MOLYBDOMENITE FROM THE RANWICK URANIUM MINE, MONTREAL RIVER AREA, ONTARIO

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Abstract

Molybdomenite has been found at the Ranwick Uranium Mine, where it occurs as small white to colourless blades closely associated with clausthalite. Molybdomenite has been synthesized and its chemical formula is established as PbSeO₃. The mineral is monoclinic with space group $P2_1$ or $P2_1/m$. Unit cell parameters are: a = 6.86, b = 5.48, c = 4.50 (all in Å), $\beta = 112^{\circ}45', Z = 2$.

INTRODUCTION

Historical review

The original description of molybdomenite given by Bertrand (1882) is quite sketchy and does not completely define the species. This was due, no doubt, to insufficient material and primitive mineralogical techniques. Goni & Guillemin (1953) restudied various selenites and selenates including molybdomenite, and published x-ray powder data and partial optical data for Bertrand's original specimens. They also proved, by qualitative tests, that molybdomenite is a lead selenite. The amount of material was too small for a quantitative analysis. Other than suggesting that molybdomenite was orthorhombic, they presented no crystallographic data.

Molybdomenite has been found at Cerro de Cocheuta, Mendoza, Argentina (the type locality); at the Trogtal, Harz, Germany (Goni & Guillemin, 1953); and at Pacajake, Bolivia, (Ahlfeld & Reyes, 1955). It may also be present in one of the deposits of the Magadanskaya oblast of the U.S.S.R. (Babkin, 1958) as reported by Sindeeva (1964). Of these reports, only the one by Goni & Guillemin contains diagnostic data for molybdomenite.

This paper records a new occurrence of molybdomenite and characterizes the species somewhat more fully than before.

During the fall of 1961, Dr. W. M. Tovell, Curator of Geology at the Royal Ontario Museum, took a collecting trip along the north shore of Lake Superior. One of the numerous mines which he visited was the Ranwick Uranium Mine, now operated as a tourist attraction by Mr. R. R. Ranson of Sault Ste Marie, Michigan. Mr. Ranson kindly presented several specimens from the mine to Dr. Tovell, who subsequently turned them over to the writer. A preliminary examination of the specimens revealed the following main minerals: pitchblende; a galena-like mineral which was later identified as clausthalite; and calcite. In addition, a few white to colourless, adamantine crystals were observed in close association with the clausthalite. Thinking this material might be one of the rare lead selenites or selenates, the writer began a detailed study. Data from an x-ray powder pattern resembled the data given by Goni & Guillemin (1953) for molybdomenite except that the powder photograph of the Ranwick mineral contained many more lines.

It was apparent that additional material would be needed in order to continue the study. Mr. Ranson offered Dr. Tovell and the writer the run of his property, and we visited the mine in August, 1962 and 1963.

Location and geology

The Ranwick Uranium Mine is north of Sault Ste Marie (near Mile 72) on Highway 17 (Trans-Canada Highway). The mine is readily visible on the south side of the highway which here runs east-west. Access to the property is extremely easy and there are ample parking facilities.

The geology of the Montreal River Area has been studied in detail by Nuffield (1956). Most of the material in this section has been abstracted from Nuffield's report. Additional information has been supplied by Mr. Ranson. The area is underlain by Algoman (?) granite, "granite" gneiss, and minor syenite, which are cut by Lower Keweenawan diabase dikes. These dikes are numerous and may make up as much as 5 per cent of the rock. The dikes range in width from less than a foot to 600 feet. The longest dikes strike west to northwest and dip steeply to the north. Some of these have been traced for over two miles. Other dikes strike northeast and dip vertically or steeply to the southeast, while a few strike north.

The Ranwick Uranium Mine is situated on a northwest-striking, northdipping diabase dike. The dike, now known as the Ranson dike, is about two miles long and the average width is 70 feet. Pitchblende was discovered near the south contact of the dike by Mr. Ranson in November of 1948. Since then, much trenching has been done across the dike and an adit was driven 1,049 feet along the footwall. From the adit, two crosscuts were driven 70 feet across the dike. Five other crosscuts of only a few feet in length were also driven. Pitchblende occurs in calcitehematite veins in fractures near the diabase-granite contact. According to Mr. Ranson, most of the selenium-bearing samples came from a surface pit about 80 feet above the adit. The writer visited this pit with Mr. Ranson and, although the pit was partially filled with water, several pitchblende-selenide specimens were collected. Only one of these contained molybdomenite. Fortunately, Mr. Ranson had collected a larger number of specimens from the pit before it was flooded and he kindly turned these over to the writer.

