INVESTIGATIONS OF CRYSTAL-CHEMICAL VARIABILITY IN LEAD URANYL OXIDE HYDRATES. II. FOURMARIERITE

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ABSTRACT

Structures have been refined for twelve crystals of fourmarierite from various localities in the Democratic Republic of Congo, and one synthetic crystal. Single-crystal diffraction data were collected using MoKα X-radiation and a CCD-based detector mounted on a Bruker three-circle diffractometer. All crystals have orthorhombic symmetry, space group Bb21m. The natural crystals have similar unit-cell parameters: a 14.00–14.03, b 16.40–16.48, c 14.32–14.38 Å. The parameters for the synthetic crystal, which contains less Pb than the natural crystals, are a 13.938(2), b 16.638(3), c 14.672(2) Å. The structures refined to agreement indices (R1) in the range of 3.6 to 6.2%. The structure of fourmarierite contains uranyl pentagonal bipyramids that share edges and corners to form sheets oriented parallel to (001). There are two distinct Pb²⁺ cations and eight H₂O groups located in the interlayer. On the basis of the structure refinements, the site occupancy of Pb(1) in the natural crystals is typically deficient, ranging from 71 to 100%, whereas the site occupancy of Pb(2) is ~100% in the crystals studied. In the synthetic crystal the Pb(1) and Pb(2) occupancies are 13% and 87%, respectively. Increase of the Pb content is correlated with a minor increase in the a unit-cell parameter, as well as large decreases in the b and c unit-cell parameters, and the unit-cell volume. The substitution O²⁻ ↔ (OH)⁻ at two anion sites [O(12) and O(15)] in the sheet of uranyl polyhedra provides the charge-balance mechanism associated with interlayer cation variation. The structural formula for fourmarierite may be written as Pb₁₋ₓ[(UO₂)₄O₃₋₂ₓ(OH)₄₊₂ₓ](H₂O)₄, Z = 8, with the constituents of the sheet of uranyl polyhedra enclosed in square braces.

Keywords: fourmarierite, uranyl mineral, structure determination, crystal chemistry.

Nous avons affiné la structure de douze cristaux de fourmariérite provenant de diverses localités de la République Démocratique du Congo, et d’un cristal synthétique. Les données ont été prélevées sur cristal unique avec rayonnement MoKα et un diffractomètre à trois cercles Bruker muni d’un détecteur de type CCD. Tous les cristaux possèdent une symétrie orthorhombique, groupe spatial Bb21m. Les cristaux naturels ont des paramètres réticulaires semblables: a 14.00–14.03, b 16.40–16.48, c 14.32–14.38 Å. Les paramètres du cristal synthétique, qui contient moins de Pb que les cristaux naturels, sont a 13.938(2), b 16.638(3), c 14.672(2) Å. Les structures ont été affinées jusqu’à un résidu R1 entre 3.6 et 6.2%. La structure de la fourmariérite contient des bipyramides à uranyle pentagonales qui partagent arêtes et coins pour former des feuillots parallèles à (001). Il y a deux sites distincts occupés par des cations Pb²⁺ et huit groupes H₂O situés dans l’interfeuillet. D’après les résultats des affinements, le site Pb(1) dans les cristaux naturels est typiquement déficitaire, contenant entre 71 et 100%, tandis que le taux d’occupation de Pb(2) est environ 100% dans les cristaux étudiés. Dans le cristal synthétique que nous avons caractérisé, les sites Pb(1) et Pb(2) ont une occupation de 13 et 87%, respectivement. L’augmentation de la teneur en Pb mène à une légère augmentation du paramètre a, et une diminution beaucoup plus importante des paramètres b et c, ainsi que du volume de la maille. La substitution O²⁻ ↔ (OH)⁻ à deux sites anioniques [O(12) et O(15)] dans le feuillet de polyédres à uranyle assure un équilibre des charges en conséquence de la variation de la teneur des cations dans l’interfeuillet. La formule structurale de la fourmariérite peut s’écrire Pb₁₋ₓ[(UO₂)₄O₃₋₂ₓ(OH)₄₊₂ₓ](H₂O)₄, Z = 8, les composants du feuillet de polyédres à uranyle étant entre crochets.

