AN X-RAY STUDY OF A SINGLE CRYSTAL OF KAOLINITE

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There are several papers on x-ray studies of kaolinite [1-4], but the structure has not yet been established in that way, because in every case powder patterns have been used. The reason is, of course, that kaolinite occurs only in a very finely divided form.

Recently, the structure was established by means of electron diffraction [5]. X-ray studies on a single crystal remain of interest, though; we report here on such studies.

B. B. Zvyagin pointed out to us that a single crystal suitable for use in x-ray diffraction might be found by examining the deposit from a very dilute suspension, as has been done for another mineral [6].

In this way we found a single crystal about $0.2 \times 0.2 \times 0.01$ mm in size. The patterns were recorded with copper radiation in a KRON-2 camera. Oscillation and rotation photographs on the <u>a</u> and <u>b</u> axes were recorded; these showed 112 spots of hol and 0kl types. The intensities were estimated visually. The F₀ were corrected for the Lorentz and polarization factors; absorption was neglected as the crystal was so small. We used Bragg's [7] atomic-scattering factors in the calculation.

The Laue patterns (recorded along c^{\bullet}) showed that the spots did not display hexagonal symmetry (although they appeared to form a hexagonal array), because their positions and intensities were inappropriate to that symmetry; kaolinite is triclinic. The cell parameters calculated from powder patterns agree within the errors of experiment with those given in [3], namely a = 5.14, b = 8.93, c = 7.37 A, $\beta = 104.5^{\circ}$ and $\alpha = 91.8^{\circ}$.

The oscillation pattern gave c as about 7.4 A. This means that the unit cell contains one layer.

The oscillation and rotation patterns on <u>a</u> and <u>b</u> were found to have spots with h + k = 2n only; the cell is base-centered. As our initial model we assumed one consisting of octahedral and tetrahedral grids, with the latter displaced by -b/3 relative to the former.

The agreement with the measured intensities was best for that form of the structure in which the triangles at the tops of the octahedra lie opposite the bases of the tetrahedra above them along \underline{a} (if we use the regular figures proposed in [8]).

Similar deductions have been made elsewhere [3, 5] about the orientations of the layers.

The F for the ideal model bore a relation to the F_0 such that R was large, so the structure had to be revised by determining the coordinates of the atoms more precisely. Successive Fourier syntheses were used in that revision. Appropriate difference syntheses were used for overlapping atoms; the F were also calculated for various relative positions of the overlapping atoms, and those giving the best R were chosen. The structure factors were calculated from sums taken with respect to 13 independent atoms, whose coordinates were reckoned relative to an origin taken at the center of an empty octahedron. The projection along <u>a</u> (Fig. 1) gave us the <u>y</u> and <u>z</u> for all atoms except O₄ and O₅, which overlap in that projection. Difference syntheses were used for those atoms. It was very much more difficult to find the coordinates from the projection along <u>b</u> (Fig. 2), because that projection shows no isolated atoms. These coordinates were found by means of difference syntheses and by minimizing R. Table 1 gives the values finally adopted. Vainshtein's formulas [9] for the errors in the results are (in A) Si ± 0.014, A1 ± 0.015, and O ± 0.020.

TABLE 1

Coordinates of the Atoms in the Structure of Kaolinite

Group [.]	x/a	y/b	z/c	Group	x/a	1 <i>y/b</i>	z/c
$\begin{array}{c} OH_1\\ OH_2\\ OH_3\\ OH_4\\ O_1\\ O_2\\ O_3\\ O_8 \end{array}$	$\begin{array}{c} 0.778 \\ 0.278 \\ 0.316 \\ 0.248 \\ 0.754 \\ 0.690 \\ 0.791 \end{array}$	$\begin{array}{c} 0.180\\ 0.320\\0.008\\ 0.184\\ 0.315\\ 0.004\\ 0.165\end{array}$	$\begin{array}{c} -0.140 \\ -0.138 \\ -0.136 \\ 0.155 \\ 0.155 \\ 0.157 \\ 0.482 \end{array}$	$\begin{array}{c} O_4\\ O_5\\ Al_1\\ Al_2\\ Si_1\\ Si_2 \end{array}$	$\begin{array}{c} 0.612 \\ 0.108 \\ 0.502 \\ 0.002 \\ 0.800 \\ 0.800 \end{array}$	$\begin{array}{c} -0.120 \\ -0.058 \\ 0.172 \\ 0.330 \\ 0.322 \\ 000 \end{array}$	$\begin{array}{c} 0.455 \\ 0.455 \\ 0.003 \\ 0.002 \\ 0.382 \\ 0.385 \end{array}$

