S4	H bonds	Σ
01		0.16			0.04				1.57				0.13	1.90
O2			0.15	0.11			$0.21^{\times 2\downarrow}$		1.54					2.01
O3	0.08								1.54				0.16,0.18	1.96
O4	0.08							0.57	1.39					1.96
O5		0.11								1.60			0.12	1.83
O6	0.14		0.15			$0.15^{\times 2\downarrow}$				1.57				2.01
O7			0.13	0.15						1.54			0.17	1.99
O8						$0.08^{ imes 2\downarrow}$		0.61		1.39				2.08
O9	0.21					$0.06^{\times 2\downarrow}$					1.61			1.88
O10		0.12			0.16						1.55		0.16	1.99
O11			0.16	0.19							1.55		0.15	2.05
O12			0.10			$0.15^{\times 2\downarrow}$		0.62			1.33			2.20
O13		0.19			0.13							1.62		1.94
O14			0.23				$0.24^{\times 2\downarrow}$					1.61		2.08
O15	0.17							0.48				1.41		2.06
O16				0.20				0.46				1.39		2.05
O17		0.18						1.71					0.15	2.04
O18	0.14							1.70						1.84
OW1	0.18				0.14								-0.12,-0.15	0.05
OW2			0.12	0.20			$0.04^{\times 2\downarrow}$						-0.15,-0.16	0.05
OW3		0.19			0.16								-0.18, -0.17	0.00
OW4				0.18	0.18								-0.13,-0.16	0.07
Σ	1.00	0.95	1.04	1.03	0.81	0.88	0.98	6.15	6.04	6.10	6.04	6.03		

TABLE 7. Bond-valence analysis for klaprothite, péligotite and ottohahnite. Values are expressed in valence units.*

Péligotite	Na1	Na2	Na3	Na4	Na5	Na6	Na7	U	S1	S2	S3	S4	H bonds	Σ
	1141	1442	INdo	INAT	INdo	1440	1447	0	51	52	55	54	11 001103	
01					0.12				1.58				0.12	1.83
02			0.15	0.13			$0.20^{\times 2\downarrow}$		1.56					2.03
O3	0.04						$0.07^{ imes 2 \downarrow}$		1.55				0.15,0.17	1.98
O4	0.03							0.60	1.44					2.07
05		0.17								1.59			0.14	1.90
O6	0.13		0.19			$0.12^{\times 2\downarrow}$				1.57				2.01
07			0.04	0.16						1.55			0.15	1.90
08						0.11 ^{×2↓}		0.63		1.35				2.09
09	0.23					$0.12^{\times 2\downarrow}$					1.58			1.93
O10		0.22,0.16									1.57			1.96
O11			0.13	0.21							1.55		0.14	2.03
012			0.10			$0.14^{\times 2\downarrow}$		0.58			1.33			2.15
O13			0.22		0.25,0.17							1.61		2.04
O14							$0.22^{\times 2\downarrow}$					1.64		2.08
O15	0.19							0.48				1.42		2.09
O16				0.17				0.47				1.38		2.01
O17		0.17			0.13			1.67						1.97
O18	0.16							1.72						1.88
OW1	0.21	0.11											-0.12	0.19
OW2			0.15	0.15									-0.14,-0.15	0.01
OW3		0.19			0.21								-0.17,-0.15	0.08
OW4				0.20	0.08								-0.14	0.14
Σ	1.00	1.03	0.98	1.00	0.98	0.98	0.97	6.15	6.13	6.06	6.04	6.04		

TABLE 7.	(contd.)	

