# DTA, TG, AND XRD STUDIES OF STURMANITE AND ETTRINGITE

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### ABSTRACT

Thermal analyses (DTA and TG) were carried out on sturmanite and ettringite from the Kalahari Manganese Field, South Africa. The TG trace for sturmanite, approximately  $Ca_6(Fe^{3+}_{1.5}Al_{0.3}Mn^{2+}_{0.2})_{\Sigma 2.0}\{[B(OH)_4]_{1.2}(SO_4)_{2.3}\}_{\Sigma 3.5}(OH)_{12}\bullet 25H_2O$ , indicates that  $H_2O(g)$  is lost at about 135°C,  $SO_3(g)$  is lost at about 1349°C, and the residue melts at about 1154°C. In sturmanite, a polymorphic transition occurs at about 627°C. For ettringite, approximately  $Ca_6Al_2(SO_4)_3(OH)_{12}\bullet 26H_2O$ , the  $H_2O(g)$  and  $SO_3(g)$  are lost at about 149 and 753°C, respectively, and the residue melts at about 1176°C. Ettringite melts after the liberation of both  $H_2O(g)$  and  $SO_3(g)$ , whereas sturmanite melts after the liberation of  $H_2O(g)$ . The loss of  $SO_3(g)$  occurs at a considerably lower temperature in ettringite than in sturmanite. Using powder X-ray diffraction, the unit-cell parameters for sturmanite are *a* 11.227(1), *c* 21.846(3) Å, *V* 2355.2(8) Å<sup>3</sup> for the hexagonal supercell, and *a* 11.147(3), *c* 10.918(5) Å, *V* 1174.9(9) Å<sup>3</sup> for the subcell. The unit-cell parameters for ettringite are *a* 11.223(1), *c* 21.474(2) Å, *V* 2342.2(5) Å<sup>3</sup> for the hexagonal supercell, and *a* 11.229(1), *c* 10.732(2) Å, *V* 1171.9(3) Å<sup>3</sup> for the subcell. The volume of sturmanite is only slightly larger than that of ettringite.

Keywords: sturmanite, ettringite, differential thermal analysis, thermogravimetric analysis, X-ray diffraction.

### Sommaire

Nous avons effectué des analyses thermiques différentielles et thermogravimétriques de la sturmanite et de l'ettringite provenant du champ minéralisé en manganèse de Kalahari, en Afrique du Sud. Le tracé thermogravimétrique de la sturmanite, dont la composition est proche de  $Ca_6(Fe^{3+}_{1.5}Al_{0.3}Mn^{2+}_{0.2})_{\Sigma 2.0}$ { $[B(OH)_4]_{1.2}(SO_4)_{2.3}\}_{\Sigma 3.5}(OH)_{12}$ •25H<sub>2</sub>O, montre qu'il y a perte de H<sub>2</sub>O(g) à environ 135°C, et de SO<sub>3</sub>(g) à environ 1349°C, et que le résidu passe à l'état fondu à environ 1154°C. La sturmanite subit une transition polymorphique à environ 627°C. Dans le cas de l'ettringite, dont la formule est proche de  $Ca_6Al_2(SO_4)_3$  (OH)<sub>12</sub>•26H<sub>2</sub>O, les fractions H<sub>2</sub>O(g) et SO<sub>3</sub>(g) sont libérées à environ 149 et 753°C, respectivement, et le résidu fond à environ 1176°C. L'ettringite fond après la libération de H<sub>2</sub>O(g) et SO<sub>3</sub>(g), tandis que la sturmanite fond après la libération de H<sub>2</sub>O(g) esul. La fraction SO<sub>3</sub>(g) est libérée à une température considérablement plus faible dans le cas de l'ettringite que pour la sturmanite. D'après les données de diffraction X obtenues sur poudre, les paramètres réticulaires de la sturmatite sous-maille. Par contre, les paramètres réticulaires de l'ettringite sont *a* 11.223(1), *c* 21.474(2) Å, *V* 2342.2(5) Å<sup>3</sup> pour la supermaille hexagonale, et *a* 11.223(1), *c* 21.474(2) Å, *V* 2342.2(5) Å<sup>3</sup> pour la supermaille hexagonale, et *a* l'ettringite.