Mots-clés: fourmariérite, minéral à uranyle, détermination de la structure, chimie cristalline.
Renewed interest in uranyl minerals results in part from their significance in environmental issues (Finn et al. 1996, Wronkiewicz et al. 1996, Buck et al. 1998, Abdelouas et al. 1999). Pb uranyl oxide hydrates as a group form from the oxidation of geologically old uraninite, because of the buildup of radiogenic Pb (Frondel 1958, Finch & Ewing 1992). Seven Pb uranyl oxide hydrate minerals have been described (Table 1). Of these, fourmarierite is of particular interest because of its structural similarity to schoepite (Finch et al. 1996), which is present in the oxidation products of primary uraninite (Finch & Ewing 1992), and may also form by alteration of nuclear waste in a geological repository.

The understanding of uranyl structures has been enhanced by the recent application of CCD-based area detectors of X-rays to mineralogical research (Burns 1997, 1998a, b, c, 1999, 2000a, b, Burns & Finch 1999, Burns & Hanchar 1999, Hill & Burns 1999, Burns & Hill 2000a, b, Burns et al. 2000). This approach has been particularly successful in the case of Pb uranyl oxide hydrates, with the solutions of the structures of masuyite (Burns & Hanchar 1999), vandendriesscheite (Burns 1997), richetite (Burns 1998a), and wölsendorfite (Burns 1999). Previously, the structures were reported for fourmarierite (Piret 1985), sayrite (Piret et al. 1983) and curite (Taylor et al. 1981).

Natural Pb uranyl oxide hydrates possess fascinating structures composed of sheets of uranyl polyhedra; they exhibit variable contents of the interlayer, perhaps depending upon the age of the primary uraninite and the nature of the alteration processes. A thorough understanding of the crystal chemistry of these minerals has not yet been attained, but such details are essential to establish an understanding of the relationships between the structures of these minerals and their paragenesis. We have undertaken a systematic study of the crystal chemistry of Pb uranyl oxide hydrates. In the first paper in a series, we report the results on our studies of curite (Li & Burns 2000a). In this contribution, the second in the series, we present data for thirteen crystals of fourmarierite. The primary motivation of the current work is to determine the extent of Pb variation in fourmarierite, and to develop an understanding of the crystal-chemical mechanisms by which such variation occurs.

The structure of fourmarierite was reported for a crystal from Shinkolobwe by Piret (1985). On the basis of the structure determination, the formula of fourmarierite was given as Pb[\((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_4. The structure contains two symmetrically independent Pb sites, both of which were reported to be fully occupied.

**EXPERIMENTAL METHODS**

Specimens containing fourmarierite were provided by the Canadian Museum of Nature, the Royal Ontario Museum, and Prof. Rodney C. Ewing, University of Michigan. In addition, we were able to synthesize crystals of fourmarierite suitable for single-crystal study.

**Synthesis of crystals**

Crystals of fourmarierite were synthesized using hydrothermal techniques. A solution was prepared with 0.4 M uranyl nitrate and PbO, with a Pb:U ratio of 1:8. The pH of the solution was adjusted using 1 M NaOH solution to 3.2. The reactants were transferred to a Teflon-lined Parr bomb that was placed in an oven and heated at 200°C for 14 days. The resulting precipitate was recovered by filtration, and contained crystals of fourmarierite and a new Pb uranyl oxide hydrate that is described by Li & Burns (2000b).

**X-ray diffraction**

Well-formed platy crystals of fourmarierite were chosen for X-ray-diffraction experiments. Single-crystal diffraction data were collected for each crystal using MoKα X-radiation and a CCD-based detector mounted on a Bruker three-circle diffractometer. Localities and sizes of the crystals, as well as details of the data collections, are listed in Table 2. The data were integrated and corrected for Lorentz, polarization, and background.