Although Nuffield (1956) reports only pitchblende, calcite, and hematite from the Ranwick mine, later studies by Nuffield (personal communication) disclosed the following species: klockmannite, native selenium, and malachite. In addition to the foregoing minerals identified by Nuffield and the molybdomenite described in this paper, the writer has identified only clausthalite and a few secondary uranium minerals.

MINERALOGY OF MOLYBDOMENITE

General features and appearance

The Ranwick molybdomenite occurs as flat, bladed crystals up to 0.5 mm in length. The crystals are colourless to white and have a pearly to adamantine lustre. Molybdomenite is translucent to transparent. It is invariably associated with clausthalite, from which it is derived.

Crystallographic characteristics

No euhedral crystals were available, so the entire crystallographic study was carried out by means of *x*-rays.

X-ray data. A crystal was mounted on a Buerger precession camera and oriented with the blade axis parallel to the spindle axis. After several films had been taken, it became apparent that the mineral was not orthorhombic as deduced by Goni & Guillemin (1953), but rather, monoclinic. From numerous films taken at small angular intervals, provisional a^* and c^* axes were located normal to the spindle axis. Using these axes, 0-, 1-, and 2- level films were obtained. The crystal was then reoriented with the *b*-axis as the precession axis and 0-, 1-, and 2- level photographs were taken. Measurements from these films yielded unit cell parameters which were used to index the *x*-ray powder data. More precise cell parameters were then calculated from the indexed powder data. The only systematic extinctions observed from the singlecrystal films were 0k0 where *k* was odd.

The unit cell of molybdomenite is defined as follows:

Space group
$$P2_1$$
 or $P2_1/m$
 $a = 6.86 \text{ Å}$
 $b = 5.48 \text{ Å}$
 $c = 4.50 \text{ Å}$
 $\beta = 112^{\circ}45^{1}$
 $a;b;c = 1.252:1:0.821$

Using a cell volume of 156 Å³ and a measured specific gravity of 7.07, Z = 2 (1.99 by actual calculation).

The indexed x-ray powder data for molybdomenite are given in Table 1 All d-spacings permitted by the space group were calculated down to d = 1.554 Å. The pattern was produced in a camera of 114.6 mm diameter using Ni-filtered Cu radiation. Also included in Table 1 are

	bdomenite Montreal R Ontario, ROM, M This s Cam diam	Canada 126131 study era	nine,	Cacheuta Type spo National This stu Camer diamet	, Meno ecimen d'Hist No. 9 dy* a er	ite Cerro d doza, Arge from Mus toire Natur 9,561 Goni Guillen (1953 Came diamet	ntina éum celle, & iin) ra	Synthe molybdon This st Came diamet	nenite udy ra
	114.6 Radia Cu/	tion		57.3 m Radiati Cu/N	ion	? Radiat Cu/N		114.6 n Radiat Cu/N	ion
kkl	d(calc)	d(obs)	Ι	$d(\mathrm{obs})$	I	d(obs)	I	d(obs)	I
$\begin{array}{c} 100\\ 101 \end{array}$	6.33 Å 4.318	6.29 Å 4.31	1 <u>121</u> 2	${f 6.37}{ m \AA}\ {f 4.35}$	12 1	8		$\begin{array}{c} 6.22 \ \mathrm{\AA} \\ 4.32 \end{array}$	1212
$\begin{array}{c} 001 \\ 110 \end{array}$	$4.150 \\ 4.143$	4.13	6	4.15	6	4.16 Å 3.59	$\frac{4}{2}$	4.12	3
111	3.392	3.40	7	3.43	7	3.40	$\tilde{7}$	3.39	7
01 <u>1</u>	3.308	3.31	9	3.32	9	3.32	7	3.30	8
$20\overline{1}$	$\left. \begin{array}{c} 3.176 \\ 2.165 \end{array} \right\}$	3.16	8	3.17	7	3.17	4	3.14	9
200 101	$3.165 \atop 2.982$	2.98	3	2.99	5	3.00	2	2.97	1
211	2.746	2.00	0	4.00	0	0.00	-	2.01	-
020	2.740)	2.741	10	2.756	10	2.75	10	2.735	10
$\begin{array}{c} 210\\111 \end{array}$	$2.739 \atop 2.619$	2.618	$\frac{10}{2}$	2.636	2	2.62	2	2.616	10
$ \begin{array}{r} 111 \\ 120 \\ 121 \\ 021 \\ 301 \\ \end{array} $	2.019 2.514 2.313 2.286 2.268	2.018	Z	2.030	2	2.02	2	2.010	T
$102 \\ 202 \\ 201 \\ 300$	2.203 2.244 2.159 2.148 2.110	2.248	4	2.264	5	2.26	3	2.246	4
$31\overline{1} \\ 11\overline{2} \\ 002$	2.095 2.077 2.075	2.093	2	2.090	6 b			2.097	1 2
$22\overline{1} \\ 220 \\ 121$	2.075 2.072 2.018	2.071	5			2.08	2	2.071	6
$21\overline{2}$	2.009	2.013	3	2.023	5	2.02	2	2.017	2
211	2.000	1.998	2					1.996	2
310	1.969	1.962	1	1.965	2			1.960	$2 \\ \frac{1}{\frac{1}{2}} \\ 1$
$\begin{array}{c} 012\\ 302 \end{array}$	$1.940 \\ 1.888$	$\substack{1.943\\1.888}$	$\frac{1}{2}$	1.892	3	1.86	2	$\substack{1.945\\1.887}$	2 1

