

Hazenite, $\text{KNaMg}_2(\text{PO}_4)_2 \cdot 14\text{H}_2\text{O}$, a new biologically related phosphate mineral, from Mono Lake, California, U.S.A.

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ABSTRACT

A new biologically related, struvite-type phosphate mineral, hazenite, ideally $\text{KNaMg}_2(\text{PO}_4)_2 \cdot 14\text{H}_2\text{O}$, has been found in and/or on completely dried-out or decomposed cyanobacteria on porous calcium-carbonate (mainly calcite and aragonite) substrates in Mono Lake, California. The mineral occurs as radiating clusters of prismatic crystals and is colorless, transparent with white streak and vitreous luster. It is brittle, with the Mohs hardness of 2–2.5; cleavage is good on {001} and no twinning was observed. The measured and calculated densities are 1.91(3) and 1.88(2) g/cm³, respectively. Hazenite is biaxial (+), with $n_\alpha = 1.494(1)$, $n_\beta = 1.498(1)$, $n_\gamma = 1.503(1)$, $2V_{\text{meas}} = 41(2)^\circ$, $2V_{\text{calc}} = 42^\circ$, $X = b$, $Y = c$, $Z = a$, and does not fluoresce under long- or short-wave ultraviolet rays. The dispersion is strong with $r < v$. It is soluble in water. The electron microprobe analysis yielded an empirical formula of $\text{K}_{0.97}(\text{Na}_{0.96}\text{Ca}_{0.02})\text{Mg}_{2.07}[(\text{P}_{0.98}\text{S}_{0.02})\text{O}_4]_2 \cdot 13.90\text{H}_2\text{O}$. Hazenite is orthorhombic with space group *Pmnb* and unit-cell parameters $a = 6.9349(4)$ Å, $b = 25.174(2)$ Å, $c = 11.2195(8)$ Å, and $V = 1958.7(3)$ Å³. There are many structural similarities between hazenite and struvite, as also revealed by their Raman spectra. The hazenite structure contains six symmetrically independent non-hydrogen cation sites, two for Mg^{2+} (Mg1 and Mg2), two for P^{5+} (P1 and P2), one for Na^+ , and one for K^+ . It can be viewed as three types of layers stacking along the **b**-axis, in a repeating sequence of ABCBACB..., where layer A consists of $\text{Mg}1(\text{H}_2\text{O})_6$ octahedra and NaO_6 trigonal prisms, layer B of $\text{P}1\text{O}_4$ and $\text{P}2\text{O}_4$ tetrahedra, and layer C of $\text{Mg}2(\text{H}_2\text{O})_6$ octahedra and very irregular KO_6 polyhedra. These layers are linked together by hydrogen bonds, plus the K-O bonds between layers B and C (K-O5-P2). Interestingly, the combination of layers B and C in hazenite exhibits a configuration analogous to the struvite-(K) structure. Hazenite is believed to form in high pH environments through the involvement of cyanobacterial activities. To our knowledge, hazenite is the first struvite-type compound that contains two structurally distinct monovalent cations (K and Na), pointing to an exclusive role of biological activity in the mineralization process.

Keywords: Hazenite, struvite-type materials, phosphates, biomineral, crystal structure, X-ray diffraction, Raman spectra

INTRODUCTION

A new biologically related, struvite-type phosphate mineral, hazenite, $\text{KNaMg}_2(\text{PO}_4)_2 \cdot 14\text{H}_2\text{O}$, has been found on the shoreline of Mono Lake, California. Mono Lake, located in the hydrologically closed Mono basin, eastern California, U.S.A., is known for its outstanding biological and geochemical features (Wiens et al. 1993 and references therein). This terminal lake currently consists of a hypersaline (84–92 g/L), alkaline (pH = 9.8) $\text{NaCO}_3\text{-Cl-SO}_4$ brine, resulting from evaporative concentration of inflow water, rock weathering, and mineral precipitation (e.g., Garrels and Mackenzie 1967; Li et al. 1997; Oremland et al. 2004). The lake is endorheic and the basin's volcanic setting includes numerous hydrothermal springs that contribute to its unique water chemistry (Bischoff et al. 1991; Council and Bennett 1993; Budinoff and Hollibaugh 2007). High concentrations of boron (~34 mmol/L) and minor amounts of K, Mg, and Ca are

found in the lake (e.g., Bischoff et al. 1991). Dissolved organic carbon concentration is ~6.7 mmol/L (Melack 1983), and phosphate is as high as 1.0 mmol/L (Mono Basin Ecosystem Study Committee 1987). The food web of Mono Lake is relatively simple, consisting of bacteria, phytoplankton, brine flies *Ephedra hians*, the brine shrimp *Artemia monica*, and migrating and nesting birds (Jørgensen et al. 2008 and references therein). Soda lakes, such as Mono Lake, are good analogs of putative ancient Martian and Archaean terrestrial aquatic biomes and represent environmental extremes in terms of their high pH and salinities. They can also be sensitive indicators of local climate changes and provide records of paleoclimate (Stine 1994; Jellison et al. 1996; Benson et al. 1998; Hollibaugh et al. 2001). Nevertheless, Mono Lake is subject to recurrent periods of meromixis (persistent chemical stratification) as a consequence of natural and anthropogenic alterations of freshwater flow into the lake, with the most recent episodes persisting from 1995 until 2003, and from 2005 until 2007 (Budinoff and Hollibaugh 2007).