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**TABLE 1. CRYSTALLOGRAPHIC AND COMPOSITIONAL DATA OF Pb URANYL OXIDE HYDRATE MINERALS**

<table>
<thead>
<tr>
<th>Pb-UO₂ minerals</th>
<th>Structural formula</th>
<th>S.G.</th>
<th>a (Å)</th>
<th>b (Å)</th>
<th>c (Å)</th>
<th>α (°)</th>
<th>β (°)</th>
<th>γ (°)</th>
<th>Pb/U</th>
</tr>
</thead>
<tbody>
<tr>
<td>wölsendorfite (1)</td>
<td>Pb₆[((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>Cmcm</td>
<td>14.131</td>
<td>13.885</td>
<td>55.969</td>
<td>1:2.15</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>sayrite (2)</td>
<td>Pb₂((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>P2_1/c</td>
<td>10.704</td>
<td>9.690</td>
<td>14.533</td>
<td>90.402</td>
<td>2:2.5</td>
<td></td>
<td></td>
</tr>
<tr>
<td>curite (3)</td>
<td>Pb₂((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>Pna2_1</td>
<td>12.551</td>
<td>13.003</td>
<td>8.390</td>
<td>1:6.27</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>masuyite (4)</td>
<td>Pb₂((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>Pn</td>
<td>12.241</td>
<td>7.008</td>
<td>6.983</td>
<td>90.402</td>
<td>3:1</td>
<td></td>
<td></td>
</tr>
<tr>
<td>fourmarierite (5)</td>
<td>Pb₁₄((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>Bb₂m</td>
<td>13.986</td>
<td>16.400</td>
<td>14.29</td>
<td>1:4.0</td>
<td></td>
<td></td>
<td></td>
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<tr>
<td>richetite (6)</td>
<td>Pb₁₄((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>P1</td>
<td>20.939</td>
<td>12.100</td>
<td>16.345</td>
<td>103.87</td>
<td>115.37</td>
<td>90.27</td>
<td>1:4.15</td>
</tr>
<tr>
<td>vandendriesscheite (7)</td>
<td>Pb₁₄((\text{UO}_2)\text{O}_4\text{O}_3\text{O}}_4\text{H}_4\text{H}_2\text{O}]_2</td>
<td>Pbcn</td>
<td>14.1165</td>
<td>41.478</td>
<td>14.5347</td>
<td>1:6.36</td>
<td></td>
<td></td>
<td></td>
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</tbody>
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effects using the Bruker program SAINT. Each dataset was corrected for absorption. Where well-developed faces bounded the crystals, Gaussian interpolation was used, but where crystal faces were less well developed, semi-empirical corrections were done on the basis of equivalent reflections. The structures were refined on the basis of $F^2$ for all unique data in space group $\text{Bb}21 \text{m}$ using the SHELXTL software package. Refinement began with isotropic displacement parameters for all atoms, followed by conversion to anisotropic displacement parameters for U and Pb atoms. The final cycles of refinement included the Pb site-occupancies, and resulted in agreement indices ($R_1$) ranging from 3.6% to 6.2%, calculated for observed ($|F_o|^2$) reflections (Table 3). The highest $R_1$ was obtained for the synthetic crystal, which was substantially smaller than the others (Table 2). Unit-cell parameters and Pb occupancies derived from site-scattering refinement for each crystal are given in Table 3, together with the corresponding final agreement index ($R_1$). Final atomic coordinates are in Table 4, and selected interatomic distances are in Table 5.

**Electron-microprobe analyses**

With the exception of the synthetic crystal, the fourmarierite crystals selected for our X-ray-diffraction investigation were mounted at the center of hollow aluminum tubes using epoxy. Each crystal was hand-polished and coated with carbon. The elemental concentrations were analyzed using an electron microprobe (JEOL Superprobe 733 at the Canadian Museum of Nature) equipped with four wavelength-dispersion spectrometers and operated at 15 kV. A beam current of 20 nA and a beam diameter of 20 μm were used. Synthetic UO$_2$ and crocoite were used as standards for U and Pb, respectively. We sought Na, K, Ca, Sr, Ba, and Th, but no significant concentrations were found. The analyses were done with 25 s spent counting for each element, using the Tracer Northern Program 5500 and 5600 software, and the ZAF correction using the PAP correction program (C. Davidson, CSIRO, pers. commun.). Several different points were analyzed for each crystal, depending on the size of the crystal examined. No significant chemical zoning was observed in electron-back-scatter images of any crystals of fourmarierite examined. The results of the chemical analyses are given in Table 6, with the proportion of H$_2$O assumed on the basis of stoichiometry.