Ottohahnite	Na1	Na2	Na3	Na4	Na5	Na6	U1	U2	S1	S2	S3	S4	S5	H bonds	Σ
	INdI	INdZ	INdS	INdH		INdO	01	02		52	33	54	35		
O1					0.21,0.08				1.60					+0.05	1.94
O2			0.11	0.09		0.17			1.59						1.96
O3		0.20	0.04						1.58					+0.08	1.90
O4					0.05			0.64	1.34						2.03
O5			0.18,0.05			0.09				1.62					1.94
O6		0.16,0.14								1.57				+0.12	1.99
O7						0.04		0.56		1.44					2.04
08			0.04				0.60			1.44					2.08
09		0.20									1.61			+0.17, +0.05	2.03
O10	0.05							0.54			1.50				2.09
011	0.05						0.55				1.49				2.09
012							0.52				1.49				2.01
013					0.13							1.71		+0.20	2.04
014				0.17,0.09								1.60		+0.14	2.00
015			0.18				0.49					1.42			2.09
O16	0.19						0.49					1.36			2.04
O17	0.20												1.64	+0.20	2.04
O18				0.20		0.05							1.59	+0.17	2.01
O19					0.17			0.48					1.41		2.06
O20	0.16							0.47					1.40		2.03
O21							1.75							+0.12	1.87
O22		0.10					1.73							+0.10	1.93
O23								1.72						+0.14	1.86
O24		0.15						1.69							1.84
OW1				0.17	0.19									-0.14,-0.10	0.12
OW2			0.17			0.23								-0.12, -0.17	0.11
OW3	0.18			0.18										-0.19,-0.19	-0.02
OW4	0.16					0.18								-0.17,-0.21	-0.04
OW5			0.19											+0.17, -0.17,	0.07
														-0.12	
OW6						0.23								+0.19, -0.20,	0.08
														-0.14	

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Ottohahnite	Nal	Na2	Na3	Na4	Na5		U1	Na6 UI U2 SI S2 S3 S4 S5	S1	S2	S3	$\mathbf{S4}$	S5	H bonds	Σ
OW7				-	0.14,0.03									+0.19,-0.25,	0.02
0W8														+0.21,+0.16,	0.02
6M0														+0.13,+0.15	0.02
Σ	0.99	0.95	0.96	0.90	1.00	0.99	6.13	6.13 6.10 6.11 6.07 6.09	6.11	6.07	6.09	6.09	6.04	-0.10,-0.10	
*S ⁶⁺ -O bond strength from Brown hydrogen-bond strengths based on	strength fruid strengths	om Brown at based on O-	and Altermatt (1985); Na ⁺ –O bond strength from Wood an O–O bond lengths from Ferraris, G. and Ivaldi, G. (1988).	1985); Na ⁺ –C hs from Ferra) bond strer tris, G. and	ngth fron Ivaldi, G	1 Wood a	and Pale	nik (199	9); U ⁶⁺ -	-O bond	strength	from B	$*S^{6+}$ -O bond strength from Brown and Altermatt (1985); Na ⁺ -O bond strength from Wood and Palenik (1999); U ⁶⁺ -O bond strength from Burns <i>et al.</i> (1997); hydrogen-bond strengths based on O–O bond lengths from Ferraris, G. and Ivaldi, G. (1988).	

TABLE 7. (contd.)

et al., 1985; Plášil *et al.*, 2010, 2015). Weak Raman bands at ~150 cm⁻¹ are attributed to $-O_{eq}-U-O_{ea}$ -bending vibrations (Ohwada, 1976).

Vibrations of Na–O bonds and lattice modes

Weak Raman bands with shoulders at ~280 cm⁻¹ could be related to Na–O stretching vibrations or to the v (U–O_{*ligand*}) stretching vibration (Plášil *et al.*, 2010, 2015). The bands located below 100 cm⁻¹ are related to the lattice modes.

Composition

Chemical analyses of klaprothite, péligotite and ottohahnite were performed using a CamScan 4D electron microprobe in energy-dispersive spectroscopy (EDS) mode (20 kV, 500 pA and beam rasterized on an area 8 μ m × 8 μ m in order to minimize sample damage). The EDS mode with the above experimental conditions, rather than wavelength-dispersive spectroscopy (WDS) mode, was chosen because of the instability of these minerals under the electron beam caused by the high contents of both Na and H₂O. Attempts to use the WDS mode with higher beam current were unsuccessful because of significant decomposition of the minerals after several seconds under the electron beam. Standards used were chkalovite for Na, synthetic UO₂ for U and synthetic ZnS for S. H₂O was not determined directly because of the scarcity of pure material. The H₂O content was calculated by stoichiometry on the basis of the number of O apfu from crystal-structure determinations. The Raman spectra confirm the presence of H₂O groups and the absence of B-O, C-O and N-O bonds in the minerals. No other elements with atomic numbers higher than 8 were observed. Analytical data are given in Table 2.