Several unusual biologically related phosphate miner-

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als have been found from Mono Lake, including newberyite $\text{MgHPO}_4 \cdot 3\text{H}_2\text{O}$ (Cohen and Ribbe 1966; Ribbe and Cohen 1966), monetite CaHPO_4 (Cohen and Ribbe 1966; Ribbe and Cohen 1966), struvite $(\text{NH}_4)\text{MgPO}_4 \cdot 6\text{H}_2\text{O}$ (Cooper and Dunning 1969), and possibly brushite $\text{CaHPO}_4 \cdot 2\text{H}_2\text{O}$ (Walker 1988). In addition, ikaite, $\text{CaCO}_3 \cdot 6\text{H}_2\text{O}$, not previously observed in lake environments, precipitates seasonally along the shore of Mono Lake (Shearman et al. 1989; Council and Bennett 1993; Bischoff et al. 1993). Another relatively rare mineral gaylussite, $\text{Na}_2\text{Ca}(\text{CO}_3)_2 \cdot 5\text{H}_2\text{O}$, was also found in Mono Lake, which is believed to presently precipitate in a year-round process that probably only began after 1970 when the salinity of the lake exceeded 80‰ (Bischoff et al. 1991), owing to the diversion of tributary streams by the Los Angeles Department of Water and Power. The salinity of Mono Lake has steadily increased since 1941 from 50‰ to the current ~90‰ (Bischoff et al. 1991; Budinoff and Hollibaugh 2007; Jørgensen et al. 2008). In this paper, we report a new biologically related phosphate mineral, hazenite $\text{KNaMg}_2(\text{PO}_4)_2 \cdot 14\text{H}_2\text{O}$, discovered at this locality. The new mineral is named after Robert Miller Hazen, a staff scientist at the Geophysical Laboratory, Carnegie Institution of Washington, for his outstanding contributions to Earth sciences in general, and mineralogy in particular, in both research and education, as well as for his contributions to understanding interactions between minerals and organic molecules. Part of the cotype sample has been deposited at the Mineral Museum of the University of Arizona (Catalog no. 18812). The new mineral and its name have been approved by the Commission on New Minerals and Nomenclature and Classification (CNMNC) of the International Mineralogical Association (IMA2007-061).

SAMPLE DESCRIPTION AND EXPERIMENTAL METHODS

Occurrence, physical, and chemical properties, and Raman spectra

Hazenite was found in and/or on completely dried-out or decomposed green algae (cyanobacteria) on porous calcium-carbonate (mainly calcite and aragonite) substrates or tufas on the south shore of Mono Lake (lat. $37^\circ 56' 31''\text{N}$, long. $119^\circ 1' 42''\text{W}$). It is observed as radiating clusters or tufts of elongated (along [100]) prismatic crystals up to $0.40 \times 0.12 \times 0.06$ mm (Fig. 1). The mineral is colorless, transparent with white streak and vitreous luster. It is brittle, with Mohs hardness of 2–2.5; cleavage is good on {001} and no twinning was observed. The measured



FIGURE 1. Photograph of hazenite crystals.

and calculated densities are 1.91(3) and 1.88(2) g/cm³, respectively. Hazenite decomposes easily in 10% HCl at room temperature and dissolves slowly in water. Optically, it is biaxial (+), with $n_\alpha = 1.494(1)$, $n_\beta = 1.498(1)$, $n_\gamma = 1.503(1)$, $2V_{\text{meas}} = 41(2)^\circ$, $2V_{\text{calc}} = 42^\circ$, $X = b$, $Y = c$, $Z = a$, and does not fluoresce under long- or short-wave ultraviolet light. The dispersion is strong with $r < v$.

The chemical composition was determined with a CAMECA SX50 electron microprobe at 10 kV and 5 nA (<http://ruff.info>), yielding an average composition (wt%) (20 points) of K_2O 8.27(28), Na_2O 5.35(26), MgO 15.09(41), CaO 0.21(26), P_2O_5 25.27(43), SO_3 0.54(50), and H_2O (by difference, to sum up to 100%) 45.27. Based on the number of oxygen atoms bonded to P^{5+} , which is 8 (see below), we obtain an empirical formula of $\text{K}_{0.97}(\text{Na}_{0.96}\text{Ca}_{0.02})\text{Mg}_{2.07}(\text{P}_{0.98}\text{S}_{0.02}\text{O}_4)_2 \cdot 13.90\text{H}_2\text{O}$, which can be simplified ideally as $\text{KNaMg}_2(\text{PO}_4)_2 \cdot 14\text{H}_2\text{O}$.