TABLE 2

Values of F and F_0

hkl	F	F	hkl	F	F _{0.}	hkl	F	F	hkl	F	Fo	hk l	F	F ₀
004	05 0	07.0	037	2.0		004						=		
001	20.2	21.2	027	5.2	5.2	004	7.8	9.4	0101	6.3	6.3	209	18.9	16.9
002 600	42.7	40.1	028	0.4	4.3	005	11.4	15.1	0102	5.6	5.8	= 400	29.4	31.8
003	20.0	30.2	1040	0.1	0.8	000	18.7	15.1	•0103	8.5	7.3	401	11.5	15.4
004	21.8	29.2	041	8.6	10.8	067	7.8	5.1	0101	7.3	7.2	402	26.1	22.0
005	21.4	22.5	042	13.9	13.6	061	13.7	20.8	0102	8.3	8.2	403	3.0	3.9
006	21.1	21.0	043	12.8	13.2	062	20.6	27.2	0103	1.7	3.8	404	9.9	11.3
007	15.0	15.1	044	20.9	18.5	063	9.7	13.2	# 200	32.1	30.7	405	9.3	7.7
008	19.3	15.4	045	8.6	9.2	064	11.1	13.0	201	17.3	20.2	406	7.5	5.8
* 020	17.3	15.3	046	12.7	12.1	065	19.2	15.4	202	18.2	22.9	*401	25.2	27.0
021	18.2	16.4	047	7.9	8.1	066	15.1	12.2	203	28.1	26.0	* 402	15.4	21.0
022	22.4	22.5	048	7.5	6.9	067	12.2	10.8	204	42.2	34.2	403	17.1	19.4
023	5.2	5.1	041	4.4	5.1	*080	11.5	14.7	205	2.9	3.9	404	14.3	13.4
024	14.7	11.7	042	17.1	14.2	081	8.9	6.9	206	18.4	15.0	405	18.9	16.5
025	6.3	7.1	043	10.4	10.7	082	9.7	8.8	207	7.5	6.2	406	27.9	19.2
026	8.2	7.7	044	17.9	12.8	083	3.1	5.4	208	10.8	7.6	407	4.6	6.0
027	8.3	9.4	045	15.2	14.9	084	5.1	5.2	≠ 201	26.2	26.1	408	14.2	9.2
028	4.2	4.5	046	9.8	10.8	085	5.5	6.1	* 202	39.4	34.8	600	15.4	11.1
021	27.8	25.9	047	10.6	9.9	081	11.1	7.7	203	35.7	33.0	601	12.2	8.1
$0\overline{2}2$	7.5	7.9	048	4.0	3.8	082	5.0	5.4	204	34.9	37.0	601	8.0	12.8
0 2 3	16.1	12.3	*060	45.3	46.2	083	4.6	5.4	205	16.3	15.2	+802	21.7	19.5
024	10.9	10.5	061	17.0	13.4	084	7.2	7.5	208	15.0	17.8	R 03	5.5	5.8
025	7.9	7.9	062	13.6	15.7	085	5.4	5.8	207	15 3	14 5	80%	10 0	0.0
026	12.3	13.2	063	10.9	14.8	*0.10.0	6.9	9.3	208	7.8	3.3	004	10.0	0.0

TABLE 3

Interatomic Distances (A) for Kaolinite

$\begin{array}{c c c c c c c c c c c c c c c c c c c $	$\begin{array}{c c c c c c c c c c c c c c c c c c c $
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Fig. 1. The $\sigma(yz)$ projection of the electron-density (only half of the projection is shown, because the other half is translationally the same).



Fig. 3. Tetrahedral grid in kaolinite.



Fig. 2. The $\sigma(xz)$ projection (half is shown).



Fig. 4. Octahedral grid in kaolinite.

Table 2 gives the F and F_0 . These have been compiled by means of two normalizations (as in [5]); the reflections were divided into two groups, to each of which we assigned a distinct factor to convert the F_0 to absolute values. One group differed from the other in that those in the first (those values indicated by• in Table 2) are single, whereas those in the other appear as doublets. Those latter are much weaker, which accounts for the need for two normalizations. The implication is that the sequence of atoms in a single layer does not apply strictly to the whole crystal [5]. Here R takes the value 15.4%.

The strengths of those reflections not visible on the films were taken as being half the strength of the weakest reflection.

The atoms are displaced from the positions corresponding to ideal cubic packing of the anions, which deviate from their ideal positions in a tetrahedral grid [10]. The tetrahedra are rotated in such a way that their bases form a ditrigonal pattern (Fig. 3). The average angle of that rotation is 21°. The top and bottom ends of the octahedra also form a ditrigonal array, but here the angles are much smaller (Fig. 4), being 6.5 and 4°, respectively. The interatomic distances (Table 3) agree well with the standard values. The common edges of the octahedra are shorter than the others. The successive layers are stacked in a way such that the oxygen atoms and OH groups in adjacent layers come together in pairs; the distances vary somewhat, the mean being 2.93 A.

These results do not differ essentially from those obtained by means of electron diffraction [5], although the coordinates and distances are by no means always in exact agreement.

The distortions resemble very closely those found in dickite, a polymorphic form of kaolinite [7].

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