The empirical formula of klaprothite, calculated as the mean of 11 representative spot analyses, is $Na_{6.01}(U_{1.03}O_2)(S_{0.993}O_4)_4(H_2O)_4$ and that of péligotite, calculated as the mean of 8 analyses, is $Na_{5.82}(U_{1.02}O_2)(S_{1.003}O_4)_4(H_2O)_4$, both based on 22 O apfu). The ideal formula for both minerals is $Na_6(UO_2)(SO_4)_4(H_2O)_4$ which requires Na_2O 21.51, UO_3 33.10, SO_3 37.05 and H_2O 8.34, total 100 wt.%. The empirical formula of ottohahnite, calculated as the mean of 12 analyses, is $Na_{5.88}(U_{0.99}O_2)_2(S_{1.008}O_4)_5(H_2O)_{8.5}$ based on 32.5 O apfu. Its ideal formula is $Na_6(UO_2)_2(SO_4)_5(H_2O)_7$. 1.5H₂O, which requires Na_2O 14.18, UO_3 43.62, SO_3 30.52 and H_2O 11.68, total 100 wt.%.

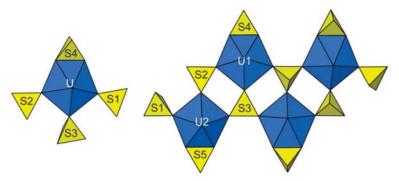


FIG. 7. (Left) The $[(UO_2)(SO_4)_4]^{6-}$ cluster in the structures of klaprothite and péligotite. (Right) The $[(UO_2)_4(SO_4)_{10}]^{12-}$ cluster constructed of four linked $[(UO_2)(SO_4)_4]^{6-}$ clusters.

The Gladstone-Dale compatibility index 1 – (K_P/K_C) for the empirical formulas of klaprothite, péligotite and ottohahnite are –0.008 (superior), – 0.015 (superior) and –0.021 (excellent), respectively (Mandarino, 2007). As noted by Kampf *et al.* (2015*c*), Gladstone-Dale calculations for uranyl sulfates should utilize $k(UO_3) = 0.118$, as provided by Mandarino (1976).

X-ray crystallography and structure refinement

Powder X-ray studies were carried out using a Rigaku R-Axis Rapid II curved imaging plate microdiffractometer, with monochromatized MoK α radiation ($\lambda = 0.71075$ Å). A Gandolfi-like motion on the ϕ and ω axes was used to randomize the orientation of the sample and observed *d*-values and intensities were derived by profile fitting using JADE 2010 software (Materials Data, Inc.). The powder data presented in Table 3 show good agreement with the patterns calculated from the structure determinations. Unit-cell parameters refined from the powder data using JADE 2010 with whole pattern fitting are for klaprothite: a = 9.812(4), b =9.739(4), c = 20.855(4) Å, $\beta = 98.743(8)^{\circ}$ and V =1969.7(12) Å³; for péligotite: a = 9.811(9), b =9.928(10), c = 10.635(10) Å, $\alpha = 88.79(3)^{\circ}$, $\beta =$ 74.00(3)°, $\gamma = 89.27(2)^{\circ}$ and V = 995.5(17) Å³: and for ottohahnite: a = 9.979(8), b = 11.660(8), c = 14.277(8) Å, $\alpha = 113.48(3)^{\circ}$, $\beta = 104.29(2)^{\circ}$, $\gamma = 91.35(2)^{\circ}$ and $V = 1462.6(18) \text{ Å}^3$.

The single-crystal structure data for klaprothite, péligotite and ottohahnite were collected at room temperature on a Rigaku R-Axis Rapid II curved imaging plate microdiffractometer, also using monochromatized Mo $K\alpha$ radiation. The data were

processed using the Rigaku CrystalClear software package and empirical (multi-scan) absorption corrections were applied using the ABSCOR program (Higashi, 2001) in the CrystalClear software suite. The structures were solved by direct methods using SIR2011 (Burla et al., 2012). SHELXL-2013 (Sheldrick, 2008) was used for the refinement of the structures. Difference-Fourier syntheses located all H atom positions in the structures of klaprothite and péligotite; however, H atom positions could not be located in the ottohahnite structure. For klaprothite and péligotite, H atom positions were refined with soft restraints of 0.82(3) Å on the O-H distances and 1.30(3) Å on the H-H distances and with the $U_{\rm eq}$ of each H set to 1.2 times that of its donor O atom.