TABLE 1. X-RAY POWDER DATA FOR MOLYBOOMENITE.

hkl	d(calc)	d(obs)	I	d(obs)	Ι	d(obs)	Ι	d(obs)	Ι
$31\overline{2}$ 102 130	$\begin{array}{c}1.785\\1.779\\1.755\end{array}$	1.783	2	1.791	4	1.76	2	1.783	2
$130 \\ 321 \\ 122 \\ 122 \\ 123 $	$\begin{array}{c}1.747\\1.736\end{array}$	1.738	$2 \\ 1$					1.742	${f 2} {1}$
$40\overline{1}$ $22\overline{2}$	1.716 1.696	1.714	1					1.710	
$112 \\ 221 \\ 131$	$1.692 \\ 1.690 \\ 1.682$	1.691	2	1.695	2vb)		1.691	3
031 320 022	$\begin{array}{c} 1.672 \\ 1.672 \\ 1.672 \\ 1.654 \end{array}$	1.672	1					1.676	1 2
$301 \\ 411 \\ 402$	1.642 1.637 1.588	1.640	2	1.647	1			1.638	2
$23\overline{1} \\ 400 \\ 230$	$\begin{array}{c}1.584\\1.582\\1.582\\1.582\end{array}$	1.582	2	1.588	2			1.585	2
$\begin{array}{c} 311 \\ 131 \end{array}$	$1.573^{'} \\ 1.558^{'}$	1.554	$2 \\ 3 \\ \frac{1}{2} \\ 2 \\ 1$	1.563	$\begin{array}{c} 2\\ 2\\ \frac{12}{3}\\ 2\end{array}$			1.558	$ \begin{array}{c} 2 \\ 4 \\ \frac{1}{2} \\ 3 \\ 1 \\ 2 \end{array} $
		1.519	3	1.526	2			1.517	4
		1.491	2	1.497	2			1.494	2
		1.450	2	1.457	3			1.452	3
		1.426	1	1.435	2			1.430	1
		1.407	1					1.408	2
				1.401	1				
		1.394	1					1.395	1
		1.370	1	1.375	12			1.376	1
		1.343	$<\frac{1}{2}$	1.351	1/2			1.344	늘
		1.314	ī	1.316	ĩ			1.316	1
		1.294	ĩ	1.299	2			1.294	2
		1.276	1	1.284	1			1.278	- 3
		1.262	붊	1.264	1			1.262	2
		1.236	1	1.241	1313121111131			1.240	$ \frac{1}{1} \frac{1}{2} \frac{1}{2} \frac{1}{2} \frac{1}{2} \frac{1}{2} \frac{1}{2} \frac{1}{2} \frac{1}{5} 1$
		1.224	î	1.230	ĩ			1.226	5

X-RAY POWDER DATA FOR MOLYBDOMENITE. contd.

