# Article



# Bernardevansite, $Al_2(Se^{4+}O_3)_3 \cdot 6H_2O$ , dimorphous with alfredopetrovite and the Al-analogue of mandarinoite, from the El Dragón mine, Potosí, Bolivia

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### Abstract

A new mineral species, bernardevansite (IMA2022-057), ideally  $Al_2(Se^{4+}O_3)_3 \cdot 6H_2O$ , has been discovered from the El Dragón mine, Potosí Department, Bolivia. It occurs as aggregates or spheres of radiating bladed crystals on a matrix consisting of Co-bearing kruťaite-penroseite. Associated minerals are Co-bearing kruťaite-penroseite, chalcomenite and 'clinochalcomenite'. Bernardevansite is colourless in transmitted light, transparent with white streak and vitreous lustre. It is brittle and has a Mohs hardness of  $2\frac{1}{2}-3$ . Cleavage is not observed. The measured and calculated densities are 2.93(5) and 2.997 g/cm<sup>3</sup>, respectively. Optically, bernardevansite is biaxial (+), with  $\alpha = 1.642(5)$ ,  $\beta = 1.686(5)$  and  $\gamma = 1.74(1)$  (white light). An electron microprobe analysis yielded an empirical formula (based on 15 O apfu) ( $Al_{1.26}Fe_{0.82}^{3+})_{\Sigma 2.08}(Se_{0.98}O_3)_3 \cdot 6H_2O$ , which can be simplified to ( $Al_*Fe^{3+})_2(SeO_3)_3 \cdot 6H_2O$ .

Bernardevansite is the Al-analogue of mandarinoite,  $Fe_2^{3+}(SeO_3)_3 \cdot 6H_2O$  or dimorphous with  $P\bar{6}2c$  alfredopetrovite. It is monoclinic, with space group  $P2_1/c$  and unit-cell parameters a = 16.5016(5), b = 7.7703(2), c = 9.8524(3) Å,  $\beta = 98.258(3)^\circ$ , V = 1250.21(6) Å<sup>3</sup> and Z = 4. The crystal structure of bernardevansite consists of a corner-sharing framework of  $M^{3+}O_6$  (M = Al and Fe) octahedra and Se<sup>4+</sup>O<sub>3</sub> trigonal pyramids, leaving large voids occupied by the H<sub>2</sub>O groups. There are two unique  $M^{3+}$  positions: M1 is octahedrally coordinated by  $(4O + 2H_2O)$  and M2 by  $(5O + H_2O)$ . The structure refinement indicates that Al preferentially occupies M1 (= 0.692Al + 0.308Fe) over M2 (= 0.516Al + 0.484Fe). The substitution of the majority of Fe in mandarinoite by Al results in a significant reduction in its unit-cell volume from 1313.4 Å<sup>3</sup> to 1250.21(6) Å<sup>3</sup> for bernardevansite. The discovery of bernardevansite begs the question whether the Fe<sup>3+</sup> end-member, Fe<sub>2</sub><sup>3+</sup>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O, has two polymorphs as well, one with  $P2_1/c$  symmetry, as for mandarinoite and the other  $P\bar{6}2c$ , as for alfredopetrovite.

Keywords: bernardevansite, alfredopetrovite, mandarinoite, new mineral, crystal structure, Raman, El Dragón mine, Bolivia

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## Introduction

Bernardevansite, ideally  $Al_2(Se^{4+}O_3)_3 \cdot 6H_2O$ , is a new mineral species from the El Dragón mine, Antonio Quijarro Province, Potosí Department, Bolivia. It is named in honour of Dr Bernard W. Evans (b. 1934, Fig. 1), an Emeritus Professor in Mineralogy and Petrology at the University of Washington in Seattle, Washington, USA. Bernard received his B.Sc. from the University of London King's College, England, 1955 and Ph.D. from the University of Oxford, England in 1959. He was an Assistant and Associate Professor at the University of California in Berkeley from 1965–1969 and a Professor at the University of Washington in Seattle from 1969–2001. Bernard's major research interests included petrology, mineralogy, geochemistry and

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Fig. 1. A portrait of Dr Bernard W. Evans in 2008.

diffraction data, illustrating its structural relationships with mandarinoite and alfredopetrovite.