The Raman spectra of hazenite were collected on a randomly oriented crystal from 9 scans at 30 s and 100% power per scan on a Thermo Almega microRaman system, using a solid-state laser with a wavelength of 532 nm and a thermoelectrically cooled CCD detector. The laser is partially polarized with 4 cm^{-1} resolution and a spot size of 1 μm .

X-ray crystallography

Both powder and single-crystal X-ray diffraction data of hazenite were collected on a Bruker X8 APEX2 CCD X-ray diffractometer equipped with graphite-monochromatized $\text{MoK}\alpha$ radiation. Based on orthorhombic symmetry (see below), the unit-cell parameters determined from the powder X-ray diffraction data (Table 1)

TABLE 1. Powder X-ray diffraction data for hazenite

Intensity	d_{meas}	d_{calc}	<i>h k l</i>
6	5.899(5)	5.903	1 0 1
17	5.606(6)	5.600	0 0 2
11	5.485(5)	5.480	0 4 1
4	5.120(3)	5.123	0 2 2
7	4.824(3)	4.827	1 3 1
30	4.659(3)	4.656	1 4 0
100	4.302(5)	4.298	1 4 1
22	4.184(9)	4.175	0 4 2
11	3.744(1)	3.744	0 5 2
3	3.468(8)	3.460	2 0 0
20	3.262(6)	3.268	1 1 3
6	3.081(3)	3.078	2 3 1
6	3.027(9)	3.037	0 8 1
11	2.931(1)	2.930	2 4 1
32	2.803(3)	2.805	0 0 4
43	2.786(5)	2.790	0 1 4
51	2.767(2)	2.766	2 5 1
48	2.742(1)	2.743	0 8 2
51	2.670(1)	2.671	2 4 2
2	2.560(4)	2.556	0 4 4
3	2.529(2)	2.531	2 1 3
7	2.357(1)	2.358	2 4 3
6	2.329(3)	2.327	2 8 0
4	2.209(1)	2.210	1 6 4
9	2.180(3)	2.177	2 0 4
10	2.172(1)	2.172	2 1 4
10	2.151(1)	2.151	2 8 2
11	2.075(3)	2.074	2 7 3
6	2.000(2)	2.003	2 5 4
3	1.914(6)	1.908	3 3 3
8	1.882(1)	1.883	2 1 1
7	1.881(6)	1.874	2 9 3
7	1.831(1)	1.830	3 5 3
7	1.804(2)	1.802	2 4 5
7	1.794(1)	1.793	2 1 20
7	1.749(1)	1.748	0 9 5
5	1.745(1)	1.744	3 3 4
11	1.734(1)	1.733	4 0 0
5	1.709(2)	1.711	2 1 22
3	1.667(2)	1.666	3 8 3
4	1.614(1)	1.614	2 3 6
2	1.591(1)	1.590	2 4 6
2	1.466(1)	1.466	4 8 2
2	1.371(2)	1.373	4 9 3
2	1.355(1)	1.355	5 4 0
3	1.345(1)	1.345	5 4 1
2	1.319(2)	1.321	3 8 6
3	1.291(1)	1.291	3 1 61
2	1.258(1)	1.258	2 5 8

are $a=6.936(2)$, $b=25.15(1)$, $c=11.212(4)$ Å, and $V=1956(1)$ Å³. For comparison, included in Table 1 are also the powder X-ray diffraction data calculated using the program XPOW (Downs et al. 1993) based on the determined structure.

Single-crystal X-ray diffraction data of hazenite were collected from a nearly equi-dimensional crystal ($0.05 \times 0.05 \times 0.06$ mm) with frame widths of 0.5° in ω and 30 s counting time per frame. All reflections were indexed on the basis of an orthorhombic unit-cell (Table 2). The intensity data were corrected for X-ray absorption using the Bruker program SADABS. The systematic absences of reflections suggest possible space group $Pmnb$ (62) or $P2_1nb$ (33). The crystal structure was solved and refined using SHELX97 (Sheldrick 2008) based on the space group $Pmnb$, because it yielded the better refinement statistics in terms of bond lengths and angles, atomic displacement parameters, and R factors. The detailed structure refinement procedures were similar to those described by Yang and Sun (2004). The positions of all atoms were refined with anisotropic displacement parameters,

except for H atoms, which were refined with a fixed isotropic displacement parameter ($U_{eq} = 0.04$). Final coordinates and displacement parameters of atoms are listed in Table 3, and selected bond distances in Table 4.