For the structures of both klaprothite and péligotite, all non-hydrogen atoms refined to full occupancy, although the Na7 sites, at centres of symmetry in both structures, exhibited significant displacement along a. Refining these as split sites improved the R-factors slightly, but it was ultimately decided to report them as unsplit sites. The same high displacement along **a** is noted for this site in the structure of the synthetic phase equivalent to klaprothite (Plášil et al., 2015). For the structure of ottohahnite, all sites were refined with full occupancy except for three O sites in relatively close proximity: The occupancies of OW7a and OW7b (1.18 Å apart and coordinated to Na5) refined to 0.74 and 0.37. The occupancy of OW9 (1.77 Å from OW7b and not coordinated to a cation) refined to 0.48. In the final refinement, the OW7a and OW7b sites were constrained to a total occupancy of 1 and the OW9 site was assigned an occupancy of 0.5.

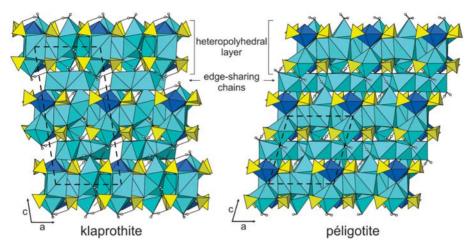


FIG. 8. The structures of klaprothite and péligotite viewed along [010]. Na polyhedra are turquoise blue. Unit-cell outlines are shown as dashed lines.

Data collection and refinement details are given in Table 4, atom coordinates and displacement parameters in Table 5, selected bond distances in Table 6 and bond-valence analyses in Table 7.

Description and discussion of the structures

In the structures of klaprothite, péligotite and ottohahnite, the U sites are surrounded by seven O atom sites forming squat pentagonal bipyramids. This is the most typical coordination for U^{6+} ,

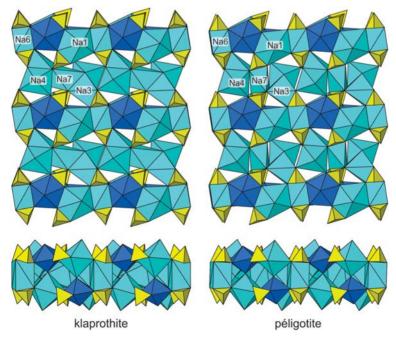


FIG. 9. Heteropolyhedral layers in the structures of klaprothite and péligotite viewed perpendicular to {001} (upper images) and along [100] (lower images); [010] is horizontal in all cases.

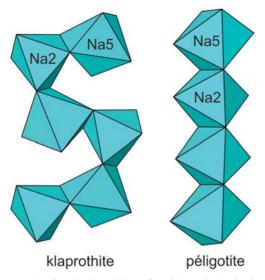


FIG. 10. Edge-sharing chains of NaO6 octahedra in the structures of klaprothite and péligotite. The chain in klaprothite is along [010] and that in péligotite is along [110].

particularly in uranyl sulfates (Krivovichev, 2010, 2013), where the two short apical bonds of the bipyramid constitute the uranyl group. The bond lengths in U⁶⁺-coordination polyhedron in all three structures (klaprothite: <U1–O_{ap}>=1.775, <U1–O_{eq}>=2.368 Å; péligotite: <U1–O_{ap}>=1.777,

 $<U1-O_{eq}>=2.364$ Å; ottohahnite: $<U1-O_{ap}>=1.765$, $<U1-O_{eq}>=2.381$, $<U2-O_{ap}>=1.775$, $<U2-O_{eq}>=2.377$ Å) are consistent with the most typical lengths observed for [7]-coordinated U⁶⁺ (Burns et al., 1997). In all three structures, all five of the equatorial O atoms also participate in SO_4 tetrahedra and, in each case, three of the SO_4 tetrahedra share single equatorial O atoms and one shares two equatorial O atoms, i.e. a shared polyhedral edge. Such a bidentate linkage between a UO₇ pentagonal bipyramid and a SO₄ tetrahedron has previously been reported in synthetic phases, but never in minerals. The $[(UO_2)(SO_4)_4]^{6-}$ uranyl sulfate cluster (Fig. 7) is the fundamental building block in the structures of klaprothite and péligotite; while in the structure of ottohahnite, four of these uranyl-sulfate clusters are combined through shared SO4 tetrahedra to form a larger $[(UO_2)_4(SO_4)_{10}]^{12-}$ cluster (Fig. 7). Among synthetic Na uranyl sulfates, the structure of the synthetic equivalent of klaprothite was reported by Plášil et al. (2015) and the same $[(UO_2)(SO_4)_4]^{6-}$ cluster with a bidentate U-S linkage was reported in the structures of $Na_{10}[(UO_2)(SO_4)_4](SO_4)_2 \cdot 3H_2O$ (Burns and KNa₅[(UO₂)(SO₄)₄](H₂O) Hayden, 2002), (Hayden and Burns, 2002a) and $Na_6(UO_2)$ $(SO_4)_4(H_2O)_2$ (Hayden and Burns, 2002b). The $[(UO_2)_4(SO_4)_{10}]^{12-}$ cluster in the structure of ottohahnite has not previously been reported in the structure of any mineral or synthetic phase.