### Sample description and experimental methods

#### Occurrence

Bernardevansite was found on a specimen (Fig. 2) collected from the El Dragón mine (19°49'15"S, 65°55'00"W), Antonio Quijarro Province, Potosí Department, Bolivia. Associated minerals are



Fig. 3. A microscopic view of aggregates or spheres of pale grey to colourless, radiating bladed bernardevansite crystals (R210010).

Co-bearing krut'aite-penroseite (matrix), chalcomenite and 'clinochalcomenite' (not IMA-approved). Detailed descriptions on the geology and mineralogy of the El Dragón mine have been given by Grundmann et al. (1990, 2007) and Grundmann and Förster (2017). This mine exploited a telethermal deposit consisting of a single selenide vein hosted in sandstones and shales. The major ore mineral is kruťaite, CuSe<sub>2</sub>, varying in composition to penroseite, NiSe<sub>2</sub>. Later solutions rich in Bi, Pb and Hg resulted in the crystallisation of minerals such as clausthalite, petrovicite, watkinsonite, and the recently described minerals eldragónite, Cu<sub>6</sub>BiSe<sub>4</sub>(Se<sub>2</sub>) (Paar et al., 2012), grundmannite, CuBiSe<sub>2</sub> (Förster et al., 2016), hansblockite, (Cu,Hg)(Bi,Pb)Se<sub>2</sub> (Förster et al., 2017), cerromojonite, CuPbBiSe<sub>3</sub> (Förster et al., 2018) and nickeltyrrellite, CuNi<sub>2</sub>Se<sub>4</sub> (Förster et al., 2019). Oxidation produced a wide range of secondary Se-bearing minerals, such as favreauite, PbBiCu<sub>6</sub>O<sub>4</sub>(SeO<sub>3</sub>)<sub>4</sub>(OH)·H<sub>2</sub>O (Mills et al., 2014), alfredopetrovite, Al<sub>2</sub>(Se<sup>4+</sup>O<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O (Kampf et al., 2016a), petermegawite  $Al_6(Se^{4+}O_3)_3[SiO_3(OH)](OH)_9\cdot 10H_2O$  (Yang et al.,



Fig. 2. The specimen on which the new mineral bernardevansite, indicated by the blue arrow, was found (R210010).



Fig. 4. A back-scattered electron image of aggregates of radiating bladed bernardevansite crystals (R210010).



Fig. 5. A back-scattered electron image of aggregates of bladed bernardevansite crystals (R210010).

Table 1. Chemical compositions (in wt.%) of bernardevansite.\*

Constituent	Mean	Range	S.D.	Probe standard
$Al_2O_3$ $Fe_2O_3$	11.38 11.60	10.05-12.94 9.03-12.86	1.03 1.60	$Al_2O_3$ (synthetic) $Fe_2O_3$ (synthetic)
H <sub>2</sub> O Total*	19.14 99.83	99.70-100.08	0.18	Added in ideal value

\*Bernardevansite is prone to the electron beam damage, however this did not seem to affect the relative proportions of cations. The large variations in the  $Al_2O_3$  and  $Fe_2O_3$  contents result from the strong correlation between the two components. S.D. – standard deviation 2022a), franksousaite  $PbCu(Se^{6+}O_4)(OH)_2$  (Yang *et al.*, 2022b) and the new mineral bernardevansite, described herein.