DISCUSSION

Crystal structure

The structure of hazenite is identical to that of the bacteria-formed phosphate reported by Yang and Sun (2004), who have given a detailed structural description of this compound, including its hydrogen bonding and structural relationships with struvite-type materials. However, for a direct comparison with struvite-type materials, we have adapted a different unit-cell setting for hazenite (Table 5). Specifically, the a and b axes for the phosphate compound studied by Yang and Sun (2004) are

TABLE 2. Summary of crystallographic data and refinement results for hazenite from single-crystal X-ray diffraction

Ideal structural formula	KNaMg ₂ (PO ₄) ₂ ·14H ₂ O
Space group	<i>Pmnb</i> (no. 62)
a (Å)	6.9349(4)
b (Å)	25.1737(15)
c (Å)	11.2195(8)
V (Å ³)	1958.7(3)
Z	4
ρ_{calc} (g/cm ³)	1.875
λ (Å)	0.71069
μ (mm ⁻¹)	0.62
θ range for data collection	1.62 to 28.11
No. of reflections collected	11511
No. of independent reflections	2416
No. of reflections with $I > 2\sigma(I)$	2339
No. of parameters refined	198
R_{int}	0.035
Final R factors [$I > 2\sigma(I)$]	$R1 = 0.042$, $wR2 = 0.090$
Final R factors (all data)	$R1 = 0.045$, $wR2 = 0.093$
Goodness-of-fit	1.21

TABLE 4. Selected non-hydrogen bond distances (Å) in hazenite

P1-O1	1.544(3)	P2-O4	1.535(3)
P1-O2	1.540(3)	P2-O5	1.534(3)
P1-O3 (x2)	1.530(2)	P2-O6 (x2)	1.543(2)
Avg.	1.536	Avg.	1.539
Mg1-OW7	2.032(3)	Mg2-OW11 (x2)	2.060(2)
Mg1-OW8	2.040(3)	Mg2-OW12 (x2)	2.074(2)
Mg1-OW9 (x2)	2.066(2)	Mg2-OW13	2.093(3)
Mg1-OW10 (x2)	2.083(2)	Mg2-OW14	2.109(3)
Avg.	2.062	Avg.	2.078
K-O5	2.672(3)	Na-OW10 (x2)	2.563(3)
K-OW11 (x2)	2.980(2)	Na-OW15 (x2)	2.425(3)
K-OW12 (x2)	2.908(2)	Na-OW15 (x2)	2.472(3)
K-OW13	3.064(3)		
Avg.	2.919	Avg.	2.487