(Traduit par la Rédaction)

Mots-clés: sturmanite, ettringite, analyse thermique différentielle, analyse thermogravimétrique, diffraction X.

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## INTRODUCTION

Ettringite, approximately Ca<sub>6</sub>Al<sub>2</sub>(SO<sub>4</sub>)<sub>3</sub>(OH)<sub>12</sub>• 26H<sub>2</sub>O, is an important industrial mineral because of its formation as a product of hydration in Portland and super-sulfated cement, and its use as satin white as a coating material for paper (Moore & Taylor 1970). An exact chemical composition cannot be obtained for ettringite because the H<sub>2</sub>O content is variable (McConnell & Murdoch 1962, Moore & Taylor 1970). Sturmanite, approximately Ca<sub>6</sub>(Fe<sup>3+1</sup> 5Al<sub>0</sub> 3Mn<sup>2+0</sup> 2)52.0 {[B(OH)4]12  $(SO_4)_{2,3}$   $\sum_{3,5} (OH)_{1,2} \cdot 25H_2O_{2,5}$  is a ferric iron, boron-containing analogue of ettringite; an exact formula for sturmanite is tentative because of the ambiguity in the number of S and B atoms (Peacor et al. 1983). Other minerals in the ettringite group are bentorite [Ca<sub>6</sub>Cr<sub>2</sub>  $(SO_4)_3(OH)_{12} \bullet 26H_2O, a 11.210(15), c 21.48(3) Å, space$ group P31c, and Z = 2: Gross 1980], jouravskite  $[Ca_6Mn_2\{(CO_3)_2(SO_4)_2\}_{\Sigma_4}(OH)_{12} \cdot 24H_2O, a 11.06, c$ 10.50 Å, space group  $P6_3/m$ , and Z = 1: Gaudefroy & Permingent 1965], and charlesite  $[Ca_6(Al,Si)_2\{(B(OH)_4)\}$  $(SO_4)_2$   $\{_{\Sigma_3}(OH)_{12} \bullet 26H_2O, a 11.16(1), c 21.21(2) Å,$ space group P31c, and Z = 2: Dunn *et al.* 1983].

Thermal studies have been done on ettringite, but not on sturmanite (e.g., Hall et al. 1996, Zhou & Glasser 2001, Wieczorek-Ciurowa et al. 2001, Shimada & Young 2001). However, most of these studies pertain to synthetic ettringite and were done in the low-temperature region, <200°C. The chemical compositions of sturmanite and ettringite indicate that H<sub>2</sub>O and SO<sub>3</sub> are important volatile constituents that may be liberated on heating. This study was carried out to determine what chemical constituents are liberated on heating, to measure the temperature where changes take place, and to compare the results for ettringite and sturmanite, using differential thermal analyses (DTA) and thermogravimetric (TG) analyses to about 1450°C. Unit-cell parameters at room temperature, before heating, were also determined using powder X-ray diffraction (XRD).

### **BACKGROUND INFORMATION**

The unit cell of ettringite is *a* 11.26, *c* 21.48 Å, Z = 2 for the supercell (Bannister *et al.* 1936). The space group for ettringite is *P*31*c*, and any apparent hexagonal symmetry was attributed to twinning or disorder (Moore & Taylor 1968, 1970, Courtois *et al.* 1968). Peacor *et al.* (1983) suggested that the space group for sturmanite is *P*31*c*, with unit-cell parameters *a* 11.16(3), *c* 21.79(9) Å for the supercell. They also suggested that the diffraction data for sturmanite are similar to those for the ettringite-group minerals in that sturmanite has a pronounced subcell having parameters A = a and C = c/2, *i.e.*, all reflections having l = 2n + 1 are very weak, and extinctions are present for reflections *hhl*, l = 2n + 1, which is consistent with a glide plane. Ettringite also has a similarly strong subcell to supercell relationship.