#### Physical and chemical properties and Raman spectra

Bernardevansite occurs as aggregates or spheres of radiating bladed crystals (Figs 3,4,5) on a matrix consisting of Co-bearing krut'aite-penroseite. Individual crystals of bernardevansite are found up to  $0.10 \times 0.03 \times 0.01$  mm, with elongation along [001] and common crystal forms  $\{100\}, \{110\}, \overline{1}10\}$  and  $\{001\}$ . Bernardevansite is colourless in transmitted light and transparent with white streak, and has a vitreous lustre. It is brittle and has a Mohs hardness of 2<sup>1</sup>/<sub>2</sub>-3. Cleavage was not observed. The density measured by flotation in heavy liquids is 2.93(5) g/cm<sup>3</sup> and the calculated density is  $2.997 \text{ g/cm}^3$  on the basis of the empirical chemical formula and unit-cell volume from single-crystal X-ray diffraction data. Optically, bernardevansite is biaxial (+), with  $\alpha = 1.642(5), \beta = 1.686(5), \gamma = 1.74(1)$  (determined in white light), 2V (meas.) =  $84(2)^{\circ}$  and 2V (calc.) =  $87^{\circ}$ . The pleochroism is very weak, from pale grey to grey, and dispersion was not observed. The calculated Gladstone-Dale compatibility index based on the empirical formula is 0.013 (superior) (Mandarino, 1981). Bernardevansite is insoluble in water or hydrochloric acid.

The chemical composition was determined using a Shimadzu-1720 electron microprobe (WDS mode, 15 kV, 10 nA and a beam diameter of 2  $\mu$ m). The standards used for the probe analysis are given in Table 1, along with the determined compositions (11 analysis points). The resultant chemical formula, calculated on the basis of 15 O apfu (from the structure determination), is  $(Al_{1.26}Fe_{0.82}^{3+})_{\Sigma 2.08}(Se_{0.98}O_3)_3\cdot 6H_2O$ , which can be simplified to  $(Al_{5}Fe_{0.33}^{3+})_{2}(SeO_3)_{3}\cdot 6H_2O$ .

The Raman spectrum of bernardevansite (Fig. 6) was collected on a randomly oriented crystal with a Thermo Almega



Fig. 6. Raman spectra of bernardevansite, mandarinoite and alfredopetrovite.

Table 2. Powder X-ray diffraction data (d in Å, / in %) for bernardevansite.\*

I <sub>cal</sub>	I <sub>meas</sub>	$d_{\rm meas}$	$d_{calc}$	hkl	I <sub>cal</sub>	I <sub>meas</sub>	$d_{\rm meas}$	$d_{\rm calc}$	hkl
3.9	8	16.261	16.360	100	3.8	9	2.119	2.119	711
37.1	39	8.147	8.180	200	1.9	2	2.063	2.062	332
100	100	7.036	7.023	110	6.5	19	2.023	2.029	033
11	9	6.091	6.086	011	5	15	1.947	1.944	040
1.6	2	5.590	5.543	$1 \ 1 \ 1$	2.7	6	1.905	1.908	215
4.2	3	5.110	5.109	<b>2</b> 11	1.8	3	1.820	1.821	415
9.6	8	4.876	4.890	002	3	4	1.790	1.787	<u>9</u> 11
6.2	3	4.679	4.685	211	2.2	4	1.755	1.756	433
4.8	6	4.469	4.465	310	3	5	1.739	1.736	730
3	1	4.275	4.257	<u>3</u> 11	2.3	6	1.706	1.710	441
11.1	14	4.100	4.090	400	2.6	3	1.691	1.686	<b>4</b> 42
1.7	3	3.899	3.901	112	2.1	3	1.655	1.654	724
3.9	1	3.629	3.613	021	2.8	3	1.631	1.628	902
43	64	3.515	3.512	220	1.1	4	1.563	1.567	235
26.6	47	3.385	3.388	40 <b>2</b>	1.2	2	1.549	1.547	542
6.5	15	3.252	3.261	411	3.8	11	1.503	1.503	026
4.9	4	3.127	3.118	312	5.3	13	1.463	1.463	724
13.5	29	3.013	3.006	013	1.7	4	1.377	1.374	452
31.7	80	2.943	2.940	2 2 <b>2</b>	1.1	3	1.331	1.331	536
17.2	37	2.769	2.772	222	1.6	2	1.315	1.312	054
9.7	15	2.719	2.707	213	2.2	2	1.293	1.291	11 2 2
5.2	17	2.562	2.561	322	1.1	4	1.252	1.248	942
3	2	2.505	2.503	520	1.2	2	1.235	1.235	208
6.5	10	2.423	2.418	ō12					
2.6	5	2.356	2.355	114					
6.7	20	2.247	2.251	$\bar{7}$ 1 1					
6.1	10	2.172	2.165	332					