TABLE 3. Coordinates and displacement parameters of atoms for hazenite

Atom	x	y	z	U_{11} (Å ²)	U_{22} (Å ²)	U_{33} (Å ²)	U_{23} (Å ²)	U_{13} (Å ²)	U_{12} (Å ²)	U_{eq} (Å ²)
K	3/4	0.2216(1)	0.5014(1)	0.0535(7)	0.0290(6)	0.0376(6)	-0.0096(4)	0	0	0.0401(4)
Na	3/4	0.0004(1)	0.9259(2)	0.0272(10)	0.0405(11)	0.0384(11)	0.0091(8)	0	0	0.0354(6)
Mg1	3/4	0.9543(1)	0.6388(1)	0.0140(7)	0.0149(6)	0.0175(6)	0.0022(4)	0	0	0.0155(4)
Mg2	3/4	0.2237(1)	0.8366(1)	0.0185(7)	0.0154(6)	0.0169(7)	0.0004(4)	0	0	0.0169(4)
P1	1/4	0.1281(1)	0.7219(1)	0.0143(5)	0.0127(5)	0.0145(5)	0.0001(3)	0	0	0.0138(3)
P2	3/4	0.1272(1)	0.2201(1)	0.0113(5)	0.0129(5)	0.0152(5)	0.0023(3)	0	0	0.0131(3)
O1	1/4	0.1358(1)	0.8584(2)	0.0256(14)	0.0311(14)	0.0225(13)	-0.0032(11)	0	0	0.0264(6)
O2	1/4	0.1828(1)	0.6608(2)	0.0287(14)	0.0159(12)	0.0277(14)	0.0033(10)	0	0	0.0241(6)
O3	0.4317(3)	0.0976(1)	0.6859(2)	0.0184(9)	0.0218(9)	0.0267(9)	-0.0006(7)	-0.0006(7)	0.0012(7)	0.0223(4)
O4	3/4	0.0673(1)	0.1945(3)	0.0229(13)	0.0157(12)	0.0357(15)	-0.0027(10)	0	0	0.0248(6)
O5	3/4	0.1380(1)	0.3546(2)	0.0223(13)	0.0258(13)	0.0172(12)	0.0001(10)	0	0	0.0217(6)
O6	0.9315(3)	0.1519(1)	0.1630(2)	0.0156(9)	0.0256(9)	0.0227(9)	0.0047(7)	0.0026(7)	-0.0021(7)	0.0213(4)
OW7	3/4	0.0095(1)	0.5068(3)	0.0214(14)	0.0303(16)	0.0410(17)	0.0174(13)	0	0	0.0309(7)
OW8	3/4	0.8944(1)	0.7614(3)	0.0138(13)	0.0440(18)	0.0510(19)	0.0293(15)	0	0	0.0363(8)
OW9	0.5349(3)	0.9161(1)	0.5446(2)	0.0220(10)	0.0333(10)	0.0211(9)	-0.0016(7)	0.0024(8)	-0.0094(8)	0.0255(4)
OW10	0.9578(3)	0.9947(1)	0.7375(2)	0.0208(10)	0.0192(9)	0.0333(10)	0.0005(8)	-0.0049(8)	-0.0017(8)	0.0244(4)
OW11	0.5359(3)	0.2493(1)	0.7231(2)	0.0301(12)	0.0212(11)	0.0412(12)	0.0051(8)	-0.0117(9)	0.0002(8)	0.0308(5)
OW12	0.9662(3)	0.1966(1)	0.9489(2)	0.0281(11)	0.0297(10)	0.0215(9)	0.0038(8)	0.0006(8)	0.0040(8)	0.0264(5)
OW13	3/4	0.1554(1)	0.7305(3)	0.0203(14)	0.0229(14)	0.0479(18)	-0.0114(12)	0	0	0.0304(7)
OW14	3/4	0.2975(1)	0.9256(3)	0.0768(25)	0.0176(14)	0.0270(16)	-0.0042(11)	0	0	0.0405(8)
OW15	0.0100(3)	0.0597(1)	0.9823(2)	0.0404(13)	0.0301(11)	0.0326(11)	-0.0035(8)	0.0081(10)	0.0002(9)	0.0344(5)
H11	0.841(5)	0.026(1)	0.484(3)							
H21	0.850(5)	0.879(1)	0.790(3)							
H31	0.454(5)	0.896(1)	0.572(3)							
H32	0.549(5)	0.909(1)	0.467(3)							
H41	0.057(6)	0.977(1)	0.758(3)							
H42	0.999(5)	0.023(2)	0.723(3)							
H51	0.450(6)	0.230(1)	0.707(3)							
H52	0.503(5)	0.278(2)	0.714(3)							
H61	0.940(5)	0.181(1)	0.017(3)							
H62	0.051(6)	0.179(1)	0.921(3)							
H71	0.847(5)	0.135(1)	0.722(3)							
H81	3/4	0.327(2)	0.890(5)							
H82	3/4	0.305(2)	0.003(5)							
H91	0.057(6)	0.080(1)	0.946(3)							
H92	0.954(5)	0.076(1)	0.039(3)							

TABLE 5. Comparison of crystallographic data for hazenite and other struvite-type materials

Chemical formula	Struvite (NH ₄)Mg(PO ₄)·6H ₂ O	Struvite-(K) KMg(PO ₄)·6H ₂ O	Synthetic NaMg(PO ₄)·7H ₂ O	Hazenite KNaMg ₂ (PO ₄) ₂ ·14H ₂ O
<i>a</i> (Å)	6.955(1)	6.903	6.731(2)	6.9349(4)
<i>b</i> (Å)	6.142(1)	6.174	6.731(2)	25.1737(15)
<i>c</i> (Å)	11.218(2)	11.164	10.982(4)	11.2195(8)
<i>V</i> (Å ³)	479.2	475.03	497.5	1958.7
Space group	<i>Pmn</i> 2 ₁ (no. 31)	<i>Pmn</i> 2 ₁ (no. 31)	<i>P</i> ₄ ₂ / <i>mmc</i> (no. 131)	<i>Pmnb</i> (no. 62)
<i>Z</i>	2	2	2	4
ρ_{cal} (g/cm ³)	1.71	1.88	1.77	1.88
Strong powder lines	4.257(100) 5.601(60) 2.919(55) 2.690(50) 2.660(45) 5.905(40) 4.139(40)	4.26(100) 3.27(90) 4.14(80) 2.650(70) 2.699(50) 2.905(50)	2.869(100) 5.739(98) 4.255(91) 6.731(66) 2.380(56) 1.798(45) 3.366(40)	4.302(100) 2.767(50) 2.670(50) 2.742(50) 2.786(40) 2.803(30) 4.659(30)
Reference	1	2	3	4

Note: References: 1 = Ferraris et al. (1986); 2 = Graeser et al. (2007); 3 = Mathew et al. (1982); 4 = This study.

switched in hazenite.