The crystal structure of ettringite was studied by several investigators (e.g., Bannister et al. 1936, Courtois et al. 1968, Moore & Taylor 1968, 1970). Taylor (1973) reviewed the crystal chemistry of the ettringite-group minerals. The ettringite structure consists of columns and channels that are parallel to the c axis (Fig. 1; Taylor 1973). The columns contain  $[Ca_6[Al(OH)_{12}]$ •  $24H_2O^{6+}$ , and the channels contain  $[(SO_4)_3 \cdot 2H_2O^{6-}]$ . per half unit cell. Each of the columns contains a chain of polyhedra, including one of Al and three of Ca. The  $[Al(OH)_6]^{3-}$  octahedra are linked together through three Ca<sup>2+</sup> ions, which are eightfold co-ordinated by four hydroxyl groups and four H<sub>2</sub>O molecules. The [Al(OH)<sub>6</sub>]<sup>3-</sup> octahedron is further co-ordinated by H<sub>2</sub>O molecules. The Ca<sup>2+</sup> polyhedra are trigonal prisms with the axis parallel to c. The Ca<sup>2+</sup> ion is eightfold coordinated by four H<sub>2</sub>O molecules (labeled A and B, Fig. 1) and four OH groups. The A  $H_2O$  molecule has nearly the same Z coordinate as the Ca<sup>2+</sup> ion. Each of these prisms shares two edges with adjacent [Al(OH)<sub>6</sub>]<sup>3-</sup> octahedra. Channels between the columns are occupied by H2O molecules and SO<sub>4</sub><sup>2-</sup>, CO<sub>3</sub><sup>2-</sup>, or B(OH)<sub>4</sub> groups. For example, the SO<sub>4</sub><sup>2-</sup> ions occur on axes along the lines  $(\frac{1}{3}, \frac{2}{3}, z)$  and  $(\frac{2}{3}, \frac{1}{3}, z)$ , in four positions. Statistically, three of these sites are occupied by three  $SO_4^{2-}$  groups, and the fourth site by two H2O molecules. A maximum of four anion groups can occupy these sites, with a corresponding reduction of the two H<sub>2</sub>O molecules to zero. If the total number of anion groups is at the maximum number of four, then the number of H<sub>2</sub>O molecules is reduced from 26 to 24 in the formula unit. The repeat distance along a column is C = c/2 = 10.7 Å, the prominent translation of the substructure, with the superstructure arising from ordering of anionic groups in the sites between the columns. The  $Ca^{2+}$  ions may be replaced by ions such as Pb<sup>2+</sup>, and the Al<sup>3+</sup> by Fe<sup>3+</sup>, Mn<sup>3+</sup>, Cr<sup>3+</sup>, etc. (Peacor et al. 1983, Wieczorek-Ciurowa et al. 2001).

Several thermal studies are available for synthetic ettringite, including some recent investigations (e.g., Hall et al. 1996, Zhou & Glasser 2001, Wieczorek-Ciurowa et al. 2001, Shimada & Young 2001). In a study of a synthetic ettringite using TG, XRD, and <sup>27</sup>Al NMR, Shimada & Young (2001) heated ettringite at various temperatures up to 200°C for a period of up to 7 h. The structure maintains some long-range order until the coordination number of Ca changes to 5 by dehydration of 12 H<sub>2</sub>O molecules from the channels and columns with heat treatment at 70°C. After 7 hours at 70°C, the short-range order is disrupted, and ettringite becomes XRD-amorphous. Thereafter, the rest of the H<sub>2</sub>O molecules in the columns and bridging OH groups in the Ca polyhedra are removed, and the framework of the columns is destroyed. This step is accompanied by changes in the coordination number of Al from 6 to 4 (Shimada & Young 2001).

### EXPERIMENTAL

The samples used in this study are from the Kalahari Manganese Field, South Africa; ettringite is from the N'chwaning II mine, and sturmanite is from the Wessels mine. The samples were coarsely crushed, and pure crystals were handpicked under a binocular microscope. The pure crystals were then crushed to a powder using an agate mortar and pestle. Portions of the powder were used for DTA, TG, XRD, and electron-microprobe analyses.

A weighed amount of powdered sample was placed into an  $Al_2O_3$  crucible for thermal analyses. For sturmanite, a fully computerized, Netzsch STA 409 EP/ 3/D Simultaneous TG–DTA equipment was used. For ettringite, a Shimadzu Thermal System 50 (TG 50 and DTA 50) was used because the first equipment was in need of repairs. Sturmanite was heated at a constant rate of 5°C/min in a static air environment. Ettringite was heated at a rate of 10°C/min. in a dynamic air environment where the flow rate of air was 60 mL/min. Thermal data were analyzed using software programs supplied with the instruments. A detailed experimental procedure is given in Hassan (1996).