\*The strongest lines are given in bold

microRaman system, using a solid-state laser with a wavelength of 532 nm at 75 mW power and a thermoelectric cooled CCD detector. The laser is partially polarised with  $4 \text{ cm}^{-1}$  resolution and a spot size of 1  $\mu$ m.

#### X-ray crystallography

Both the powder and single-crystal X-ray diffraction data for bernardevansite were collected on a Rigaku Xtalab Synerg D/S 4-circle diffractometer equipped with  $CuK\alpha$  radiation. Powder X-ray diffraction data were collected in the Gandolfi powder mode at 50 kV and 1 mA (Table 2) and the unit-cell parameters were refined using the program by Holland and Redfern (1997): a = 16.535(1), b = 7.7762(5), c = 9.8841(6) Å,  $\beta = 98.337(7)^{\circ}$  and V = 1257.5(5) Å<sup>3</sup>.

All bernardevansite crystals examined are pervasively twinned on (100) with a twin law ( $\overline{1}$  0 ½, 0  $\overline{1}$  0, 0 0 1). Single-crystal X-ray diffraction data were collected from a 0.03 × 0.02 × 0.01 mm fragment. The systematic absences of reflections suggest the unique space group *P*2<sub>1</sub>/*c*. The structure was solved and refined using *SHELX2018* (Sheldrick, 2015a, 2015b). No H atoms were located through the difference-Fourier syntheses. The refined Al/Fe ratios at the octahedral *M*1 and *M*2 sites are (0.692Al + 0.308Fe) and (0.516Al + 0.484Fe), respectively, yielding a total Al/Fe ratio of 1.208/0.792, which is very close to that (1.211/0.789, normalised)

Table 3. Summary of crystallographic data and refinement results for bernardevansite, mandarinoite and alfredopetrovite.

	Alfredopetrovite	Bernardevansite	Mandarinoite
Ideal formula	$Al_2(Se^{4+}O_3)_3 \cdot 6H_2O$	Al <sub>2</sub> (Se <sup>4+</sup> O <sub>3</sub> ) <sub>3</sub> ·6H <sub>2</sub> O	Fe <sub>2</sub> <sup>3+</sup> (Se <sup>4+</sup> O <sub>3</sub> ) <sub>3</sub> ·6H <sub>2</sub> O
Crystal symmetry	Hexagonal	Monoclinic	Monoclinic
Space group	P62c	P21/c	P21/c
a (Å)	8.818(3)	16.5016(5)	16.810(4)
b (Å)	8.818(3)	7.7703(2)	7.880(2)
c (Å)	10.721(2)	9.8524(3)	10.019(2)
α (°)	90	90	90
β (°)	90	98.258(3)	98.26(2)
γ (°)	120	90	90
V (Å <sup>3</sup> )	722.0(5)	1250.21(6)	1310.4
Z	2	4	4
$\rho_{cal}$ (g/cm <sup>3</sup> )	2.50		2.98
2θ range for data collection (°)	≤40 (Mo <i>K</i> α)	<130.16 (Cu <i>K</i> α)	<60 (Mo <i>K</i> α)
No. of reflections collected	1817	9235	4658
No. of independent reflections	246	3407	
No. of reflections with $l > 2\sigma(l)$	240	3156	2101
No. of parameters refined	40	184	
R(int)	0.064	0.037	
Final $R_1$ , $wR_2$ factors $[I > 2\sigma(I)]$	0.027, 0.063	0.041, 0.110	0.064, 0.084
Goodness-of-fit	1.07	1.03	,
Reference	Kampf <i>et al.</i> (2016a)	This study	Hawthorne (1984)

Table 4. Fractional atomic coordinates and equivalent isotropic displacement parameters  $(\mathring{A}^2)$  for bernardevansite.