There are many structural similarities between hazenite and struvite-type materials, which have been of great interest because of their broad and important biological, agricultural, and industrial implications (e.g., Dickens and Brown 1972; Banks et al. 1975; Angoni et al. 1998; Doyle and Parsons 2002; Trobajo et al. 2007; Haferburg et al. 2008; Weil 2008; Chen et al. 2010). For example, Mg²⁺ cations in these compounds are all coordinated octahedrally by six H₂O molecules whose H atoms are strongly bonded to oxygen atoms in (PO₄)³⁻ groups. No H₂O molecules are shared between Mg(H₂O)₆ octahedra. Excluding H⁺, there are six symmetrically nonequivalent cation sites in hazenite, two for Mg²⁺ (Mg1 and Mg2), two for

P⁵⁺ (P1 and P2), one for Na⁺, and one for K⁺. Viewed along *c*, the hazenite structure can be considered as stacking of three types of layers along the *b*-axis, in a repeating sequence of ABCBABC... (Fig. 2a), where layer A consists of Mg1(H₂O)₆ octahedra and NaO₆ trigonal prisms, layer B of P1O₄ and P2O₄ tetrahedra, and layer C of Mg2(H₂O)₆ octahedra and very irregular KO₆ polyhedra. These layers are linked together by hydrogen bonds, plus the K-O bonds between layers B and C (K-O5-P2). Interestingly, the combination of layers B and C in hazenite exhibits a configuration directly comparable with the struvite-(K) structure (Mathew and Schroeder 1979; Graeser et al. 2008) (Fig. 2b). In fact, the struvite-(K) structure can be readily achieved from the hazenite structure with a

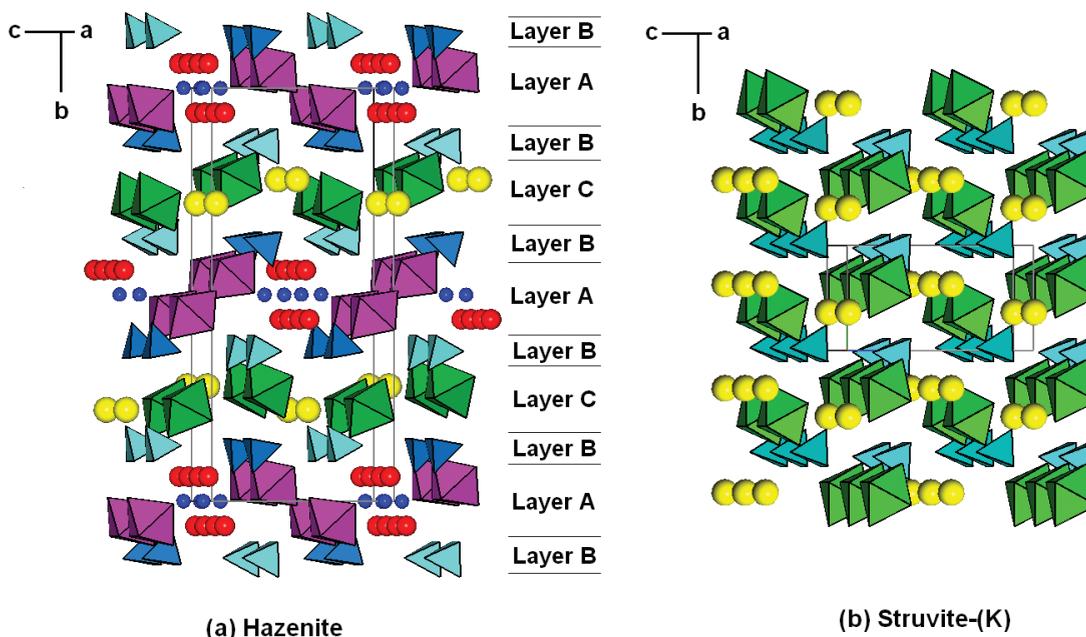


FIGURE 2. (a) Crystal structure of hazenite. Tetrahedra = PO₄³⁻ groups and octahedra = Mg(H₂O)₆. The largest, medium, and smallest spheres represent K, Ow9 (the water molecule bonded to Na only), and Na, respectively. See the text for the definition of layers A, B, and C. (b) Crystal structure of struvite-(K). Tetrahedra, octahedra, and spheres represent PO₄³⁻ groups, Mg(H₂O)₆, and K, respectively.

substitution of layer C for layer A, although the average K-O bond length (2.919 Å) in hazenite is slightly shorter than that (3.034 Å) in struvite-(K). However, the combination of layers A and B in hazenite does not show much resemblance to the $\text{NaMgPO}_4 \cdot 7\text{H}_2\text{O}$ structure determined by Mathew et al. (1982). In particular, as noted by Yang and Sun (2004), Na^+ in hazenite is situated in a six-coordinated trigonal prism, with four H_2O molecules bonded to Na^+ cations only, and each prism shares two edges with two neighboring prisms to form zigzag chains parallel to **a**, accounting for its elongated morphology along [100]. In contrast, Na^+ in $\text{NaMgPO}_4 \cdot 7\text{H}_2\text{O}$ is in an elongated octahedron, with two H_2O molecules bonded to Na^+ only, and each octahedron sharing two corners with two other $\text{Na}(\text{H}_2\text{O})_6$ octahedra to form straight chains along **c**.