**RESULTS**

**Cation polyhedra**

The structure of fourmarierite contains four symmetrically unique U sites. All of the U$^{6+}$ cations are strongly bonded to two atoms of O (designated O$_{\text{Ur}}$),...
resulting in approximately linear uranyl ions with $\langle U-O_{\text{Ur}} \rangle \approx 1.8$ Å, as is typically observed in the structure of U$^{6+}$-bearing phases (Burns et al. 1997a). Each U$^{6+}$ cation is coordinated by five additional anions, forming $\text{UO}_2^{2+}$ uranyl ion, $\text{f}$: unspecified ligand]. For the natural crystals, $\langle U(1)-5f_{\text{eq}} \rangle$, $\langle U(2)-5f_{\text{eq}} \rangle$, $\langle U(3)-5f_{\text{eq}} \rangle$, and $\langle U(4)-5f_{\text{eq}} \rangle$ range from 2.35 to 2.38, 2.35 to 2.37, 2.37 to 2.40, and 2.35 to 2.37 Å, respectively. In the case of the synthetic crystal, which has much less Pb than the natural crystals, the average bond-lengths of the uranyl polyhedra are 2.36, 2.43, 2.40 and 2.38 Å for the U(1), U(2), U(3) and U(4) polyhedra, respectively.

There are two distinct Pb sites, and eight symmetrically distinct H$_2$O sites in the interlayer of the structure of fourmarierite. In the natural crystals, Pb(1) is coordinated by four O$_{\text{Ur}}$ atoms and three H$_2$O groups, with $\langle \text{Pb(1)}-f_{\text{eq}} \rangle$ ranging from 2.68 to 2.72 Å. Pb(2)
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is coordinated by six $O_{ur}$ atoms and three $H_2O$ groups, with $<Pb(2)-O>$ ranging from 2.72 to 2.75 Å (Table 5, Fig. 1). The coordination polyhedron about the Pb(1) site in the synthetic crystal is similar to those in the natural crystals, although it is only 13% occupied, with $<Pb(1)-f>$ bond length of 2.62 Å. The coordination polyhedron about Pb(2) in the synthetic crystal, which is 87% occupied, is not identical to those in the natural crystals. The $H_2O(19)$ and $H_2O(21)$ groups assume different positions in the interlayer of the synthetic crystal, resulting in the addition of $H_2O(19)$ and removal of $H_2O(21)$ from the Pb(2) coordination polyhedron.

**Structural connectivity**

The $Ur_f$ pentagonal bipyramids share edges and vertices, forming sheets oriented parallel to (100) (Fig. 2). Symmetrically equivalent sheets are located at $c \approx 0.25$ and $c \approx 0.75$ (Fig. 3). In the interlayer, the $Pb(1)f_f$ and $Pb(2)f_f$ polyhedra share an edge, forming dimers (Fig. 4). Of the eight $H_2O$ groups, four [$O(16), O(17), O(18)$ and $O(19)$] are bonded to $Pb$ in the interlayer, and four [$O(20), O(21), O(22)$ and $O(23)$] are held in the structure by hydrogen bonds only (Fig. 3). Adjacent sheets are connected through bonds to the $Pb^{2+}$ cations and the $H_2O$ groups (Fig. 3).
Variation of Pb content and unit-cell parameters

The structural studies of 13 crystals of fourmarierite demonstrate significant variations in the Pb content, with total Pb ranging from 0.86 to 1.02 apfu (atoms per formula unit) for the natural crystals, and being 0.50 apfu for the synthetic crystal (Table 3). The Pb(1) site contains from ~70 to 100% Pb in the natural crystals, but only 13% Pb in the synthetic crystal. The Pb(2) site is close to fully occupied in the natural crystals, and contains 87% Pb in the synthetic crystal.

The variation of the unit-cell parameters as a function of Pb content is shown in Figure 5. The unit-cell volume (Fig. 5d) decreases with increasing Pb content, presumably because of increased bond-strengths between the structural sheets and Pb in the interlayer.
The a dimension, which is parallel to the sheet of polyhedra, is relatively weakly affected by changes in the Pb content (Fig. 5a). Both the b and c dimensions show a large systematic decrease with increasing Pb (Figs. 5b, c). On the basis of the trends in Figure 5, the cell parameters given by Piret (1985) for fourmarierite are consistent with both Pb sites being close to 100% occupied by Pb, in agreement with his interpretation.