Atom	x/a	y/b	z/c	U <sub>eq</sub>	Occ. (<1)
M1Al	0.39931(11)	0.74117(17)	0.9075(2)	0.0155(7)	0.692(14)
M1Fe	0.39931(11)	0.74117(17)	0.9075(2)	0.0155(7)	0.308(14)
M2Al	0.90506(9)	0.24015(16)	0.04690(19)	0.0141(6)	0.516(14)
M2Fe	0.90506(9)	0.24015(16)	0.04690(19)	0.0141(6)	0.484(14)
Se1	0.98068(4)	0.03893(9)	0.31646(7)	0.0160(2)	
Se2	0.22324(5)	0.72432(9)	0.72880(8)	0.0163(2)	
Se3	0.52831(4)	0.48070(9)	0.81382(7)	0.0164(2)	
01	0.9810(3)	0.2129(7)	0.4212(6)	0.0205(10)	
02	0.0775(3)	-0.0346(6)	0.3679(5)	0.0198(10)	
03	0.9975(3)	0.1369(7)	0.1690(4)	0.0207(10)	
04	0.3162(3)	-0.1799(6)	0.7648(5)	0.0200(10)	
05	0.1905(3)	-0.1703(6)	0.5782(5)	0.0198(10)	
06	0.1688(3)	-0.1636(6)	0.8314(5)	0.0219(10)	
07	0.5208(3)	0.3791(7)	0.6612(5)	0.0216(10)	
08	0.4297(3)	0.5494(6)	0.7990(5)	0.0182(10)	
09	0.4774(3)	0.6800(6)	0.0686(5)	0.0196(10)	
010	0.3191(3)	-0.4257(6)	0.9641(5)	0.0205(10)	
011	0.3609(3)	-0.0850(6)	0.0260(5)	0.0240(10)	
012	0.8865(3)	0.0117(6)	-0.0434(5)	0.0221(10)	
013	0.6404(4)	0.7624(6)	0.0754(8)	0.0311(14)	
014	0.2527(4)	0.3132(9)	0.1937(7)	0.0468(16)	
015	0.1770(4)	0.2449(6)	0.4252(8)	0.0351(17)	

**Table 5.** Atomic displacement parameters  $(\mathring{A}^2)$  for bernardevansite.

Atom	U <sup>11</sup>	U <sup>22</sup>	U <sup>33</sup>	U <sup>12</sup>	U <sup>13</sup>	U <sup>23</sup>
M1	0.0121 (10)	0.0190(10)	0.0155(11)	0.0002(5)	0.0025(7)	0.0015(5)
M2	0.0103(9)	0.0189(9)	0.0134(9)	-0.0004(5)	0.0031(6)	0.0008(4)
Se1	0.0132(4)	0.0198(4)	0.0149(4)	-0.0004(2)	0.0016(3)	0.0013(3)
Se2	0.0132(4)	0.0199(3)	0.0158(4)	-0.0005(3)	0.0023(3)	0.0006(3)
Se3	0.0135(4)	0.0198(4)	0.0160(4)	0.0002(3)	0.0022(3)	0.0008(3)
01	0.015(2)	0.024(2)	0.024(3)	-0.002(2)	0.006(2)	-0.005(2)
02	0.013(2)	0.021(2)	0.024(2)	0.0007(18)	-0.0010(19)	0.0049(19)
03	0.015(2)	0.032(3)	0.015(2)	-0.002(2)	0.0026(17)	0.004(2)
04	0.016(2)	0.026(3)	0.018(2)	-0.003(2)	0.0004(18)	0.000(2)
05	0.015(2)	0.030(3)	0.015(2)	-0.002(2)	0.0039(18)	0.0007(19)
O6	0.017(2)	0.028(3)	0.021(2)	0.002(2)	0.0059(19)	0.001(2)
07	0.018(2)	0.029(3)	0.017(2)	0.003(2)	0.0021(18)	-0.002(2)
O8	0.016(2)	0.022(2)	0.016(2)	0.0052(19)	0.0016(18)	0.0011(19)
O9	0.019(2)	0.022(2)	0.018(2)	0.002(2)	0.0017(18)	0.006(2)
O10	0.016(2)	0.022(2)	0.024(2)	0.003(2)	0.0044(19)	0.0028(19)
011	0.027(3)	0.021(2)	0.023(2)	-0.002(2)	0.000(2)	0.003(2)
012	0.018(2)	0.027(3)	0.022(2)	0.003(2)	0.006(2)	-0.002(2)
013	0.027(3)	0.024(3)	0.043(4)	0.000(2)	0.010(3)	0.005(2)
014	0.065(4)	0.036(4)	0.041(4)	0.007(3)	0.012(3)	0.003(3)
015	0.023(3)	0.036(4)	0.047(4)	-0.006(2)	0.008(3)	-0.012(2)