XRD data were obtained using a fully computerized Siemens D5000 Diffractometer. The XRD data were obtained with the diffractometer operating in the vertical position and in the  $\theta$ - $\theta$  operating mode. We used Ni-filtered CuK $\alpha$  radiation in conjunction with a position-sensitive detector. Data were collected at room temperature for the 2 $\theta$  range of 6 to 110°. A continuous scan was used with a step size of 0.015° and step time of 20.0 s. The unit-cell parameters were obtained by least-squares refinement using the program WIN-MET-RIC. The zero-shift of the diffractometer was determined by maximizing  $F_N$  (a figure-of-merit for all reflections) in the refinement procedure (de Wolff 1968).

The sturmanite sample was used up for thermal analyses. However, for ettringite, we used the available sample for chemical analysis using the electron microprobe (EMP), and for additional XRD runs after heating the sample to 260°C, and then quenching the sample to room temperature. At 260°C, the thermal analyses indicated that all the H<sub>2</sub>O molecules were liberated.

The chemical analysis for ettringite was done using a Cameca Camebax electron microprobe using the operating program MBX (copyright by Carl Henderson, University of Michigan) and the correction was done using Cameca's PAP program. The analytical conditions were 15 kV and 9.2 nA beam current. Natural minerals were used as standards: microcline (Si $K\alpha$ , K $K\alpha$ ), albite (Na $K\alpha$ ), forsterite (Mg $K\alpha$ ), "apatite" (PK $\alpha$ ), anorthite (Al $K\alpha$ , Ca $K\alpha$ ), and gypsum (S $K\alpha$ ). The oxide weight



FIG. 1. The general features of the structure of ettringite. (a) Projection along [001] showing the polygons C that represents the columns of [Ca<sub>3</sub>Al(OH)<sub>6</sub>•12H<sub>2</sub>O]<sup>3+</sup>, and the triangles S that represents the SO<sub>4</sub><sup>2-</sup> ions and H<sub>2</sub>O molecules in the channels. (b) Part of a column projected on (110). The A and B circles represent H<sub>2</sub>O molecules, but those attached to the Ca ions lying in the central vertical line of the figure are omitted, as are all the H atoms (from Taylor 1973).

percentages resulting from the EMP analyses are given in Table 1. The H<sub>2</sub>O content for ettringite was obtained by subtraction. Ettringite damages quite easily in the EMP with loss of SO<sub>3</sub> and H<sub>2</sub>O, so a diffuse electron beam was used for the analysis. The results obtained for ettringite are comparable to those in the literature (Table 1).

## **RESULTS AND DISCUSSION**

### DTA and TG

Using about 76 mg powder, the TG and DTA curves, and their corresponding derivative curves (DTG and DDTA, respectively) were obtained for sturmanite (Fig. 2). The DTG and DDTA curves were obtained from the corresponding raw data using a narrow window for filtering the measured raw data. The differentiation was done by using a modified Golav-Savitzky algorithm of second order. The characteristic data for sturmanite obtained from these curves are summarized (Table 2). Four peaks are observed in the DTA curve; peaks 1, 2, and 3 are well defined in both the DTA and DDTA curves, but peak 4, although visually detectable, is less obvious, but is clearly seen in the TG and DTG traces (Fig. 2). Peaks 2 and 3 occur as discontinuities only in the DTA trace where there is no loss in weight, so they are attributed to polymorphic phase-transitions. Peak 3 is related to melting of the sturmanite residue because it occurs at a higher temperature compared to peak 2, which is related to a polymorphic transition. A brownish black "melt" was observed in the crucible after the experiment.