It is possible that a complete range of compositions between schoepite and fourmarierite occurs, although the fourmarierite crystals examined in the current study do not approach the full range. Such a series would have to involve a phase transition from space group $Bb\overline{2}1m$ (fourmarierite) to $P2_1ca$ (schoepite). Linear regression of the unit-cell data for the natural crystals presented in Figure 5 indicates that the b and c parameters in the absence of Pb would be 16.86(8) and 14.67(4) Å (standard errors in brackets), which are similar to the values 16.813(5) and 14.731(4) Å obtained for schoepite by Finch et al. (1996). Inclusion of the data for the synthetic crystal results in intercepts of 16.86(3) and 14.99(4) Å for the b and c parameters, respectively, suggesting that a discontinuity occurs in the c parameter as a function of Pb content (Fig. 5c). An estimate of the Pb

![Coordination polyhedra about the Pb(1) and Pb(2) sites in the natural crystals, which link to form a dimer of composition Pb$_2$O$_{14}$.](image)
content of crystals of compositions intermediate between schoepite and fourmarierite may be obtained from the unit-cell parameters.

The total Pb contents of the natural crystals as derived by structure refinement and by analysis using an electron microprobe are compared in Figure 6. The microprobe analyses gave total Pb contents consistently ~8 mol.% lower than the structure refinements. The reason for this discrepancy is unknown, but may originate from factors such as crystal instability under the electron beam and the difficulty of identifying suitable standards for the analysis of these complex minerals.

**Charge-balance mechanism**

Variation of the Pb content of the interlayers of the structure of fourmarierite must be charge-balanced locally in the structure. Several possibilities exist to provide for the charge balance. Charged species other than Pb$^{2+}$ in the interlayer could be involved, but careful inspection of difference-Fourier maps did not reveal additional sites in the interlayer, nor did the electron-microprobe analyses indicate constituents other than U and Pb. However, because of the dominance of U and Pb in the scattering of X-rays, small impurities of additional species in the interlayer cannot be ruled out. Another possibility is the substitution (OH)$^-$ → H$_2$O at any of the sites from O(16) to O(23) in the interlayer. On the basis of the bond-valences incident upon these H$_2$O sites, substitution of (OH)$^-$ would lead to serious underbonding for the (OH)$^-$ group, thus this charge-balancing mechanism seems unlikely.

The substitution (OH)$^-\rightarrow$O$^{2-}$ within the sheet of uranyl polyhedra, at the equatorial vertices of the $Urdf_5$ pentagonal bipyramids, could provide a local charge-balancing mechanism. Such a substitution would be reflected by the bond-valence sums at the anion sites within the sheets; replacement of (OH)$^-$ by O$^{2-}$ should result in a shortening of the U–O bond lengths associated with the site. We examined the bond-valence sums at the anion sites for each crystal and found a relationship between the total Pb content of the crystal and the bond-valence sums at the O(12) and O(15) sites (Fig. 7), both of which are occupied mainly by (OH)$^-$. Note that neither of these sites is bonded to Pb, so that observed changes in the bond-valence sums may only be attributable to changes in U–O bond lengths. Other anion sites do not exhibit systematic changes in bond-valence sums with Pb content.

**Fig. 2.** Sheet of uranyl polyhedra in the structure of fourmarierite projected onto (001). The $Urdf_5$ pentagonal bipyramids are shown shaded with crosses.
Fig. 3. The structure of fourmarierite projected onto (100). Large and small circles in the interlayer represent Pb$^{2+}$ cations and H$_2$O groups, respectively. Pb–bonds are shown by lines.

Fig. 4. Interlayer constituents projected onto the sheet of uranyl polyhedra at $z \approx 0.25$. Large solid black circles represent Pb cations at $z \approx 0.5$, small solid black circles represent H$_2$O groups bonded to Pb at $z \approx 0.5$, shaded circles represent H-bonded H$_2$O groups at $z \approx 0.5$ in the interlayer, and open circles represent O$_{Ur}$ bonded to Pb at $z \approx 0.75$. 
In the case of the O(15) site, the sum of bond valences [attributable to the U–O(15) bonds] ranges from 1.45 to 1.51 \( v_u \) (valence units) in the natural crystals, and is 1.24 \( v_u \) in the synthetic crystal. The sum increases steadily with increasing Pb content in the crystals (Fig. 7b). This increase indicates that the U–O(15) bonds get stronger with increasing Pb content, consistent with an increase in the amount of \( O^2^- \leftrightarrow (OH)^- \) substitution at the O(15) site. The situation is similar for the O(12) site (Fig. 7a). There, the bond-valence sums incident upon the anion increase from 1.34 in the synthetic crystal to ~1.65 in the natural crystals. These trends demonstrate that \( O^2^- \leftrightarrow (OH)^- \) substitution within the sheets of uranyl polyhedra is the mechanism that provides charge balance as Pb varies in the interlayer.