measured from the electron microprobe analysis. The structure was refined as a 2-component twin with a twin ratio of 0.81/0.19. Final refinement statistics for bernardevansite are listed in Table 3. Atomic coordinates and displacement parameters are given in Tables 4 and 5, respectively. Selected bond distances are presented in Table 6. The bond-valence sums were calculated using the parameters given by Brese and O'Keeffe (1991) (Table 7). The crystallographic information file has been deposited with the Principal Editor of *Mineralogical Magazine* and is available as Supplementary material (see below).

#### Crystal structure description and discussion

Bernardevansite, Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O, is isostructural with mandarinoite, Fe<sub>2</sub><sup>3+</sup>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O (Hawthorne, 1984), rather than with the Al end-member  $P\bar{6}2c$  alfredopetrovite, Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O (Morris *et al.*, 1992; Kampf *et al.*, 2016a). In other words, it is dimorphous with alfredopetrovite. The crystal structure of bernardevansite consists of a corner-sharing framework of  $M^{3+}O_6$  (M = Al and Fe) octahedra and Se<sup>4+</sup>O<sub>3</sub> trigonal pyramids, leaving large voids that are occupied by the H<sub>2</sub>O groups (Fig. 7). There are three unique Se positions in bernardevansite, each of which is coordinated to three O atoms to form characteristic SeO<sub>3</sub> trigonal pyramids. There are two unique  $M^{3+}$  positions: M1 is octahedrally coordinated by  $(4O + 2H_2O)$  and M2 by  $(5O + H_2O)$ . The structure refinement indicates that Al preferentially occupies M1 (= 0.692 Al + 0.308 Fe) over M2 (= 0.516 Al + 0.484 Fe). There are three distinct H<sub>2</sub>O molecules (O13, O14 and O15) in the structure that are not bonded to any non-H cation (Table 7), in addition to three H<sub>2</sub>O molecules (O10, O11 and O12) bonded to M cations. Although our structure determination failed to locate H atoms, all O-O distances for H-bonding in bernardevansite are consistent and comparable with those found in mandarinoite (Hawthorne, 1984) (Table 6).

Table 6. Selected bond distances (Å) for bernardevansite, Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O and mandarinoite, Fe<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O.\*