*Note: The following additional lines were observed and are probably due to chalcomenite: $5.37(\frac{1}{2})$, $4.93(\frac{1}{2})$, $3.85(\frac{1}{2})$, 3.55(1), 2.550(1), $2.380(\frac{1}{2})$, $2.309(\frac{1}{2})$, and 1.753(4).

the data given by Goni & Guillemin (1953) for type molybdomenite and data obtained by the writer for both type molybdomenite and synthetic molybdomenite. The type specimen was made available by Mlle. S. Cailliere, Muséum National d'Histoire Naturelle, Paris.

Physical and optical properties

The natural crystals of molybdomenite were too small for specific gravity determinations. We did, however, measure the specific gravity of eleven fragments of synthetic molybdomenite by means of a Berman balance. Under a binocular microscope, three of the fragments revealed numerous pores caused by the fusion process used for synthesizing the compound. These porous fragments gave consistently lower specific gravities than the others. Consequently, they were eliminated from consideration. The average value for the remaining eight fragments is 7.07 ± 0.05 . The specific gravity calculated from the unit cell data is 7.12.

Molybdomenite has two cleavages parallel to the length of the crystals. A good cleavage lies in the plane of the blades, parallel to $\{001\}$. The orientation of the other cleavage could not be determined but it is parallel to some plane in the $\{h0l\}$ family.

The writer can add very little to the optical properties of molybdomenite. His highest refractive index liquid (2.10) was lower than the refractive indices of the Ranwick molybdomenite. The orientation of the optical indicatrix could not be determined because the material was too fibrous, nor was it possible to resolve the difference in optic sign previously recorded for this material: optically positive by Bertrand (1882) and optically negative by Goni & Guillemin (1953). The modern techniques available to Goni & Guillemin make the writer more inclined to accept their data, which follow:

> Biaxial negative parallel extinction (presumably for sections parallel to the *b*-axis) sign of elongation positive 2V near 80° $n_x = 2.12, n_g = 2.14$

Assuming a composition of $PbSeO_3$ and a calculated specific gravity of 7.12, the Gladstone-Dale rule was applied to the mineral, using the revised values for the specific refractive energies given by Mandarino (1964). An "average" refractive index of 2.11 was obtained. The observed average was 2.13.

Chemical characteristics

Again, material was insufficient for a quantitative analysis. Two methods were explored to obtain the chemical composition of molyb-domenite: (1) x-ray spectrographic determination of the Pb/Se ratio and (2) synthesis.

X-ray spectrographic determination. Mr. Donald C. Harris, Assistant Curator of Mineralogy at the Royal Ontario Museum, determined the ratio of lead to selenium in the Ranwick molybdomenite by using the following samples as standards:

1. synthetic PbSeO₃ (Levy)

2. synthetic PbSeO₃ (Mrose)

3. synthetic PbSeO₃ (Mandarino)

Samples 1 and 2 were kindly supplied by Miss Mary Mrose of the United States Geological Survey; sample 3 was synthesized at the Royal Ontario Museum.

The analyses were made with a Norelco x-ray spectrograph equipped with a vacuum sample chamber (in which the sample is inverted), a pulse height analyser, and a scintillation counter.

One to two milligrams of each sample were concentrated inside a sample holder on the centre of the Mylar. Mr. Harris determined the peak intensities of the PbL α and SeK α lines, and then calculated the ratios of the intensities of PbL α to SeK α (see Table 2).

Sample No.	Material	$PbL\alpha/SeK\alpha$ ratio corrected for background
1	PbSeO ₃ synth. (Levy)	0.915
2	PbSeO ₃ synth. (Mrose)	0.928
3	PbSeO ₃ synth. (Mandarino)	0.927
4	Molybdomenite, Ranwick mi	ne 0.932

TABLE 2. X-RAY SPECTROGRAPHIC DETERMINATION OF $PbL\alpha/SeK\alpha$ ratios

Since the ratio of the intensity of the $PbL\alpha$ line to that of the Se $K\alpha$ line is practically constant, it follows that the ratio of the Pb to Se is the same (1:1) in all the samples. This fact, however, does not conclusively define molybdomenite. We cannot assign a formula of PbSeO₃ to molybdomenite from these data alone because we may be dealing with a "basic" selenite, such as Pb₂Se₂O₇, in which Pb/Se is also 1:1.

Synthesis. Mellor (1930) lists various procedures for synthesizing PbSeO₃ and Pb₂Se₂O₇. These include:

- 1. Precipitation by adding selenious acid or an alkali selenite to a solution of lead chloride or lead nitrate.
- 2. Fusion of lead selenate.
- 3. Treatment of lead dioxide with selenious acid (a small amount of lead selenate is also produced).