**Formula of fourmarierite**

The structure determinations indicate that there are 32 U\(^{6+}\) cations in the unit cell. Most O atoms can be designated as \( O^2^- \), (OH\(^-\) or \( H_2O \) groups on the basis of the bond-valence analysis (Table 7), except O(12) and O(15), which have bond-valence sums that range considerably in the crystals studied. The variability in these sums (Fig. 7) reflects the substitution of \( O^2^- \leftrightarrow (OH)^- \) at these two sites. If \( y \) represents the total number of \( O^2^- \) in sites O(12) and O(15), then there is \( (2-y) \) (OH\(^-\) in these sites. All O\(^2^-\) and (OH\(^-\)) are located in the sheets of uranyl polyhedra, which therefore have the

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**Fig. 5.** Unit-cell parameters versus total Pb (apfu) obtained from crystal-structure refinement of thirteen single crystals of fourmarierite. The lines were obtained from linear regression of the data. In (c), the broken line was derived using data from the natural crystals only.

**Fig. 6.** Comparison of the proportion of Pb (apfu) determined by X-ray diffraction and electron-microprobe analysis. The broken line represents an ideal, iso-lead (apfu) segment from the two experimental methods.
composition \[ (UO_2)_{4-y}O_2^{2+} (OH)_{5-y}^{-} \]. In order for the structure to be charge-balanced, there must be \((1 + y)\) \(\text{Pb}\) \(\text{apfu}\). The formula of fourmarierite may be written as \(\text{Pb}_{1-x} [(UO_2)_{4-x}O_3^{2-} (OH)_{4+x}] (H_2O)_4 \), \(Z = 8\). Where \(x = 0\), this formula is identical to that given by Piret (1985). Where \(x = 1\), the composition of the schoepite sheet results (Finch et al. 1996).

**Interlayer of fourmarierite**

For a given sheet-topology, the distribution of interlayer constituents is closely related to the type of interlayer species and the arrangement of the sheets of uranyl polyhedra on either side. The hydrogen-bond network in the interlayer of schoepite (Finch et al. 1996) may be considered to be interrupted by \(\text{Pb}^{2+}\) cations in the interlayer of fourmarierite, resulting in isolated \(\text{Pb}\) and \(\text{H}_2\text{O}\) clusters (Fig. 4). In the structure of fourmarierite, \(\text{U}\) sites of adjacent sheets are aligned, and if viewed perpendicular to the sheet, both of the \(\text{Pb}(1)\) and \(\text{Pb}(2)\) sites reside in triangular prisms defined by \(\text{O}_{\text{Ur}}\) atoms (Fig. 4). Similar configurations are observed for interlayer \(\text{K}\) cations between the \(\alpha-\text{U}_3\text{O}_8\)-type sheets in the structure of compreignacite (Burns 1998b), for interlayer \(\text{Pb}\) between the \(\alpha-\text{U}_3\text{O}_8\)-type sheets in the structure of masuyite (Burns & Hanchar 1999), and for interlayer \(\text{Ba}\) between the \(\alpha-\text{U}_3\text{O}_8\)-type sheets in the structure of protasite (Pagoaga et al. 1987).

Because of the strong bonds within the uranyl ions, the \(\text{O}_{\text{Ur}}\) atoms receive \(~1.7\) \(\text{vu}\) from the central \(\text{U}^{6+}\) cation. An important role of the interlayer constituents is to contribute valence to \(\text{O}_{\text{Ur}}\) atoms. There is a complementary relationship between the distribution of interlayer cations and hydrogen-bonded \(\text{H}_2\text{O}\) groups; hydrogen-bonded \(\text{H}_2\text{O}\) groups are usually present in the regions where there are no interlayer cations and provide hydrogen bonds to \(\text{O}_{\text{Ur}}\) atoms. Reduction of the amount of \(\text{Pb}\) in the \(\text{Pb}(1)\) and \(\text{Pb}(2)\) sites in fourmarierite impacts upon the sums of bond valences incident upon some of the \(\text{O}_{\text{Ur}}\) atoms, as these are bonded to \(\text{Pb}^{2+}\). It is possible that partial replacement of \(\text{Pb}^{2+}\) by \(\text{H}_2\text{O}\) groups occurs, thus providing for hydrogen bonds that may be accepted by the \(\text{O}_{\text{Ur}}\) atoms.