Ref:	Mandarinoite (1)	Bernardevansite (2)		Mandarinoite (1)	Bernardevansite (2)		Mandarinoite (1)	Bernardevansite (2)
M1-04	1.980(11)	1.918(5)	Se1-01	1.699(10)	1.700(5)	010-05	2.629	2.647
<i>M</i> 1–07	1.956(12)	1.897(5)	Se1-02	1.717(11)	1.704(5)	010-013	2.768	2.742
<i>M</i> 1–08	2.007(10)	1.941(5)	Se1-03	1.699(10)	1.698(4)			
<i>M</i> 1–09	2.032(10)	1.954(5)	<se1-0></se1-0>	1.705	1.701	011-08	2.810	2.779
M1-010	2.081(13)	1.989(5)				011-013	2.711	2.697
M1-011	2.066(12)	1.950(5)	Se2-04	1.699(11)	1.696(5)			
<m1-0></m1-0>	2.021	1.942	Se2-05	1.728(11)	1.712(5)	012-01	2.770	2.699
			Se2-06	1.694(10)	1.687(5)	012-03	2.886	2.753
M2-01	1.967(10)	1.919(6)	<se2-0></se2-0>	1.707	1.699	012-015	2.709	2.666
M2-02	2.011(11)	1.945(5)						
M2-03	2.039(13)	1.973(5)	Se3-07	1.698(9)	1.687(5)	013-09	2.801	2.757
M2-05	2.042(12)	1.983(5)	Se3-08	1.710(10)	1.699(5)	013-014	2.715	2.700
M2-06	1.994(11)	1.923(5)	Se3-09	1.717(10)	1.715(5)			
M2-012	2.074(12)	1.989(5)	<se3-0></se3-0>	1.708	1.701	014-015	2.815	2.804
<m2-0></m2-0>	2.021	1.955				014-04	3.086	3.081
						015-02	2.729	2.732
						015-014	2.815	2.795

\*Notes: *M* = Fe and (Al,Fe) for mandarinoite and bernardevansite, respectively. References: (1) Hawthorne (1984); (2) this study.

Table 7. Bond-valence sums for bernardevansite.\*

	М1	М2	Se1	Se2	Se3	Sum
01		0.544	1.348			1.892
02		0.507	1.336			1.843
03		0.470	1.358			1.828
04	0.509			1.365		1.874
05		0.458		1.305		1.763
06		0.538		1.397		1.935
07	0.539				1.397	1.936
08	0.479				1.352	1.831
09	0.463				1.295	1.758
010	0.421					0.421
011	0.468					0.468
012		0.449				0.449
013						-
014						-
015						-
Sum	2.878	2.967	4.042	4.068	4.044	

\*Note: The bond valence sums for M1 and M2 were calculated based on (0.692 Al + 0.308 Fe<sup>3+</sup>) and (0.516 Al + 0.484 Fe<sup>3+</sup>), respectively.

The substitution of the majority of Fe in mandarinoite by Al in bernardevansite results in a significant reduction in unit-cell volume from 1313.4 Å<sup>3</sup> to 1250.21(6) Å<sup>3</sup>, which motivated this investigation. Compared to mandarinoite, which has the identical average  $\langle M$ -O> bond distances (2.021 Å) for the two octahedral sites (Hawthorne, 1984), the  $\langle M$ -O> distance for the M1 site (1.942 Å) in bernardevansite is shorter than that for the M2 site (1.955 Å), consistent with the preference of Al at M1 over M2 (0.692 vs. 0.516), as the ionic radius of <sup>VI</sup>Al<sup>3+</sup> (0.535 Å) is smaller than that of <sup>VI</sup>Fe<sup>3+</sup> (0.645 Å) (Shannon, 1976). A survey of the literature appears to suggest that, for a structure with two or more octahedral sites,  $Al^{3+}$  is likely to be favoured by the site coordinated with more H<sub>2</sub>O molecules. This is indeed the case for bernardevansite, as the *M*1 site is coordinated by (4O + 2H<sub>2</sub>O) and *M*2 by (5O + H<sub>2</sub>O). Another typical example is coquimbite, which contains three distinct octahedral sites (*M*1, *M*2 and *M*3), with *M*1 coordinated by (6H<sub>2</sub>O), *M*2 by (6O<sup>2-</sup>) and *M*3 by (3H<sub>2</sub>O + 3O<sup>2-</sup>). All structure determinations on coquimbite have shown that Al<sup>3+</sup> is predominately or exclusively ordered into the *M*1 site (e.g. Demartin *et al.*, 2010; Yang and Giester, 2018; Mauro *et al.*, 2020 and references therein).