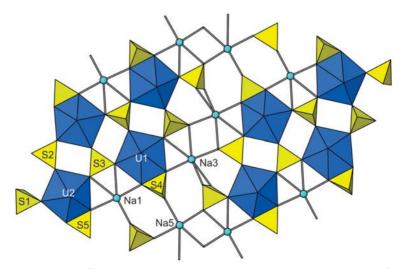


FIG. 11. One slice parallel to $\{11\overline{2}\}$ in the structure of ottohahnite containing the $[(UO_2)_4(SO_4)_{10}]^{12}$ -cluster and Na1, Na3 and Na5 atoms. Na–O bonds are shown as sticks.

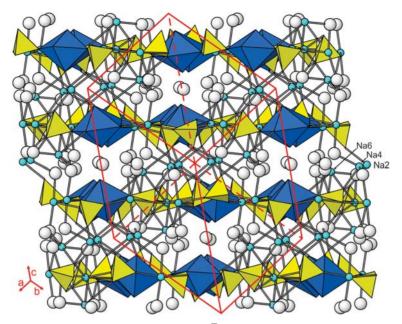


Fig. 12. The structure ottohahnite viewed down [111] with $\{11\overline{2}\}$ horizontal. O atoms of H₂O groups are shown as white spheres. The unit cell outline is shown in red.

In the structures of klaprothite and péligotite, seven Na sites, in 6-, 7- and 8-coordinations, link to the O atoms in the clusters and to four H₂O groups, vielding relatively loosely bonded frameworks (Fig. 8). In both structures, the Na1O₇, Na3O₇, Na4O₆, Na6O₈ and Na7O₆ polyhedra combine with the uranyl sulfate clusters forming more tightly bonded heteropolyhedral layers parallel to {001} in which the Na polyhedra share some faces and edges with one another, and edges and corners with the UO₇ and SO₄ polyhedra of the uranyl sulfate clusters. The polyhedral layers in the two structures (Fig. 9) are topologically identical except for a minor difference in the Na7O₆ coordination. The principal difference between the structures is the configuration of the Na2O6 and Na5O6 octahedra between the layers. These octahedra form a zig-zag edge-sharing chain parallel to [010] in the structure of klaprothite, while in the structure of péligotite they form a straight edge-sharing chain parallel to [110] (Fig. 10).

In the structure of ottohahnite, the $[(UO_2)_4(SO_4)_{10}]^{12-}$ clusters are linked to one another *via* bonds to six different Na atoms. The Na atoms, with coordinations ranging from 6 to 8, bond to O atoms in the clusters and to H₂O groups. There are seven different H₂O groups (one of which is split) bonded to Na atoms and two H₂O groups,

one of which is half occupied, bonded only though hydrogen bonds. The plane of the $[(UO_2)_4(SO_4)_{10}]^{12-}$ cluster is parallel to $\{11\overline{2}\}$. Na1, Na3 and Na5 are in approximately the same plane with the cluster. Figure 11 shows the linkages between the clusters and Na1O₇, Na3O₈ and Na5O₇ coordinations within one slice through the structure parallel to $\{11\overline{2}\}$. The Na2, Na4 and Na6 atoms are placed between these slices linking them together via Na–O bonds and yielding Na2O₆, Na4O₆ and Na6O₇ coordinations. The isolated H₂O groups (OW8 and OW9) are also sandwiched in between the slices (Fig. 12).

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