There are two well-defined DTG peaks 1 and 4 (Fig. 2a). The sharp peak 1 corresponds to a net loss in weight of 41.6% (Table 2). This weight loss is attributed to the loss of  $H_2O(g)$ . According to the chemical

TABLE 1. CHEMICAL COMPOSITION OF STURMANITE AND ETTRINGITE

	Sturmanite+	Ideal	Ettringite;	[deal	ЕМР
CaO wi %	aOwt% 25.67		26.6	26.81	26 70
Na <sub>2</sub> O		40.00	40.0		0.01
Fe-O,	9.11	12,09			
MnO	1.16				
P.O.					0.20
Al <sub>i</sub> O,	1.08		7,0	8 12	9.04
sio.					0.72
SO.	14.02	12.13	18.8	19 14	18.67
B.Ó,	3.18	5.27			
H O	45.83	45.02	46.3	45,93	44.66*
others			0.8		
Total	100.00	100	99.5	100	100

\* Composition reported by Peacor et al. (1983), which leads to the formula  $Ca_{4}(Fe^{3} {}_{1.5}Al_{0.3}Mn^{2} {}_{0.2})_{2270} \{ |B(OH)_{4}]_{1.2}(SO_{4})_{2.3} \}_{\Sigma 3.5} (OH)_{12} {}^{+}25H_{2}O_{10}$ Composition reported by McConnell & Murdoch (1962).

\* ILO obtained by subtraction. EMP: electron-microprobe data

data on sturmanite, H<sub>2</sub>O constitutes 46.7 wt.% (Table 1), which is more than the weight loss for peak 1. Peak 4 corresponds to a weight loss of about 12.3%, which is comparable to the 14.2 wt.% of SO<sub>3</sub> shown by the chemical analyses (Table 1). Peak 4 is thus attributed to the loss of  $SO_3(g)$ . The loss of weight is more gradual for peak 4 than for peak 1.

Mn is held to be in the 2+ oxidation state in sturmanite (Peacor et al. 1983). No oxidation of Mn<sup>2+</sup> was detected in the DTA and TG analyses of sturmanite, as there was no weight gain. Mn<sup>2+</sup> should have oxidized to Mn<sup>3+</sup> starting at about 700°C, as was observed in the DTA and TG analyses of helvite and danalite (Antao & Hassan 2002). The amount of Mn<sup>2+</sup> present may be too small to be detected by the DTA-TG technique. As an alternative, the Mn may be already trivalent, thus was not oxidized further. Moreover, the presence of Fe<sup>3+</sup>and Al<sup>3+</sup> suggests that Mn may be present as Mn<sup>3+</sup> (Peacor et al. 1983); however, they tentatively assigned the Mn to Mn<sup>2+</sup>. In helvite and danalite, the oxidized Mn<sup>3+</sup> cation undergoes further oxidation to Mn<sup>4+</sup> from about 1300°C (Antao & Hassan 2002), which was also not observed for sturmanite. Jouravskite, an isostructural phase, contains Mn<sup>4+</sup>, and sturmanite could contain Mn in the tetravalent state as well (Peacor et al. 1983). In

TABLE 2. DATA FROM THE THERMAL ANALYSIS OF STURMANITE AND ETTRINGITE

		Sturmanite			Etrringite				
	ΤG	DTG	DTA	DDTA	TG	DTC	i DT/	A DDTA	Changes
				Pea					
Onset-T (°C) Peak-T (°C) End-T (°C) % Wt. Loss Enthalpy (J/g)	102  167 41.6	89 135 157	92 139 181 1660	125 151	68 249 40,4	64 149 201	93 145 196 641	128 169	Loss of H <sub>i</sub> O (and OH)
				Per	ak Z <sup>:</sup>				
Onset-T (°C) Peak-T(°C) End-T (°C)			-627	 635					Poly- morphic transition
				Pea	nik 3 <sup>‡</sup>				
Onset-T (°C) Peak-T(°C) End-T (°C)			1154	1128 1156			1122 1176 1209	1164 1188	Melting of residue
				Pe	ık 4 <sup>5</sup>				
Onset-T (°C) Peak-T(°C) End-T (°C) % Wt. Loss	1274 1422 12 3	1266 1349 1393			658  968 15.4	699 753 848	708 765	732	Loss of SO <sub>5</sub> (g)

\* exothermic peaks (peaks 1, 2, and 3); \* peak 4 is exothermic for sturmanite but endothermic for ettringite. Total % wt. loss for peaks 1 and 4 = -53.9 in sturmanite and -55 8 in entringite. Total wt. loss from start to finish is -56.9% in sturmanite and -59.7% in ettringite.