According to the Raman spectroscopic studies on hydrous materials containing  $(\text{SeO}_3)^{2-}$  (e.g. Wickleder *et al.*, 2004; Frost *et al.*, 2006; Frost and Keeffe, 2008; Djemel *et al.*, 2013; Wolak *et al.*, 2013; Kasatkin *et al.*, 2014; Mills *et al.*, 2014; Kampf *et al.*, 2016b), we made the following tentative assignments of major Raman bands for bernardevansite. The broad bands between 2900 and 3500 cm<sup>-1</sup> and those between 1500 and 1750 cm<sup>-1</sup> are due to the O–H stretching and H–O–H bending modes in H<sub>2</sub>O groups, respectively. The bands at 844 and 685 cm<sup>-1</sup> are ascribable to the Se<sup>4+</sup>–O symmetric and antisymmetric stretching vibrations, respectively, within the Se<sup>4</sup> +O<sub>3</sub> groups, whereas those from 320 to 570 cm<sup>-1</sup> originate from the O–Se<sup>4+</sup>–O bending modes. The bands below 320 cm<sup>-1</sup> are mainly associated with the rotational and translational modes of Se<sup>4+</sup>O<sub>3</sub> groups, as well as the  $M^{3+}$ –O interactions and lattice vibrational modes.

For comparison, the Raman spectra of alfredopetrovite,  $Al_2(Se^{4+}O_3)_3 \cdot 6H_2O$  and mandarinoite,  $Fe_2^{3+}(Se^{4+}O_3)_3 \cdot 6H_2O$ , from the RRUFF Project (http://rruff.info/R210014 and http://rruff.info/R140742, respectively) are also plotted in Fig. 6. Evidently, the spectrum of bernardevansite is more similar to that of



Fig. 7. Crystal structure of bernardevansite. Green, yellow and grey polyhedra represent M10<sub>6</sub>, M2O<sub>6</sub> and SeO<sub>3</sub> groups, respectively. Purple and aqua spheres represent Se (Se1, Se2 and Se3) atoms and H<sub>2</sub>O (O13, O14 and O15) groups that are not bonded to any non-H cation, respectively.

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Fig. 8. A concept of classification of  $M^{3+}(SeO_3)_3 \cdot 6H_2O$  minerals (M = AI and Fe).

mandarinoite than to that of alfredopetrovite, pointing to the structural similarities between bernardevansite and mandarinoite.

Although the bernardevansite sample we studied here,  $(Al_{0.61}Fe_{0.39}^{3+})_2(SeO_3)_3 \cdot 6H_2O_5$ , is isostructural with mandarinoite, its chemistry is closer to that of alfredopetrovite, the Al endmember of the Fe<sub>2</sub><sup>3+</sup>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O-Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O system, as illustrated in Fig. 8. This raises an interesting question about its ideal chemical formula. Should it be expressed as (1) an Fe-bearing formula,  $(Al_{1-x}Fe_x)_2(SeO_3)_3 \cdot 6H_2O$ , where 0 < x < 0.5, or (2) an Fe-free end-member formula, Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O. The first Fe-bearing formula requires that Fe is essential to stabilise the  $P2_1/c$  mandarinoite-type structure and there is no complete solid-solution Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O series between and  $Fe_2(SeO_3)_3$ ·6H<sub>2</sub>O. This formula appears to be consistent with synthetic experiments, as several hydrothermal syntheses of Al selenites conducted thus far have revealed only the hexagonal form of Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O and no monoclinic form (Morris et al., 1992; Ratheesh et al., 1997 and references therein). In contrast, the second Fe-free formula implies that Al<sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O possesses two polymorphs: a monoclinic  $P2_1/c$  mandarinoite-type form and a hexagonal P62c alfredopetrovite form. Regardless of its ideal chemical formula, the discovery of bernardevansite begs the question whether the Fe<sup>3+</sup> end-member, Fe<sup>3+</sup><sub>2</sub>(SeO<sub>3</sub>)<sub>3</sub>·6H<sub>2</sub>O, has two polymorphs as well, one with  $P2_1/c$  symmetry, as for mandarinoite, and the other  $P\overline{6}2c$ , as for alfredopetrovite.

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