**DISCUSSION**

The demonstration that \((\text{OH})^{-} \rightarrow \text{O}^{2-}\) substitution within the sheets of uranyl polyhedra is the charge-balancing mechanism linked to \(\text{Pb}\) variation in the structure of fourmarierite is significant because this is the first demonstration of anion substitution within a sheet of uranyl polyhedra. This discovery has important structural implications because it permits a range of interlayer

\[
\begin{array}{ccccccc}
\text{O(1)} & 1.61 \\
\text{O(2)} & 1.59 \\
\text{O(3)} & 1.59 \\
\text{O(4)} & 1.59 \\
\text{O(5)} & 1.61 \\
\text{O(6)} & 1.67 \\
\text{O(7)} & 1.61 \\
\text{O(8)} & 1.50 \\
\text{O(9)} & 0.42 \\
\text{O(10)} & 0.75 \\
\text{O(11)} & 0.48 \\
\text{O(12)} & 0.48 \\
\text{O(13)} & 0.67 \\
\text{O(14)} & 0.65 \\
\text{O(15)} & 0.41 \\
\text{O(16)} & 0.14 \\
\text{O(17)} & 0.14 \\
\text{O(18)} & 0.17 \\
\text{O(19)} & 0.14 \\
\text{O(20)} & 0 \\
\text{O(21)} & 0 \\
\text{O(22)} & 0 \\
\text{O(23)} & 0 \\
\hline
\text{Total} & 6.00 & 5.94 & 5.95 & 5.77 & 1.57 & 1.80
\end{array}
\]

*Bond-valence parameters for \(\text{U}^{7+}\) from Burns et al. (1997a) and for \(\text{Pb}\) from Beese & O'Keeffe (1991).*
contents without changing the topology of the sheet. These findings also indicate that there may be a link between the total charge on the fourmarierite sheet (and therefore the interlayer composition) and the pH of the solution that was in contact with the growing crystals. The Pb content of the interlayer, as well as the amount of hydrogen in the sheet of uranyl polyhedra, may prove to be a sensitive indicator of the geochemical conditions at the time of crystal growth.

The structure of fourmarierite may be obtained from that of schoepite by replacing selected (OH)\(^-\) by O\(^2-\) in the sheet, insertion of Pb into the interlayer, and removal of specific H\(_2\)O groups. Foord \textit{et al.} (1997) reported that mineral “A” may have the schoepite-type structure, but an appreciable amount of Pb was indicated by chemical analysis. Mineral “A” was originally described by Frondel (1956), and it probably formed as an alteration product in the early stages of corrosion of uraninite. Further studies are needed to determine whether mineral “A” may be Pb-bearing schoepite, Pb-poor fourmarierite such as the synthetic crystal obtained in the current study, or a distinct structure.

The finding that (OH)\(^-\) \rightleftharpoons O\(^2-\) substitution occurs in sheets of uranyl polyhedra based upon the fourmarierite anion-topology may be important for understanding the behavior of radionuclides during the alteration of nuclear waste. Buck \textit{et al.} (1998) observed that dehydrated schoepite that had formed as an alteration product of spent fuel contains approximately 550 ppm Np, which is consistent with earlier predictions that uranyl phases can incorporate transuranic elements into their structures (Burns \textit{et al.} 1997b). However, Np incorporated into dehydrated schoepite is most likely pentavalent, and a charge-balance mechanism is required for Np\(^{5+}\) to substitute for U\(^{6+}\). In light of our findings, a possible charge-balance mechanism is (OH)\(^-\) \rightleftharpoons O\(^2-\) substitution in the sheets of uranyl polyhedra. We suggest that a considerable amount of Np could be incorporated into dehydrated schoepite without destabilizing the structure.

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References


CRYSTAL-CHEMICAL VARIABILITY IN FOURMARIERITE


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