Dickens and Brown (1972) examined several struvite-type compounds and their structural relationships and presented a general chemical formula for such materials: $M^+X^{2+}(\text{YO}_4) \cdot n\text{H}_2\text{O}$, where $n = 6-8$, $Y = \text{P}$ or As , and the radius of monovalent M^+ is greater than that of divalent X^{2+} . Banks et al. (1975) and Weil (2008) investigated several struvite analogs of the type $M^+\text{MgYO}_4 \cdot 6\text{H}_2\text{O}$ ($M^+ = \text{NH}_4, \text{K}, \text{Rb}, \text{Cs}, \text{and Tl}$). Recently, struvite analogs with $X^{2+} = \text{Co}$ and Ni have been reported (Blachnik et al. 1997; Touaïher et al. 2001; El Bali et al. 2005; Trobajo et al. 2007; Hafnerburg et al. 2008). Although Na^+ ions have been included in many laboratory synthesis preparations, no Na-analog of the struvite-type structure has been identified thus far (e.g., Banks et al. 1975; Mathew et al. 1982). It therefore has been generally agreed that the struvite-type structure is unable to accommodate monovalent M^+ cations smaller than K^+ and the incorporation of Na^+ requires extra H_2O molecules ($n > 6$) to compensate its smaller ionic size, as in $\text{NaMgPO}_4 \cdot 7\text{H}_2\text{O}$ (Mathew et al. 1982) and $\text{NaNiPO}_4 \cdot 7\text{H}_2\text{O}$ (Trobajo et al. 2007). Clearly, the determination of the hazenite structure provides further support to this statement. Furthermore, to our knowledge, hazenite is the first struvite-type compound that contains two structurally distinct M^+ cations (K and Na), pointing to the exclusive role of biological activities in its mineralization process.

Raman spectra

There have been numerous studies on struvite-related materials by means of infrared and Raman spectroscopic techniques (Banks et al. 1975; Angoni et al. 1998; Stefov et al. 2004, 2005; Frost et al. 2005; Koleva 2007). Yang and Sun (2004) measured the infrared spectrum of hazenite and showed the remarkable similarities in the IR spectra between hazenite and struvite. Here, we present our Raman spectroscopic measurements of hazenite in Figure 3, along with the Raman spectra of a struvite sample from the RRUFF project (deposition R050540: <http://rruff.info/R050540>) for comparison. A tentative assignment of major Raman bands for hazenite is given in Table 6. It is evident from Figure 3 that the PO_4 symmetric stretching (ν_1) mode at $\sim 934 \text{ cm}^{-1}$ for hazenite is shifted to lower frequency relative to that for struvite by $\sim 13 \text{ cm}^{-1}$. This shift is expected, because, as compared with the $\text{NH}_4^+\text{-O}$ bonds in struvite, the stronger M^+ ($=\text{K}, \text{Na}$)-O interactions in hazenite make the intra-molecular P-O bonds weaker, giving rise to the shift of the P-O stretching mode to lower frequency. Similar observations have been reported for struvite-type compounds by Stefov et al. (2004, 2005) and for

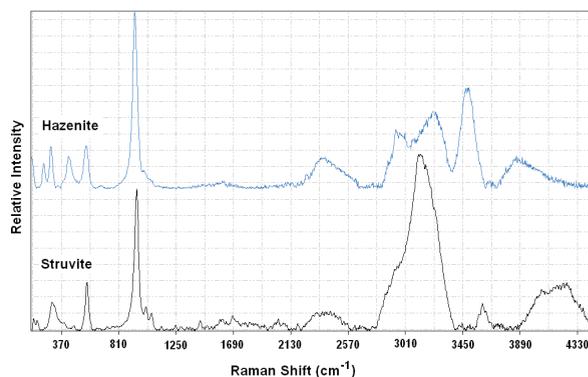


FIGURE 3. Raman spectrum of hazenite, along with that of struvite taken from the RRUFF project (see the text for details) for comparison. The spectra are shown with vertical offset for more clarity.