Mrose (personal communication) produced PbSeO₈ by the precipitation method. X-ray powder data for synthetic PbSeO₈ are given by Goni & Guillemin (1953). It is not stated how their material was synthesized, but it was probably by precipitation. A.S.T.M. card 4-0710 lists x-ray powder data for synthetic PbSeO₈ prepared by the Wyandotte Chemicals Corporation, Wyandotte, Michigan. According to Keuntzel (personal communication) no record remains of the analysis nor has any of the material been preserved. The writer synthesized PbSeO₈ by two methods. First, solutions of lead nitrate and potassium selenite were mixed to form a precipitate of PbSeO₈. In the second method, a mixture of PbO and SeO₂ in a molecular ratio of 1:1 was fused in a sealed evacuated Vycor tube and allowed to cool slowly. Because of the great affinity of SeO₂ for water, the SeO₂ was heated in an oven for several hours at a temperature above 100° C and then placed in a desiccator for several days. The

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Synthesized by		Le	Levy		Mrose Precipitation	e ition	This study Precipitation	udy trion	Corporation ASTM 4-0710	ation 4-0710	Fusion of PbO + SeO ₂	iay of SeO ₂	Molybdomenite Ranwick Mine, Ontario	nenute Mine, io
	$d(obs)^1$ 5.95	780	$d(obs)^2$ 5.91	2	$d(obs)^2$ 5.99	7 °°	$d(obs)^2$ 5.91	1 8.	d(obs)	I	$d(obs)^3$ 6.22	L Hist	$d(obs)^3$ 6.29	
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	3.89 9.76	010	3.85	7	3.88 77 77	4-	3.86	က	3.39	09	3.39	~~~	3.40	2
	o.70 3.48	0 4	3.50	ŝ	o.70 3.50	4	3.47	x) 4	3.31 3.14	04	3.30 3.14	xσ	3.31 3.16	ກແ
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	5.62 5.62	0 10 0	2.97	8b	2.95	6	2.94	× x	2.74	100	2.735	10	2.741	10
	2.83	ŝ	9, 779,	0vh	2.814	x	2.814	9			2.616		2.618	07
	2.77	ي مر			2.764	×			2.24	50	2.246	4	2.248	4
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	2.51	3		2	2.501	က	2.495	4	2.02	40	2.017	2	2.013	~~
	2.39	210	2.392	, ,	2.386	0	2.380	r-iea			1.996	3	1.998	2
	2.28	N	2.292	-	2.281		2.275				1.960		1.962	, -
	9, 14	с т	9, 137	ę	2.137	(191)-	2.191 9 139	4910	1 80	30	1.945 1.887	- \$9	1.943	c
	2.05	က	2.053	4	2.058	. 1Q	2.049	00	1.78	808	1.783	- 01	1.783	101
	2.01	C1	1		2.010	109	2.006	-	1.73	30	1.742	2	1.738	01
	1.97	m	1.997	io,	1.977	ιQ (1.973	9			1.710	i	1.714	Ч
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	1.76	ი	1.762	1)		2		8		3	100.1	1
					1.750	9	1.750	2	1.55	30	1.558	2	1.554	\$

material was weighed on a Mettler Gram-Atic balance, quickly, to prevent water from being taken up by the SeO₂. During the fusion process, no evidence of water was observed, indicating that a truly anhydrous compound was synthesized.

Partial x-ray powder data for all these synthetic products are compared in Table 3. The first three substances are clearly identical, but do not give the molybdomenite pattern. The remaining three samples (the PbSeO₃ produced by the Wyandotte Chemicals Corporation, the sample prepared by the writer and the Ranwick mineral) give essentially the same x-ray powder patterns—molybdomenite. The Ranwick molybdomenite has already been proved identical to the molybdomenite from Argentina (see Table 1).

From the above evidence, there seems to be little doubt that the chemical formula of molybdomenite is PbSeO₃. It is also clear that there are at least two polymorphs of PbSeO₃. Mrose (personal communication) has found a similar situation in samples of PbSeO₄. The writer's precipitated PbSeO₃, upon fusion, inverts to molybdomenite. A powder pattern for such a product showed all the lines of molybdomenite plus some lines from the precipitated PbSeO₃ (due to incomplete fusion).

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