FIG. 2. Sturmanite: (a) TG and DTG curves, and (b) DTA and DDTA curves. Corresponding peaks at a particular temperature are assigned the same number and are labeled on the DTA and DTG curves in this figure and Figure 3.

addition, Cr-substituted ettringite has been synthesized by Wieczorek-Ciurowa *et al.* (2001). Therefore, the present results are most consistent with the hypothesis that the Mn is trivalent.

Thermal curves were obtained for ettringite by using about 10 mg of powder (Fig. 3). Characteristic data for ettringite obtained from these curves are given in Table 2. Three peaks are observed in the DTA trace (peaks 1, 3, and 4; Fig. 3b) and are labeled to correspond to those in sturmanite. Peaks 1 and 3 are well defined in both the DTA and DDTA curves, but peak 4 is clearly seen in the TG and DTG traces (Fig. 3). Peak 3 occurs in the DTA trace, and at that temperature, there is no loss in weight, so peak 3 is attributed to the melting of the residue of ettringite.

Peaks 1 and 4 are well defined in the DTG curve for ettringite (Fig. 3a). As in sturmanite, peaks 1 and 4 correspond to the loss of  $H_2O(g)$  and  $SO_3(g)$ , respectively. The TG curve gives a loss of 40.4% over peak 1 and 15.4% over peak 4 (Table 2). According to the chemical composition of ettringite,  $H_2O$  constitutes about 44.7% and SO<sub>3</sub> constitutes 18.7% (Table 1). The weight losses obtained from the TG curve are less than those expected from the chemical composition. The rate of loss is slower over peak 4 than over peak 1 (Fig. 3a).

In both ettringite and sturmanite, the observed TG weight loss for  $H_2O(g)$  and  $SO_3(g)$  are less than those expected from the chemical compositions of the two

minerals. With regards to H<sub>2</sub>O, these differences reflect incomplete liberation of H<sub>2</sub>O, as was observed for synthetic ettringite by Shimada & Young (2001). They showed that when synthetic ettringite is heated to 120°C, the number of H<sub>2</sub>O molecules remaining in the chemical formula is 6.6 with respect to 30.9 H<sub>2</sub>O molecules at room temperature. The ideal formula of ettringite is 3CaO•Al<sub>2</sub>O<sub>3</sub>•3CaSO<sub>4</sub>•32H<sub>2</sub>O, which leads to 45.93 wt.% H<sub>2</sub>O (Table 1). The TG-established loss of H<sub>2</sub>O from ettringite was 41.40 wt.% (Table 2), which corresponds to 26.61 molecules of H<sub>2</sub>O, so 5.39 molecules of H<sub>2</sub>O remained in the sample. These results indicate that natural and synthetic ettringite do behave a little differently. A similar analysis for sturmanite, using the ideal empirical formula, 4CaO•Fe<sub>2</sub>O<sub>3</sub>•<sup>1</sup>/<sub>2</sub>B<sub>2</sub>O<sub>3</sub>•2CaSO<sub>4</sub>• 33H<sub>2</sub>O, and the TG-established wt.% loss of H<sub>2</sub>O (Tables 1, 2), indicates that 4.3 molecules of H<sub>2</sub>O remained in the sample.

A batch of ettringite powder was selected for XRD quenching experiments. An XRD trace of the sample, taken at room temperature, indicated that the sample contains a small amount of calcite as an impurity phase (Fig. 4a). This sample was then heated in an oven from room temperature to 260°C at a rate of 5°C/min. At 260°C, the thermal analyses indicated that ettringite is dehydrated. The sample was held at 260°C for one hour and then cooled to room temperature. An XRD trace of



FIG. 3. Ettringite: (a) TG and DTG curves, and (b) DTA and DDTA curves.

the sample contained peaks from ettringite and the impurity calcite phase (Fig. 4b). The same batch was again heated from room temperature to  $260^{\circ}$ C at a rate of  $5^{\circ}$ C/ min, and the sample was held at  $260^{\circ}$ C for 7 h and then cooled to room temperature. The XRD trace still showed peaks that are indicative of ettringite and the minor calcite phase (Fig. 4c). These results are in contrast to those obtained for synthetic ettringite by Shimada & Young (2001). When synthetic ettringite is heated to  $70^{\circ}$ C and held for 7 h, they showed that synthetic ettringite becomes XRD-amorphous, and their XRD trace contained a minor amount of calcite. These conflicting results indicate that synthetic and natural ettringite do behave differently.