TABLE 6. Tentative assignment of Raman bands for hazenite

Bands (cm^{-1})	Intensity	Assignment
3900–2500	Relatively strong, broad	Water ν_1 - ν_3 sym- and antisymmetric stretching
2380	Weak, broad	Water-phosphate hydrogen bonding
1620	Very weak, broad	Water bending
1100–988	Shoulder, weak	ν_3 (PO_4) antisymmetric stretching
935	Strong, sharp	ν_1 (PO_4) symmetric stretching
685	Very weak	Water-water H bonding
565	Relatively strong, sharp	ν_4 (PO_4) antisymmetric bending
430	Relatively strong, sharp	ν_2 (PO_4) symmetric bending
290	Relatively strong, sharp	Mg-O stretching
258–136	Relatively strong, multi-peaks	Lattice and Mg-O vibrational modes

ditmarite-type compounds $M^+M''\text{PO}_4 \cdot \text{H}_2\text{O}$ ($M^+ = \text{K}^+, \text{NH}_4^+$; $M'' = \text{Mn}^{2+}, \text{Co}^{2+}, \text{Ni}^{2+}$) by Koleva (2007).

Formation

Hazenite was first found to form in a microbiological culture containing the filamentous, photosynthetic cyanobacterium *Lyngbya sp.*, which was isolated from the interior of calcareous tufa deposits collected from Mono Lake (Yang and Sun 2004). The growth medium was composed of distilled water and the lake water in a 1:1 ratio, solidified with 1.2% agar. At the end of growth and as the medium in the petri dish dried up, a few tufts of elongated-tabular crystals of hazenite appeared around dead cyanobacterial colonies. Following this discovery, the subsequent field trips to Mono Lake led us to find the mineral in nature.

Struvite is a relatively rare mineral in nature and its formation by microorganisms was first documented as early as in the 19th century (Robinson 1889). It is well known now that struvite can form inside living bacteria, as a diagenetic byproduct of bird feces, in kidney stones, or in association with decaying organisms or decomposing organic matter (such as in old graveyards, under floors of stables, in guano, or wastewater treatments). One common feature shared by these environments is high alkalinity, suggesting that high pH conditions are necessary for the formation of struvite and/or similar phosphate minerals (Rivadeneira et al. 1993; Grover et al. 1997; Doyle and Parsons 2002; Chen et al. 2010). Because of the very hypersalinity and high alkalinity of the Mono Lake water, hazenite is also believed to form in high-pH conditions with the strong involvement of cyanobacte-

rial activities. Yang and Sun (2004) suggested that phosphate metabolism or phosphate-containing organic compounds could be a good source for inorganic phosphate. In the cytoplasm of a living bacterium, much of the phosphate in a cell occurs in the form of phospholipids in the membranes. Presumably, following the death of algae, the slow breakdown of phosphorus-containing organic compounds, on the one hand, would become the steady supply for phosphate. The lake water, on the other hand, provides Na, K, and Mg for the hazenite formation. In addition, the cell membranes may be energetically favored for nucleation of crystals, as they can provide structures that serve as substrates for heterogeneous crystal nucleation (with lower energy barrier than homogeneous nucleation) and stereochemical arrangements for the mineral components (Lowenstam and Weiner 1989). Noteworthy, Ben Omar et al. (1994, 1995, 1998) and Gonzalez-Munoz et al. (1993, 1996) conducted a series of experiments on struvite formation with several soil inhabiting genera, like *Azotobacter*, *Bacillus*, *Myxococcus*, and *Pseudomonas*. One of their studies (Ben Omar et al. 1994) shows that the struvite precipitation would not occur until the autolysis of *M. Xanthus* cultures and the actual physical presence of the bacteria was necessary. Ben Omar et al. (1995, 1998) also found that dead cells, disrupted cells, and membrane fractions of this microorganism induce the struvite crystallization. Evidently, their observations lend support to our proposed biomineralization process of hazenite.

It has been well documented that ikaite, $\text{CaCO}_3 \cdot 6\text{H}_2\text{O}$, precipitates seasonally in Mono Lake (Shearman et al. 1989; Council and Bennett 1993; Bischoff et al. 1993). During the winter, cold water temperatures and high concentrations of PO_4^{3-} and organic carbon inhibit calcite precipitation, allowing the metastable ikaite to form. During the spring warming, however, ikaite decomposes to CaCO_3 and H_2O , occasionally leaving pseudomorphs of the primary precipitate. Yet, it is unclear whether and how the hazenite formation in Mono Lake may be affected by the seasonal weather changes or the diversion of tributary streams (Bischoff et al. 1991). Apparently, more systematic studies, including field observations, are needed to address these questions.

In environmental mineralogy, the precipitation of struvite from sewage has been a subject of considerable investigations because it may offer a potential route for dephosphorization of wastewater from industries and to recover phosphates for recycling in the form of fertilizer. At present, there are numerous research projects focusing on the controlled precipitation of struvite and its later utilization as a fertilizer (Booker et al. 1999; Doyle and Parsons 2002; Shu et al. 2006; Machnicha et al. 2008; Forrest et al. 2008) The discovery of hazenite and a better understanding of its formation process, especially in terms of the degree of biologically controlled vs. biologically induced mineralization (Lowenstam 1981; Mann 1983), will undoubtedly provide us with additional knowledge in this regard and facilitate future work on biomineralization of struvite-type phosphate materials.

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