In general, both ettringite and sturmanite samples undergo two main well-separated weight-loss stages. The loss of  $H_2O(g)$  begins at about 68°C in ettringite and at about 102°C in sturmanite. The majority of H<sub>2</sub>O molecules escape in a single step, but a significant amount of H<sub>2</sub>O remained in the sample. The loss of SO<sub>3</sub>(g) in ettringite begins at about 658°C, but in sturmanite this loss occurs at a considerably higher temperature, about 1274°C. The SO<sub>3</sub>(g) escapes in a single step in both samples. In lazurite, the loss of  $SO_3(g)$  begins at about 1264°C and continues beyond about 1420°C in several steps (Hassan 2000). The temperature at which  $SO_3(g)$  is liberated in lazurite is comparable to that in sturmanite. The ettringite sample melts after the liberation of both  $H_2O(g)$  and  $SO_3(g)$ , whereas the sturmanite sample melts after the liberation of  $H_2O(g)$ . The residue of ettringite melts at 1176°C. For sturmanite, the residue melts at 1154°C. The two weight-loss stages begin earlier in ettringite than in sturmanite. These results indicate that the bonds are weaker in ettringite than in sturmanite, which facilitates the escape of volatiles at lower temperatures in ettringite.

## XRD data

The hexagonal unit-cell obtained in this study for sturmanite is *a* 11.157(1), *c* 21.846(3) Å, *V* 2355.2(8) Å<sup>3</sup>, with a tolerance in  $|2\theta| \le 0.018^{\circ}$  for 31 XRD peaks refined in the space group *P*31*c*. The XRD results are similar to those of Peacor *et al.* (1983): *a* 11.16(3), *c* 21.79(9) Å); however, not all their X-ray-diffraction peaks were indexed. All our XRD peaks are indexed, and in particular, a strong 223 peak was observed on the shoulder of the 216 peak; this 223 peak was not observed in the earlier study. The principal X-ray-diffraction peaks are slightly different from those obtained by Peacor *et al.* (1983); in this study they are: [*d*<sub>obs</sub> in Å(I)(*hkl*)]: 9.661(100)(100), 5.6 (79)(110), 2.774(72) (304), 2.579(52)(216), and 3.904(48)(114).

Peacor *et al.* (1983) noted that sturmanite has a very pronounced subcell with parameters A = a and C = c/2. We have refined the subcell parameters in the space group *P*31*c* and obtained *a* 11.147(3), *c* 10.918(5) Å, *V* 1174.9(9) Å<sup>3</sup>, with refinement statistics better than those for the supercell.



FIG. 4. Ettringite XRD traces: (a) room temperature, (b) heated to 260°C for 1 h, and (c) heated to 260°C for 7 h. The Miller indices of three of the ettringite peaks are labeled. The main peak of calcite is indicated.

The parameters of the hexagonal subcell of ettringite obtained in this study are *a* 11.229(1), *c* 10.732(2) Å, *V* 1171.9(3) Å<sup>3</sup>, with a tolerance in  $|2\theta| \le 0.025^{\circ}$  for 48 XRD peaks refined in space group *P*31*c*. The supercell parameters are *a* 11.223(1), *c* 21.474(2) Å, *V* 2342.2(5) Å<sup>3</sup>, refined in space group *P*6<sub>3</sub>/*mmc*. Comparable values in the literature are *a* 11.26, *c* 21.48 Å (*e.g.*, Bannister *et al.* 1936, Moore & Taylor 1970). The unit-cell volume of sturmanite is slightly larger than that of ettringite. However, volatiles escape from ettringite more easily than from sturmanite.

#### ACKNOWLEDGEMENTS

We thank Nic Beukes and Pieter De Bruyn for providing the ettringite and sturmanite samples, respectively, and D.H. Lindsley for his help with the electron-microprobe analysis, which was financially supported by a NSF grant to J.B. Parise, EAR–0125094. We thank the two anonymous reviewers, Associate Editor M.E. Gunter, and R.F. Martin for useful comments.

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- Received January 5, 2002, revised manuscript accepted July 22, 2002.